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# Severe convection in the Mediterranean from microwave observations and a convection-permitting model

Jean-François Rysman,<sup>a\*</sup> Chantal Claud,<sup>a</sup> Jean-Pierre Chaboureau,<sup>b</sup> Julien Delanoë,<sup>c</sup> and  
Beatriz M. Funatsu<sup>d</sup>

<sup>a</sup> *Laboratoire de Météorologie Dynamique, UPMC Univ. Paris 06,*

<sup>b</sup> *Laboratoire d'Aérodynamique, Université de Toulouse and CNRS, France,*

<sup>c</sup> *Université Versailles St-Quentin; CNRS/INSU, LATMOS-IPSL,*

<sup>d</sup> *CNRS, LETG-Rennes-COSTEL, UMR 6554, Université Rennes 2*

\*Correspondence to: [jfrysman@lmd.polytechnique.fr](mailto:jfrysman@lmd.polytechnique.fr)

This study investigates severe convection in the Mediterranean during the first Special Observation Period (SOP-1; 5 September-6 November 2012) of the Hydrological Cycle in the Mediterranean Experiment (HyMeX) with the objectives of providing novel information about severe convection on its vertical structure, spatio-temporal variability as well as evaluating the ability of a convection-permitting model to reproduce this variability. Two criteria, namely deep convection (DC) and convective overshooting (COV), are computed using the water vapour channels of the Microwave Humidity Sounder (MHS). Special attention is paid to the COV as it is associated with particularly severe weather. For the first time, the COV criterion was assessed in the Mediterranean, using two case studies conjointly observed by the airborne Rasta radar and MHS. COV is characterised by high ice water content (up to  $2 \text{ g.m}^{-3}$ ) in the mid and upper troposphere (up to 12.5 km in the stratosphere). During the SOP-1, DC and COV occurred about 0.1 % and 0.03 % of the total observation time, respectively. The Atlantic weather regimes appear to affect the temporal distribution of these convective events. Most of the DC and COV occurrences were found along the western coasts of Italy and Greece, mainly during the 10-15 October and 25 October-3 November episodes. These two episodes, for which severe meteorological events (e.g., tornadoes) were reported, are significant when compared with the 2002-2013 climatology (above the 75th percentile). Both criteria are also employed to assess the current ability of the Meso-NH model to forecast severe convection using a model-to-satellite approach. The forecasted DC and COV are found to be highly correlated in time with the observations, but are strongly underestimated. This suggests that the model missed a significant part of the most intense convective events and their associated hazards and underlines the need for better characterisation of model uncertainties associated with severe convection.

*Key Words:* Convection-permitting model, Convective Overshooting, Deep Convection, HyMeX SOP-1, Mediterranean region, Microwave Observations

## 1. Introduction

The Mediterranean region is frequently affected by intense meteorological events such as heavy precipitation and flash floods as well as occasional but potentially devastating tropical-like cyclones (Papagiannaki *et al.* 2013; Tous and Romero 2013; Trigo *et al.* 2002; Lionello *et al.* 2006; Claud *et al.* 2010). These events mostly occur during the autumn season over the western Mediterranean basin (Melani *et al.* 2013; Ducrocq *et al.* 2014; Raveh-Rubin and Wernli 2015). Because of the high population density (more than 500 people per  $km^2$  in some regions), they often cause significant societal damage, for example, there were more than 800 casualties during the Algiers floods in 2001. In particular, deep convective systems (DC) can accumulate considerable volumes of rain over a region in only a few hours, e.g., Genoa in October 1970 (Roth *et al.* 1996; Siccardi 1996), Vaison-la-Romaine in September 1992 (Sénési *et al.* 1996), Izmir in November 1995 (Kömüşçü 1998), Algiers in November 2001 (Argence *et al.* 2008), Gard in September 2002 (Delrieu *et al.* 2005; Ducrocq *et al.* 2008), eastern coasts of Greece in January 2004 (Lagouvardos *et al.* 2007), Draguignan in June 2010 (Fresnay 2014) and, Nice-Genoa in November 2011 (Silvestro *et al.* 2012; Rebora *et al.* 2013; Buzzi *et al.* 2014). A special class of DC events called convective overshooting (COV) has been known for a long time to be associated with particularly severe weather such as heavy rainfall, damaging winds, hail, and tornadoes in the United States (Reynolds 1980; Negri and Adler 1981; Adler *et al.* 1985; Brunner *et al.* 2007) and Europe (including the Mediterranean region; Bedka (2011)). Moreover, Hong *et al.* (2008) showed in a climatological analysis from  $30^\circ$  S to  $30^\circ$  N that COV tends to be more frequent over land (e.g., tropical Africa and South America) than over ocean and presents a strong seasonal variability. COV are associated with very intense updrafts, so that convection reaches the tropopause and penetrates into the lower stratosphere. They

occur within convective systems that can have large spatial and temporal extension (several hours and more than 100 km, Orlanski 1975). No clear consensus exists about the exact life time and size of COV but they are assumed to have a small diameter compared to convective systems in the order of several tens of kilometres and duration in the order of several tens of minutes. Within overshooting clouds, clusters <15 km in diameter of very low brightness temperatures were observed in the infrared imagery for several minutes (<15 min) (Bedka 2011; Bedka *et al.* 2012). These small phenomena (several tens of kilometers and several tens of minutes) are difficult to forecast and therefore there is a need for better detection and understanding. An improved forecast should lead to a better mitigation of the associated damage, having a high societal impact.

In this context, the international project, the Hydrological cycle in the Mediterranean Experiment (HyMeX) was launched in 2010 to gain a better understanding of hydrological cycle and heavy precipitation events in the whole Mediterranean region (Drobinski *et al.* 2014). The first Special Observation Period (SOP-1) of HyMeX took place in the north-western Mediterranean Sea and surrounding regions from 5 September 2012 to 6 November 2012 (Ducrocq *et al.* 2014). More than 150 research instruments such as radars, lidars, raingauges, electric field sensors and meteorological balloons were deployed during this major field campaign. Three aircraft also participated in the SOP-1 with on-board instruments such as microphysical in-situ probes, cloud radar and differential absorption lidar (for details see Ducrocq *et al.* 2014; Ferretti *et al.* 2014). In addition to research instruments specially dedicated to the HyMeX SOP-1, spaceborne instruments (e.g., Microwave Humidity Sounder (MHS); Bonsignori (2007), Meteosat Second Generation (MSG) Spinning Enhanced Visible and InfraRed Imager (SEVIRI); Schmetz *et al.* (2002)) and operational ground-based networks also provided information about the meteorological environment of the Mediterranean region during this period. In this study, we use observations from the passive microwave sounder MHS to identify deep convection (DC) and convective overshooting (COV) occurrences. This instrument covers a wide area with a swath of 1920 km with a fairly fine pixel resolution (16 km at nadir). Such a spatial resolution is fine enough to observe DC and COV events as shown by Hong *et al.* (2005). Yet the finest convective towers highlighted in Bedka (2011); Bedka *et al.* (2012) are probably missed. MHS

also provides observations over the sea where available data are scarce. During the HyMeX SOP-1, MHS flew on-board three polar orbiting satellites each of them flying over the Mediterranean basin twice a day with different crossing times.

Hong *et al.* (2005) developed a method to detect the occurrence of deep convection and convective overshooting over the Tropics using water vapour channels of the AMSU-B (Advanced Microwave Sounding Unit) module that is quasi-similar to the MHS module. This innovative method uses a threshold on the brightness temperature differences between the three channels to detect DC and a more stringent relation between brightness temperature differences to detect COV. These DC and COV criteria have been used so far for tropical regions (Hong *et al.* 2008; Funatsu *et al.* 2012; Aumann and Ruzmaikin 2013), while only DC criterion has been applied in the Mediterranean region (e.g., Funatsu *et al.* 2007, 2008, 2009; Claud *et al.* 2012; Alhammoud *et al.* 2014; Sanò *et al.* 2015). Here, we take advantage of the airborne radar measurements performed during the HyMeX SOP-1 to evaluate the COV criterion in the Mediterranean region for the first time. Moreover, this study provides the first-ever characterisation of COV in terms of frozen hydrometeor content and diameter. Then, we use both DC and COV criteria to provide an estimation of the spatial and temporal distribution of severe convection during the HyMeX SOP-1. This characterisation can be used for validating simulations by convection-permitting models as will be illustrated for the Meso-NH model (Lafore *et al.* 1998).

The paper is organised as follows. In the section 2, we describe the instruments and the method. Since the method was developed for tropical convective systems and for the AMSU-B radiometer, the validity of the Hong *et al.* (2005) criteria for the Mediterranean region and for the MHS radiometer is assessed in section 3. In section 4, we document the occurrence of DC and COV during the HyMeX SOP-1 and we compare this period to the previous 10 years of AMSU-B/MHS measurements. Section 5 is dedicated to the assessment of the Meso-NH model to accurately forecast severe convection. The discussion and conclusions are presented in section 6.

## 2. Data and Method

### 2.1. Microwave Humidity Sounder (MHS)

DC and COV detection and characterisation are achieved using space-borne MHS measurements. MHS is a cross-scanning radiometer primarily designed for water vapour atmospheric profile retrieving. It measures upwelling microwave radiation using five channels: two window channels at 89 and 157 GHz (not used in this study) and three channels in the water vapour absorption line (at  $183.3\pm 1$ ,  $183.3\pm 3$  and 190.3 GHz). These three channels probe more or less deeply into the atmosphere depending on their frequencies. The swath width of MHS is 1920 km, the nadir resolution is 16 km and the viewing angle ranges from  $0.6^\circ$  to  $60^\circ$ .

During the HyMeX SOP-1, three MHS radiometers on-board the polar orbiting satellites National Oceanic and Atmospheric Administration (NOAA) satellites NOAA-18, NOAA-19 and the European Operational Meteorological satellite MetOp-A respectively, were available. These satellites have synchronous orbits and fly over the Mediterranean region twice a day: 02:00-05:00 and 13:00-16:00 for NOAA-18, 23:00-02:00 and 10:00-13:00 for NOAA-19, and 07:00-10:00 and 18:00-21:00 UTC for MetOp-A.

Hong *et al.* (2005) developed a method to detect both deep convection and convective overshooting events using the AMSU-B radiometer that is quasi-similar to the MHS radiometer (for a comparison, see Claud *et al.* 2012). Using a microwave radiometer and an X-band radar on-board a scientific aircraft, they found that for clear sky conditions, brightness temperature measured by each water vapour channel is always higher for channels with frequencies further away from the centre frequency at 183.3 GHz, but within convective clouds this relationship is reversed. This is because channels further away from the centre at 183.3 GHz can probe deeper into the clouds. Consequently, those channels are more subject to scattering by hydrometeors from the middle or low layers of the deep convective clouds (Burns *et al.* 1997; Wang *et al.* 1997), which depresses the observed brightness temperatures. Moreover, Hong *et al.* (2005) showed that close to the convective core, the brightness temperatures difference between channels tends to be close to zero while

it remains higher for higher frequencies where the convective cloud penetrates the tropopause layer. The reason is that the temperature weighting functions of the three water vapour channels tend to be similar where there is a large frozen hydrometeor content (Burns *et al.* 1997). As a result, Hong *et al.* (2005) established two criteria to detect deep convection and convective overshooting using these frequencies.

