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1 **Title:** Photosynthesis and calcification in the articulated coralline alga *Ellisolandia elongata*
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3

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13

14 **Short running title:** Photosynthesis and calcification in *E. elongata*

15

16

17 **Abstract**

18

19 Calcifying coralline algae are functionally important in many ecosystems but their existence
20 is now threatened by global climate change. The aim of this study was to improve our
21 understanding of coralline algal metabolic functions and their interactions by assessing the
22 respiration, photosynthesis and calcification rates in an articulated (geniculate) coralline
23 alga, *Ellisolandia elongata*. Algae were collected from an intertidal rock-pool on the coast
24 of Brittany (France). Physiological rates were assessed in summer and in winter by
25 measurements of oxygen, dissolved inorganic carbon and total alkalinity fluxes at five
26 irradiance levels and in the dark using incubation chambers.

27 Respiration, photosynthetic and calcification rates were strongly affected by season.
28 Respiration increased with temperature, being ten-fold higher in summer than in winter.
29 Photosynthetic parameters of the photosynthesis-irradiance ($P-E$) curve, P_g^{max} , P_n^{max} and E_k ,
30 were two- to three-fold higher in summer relative to winter. Photoinhibition was observed
31 under high irradiance levels indicating an acclimation of *E. elongata* to low irradiance.
32 Parameters of the calcification-irradiance ($G-E$) curve, G^{max} and E_k , were approximately
33 two-fold higher in summer, compared to winter. Calcification in summer was more strongly
34 inhibited at high irradiance levels than photosynthesis. By inhabiting intertidal rock pools,
35 *E. elongata* exhibits tolerance to a dynamic physico-chemical environment. Information on
36 respiration, photosynthesis and calcification rates in a calcifying coralline alga inhabiting
37 such dynamic environments in terms of pH and temperature is important for better
38 understanding of how global changes, including ocean acidification and warming, will
39 affect coralline algae in the future.

40

41

42 **Introduction**

43

44 Climate change and ocean acidification are advancing as a result of anthropogenic
45 increase in carbon dioxide concentrations in the Earth's atmosphere (IPCC, 2013).
46 Investigating organism physiology in relation to environmental variability provides useful
47 information for predicting acclimation and adaptation to future conditions (Kelly &
48 Hofmann, 2013). Calcifying benthic algae are generally considered highly sensitive to
49 climate change and ocean acidification (Kroeker *et al.*, 2013) but many species are also
50 common in habitats where pH and temperature gradients are steep. Among calcifying
51 benthic algae, red coralline algae (Corallinales, Rhodophyta) are the most important taxa in
52 terms of global carbon cycling and ecosystem engineering (Nelson, 2009). They can be
53 found over a wide latitudinal range and within all depths of the photic zone (Adey &
54 Macintyre, 1973; Steneck, 1986) where they are known to provide ecosystem services such
55 as by sheltering high biodiversity and favour larval recruitment and settlement (Adey,
56 1998). Coralline algae are also of importance with respect to the carbon and carbonate cycle
57 in shallow coastal waters through high community calcium carbonate production (Martin *et*
58 *al.*, 2007). Articulated coralline algae (also referred to as geniculate coralline algae) have
59 non-calcified segments between calcified intergenicula that provide flexibility to the erect
60 algal thallus (Martone & Denny, 2008). They have been reported to support diverse
61 assemblages of macrofauna and are important providers of three dimensional frameworks
62 (Kelaher *et al.*, 2001), being considered as ecosystem engineers (Dayton, 1972; Benedetti-
63 Cecchi, 2006; Noël *et al.*, 2009). Their unpalatable calcareous structure acts as a barrier
64 against grazing, thus protecting diverse algal epiphyte assemblages (Pfister & Hay, 1988).

65 Coralline algae have been identified as among the most vulnerable calcifying marine taxa
66 threatened by ocean acidification and decreasing calcium carbonate saturation levels
67 (Kroeker *et al.*, 2013). Reliable predictions of coralline algal responses to future
68 environmental changes require knowledge of the processes involved in calcification and
69 their interaction with other metabolic processes and environmental factors. Coralline algae
70 precipitate calcite within the cell wall and can also deposit calcite within intercellular spaces
71 (Borowitzka, 1989). They precipitate high-Mg calcite (Ries, 2011) which becomes more
72 solvable with increased incorporation of Mg^{2+} into the calcite lattice. The solubility of
73 calcite exceeds that of aragonite when the Mg/Ca mol% is above approximately 12 (Morse
74 *et al.*, 2006; Andersson *et al.*, 2008) with higher values often observed in calcite produced
75 by coralline algae (Chave, 1954). An organic matrix within the cell wall is important for the
76 calcification process but the exact chemical and functional properties of the organic matrix
77 are yet not fully understood (Borowitzka, 1989; Rahman & Halfar, 2014). For example,
78 only recently the presence of chitin was discovered in a coralline alga species, a substance
79 which likely serves an important role in nucleation, growth and orientation of the mineral
80 crystals (Rahman & Halfar, 2014).