Deep convection is detected whenever:

$$\begin{aligned}
 TB_3 - TB_5 &\geq T_0 \\
 \text{and } TB_3 - TB_4 &\geq T_0 \\
 \text{and } TB_4 - TB_5 &\geq T_0
 \end{aligned} \tag{1}$$

with  $T_0 = 0.04761 - 0.01678 \theta + 0.00599 \theta^2$  where  $\theta$  is the viewing angle, and  $TB_{3-5}$  are brightness temperatures for channels 3-5 (i.e.,  $183.3 \pm 1$ ,  $183.3 \pm 3$  and  $190.3$  GHz)

Convective overshooting is detected where:

$$TB_3 - TB_5 \geq TB_3 - TB_4 \geq TB_4 - TB_5 > T_1 \tag{2}$$

Using simulations and a microwave radiative transfer model, Hong *et al.* (2005) showed that the COV criterion performs better when  $\theta \leq 35^\circ$ . For  $\theta \leq 30^\circ$ ,  $T_1 = 0$  K and for  $35^\circ \geq \theta > 30^\circ$ ,  $T_1 = 7$  K. Because COV occur within DC systems, the DC criterion is also fulfilled every time the COV criterion is reached.

It must be noted that these criteria have been established for tropical conditions and using AMSU-B channel frequencies ( $183.3 \pm 1$ ,  $183.3 \pm 3$  and  $183.3 \pm 7$  GHz) that are slightly different from MHS channels frequencies. In the Mediterranean region, Funatsu *et al.* (2007) showed that DC events correspond to a TRMM 3B42 3-hour accumulated rainfall product of at least 20 mm in 50 % of the cases. In this paper, we evaluate both the DC and COV criteria in the Mediterranean region



(10° W-35° E and 32° N-48° N) using vertical profiles of reflectivity completed during the HyMeX SOP-1 by an airborne radar.

## 2.2. *Rasta*

Rasta (RADar SysTEM Airborne) is an airborne 95 GHz Doppler cloud radar (Protat *et al.* 2004; Bouniol *et al.* 2008; Protat *et al.* 2009; Delanoë *et al.* 2013; Bousquet and Smull 2006). It measures reflectivity using six antennas: three pointing downwards in three non-collinear directions, including one nadir pointing angle, and three pointing upwards in three non-collinear directions, including one zenith pointing angle. The vertical resolution is 60 m and the temporal resolution is 1.25 seconds. The minimal detectable cloud reflectivity is about -35 dBZ depending on the antenna, and the calibration accuracy is about 1-2 dB. This instrument makes it possible to investigate microphysical properties of clouds (e.g., Delanoë *et al.* 2013) since it detects hydrometeor sizing from about a hundred microns to a few millimetres. Mean volume-weighted diameter of frozen hydrometeors and ice water content were retrieved at each radar gate (i.e., 60 m) using the radonvar method (adapted from Delanoë *et al.* 2007, 2014). The retrieval error is estimated to be within 30 %, a first evaluation having been carried out using collocated in-situ data such as particle size and bulk measurements. During the HyMeX campaign, Rasta was mounted on the Falcon 20 aircraft which carried out 18 flights (about 50 flying hours). Because the mission of the aircraft was to fly close to the convective systems, Rasta provided unique measurements to evaluate the COV criterion.

## 3. Evaluation of COV criterion

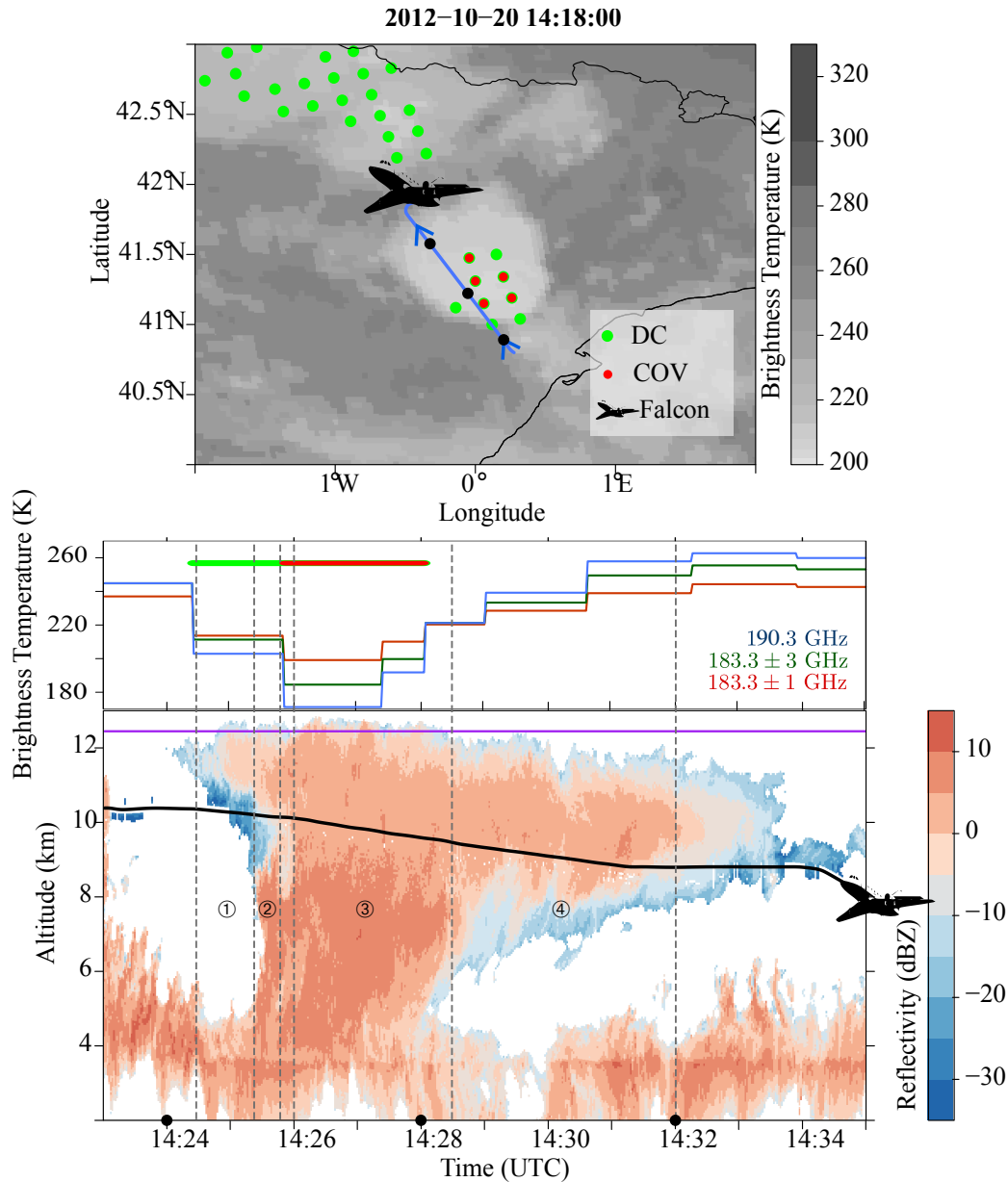
During the HyMeX SOP-1, two COV events were observed near concomitantly by MHS and Rasta instruments, with measurements within 10 minutes and 10 km apart. These nearly coincidental observations allowed us to evaluate the COV criterion for the first time in the Mediterranean. Analysis of Rasta-derived products provides unique information about the content and mean diameter of the frozen hydrometeors associated with COV. Note that a complete meteorological analysis of these two cases is beyond the scope of this paper

### 3.1. 20 October Case

The 20 October case corresponds to the Intensive observation period (IOP) 15a (Ducrocq *et al.* 2014). The synoptic situation was characterised by a north-south elongated trough located along Portugal that slowly evolved to a cut-off low. This situation triggered strong frontal convergence and warm advection over western Spain. Overall the event triggered 149 mm of rain in 24 h over Spain and caused flooding in the Ebro Basin (shown in Figure 1 (Top)).

Figure 1 shows DC and COV occurrence derived from MHS at 14:18 UTC on 20 October 2012 (Top), brightness temperatures from MHS (Middle) and reflectivity profiles from Rasta (Bottom) between 14:23 and 14:35 UTC (flight 23 of the Falcon 20 aircraft). At 14:18 UTC, the MSG 10.8  $\mu\text{m}$  channel showed brightness temperatures down to 203 K over the northeastern part of Spain, highlighting a convective region. A secondary region of convective activity with brightness temperatures down to 220 K was found along the Spanish foothills of the Pyrenees. In the following study, we focused our analysis on the deepest system, centred at 0° W-41.5° N. For this system, MHS observations highlight a deep convective region of approximately 140 km<sup>2</sup> (9 pixels), in which an area of 80 km<sup>2</sup> (5 pixels) of COV tops is detected.

At 14:25 UTC, the Falcon 20 aircraft entered the cloud system, characterised by reflectivity higher than -35 dBZ (i.e., Rasta sensitivity), and came out at 14:33 UTC. This system showed the typical anvil shape and could be separated into several segments using vertical profiles of reflectivity (Figure 1 (Bottom)), mean diameter and mass content of frozen hydrometeors (Figure 2). In the first segment, from 14:24.5 to 14:25.3 UTC (①), cloud tops exceeded 12.5 km. The reflectivity ranged from -30 dBZ to 0 dBZ at altitudes >10 km and lower than -35 dBZ below. As the aircraft moved further into the cloud (14:25.3-14:25.8 UTC (②)), the reflectivity increased up to 5 dBZ above 10 km and surpassed 10 dBZ in the layer from 4 km to 9 km. Retrievals based on Rasta indicated that the mean diameter of frozen hydrometeors ranged from 0.8 mm at 3.8 km to 0.2 mm at 12 km (Figure 2 (Top)) and ice water content had a range of approximately 0.8-1 g.m<sup>-3</sup> from 3.5 to 6 km, 1.4-1.6 g.m<sup>-3</sup> from 7 to 8 km and 0.1-0.2 g.m<sup>-3</sup> from 8 to 12 km (Figure 2 (bottom)). As a result, the larger hydrometeor loading at lower levels led to larger scattering that in turn led to more depressed brightness temperatures for channels that probed deeper into the atmosphere. As a result, MHS measurements in Figure 1 (Middle)



**Figure 1.** Case of 20 October 2012. (Top) Brightness temperatures from the MSG  $10.8 \mu\text{m}$  SEVIRI channel at 14:18 UTC are shown in grey levels, the track of the Falcon 20 is shown as a blue line, DC and COV are symbolised by green and red spots respectively. (Middle) Dark red ( $183.3 \pm 1$  GHz), dark green ( $183.3 \pm 3$  GHz) and dark blue (190.3 GHz) lines show the MHS brightness temperatures from NOAA 18 satellite at 14:18 UTC for gridpoints that are the closest to the Rasta flight track. DC and COV are symbolised by thick green and red lines respectively. (Bottom) Reflectivity from Rasta radar is shown in colour, the track of the Falcon 20 is shown as a black line, the tropopause height is displayed with a purple line. Black dashed lines have been added to highlight several segments of the cloud that are discussed in the text. Black points are displayed along flight track at 14:24, 14:28 and 14:32 UTC.

showed lower brightness temperatures (205 K) for channels further away from the centre frequency at 183.3 GHz, which was the opposite of that found for clear sky conditions. DC condition (Equation 1) was satisfied in this time window (from 14:24.5 to 14:25.8 UTC).

In the following segment spanning 14:26 to 14:28.5 UTC (③), the clouds also exceeded the 12.5 km height. Reflectivity ranged from 0 to 10 dBZ between 4 km to 12 km and was enhanced above 8 km compared with the previous segment. This cloud section corresponded to the main convective part of the cloud system. Two elevated towers observed at 14:26.5 and 14:27.5 UTC presented reflectivity up to 5 dBZ at 12-km altitude (in contrast, outside these towers reflectivity was lower than 0 dBZ at this altitude). Similarly to the previous segment, lower brightness temperatures (for example, down to 175 K for channel at 190.3 GHz) were found for channels further away from the centre frequency at 183.3 GHz. Since their differences were found to obey the differential order relation given in Equation 2, DC and COV were identified. Microphysical characteristics help to understand why COV is detected in this particular cloud section (Figure 2). The mean diameter of frozen hydrometeors was slightly higher than in the previous segment at high levels (0.8 to 0.5 mm from 3.8 to 8 km, 0.4 to 0.3 from 8 to 12 km) but the main difference was the much more substantial ice water content between 6 and 8 km (1.6 to 1.8  $\text{g}\cdot\text{m}^{-3}$ ) and especially between 10 and 12 km, where values ranged from 1 to 1.4  $\text{g}\cdot\text{m}^{-3}$ . In one elevated tower, the ice water content even reached 1.6  $\text{g}\cdot\text{m}^{-3}$ . This very high ice water content likely explains why COV is detected in this cloud section and not in the former.

Finally, from 14:28.5 to 14:32 UTC (④), the cloud tops were less elevated (cloud tops lower than 12 km) and thinner (approximately 3 km) than those within the main convective region, yet they showed significant reflectivity (up to 10 dBZ). This was due to the rather high ice water content (up to 1.2  $\text{g}\cdot\text{m}^{-3}$ ) (Figure 2). Moreover the mean diameter was larger (0.4-0.5 mm) than in previous segments at the same altitude suggesting that hydrometeors in this segment grew during a longer period of time and probably originate from the convective part of the cloud system. MHS measurements showed that the order between the three channels was reversed compared to the previous sections: lower temperatures (230-239 K) were found for channel  $183.3 \pm 1$  GHz, while higher temperatures (237-252 K) are found for channel 190.3 GHz. Thus DC and COV were not detected. The flight continued after 14:32 UTC, where it reached the base of the upper-level clouds. Once again neither DC nor COV were detected.

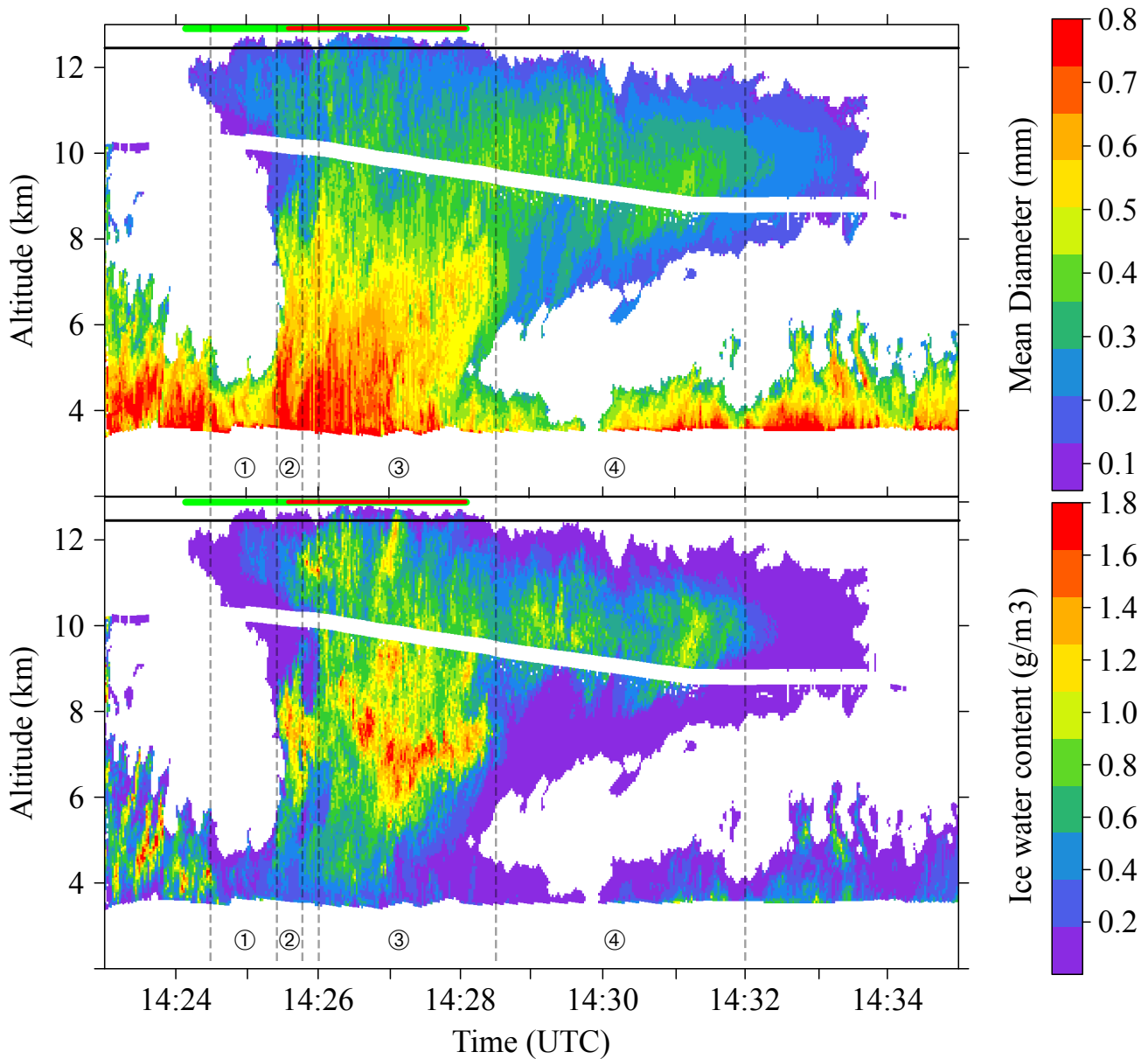
In order to further evaluate the COV detection criterion, we checked the tropopause height in Barcelona using radiosoundings. The tropopause height was determined following the World Meteorology Organisation definition of the tropopause, i.e., “The lowest level at which the lapse rate decreases to  $2^{\circ}\text{C}$  per km or less, provided that the average lapse rate between this level and all higher levels within 2 km does not exceed  $2^{\circ}\text{C}$  per km”. The tropopause height was estimated to be at approximately 12.5 km and 12.3 km respectively at 12:00 UTC and 18:00 UTC. Since the cloud system exceeded these altitudes in the COV-detected regions, the COV criterion was validated for this particular case.

### 3.2. 12 October Case

The 12 October case corresponds to the IOP 12b (Ducrocq *et al.* 2014). The synoptic situation was characterised by trough travelling from the southwestern Mediterranean that destabilised the warm and moist air present over the southern part of the Balearic Islands. In addition a low level convergence was present between Catalonia and the Balearic Islands. These meteorological conditions triggered 100 mm of rain in 24 h with wind gusts ( $100\text{ km}\cdot\text{h}^{-1}$ ), hail and local flooding in the Balearic Islands.

Figure 3 shows DC and COV occurrence derived from MHS at 13:12 UTC on 12 October 2012 (Top), brightness temperatures from MHS (Middle) and reflectivity profiles from Rasta (Bottom). Reflectivity profiles cover the period between 12:50 and 13:20 UTC on the same day for flight 19 of the Falcon 20 aircraft. A convective area is visible over the Balearic Islands, characterised by brightness temperatures down to 200 K (Figure 3 Top). In the same region, an overshooting region of  $170\text{ km}^2$  was detected from MHS, embedded within a deep convective region of approximately  $240\text{ km}^2$  (15 pixels) over the southern coast of Majorca.

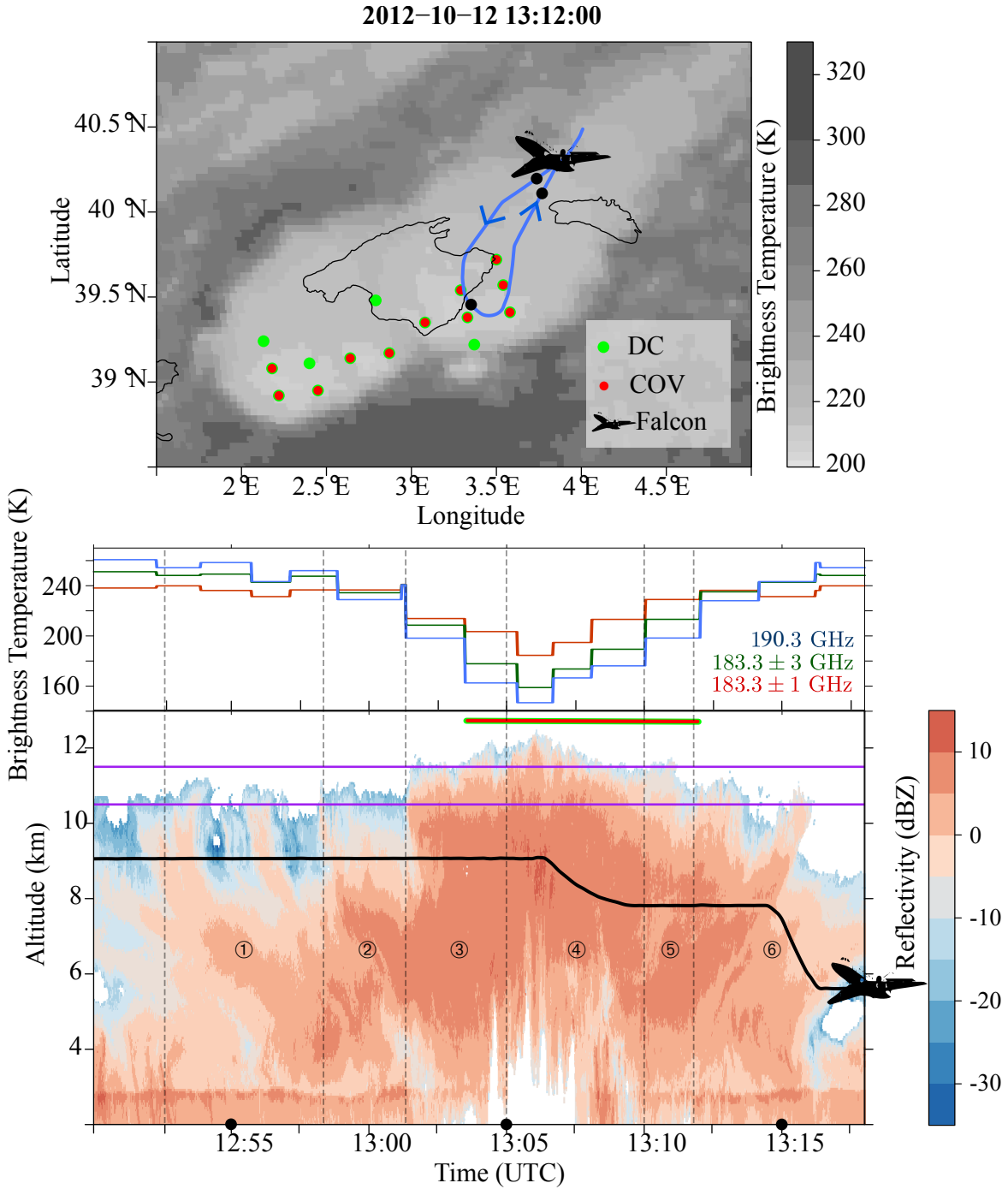
The cloud system sampled by the Falcon 20 aircraft was characterised by high reflectivities (Figure 3 (Bottom)) and again can be described in several segments. In the first segment, from 12:50 to 12:58 UTC (①), clouds reached 10 km in height and showed low reflectivities ranging from -20 to 7 dBZ. Brightness temperatures from MHS (Middle) showed lower temperatures (232-239 K) for channel at  $183.3 \pm 1\text{ GHz}$  and higher temperatures (253-261 K) for channel at



**Figure 2.** Case of 20 October 2012. Mean volume-weighted diameter of frozen hydrometeors (Top) and ice water content (Bottom) retrieved using Rasta radar. DC and COV are symbolised by green and red lines respectively and the tropopause height is displayed with a black line.

190.3 GHz. Consequently, neither COV nor DC were detected. In this cloud section, the mean diameter of the frozen hydrometeors ranged from approximately 0.4 to 0.6 mm between 3 and 6 km, and from 0.1 to 0.4 mm at higher altitudes (Figure 4). The ice water content was low ( $<0.5 \text{ g}\cdot\text{m}^{-3}$ ).

From 12:58 to 13:01.5 UTC (②), the reflectivity increased below 6-km altitude, but the top of the clouds remained unchanged.



**Figure 3.** Case of 12 October 2012. (Top) Brightness temperatures from the MSG 10.8  $\mu\text{m}$  SEVIRI channel at 13:12 UTC are shown in grey levels, the track of the Falcon 20 is shown as a blue line, DC and COV are symbolised by green and red spots respectively. (Middle) Dark red ( $183.3 \pm 1$  GHz), dark green ( $183.3 \pm 3$  GHz) and dark blue (190.3 GHz) lines show the MHS brightness temperatures from NOAA 18 satellite at 13:12 UTC for gridpoints that are the closest to the Rasta flight track. DC and COV are symbolised by green and red lines respectively. (Bottom) Reflectivity from Rasta radar is shown in colour, track of the Falcon 20 is shown as a black line, the tropopause height is displayed with purple lines. Black dashed lines highlight several segments of the cloud that are discussed in the text. Black points are displayed along flight track at 12:55, 13:05 and 13:15 UTC.

From 13:01.5 to 13:05 UTC (③) the aircraft entered into a much more active section of the cloud system: the top of the clouds reached 11.5 km and the reflectivity at mid levels (6-8 km) surpassed 10 dBZ. Although lower brightness temperatures were found for channels further away from the centre frequency at 183.3 GHz, the DC criterion (Equation 1) was not fulfilled since the viewing angle was high ( $>31^\circ$ ) while the difference between channels  $183.3\pm 1$  and  $183.3\pm 3$  GHz was small ( $\sim 5$  K). The mean diameter of frozen hydrometeors was high (0.7 mm) at low levels and moderate at mid and high levels (0.4-0.5 mm), while the ice water content was significant between 6 and 10 km ( $1.3$  to  $1.8 \text{ g}\cdot\text{m}^{-3}$ ).

The following cloud segment (13:05 to 13:10 UTC (④)) appeared to be the most active part of the convective cloud system. The density of hydrometeors (ice, graupels and perhaps liquid water) was so considerable that sometimes, below 6 km, there is complete radar signal extinction. The reflectivity was above 10 dBZ from 7-8 km and approximately 5 dBZ at higher altitudes. Moreover the cloud tops reached 12.3 km at 13:06.5 UTC. As expected, the lowest brightness temperatures ( $<150$  K) were found for channels further away from the centre frequency at 183.3 GHz. Both COV and DC were detected in this region. Looking at the microphysics (Figure 4), the mean diameters were 0.6-0.8 above 7.5 km and the ice water content was from 1.4 to  $2 \text{ g}\cdot\text{m}^{-3}$ . In the most elevated part of the cloud, a column of high ice water content was found ( $>2 \text{ g}\cdot\text{m}^{-3}$ ).

From 13:10 to 13:12 UTC (⑤), the cloud top reached 11.5 km but reflectivity in high levels was much lower (-5 to 0 dBZ) than in previous segments whereas in lower levels, the reflectivity was approximately equal to 5 dBZ (and slightly higher around 6 km). Both DC and COV were detected in this region. It was a little surprising to find COV in the less active region of this cloud system. This can be attributed to the non-optimal collocation between the aircraft track and the MHS observation (they were at about 10 km apart as shown in Figure 3 (Top)). In this segment, the mean diameter of frozen hydrometeors was 0.5-0.6 mm for 3-7.5 km and from 0.5 to 0.15 mm at higher altitudes.

After 13:12 UTC (⑥), the aircraft moved into a less active part of the cloud system with lower cloud tops (8 km) and characterised by lower reflectivities (-5-5 dBZ). The order of brightness temperatures was progressively reverse compared



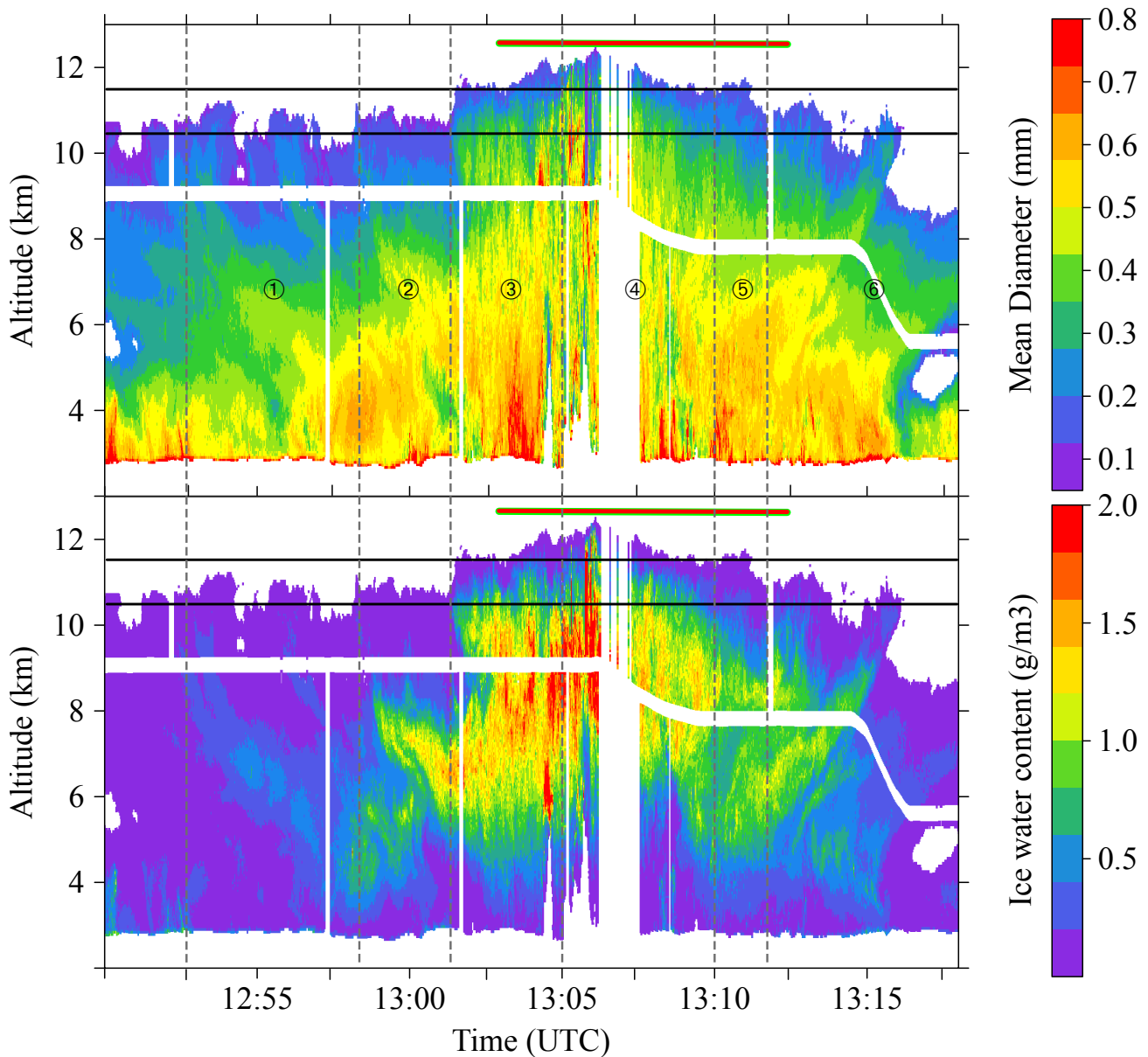
to previous segments. The mean diameter of hydrometeors ranged from 0.6 mm at low levels to 0.2 mm at high levels, and ice water content was lower than  $0.5 \text{ g.m}^{-3}$ .

Once again we checked the tropopause height to evaluate the COV assumption. The two atmospheric soundings from Majorca in Balearic Islands indicated a tropopause height of approximately 10.5 km and 11.5 km respectively at 00:00 UTC and 18:00 UTC. As a result, in this situation too, the COV criterion detected successfully the clouds of high altitude and high reflectivity.

### 3.3. Discussion

These two examples demonstrate the validity of the Hong *et al.* (2005) criterion for detecting convective overshoots in another region than the tropics. In addition, the retrieved microphysical properties of clouds allowed, for the first time, to characterize COV in terms of diameter and content of frozen hydrometeors. We found that COV is associated with a heavy loading of frozen hydrometeors, above  $1.5 \text{ g.m}^{-3}$  in the middle-to-upper troposphere and up to  $0.3 \text{ g.m}^{-3}$  above the tropopause. For the 20 October case, the intrusion into the stratosphere was observed by Rasta during approximately 3.5 minutes. Therefore, given a Falcon 20 flight speed of  $800 \text{ km.h}^{-1}$ , the COV area diameter is estimated to be 50 km. This matches well with the area deduced using MHS. Regarding the 12 October case, as the aircraft was forced to turn around as it reached the most intense part of the cloud (13:05.5), it is not possible to estimate the COV area diameter using Rasta observations.

Within regions where COV was observed, the Rasta observations also revealed several fine-scale elevated towers reaching 12-km altitude with reflectivity up to 5 dBZ, hydrometeor content of  $1.5 \text{ g.m}^{-3}$  (corresponding to a mixing ratio of  $5 \text{ g.kg}^{-1}$  given an air density of  $0.3 \text{ g.m}^{-3}$ ) and diameters up to 0.4 mm. Such towers were observed for about 30 s, or along a distance of approximately 6 km. They corresponded well to the overshooting tops detected in the infrared imagery reported in the literature. The latter appeared as small clusters (<15 km in diameter) of very low brightness



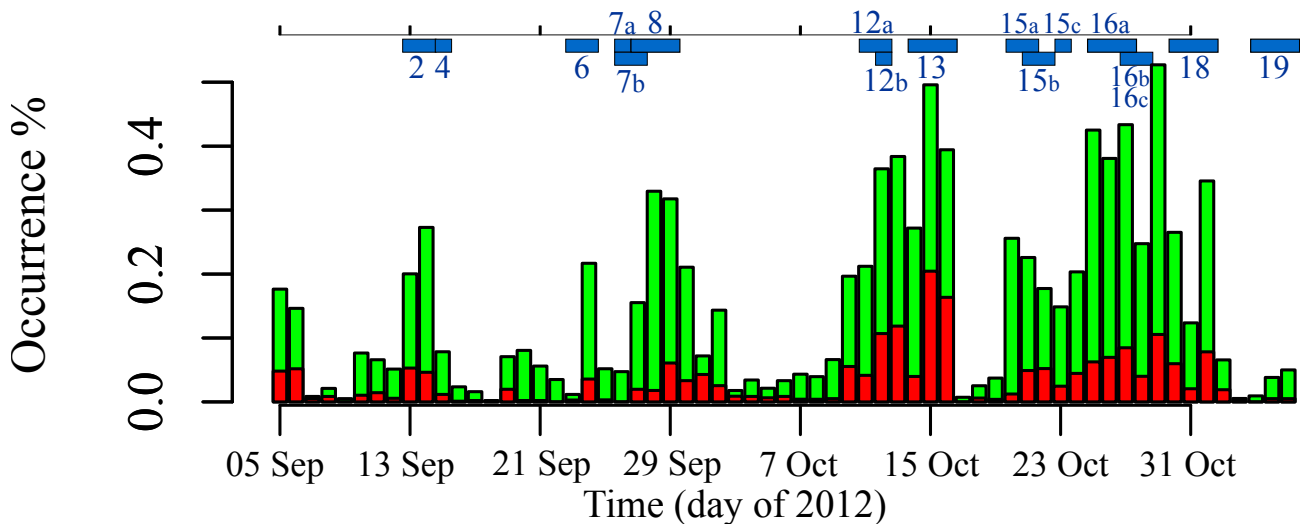
**Figure 4.** Case of 12 October 2012. Mean volume-weighted diameter of frozen hydrometeors (Top) and ice water content (Bottom) retrieved using Rasta radar. DC and COV are symbolised by green and red lines respectively and the tropopause height is displayed with a black line.

temperatures (Bedka 2011). They are explained as vertical updrafts intruding the stratosphere and cooling at a rate of 7-9  $\text{K}\cdot\text{km}^{-1}$ . This results in a significant contrast between the region of overshooting tops and the surrounding anvil clouds. Such updrafts are of great importance for troposphere-stratosphere exchange. To the best of the authors knowledge, both HyMeX cases shown here are the very first observations of the vertical structure of updrafts overshooting the tropopause reported in the Mediterranean.

#### 4. Severe convection during the HyMeX SOP-1

The HyMeX SOP-1 took place from 5 September to 6 November 2012. In this section, we study the temporal and spatial variability of DC and COV occurrences during the whole campaign in the Mediterranean region ( $10^{\circ}$  W- $35^{\circ}$  E and  $32^{\circ}$  N- $48^{\circ}$  N). To this end, we compute DC and COV occurrence for a  $0.2^{\circ}$  grid and we normalise these occurrences using the total number of satellite observations. In the following, the results are given as a percentage of observations.

##### 4.1. Temporal variability



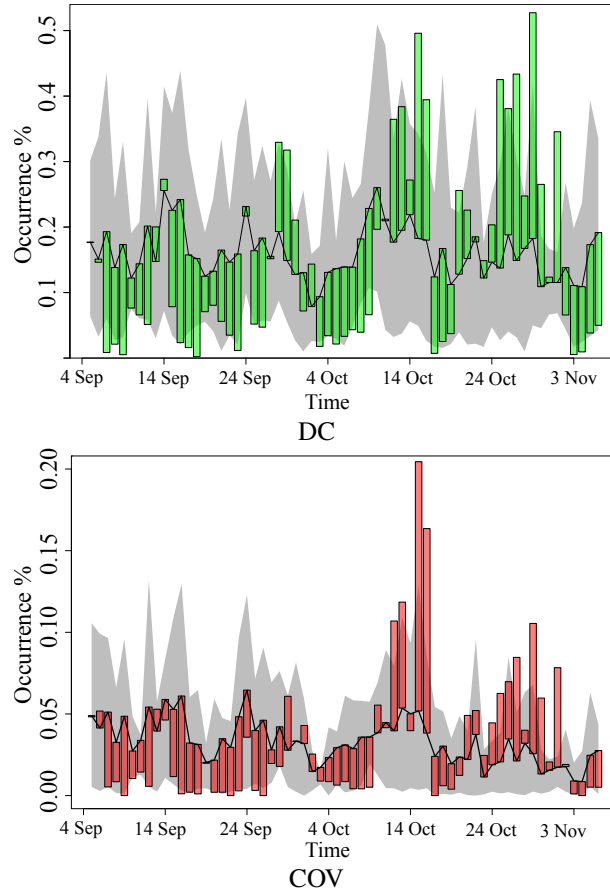
**Figure 5.** Occurrences of DC (green) and COV (red) in % during the HyMeX SOP-1. Note the occurrences are computed for the whole Mediterranean basin (not only in the western Mediterranean where the SOP-1 took place). Intensive observation periods (IOPs) for heavy precipitation events, flash-flood events and strong wind events are shown as blue boxes and labels.

Four episodes of strong convective (DC and COV) activity took place during the HyMeX SOP-1 (Figure 5): 13-14 September, 27-30 September, 10-16 October and 20 October-3 November. These episodes cover all the IOPs (Intense Observation Periods) targeted on heavy precipitation, flash-floods and strong wind events (Ducrocq *et al.* 2014). Less than 40 % of total DC events and less than 35 % of total COV events occurred during the first part of the HyMeX SOP-1 (5 September-9 October). Over the northwestern Mediterranean, supercells with heavy hail and wind gusts were reported in Catalonia and Majorca on 12-13 September. In another event, several mesoscale convective systems associated with a cut-off low resulted in a flash-flood and a tornado that caused 13 casualties and 120 millions of euros of damage in

Spain between 27-29 September. The second part of the HyMeX SOP-1 was more active. The 10-16 October episode is characterised by heavy convective activity during 0.25 % of the time on average, with COV occurrence accounting for a third of all deep convection cases. A second episode (20 October-3 November) had the highest activity in terms of convective events ( $\sim 0.3$  % of the time on average, with a peak of 0.5 % on 29 October) but showed fewer strong and elevated convective towers (0.05 %). During the last two episodes, severe weather events such as heavy rainfall, flooding, a tornado and high waves were reported over Spain, France and Italy as well as 6 casualties.

The convective activity can be related to the Atlantic weather regimes that occurred during the HyMeX SOP-1 and described by Ducrocq *et al.* (2014). Specifically, episodes with low convective activity were associated with an Atlantic ridge weather regime (5-23 September) and to a positive phase of North Atlantic Oscillation (NAO) (at the beginning of October). Conversely, the episodes of strong convective activity were associated with negative NAO phases (23-30 September and 27 October-5 November) or a blocking phase (16-26 October). These two weather regimes are usually related to unstable weather in the Mediterranean (Ducrocq *et al.* 2014). For instance Chaboureau and Claud (2006) showed that cloud systems were found more frequently during the negative NAO phase. Overall these results show a relationship between occurrence of DC and COV and large-scale weather systems and are consistent with Funatsu *et al.* (2008) who showed that an upper-level trough was present in about half of DC events in the autumn.

In order to compare the observed patterns of convective activity during the HyMeX SOP-1 with previous years, we used the same time frame (i.e., 5 September to 6 November) from 2002 to 2013 to compute the 12-year average of DC and COV occurrence (Figure 6). Overall the frequency of DC events swung around 0.2 %  $\pm 0.1$  % while the occurrence of COV events decreased with time from approximately 0.05 to 0.025 %. The inter-annual variability of DC and COV was strong, as shown by the interquartile range (shaded area). This was also true for the HyMeX SOP-1 as shown by the difference in their occurrence against the 2002-2013 period. During the HyMeX SOP-1, DC and COV activity was not notably different from the 2002-2013 period except for the episode of 11-16 October 2012 which was especially noticeable with COV occurrence 2 to 4 times above average values. During this episode, two exceptionally large convective systems

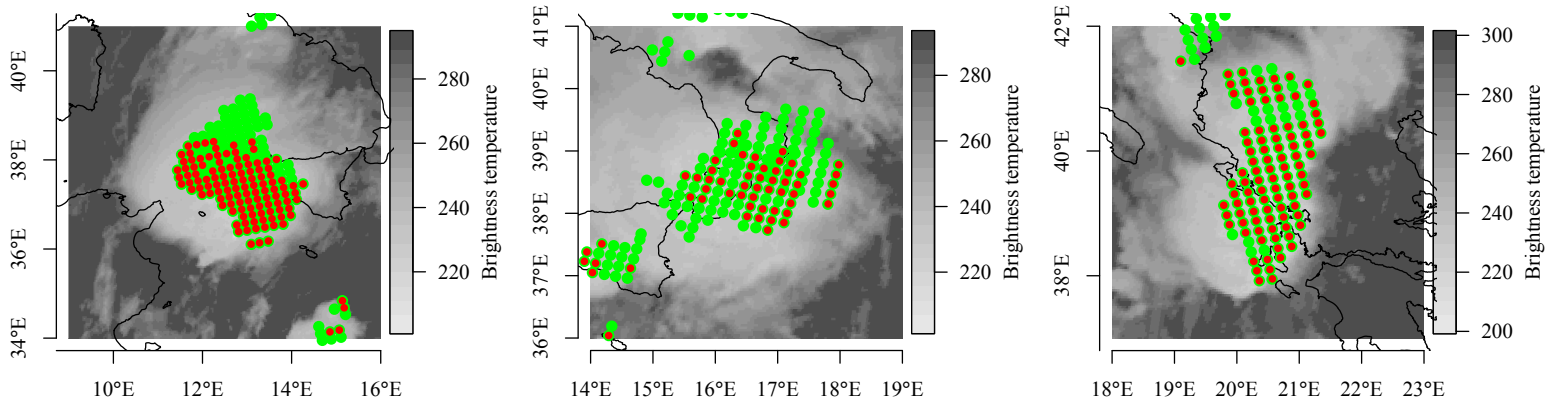


**Figure 6.** Black line indicates average DC (green, top) and COV (red, bottom) occurrence during the 2002-2013 period. Boxes indicate the difference between the HyMeX SOP-1 and the 2002-2013 period. Shaded grey shows the interquartile range of DC/COV occurrences during the 2002-2013 period.

2012-10-15 19:40:00

2012-10-16 00:45:00

2012-10-16 12:35:00



**Figure 7.** Case of 15-16 October 2012. Brightness temperature from MSG 10.8  $\mu\text{m}$  SEVIRI channel is shown in grey levels and DC and COV are symbolised by green and red spots.

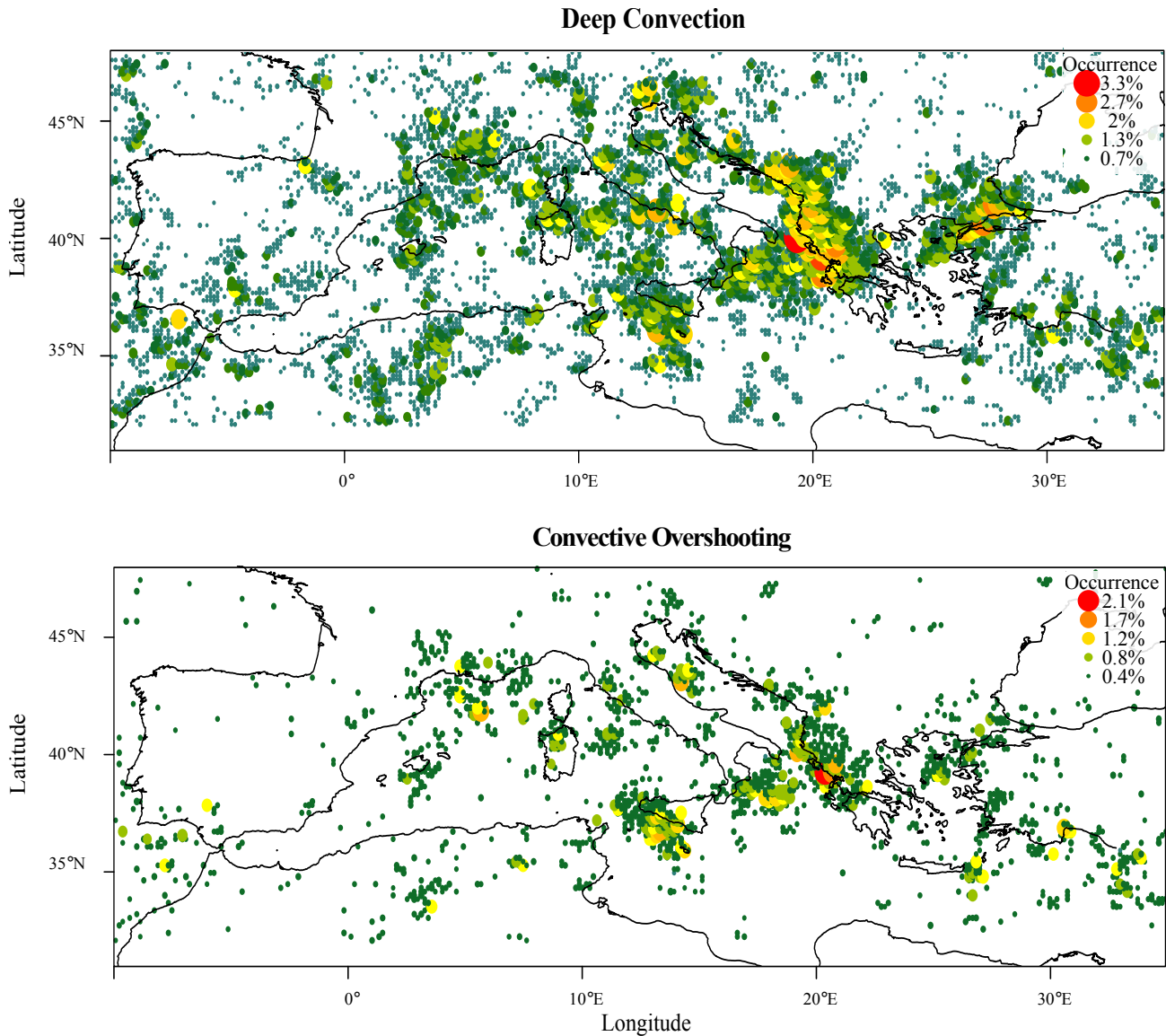
with diameters higher than 300 km travelled from Tunisia to Greece and numerous convective structures developed in the Ionian Sea. The 15-16 October event is presented in Figure 7. The large scale situation of the 15-16 October event was quite

common with a cold front from Algeria to Corsica that moved slowly eastward. Ahead of this front, a mesoscale cyclonic system developed a considerable COV region with a diameter of 200 km making this event particularly remarkable. Unfortunately no flight was possible in this area at that time. During these two days, heavy rainfall along the eastern coast of the Adriatic Sea were reported as well as severe winds between Tunisia and Sicily, large hail in Sicily and a tornado along the western coast of Greece.

Figure 6 also reveals that from 25 October to 2 November, convective activity (DC) was twice as high as the 12-year daily average value. This convective activity was related to a low pressure system that travelled slowly from western Spain to eastern Greece. This low pressure triggered frontal activity that in turn favoured convective activity ahead of the front. During this period, meteorological services reported numerous severe weather events in the Mediterranean: heavy rain and severe winds along western Italy, south-eastern France and the eastern Adriatic coast, tornadoes in northeastern Spain, Corsica, central Italy, Sicily, Malta, Croatia and western Greece.

#### 4.2. Spatial variability

The spatial distribution of DC and COV is shown in Figure 8. Areas of high convective activity were mainly located over the sea and the north-western coasts of the Mediterranean basin. In particular, Albanian and Greek western coasts showed DC occurrence as high as 2.5-3 % and Turkey, Sicily, Corsica, Sardinia, the Balearic Islands and the western and southern coasts of France, as well as the Sea of Marmara showed DC occurrences from 1.5 to 2 %. Regarding COV, the western Sicilian coast was strongly affected by these events during the HyMeX SOP-1 (frequency up to 1.5 %). Other areas with significant number of COV were the southern coast of Italy (1 %), the western coast of Albania and Greece (up to 2 %), the Gulf of Lyon (> 1 %), the Tyrrhenian Sea (> 2 %), the Ionian Sea (> 1 %) and the Straits of Sicily (> 2 %). Only a few studies provide information about the occurrence of convective activity over the sea in the Mediterranean region: Funatsu *et al.* (2009) showed a similar pattern to our results, with DC occurrence mainly over the sea, with the exception of the Spanish and Tunisian coasts where DC is more frequent in Funatsu *et al.* (2009). Similarly, Melani *et al.* (2013)

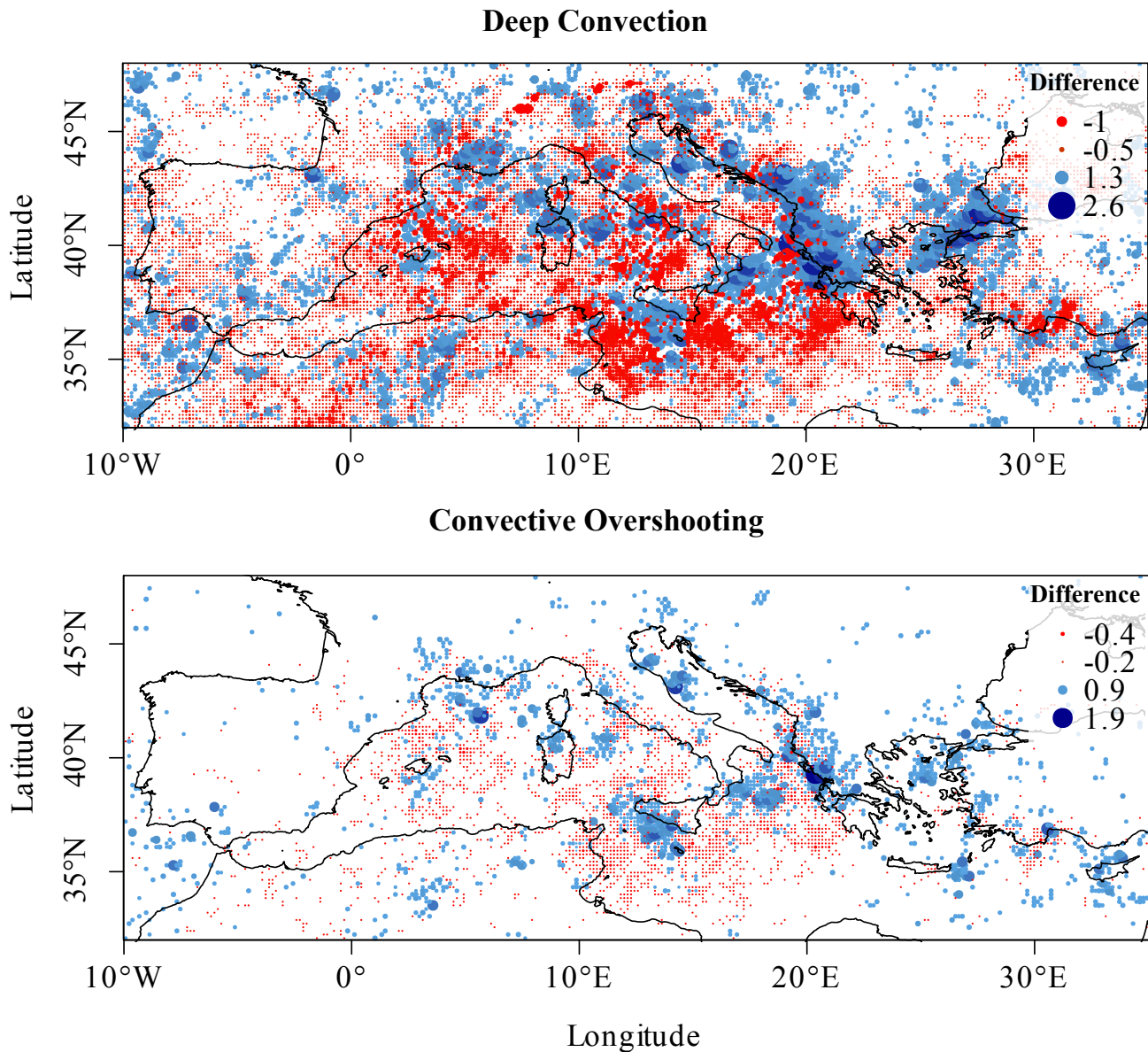


**Figure 8.** Occurrences of DC (top) and COV (bottom) during the HyMeX SOP-1. The occurrence are expressed as specific values.

showed that long lasting deep convective systems occur in regions where we identified high occurrence of DC, that is southeastern Sicily, the Balearic Islands and the Ionian and Tyrrhenian seas, while Bedka (2011) detected overshooting tops preferentially in the western Mediterranean sea and northwestern Africa using infrared imagery.

The interannual variability in the convective activity in autumn is high as noted by Funatsu *et al.* (2009); Bedka (2011) and Melani *et al.* (2013). Figure 9 shows the difference in DC and COV between the HyMeX SOP-1 and the whole 2002-2013 period. On the one hand DC and COV events are slightly below average during the HyMeX SOP-1 for the





**Figure 9.** Difference of occurrence of DC and COV between the HyMeX SOP-1 and 2002-2013 period. The difference are expressed as specific values.

south-western part of the Mediterranean region (down to -1 %). On the other hand, DC occurrence is above the mean for the western Albanian and Greek coasts ( $\sim 1.5-2.5$  %), while COV occurrence is above the mean for the western Greek and Sicilian coasts as well as southern Italy ( $\sim 1-2$  %).

Both Figures 8 and 9 reveal the existence of a DC “alley” from Tunisia to Greece during HyMeX SOP-1. This DC “alley” is also observed on average for the whole 2002-2013 period (not shown). A number of mesoscale convective systems occurred during the HyMeX SOP-1 in the Adriatic Sea and the Ionian Sea as revealed by the MSG imagery (not



shown). They likely explain the high number of DC and COV occurrences in this region. As underlined in Melani *et al.* (2013), who also observed strong convective activity over these areas based on a totally independent approach, reasons that could explain preferential areas for DC and COV occurrence remain unclear, but two factors play a significant role: first, in the Mediterranean region, coastal regions often present steep orography that triggers convergence and vertical air motion (Rotunno and Ferretti 2003; Bousquet and Smull 2006). Second, positive anomalies of the sea surface temperature might also affect the distribution of convective events (Funatsu *et al.* 2009). It is also interesting to note that the track of medicanes (Mediterranean tropical-like cyclones) often matches the DC “alley” (Claud *et al.* 2010; Chaboureau *et al.* 2012b; Tous and Romero 2013; Miglietta *et al.* 2013). Our results also show that the western coasts of land are mostly affected by DC and COV, very likely because these convective events are often associated with frontal systems that come from the west i.e., the Atlantic region (Kotroni *et al.* 1999; Funatsu *et al.* 2008).

## 5. Comparison of DC and COV occurrence between observations and the Meso-NH model

One of the main objectives of the HyMeX campaign is to improve the forecast of heavy precipitation events. To this end, about 30 atmospheric research and operational models were run during the whole SOP-1 (see Table ES4 of Ducrocq *et al.* 2014). Among them, 11 convection-permitting models provided deterministic forecasts over domains covering the western Mediterranean with grid spacing ranging from 1.5 to 3 km (see Figure 5 of Ducrocq *et al.* 2014). In this section, we evaluate the capacity of the Meso-NH model to reproduce DC and COV occurrence during the HyMeX SOP-1. As yet, no attempt has been made to calculate any categorical score for assessing the capability of Meso-NH to forecast DC and COV at the right location and at the right time.

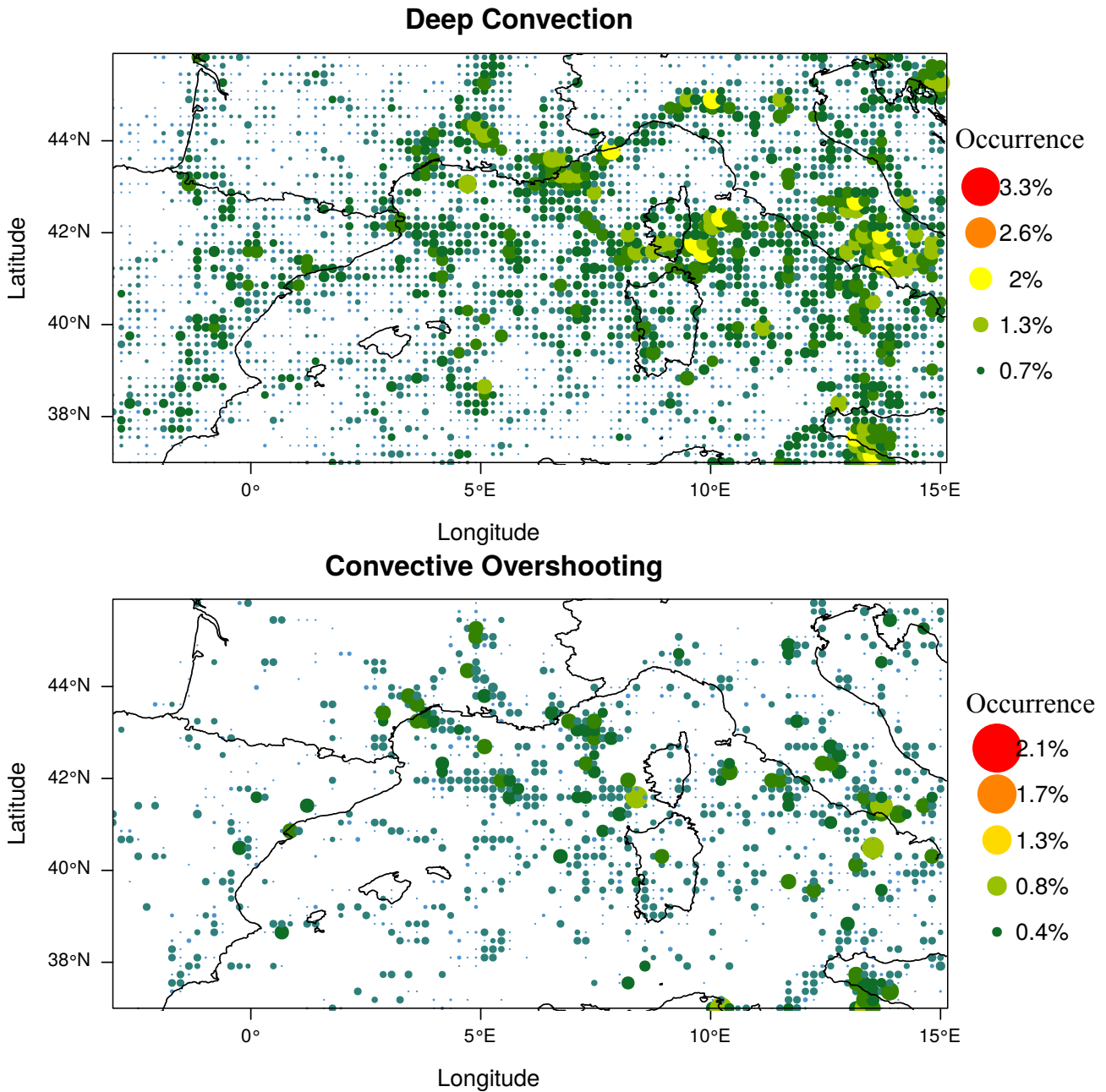
### 5.1. Meso-NH specification

During the HyMeX SOP-1, the Meso-NH model (Lafore *et al.* 1998) was run daily for 48 h from 29 August to 6 November 2012. The initial and boundary conditions were provided by European Center for Medium-range Weather Forecasts (ECMWF) analysis and forecasts issued at 00:00 UTC each day. In the case of heavy precipitation events over

the western Mediterranean, the model was run using the two-way interactive grid-nesting method (Stein *et al.* 2000) with two nested grids, a horizontal grid mesh of 15 and 2.5 km and a vertical grid with 62 levels. Otherwise, the model was run with the coarser grid only. The model includes parameterisations for radiation (Mlawer *et al.* 1997), turbulence (Cuxart *et al.* 2000), subgrid shallow convection (Pergaud *et al.* 2009), mixed-phase microphysics (Pinty and Jabouille 1998), subgrid cloud cover and condensate content (Chaboureau and Bechtold 2005) and surface exchanges (Masson *et al.* 2013). The convection scheme of Bechtold *et al.* (2001), was activated for the 15-km grid, while convection was assumed to be explicitly resolved for the 2.5-km grid (648 by 480 grid points, see the domain in Figure 8). From the model outputs available every 3 hours, brightness temperatures were computed using the radiative transfer code RTTOV (Radiative Transfer for Tiros Operational Vertical Sounder) version 11 (Saunders *et al.* 2005). The scattering properties of frozen hydrometeors were those calculated by Geer and Baordo (2014) for dendrite snowflakes using the discrete dipole approximation, while precomputed Mie tables were used for other hydrometeors. The simulation of brightness temperatures from the model outputs allowed a direct comparison with passive microwave observations. Such a comparison, the so-called model-to-satellite approach has already been performed successfully at various wavelengths and has shown the overall performance of the Meso-NH model to correctly predict cloud and rain fields (Chaboureau *et al.* 2008; Söhne *et al.* 2008; Clark and Chaboureau 2010; Chaboureau *et al.* 2012a,b, amongst others). In this study, the comparison is restricted to the inner model (2.5-km grid) where the convection is resolved. For each mesh point of the model, the MHS synthetic brightness temperatures are computed and the DC and COV criteria are applied. The normalised occurrence of DC and COV is then computed for a 0.2 ° grid to be compared with those derived from MHS observations.

## 5.2. Results

As observed, DC was mainly forecasted over the Mediterranean sea and the surrounding land, with occurrences ranging mainly from 0.7 to 1.3 % (Figure 8). Moreover, DC was present for almost every grid point in the simulations, while absent in some regions in the observations. This non-ubiquitous distribution in the observations may be attributed to a

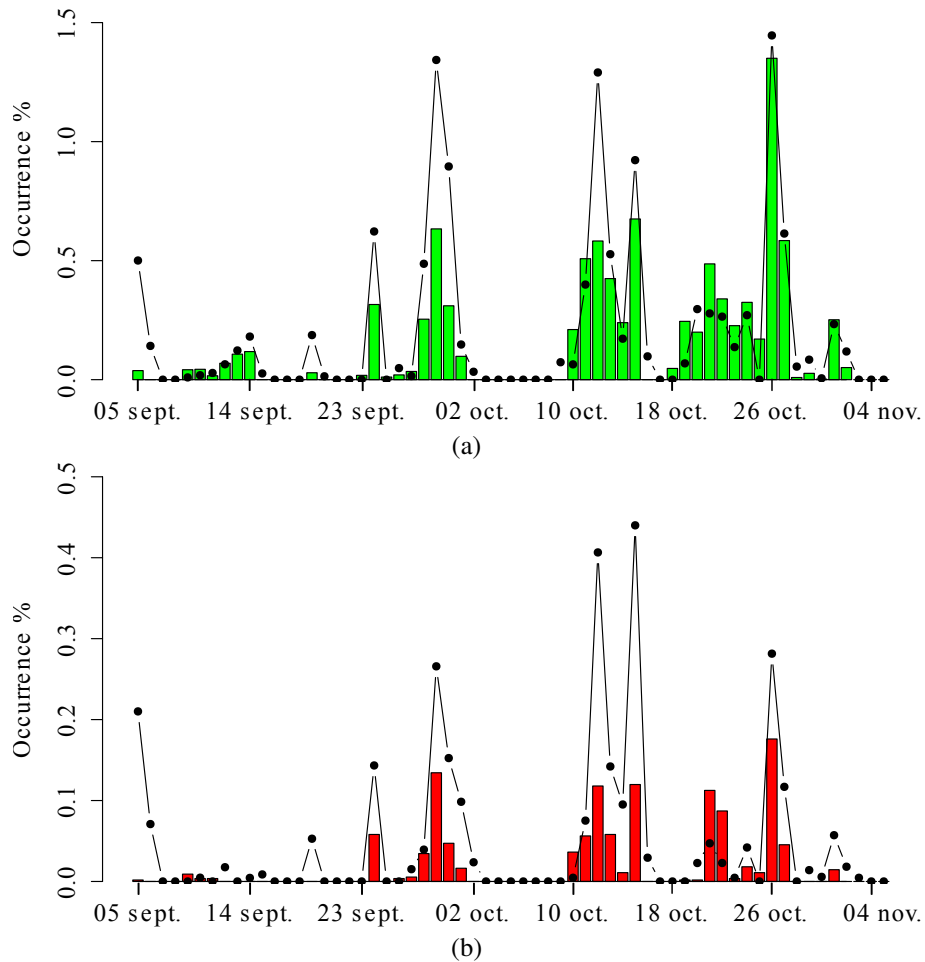


**Figure 10.** Occurrence of DC (top) and COV (bottom) during the HyMeX SOP-1 for the Meso-NH model. The occurrences are expressed as specific values.

sampling issue. Another explanation could be a lack of organised mesoscale convective systems in the forecasts. Indeed, more isolated convective storms are likely to result in DC occurrence which are scattered more widely over the domain.

However, the model shows a few areas with occurrences larger than 1.3 %: southwestern Sicily, eastern Corsica, in the vicinity of Roma, northwestern Italy and southeastern France. In particular, DC was forecasted over Sicily at the right

location and with the right occurrence. Other hot spots in the forecasts were slightly offset compared to the observations (e.g., DC was observed in western Corsica, but forecasted in eastern Corsica) and with lower occurrence (Figure 8). Again, this could be explained by the sampling issue and some model deficiencies as detailed below. The distribution of forecasted COV occurrence mimics its DC counterpart, but with a reduced number to those observed.



**Figure 11.** DC occurrence for the Meso-NH model (green) and observations (black dots) (Top). COV occurrence (red) and observations (black dots) (Bottom). Occurrences are computed for the 2.5-km grid of the Meso-NH model ( $3^{\circ}$  W- $15^{\circ}$  E and  $37^{\circ}$  N- $46^{\circ}$  N).

The comparison was further pursued by examining the time evolution (Figure 11). Periods of high DC activity (13-16 September, 27-31 September, 10-15 October and 19 October to 1 November) were well forecasted. As a result, the correlation coefficient between forecasts and observation was 0.86. Such a high value is likely due to the nature of convective systems that yield a DC signature. Indeed, as noted in the previous section, most of them were associated

with synoptic-scale troughs that are usually well described in modern operational analysis (Ducrocq *et al.* 2014; Ferretti *et al.* 2014). As a consequence, convection was triggered at the right time in the daily forecasts of the convection-permitting model. Yet, the average DC occurrence in the forecasts was two times lower than in the observations. The time evolution of COV was less accurately reproduced than for DC and the correlation coefficient was 0.74. Overall, the occurrence of COV in the forecasts was five times lower than in the observations.

Several factors could explain the discrepancies between observations and forecasts, in particular those seen in the spatial distribution. The first factor is the sampling issue. On the one hand, the sensitivity of DC and COV to the time of the day was examined in the three-hourly forecasts. It results that DC and COV are both very intermittent events. On the other hand, the three polar-orbiting satellites fly over the Mediterranean up to 6 times a day, so less regularly and frequently than the 3-hourly model outputs. However no systematic variation of DC and COV with time was found in the forecasts (not shown). This suggests that the observations should not present any systematic bias with respect to the time of the day.

A second factor explaining the discrepancy could be the frozen hydrometeors produced by the model. They have certainly not always been forecasted at the right time and/or the right place. Forecasting localised intermittent events such as thunderstorms is a big issue that is, however, outside of scope of the present study. Beyond this predictability issue, it is not clear whether the frozen hydrometeors are forecasted with correct amounts or not. In a recent model inter-comparison study focusing on an intense tropical mesoscale convective system, Varble *et al.* (2014) found a high positive bias on simulated convective areas for all the evaluated models, including Meso-NH. This bias was the result of overly intense updrafts partly due to over-strong convective dynamics and the freezing of liquid condensate in parameterised microphysics. Conversely, we show here that DC and COV were underestimated by the model, highlighting the difficulty in forecasting severe convection accurately and universally.

A final factor explaining the discrepancy is the great sensitivity of the simulated microwave signal to the assumptions made in the radiative transfer model (Saunders *et al.* 2005). In particular, Geer and Baordo (2014) found that the simulated signal varies greatly with the choice of snow particle shape and size distribution. We too found a similar sensitivity

here (not shown). A more detailed case study will allow us to disentangle the uncertainties due to the dynamics and the microphysics from those associated with the radiative properties of frozen hydrometeors.

## 6. Conclusion

In this paper, we used a method proposed by Hong *et al.* (2005) to detect Deep Convection (DC) and Convective Overshooting (COV) during the HyMeX SOP-1, based on passive microwave measurements from MHS (Microwave Humidity Sounder). In the first part of this study, we took advantage of the observations gathered during the SOP-1 to evaluate the COV method (the evaluation of DC has been achieved previously). In particular, we validated the method for two convective events sampled conjointly by MHS and the airborne Rasta radar. As expected, COV was detected for these two events where convective clouds rose beyond the tropopause. Our analyses showed that COV were associated with high reflectivity (up to 15 dBZ) and high ice water content (up to  $2 \text{ g m}^{-3}$ ) in mid and high levels of the troposphere.

In a second part, we studied the spatial and temporal variability of DC and COV during the HyMeX SOP-1. A risk for every field campaign is that the observation period does not present the expected meteorological conditions, but we showed that while the beginning of the campaign was dryer than one would have expected, the second part was impacted by a significant number of heavy precipitation events. The most active episodes of the HyMeX SOP-1 were found in the 10-16 October and the 20 October-2 November periods. The whole northern part of the Mediterranean basin was impacted by severe convection but the most affected regions were the western part of the Greek coast, the Sea of Marmara and the Sicilian Island where up to two cumulated days of DC occurred during the HyMeX SOP-1. The comparison of the HyMeX SOP-1 with the 2002-2013 period indicated that the SOP-1 falls mostly within the interquartile range, except for some instances in the second half of the SOP period. The southern part of Mediterranean basin was less impacted than average while the northern part (and in particular the Greek coast) was more impacted. Moreover, less DC and COV than average occurred during the period from 5 September to 11 October, while during the second part of the SOP-1 (11 October-6 November), convective activity was higher than average. The temporal distribution of convective events appeared to be

related to the Atlantic weather regimes. In particular, NAO- and blocking conditions were associated, for the SOP-1, with strong convective activity.

The last part of this paper was devoted to the evaluation of severe convection forecasts in the Meso-NH model through a model to satellite approach. Overall, this evaluation revealed that the forecast of DC and COV is satisfactory since their locations are rather similar to the observations, and with fairly high temporal correlation coefficient values (0.86 for DC and 0.74 for COV). Yet, the occurrences computed from the convection-permitting model tend to be lower than in the observations. These results underline the efforts needed for a better characterisation of the model uncertainties associated with the dynamics, microphysics and radiative properties of frozen hydrometeors.

In summary, this paper showed that the DC and COV criteria developed by Hong *et al.* (2005) can be applied in the Mediterranean region during the autumn. It also provided very unique observations of the vertical structure of two severe thunderstorms that could help to gain a better understanding of their underlying dynamics. Moreover, it revealed periods and regions that were the most favourable to DC and COV during the HyMeX SOP-1. An important aspect is that we detected a significant amount of DC and COV over sea where very few measurements are available. Thus there is a clear need to find innovative solutions (e.g., meteorological radar on-board ships, microwave radiometers on-board geostationary satellites) in addition to the existing ones (MHS on-board heliosynchronous platforms, SEVIRI on-board MSG) to monitor severe convection over sea.

As this study highlights the occurrences of COV events in the Mediterranean region, the question of possible troposphere-stratosphere exchanges associated with these events must be addressed, as they can significantly modify stratospheric air composition. For instance, in one of the few in-situ observations of troposphere-stratosphere exchange related to deep convection at mid-latitudes (see also Poulida *et al.* 1996; Hegglin *et al.* 2004; Homeyer *et al.* 2014), Fischer *et al.* (2003) showed that for an event in the Mediterranean, that up to 50 % of lower stratospheric air originates from boundary layer air in such convective events.

Space-borne passive microwave sounders for water vapour observation have provided data since 1999 with several different radiometers, among them MHS, AMSU-B and ATMS (Advanced Technology Microwave Sounder) are nearly equivalent instruments. In this paper, we only used the former, but we plan to use the other two to provide a climatological analysis of DC and COV over the whole period. The same analysis could also be extended to other mid-latitude regions. In the future we also plan to use the huge number of measurements acquired during the HyMeX campaign for case studies that will help to characterise small scale properties of DC and COV events better. In particular, 3D reflectivity and microphysical characteristics derived from the radar network measurements appear to be promising for a better understanding of COV (see recent works of Homeyer and Kumjian (2014)). Finally, the ability of other models that ran during the HyMeX SOP-1 to forecast DC and COV in the Mediterranean basin will be evaluated.

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