81 The equilibrium between inorganic carbon species in an aquatic system is an important
82 environmental factor affecting the processes that underlay $CaCO_3$ bioprecipitation
83 (Borowitzka, 1982; Borowitzka & Larkum, 1987; Raven, 2011; Basso, 2012). For example,
84 supersaturation of calcium carbonate is generally required for calcification to take place in
85 algae (Borowitzka, 1989). The calcium carbonate saturation state (Ω_{CaCO_3}) in seawater and
86 at the calcification site is affected by photosynthesis and respiration through the removal or
87 addition of CO_2 from or to the extracellular environment (Borowitzka, 1989). In this way,
88 calcification is induced by photosynthesis as is evident by the strong relationship between

89 the metabolic rates of these processes in all calcifying algae (Borowitzka & Larkum, 1976;
90 Pentecost, 1978; Borowitzka, 1989; Martin *et al.*, 2013a).

91 Photosynthetic rates in algae are primarily controlled by environmental factors such as
92 irradiance, nutrient concentration, temperature and the availability of inorganic carbon as
93 photosynthetic substrate (Hurd *et al.*, 2014). A useful method for investigating
94 photophysiological characteristics in algae is to look at the relationship between irradiance
95 and photosynthesis through *P* (photosynthesis) versus *E* (irradiance) curves (also termed *P-I*
96 curves) (Lüning, 1990). Photosynthetic rates increase with increasing irradiance until the
97 photosynthetic mechanism of the algae becomes saturated (Lüning, 1990). The irradiance
98 needed to saturate an alga is generally high for algae acclimated or adapted to high levels of
99 irradiance and low for algae acclimated or adapted to dim light conditions (Lüning, 1990).
100 Under higher irradiances algae may either continue at the maximum photosynthetic rate or
101 photosynthetic rates may be negatively affected above a certain irradiance level, i.e.
102 photoinhibition occurs (Lüning, 1990). Due to the close relationship between
103 photosynthesis and calcification an irradiance response curve may also be produced for
104 calcification (then termed *G-E* (or *G-I*) curve) (Chisholm, 2000; Martin *et al.*, 2006; Martin
105 *et al.*, 2013a; Martin *et al.*, 2013b). Through comparing irradiance response curves for
106 photosynthesis and calcification in algae, a better understanding of the relationship between
107 these metabolic processes may be achieved.

108 Articulated coralline algae frequently inhabit the intertidal environment and in particular
109 tidal pools (Johansen, 1981). Tidal pools are habitats where water exchange is limited or
110 temporally non-existent as a result of local tidal activity (i.e. amplitude and tidal cycles) and
111 pool locality (i.e. high or low within the intertidal zone) (Morris & Taylor, 1983). Because
112 of biological processes and physico-chemical fluxes, the carbonate chemistry within a tidal

113 pool may exhibit great periodical variations which is generally reflected by large pH
114 fluctuations (Morris & Taylor, 1983; Huggett & Griffiths, 1986).

115 *Ellisolandia elongata* (previously known as *Corallina elongata*) inhabits intertidal rock
116 pools where large fluctuations in physico-chemical variables occur diurnally (Egilsdottir *et*
117 *al.*, 2013). It has been suggested that organisms inhabiting environments where large
118 fluctuations in CO₂/pH and temperature are common will have a relatively good capacity to
119 acclimate or adapt to future ocean acidification and warming (Kelly & Hofmann, 2013;
120 Hofmann *et al.*, 2014). Indeed, studies have indicated a greater tolerance to ocean
121 acidification in *E. elongata* previously inhabiting a rock-pool environment when compared
122 to other coralline algae whose distribution is limited to more stable environments
123 (Egilsdottir *et al.*, 2013; Noisette *et al.*, 2013).

124 Although, *E. elongata* is predicted to exhibit resilience under future conditions, there have
125 been no studies on the relationship between photosynthesis and calcification in *E. elongata*
126 under natural conditions, and there is a lack of information on the seasonal difference in the
127 rates of these processes. Such data is valuable, for example for predicting the response of *E.*
128 *elongata* to future conditions. The aim of this study was to characterize the metabolic
129 activity of photosynthesis and calcification and how they interact *in situ* in the articulated
130 coralline alga *E. elongata* inhabiting tidal pools where strong physico-chemical variations
131 occur.

132

133 **Material and Methods**

134

135 *1. Collection of algae*

136

137 Individuals of *Ellisolandia elongata* (J. Ellis & Solander, 1786) K. R. Hind & G. W.
138 Saunders 2013 were collected in a shaded tidal-pool in the low intertidal level on the rocky
139 beach “Les Amiets” in Cléder on the North coast of Brittany, France (48°41’45 N, 4°7’26
140 W) on the 11th of August 2010 (summer) and on the 21st of February 2013 (winter). The
141 irradiance in photosynthetic active radiations (PAR) at the surface of the pool was measured
142 using a flat quantum sensor (LiCor, Li-190SA, LI-COR Biosciences, USA). Irradiance was
143 *ca.* 30 $\mu\text{mol photons m}^{-2} \text{ s}^{-1}$ in summer and *ca.* 15 $\mu\text{mol photons m}^{-2} \text{ s}^{-1}$ in winter under the
144 shade of rocks overhanging the tidal pool at around midday under sunny conditions.
145 Temperature in the pool at the time of collection of algae was 17 °C in summer and 10 °C in
146 winter. The physico-chemical characteristics of the rock pool are given in Table 1.
147 Individual algae with a frond length of 5-10 cm and free of epiphytic organisms were
148 selected for the experiment. They were carefully separated from their substrate to obtain
149 their encrusting base. After experiments the fresh weight (FW) of the algae was determined
150 and converted to dry weight (DW) using the FW/DW relationship established by Egilsdottir
151 *et al.* (2013).

152

153 *2. Incubation procedure and physiological measurements*

154

155 *E. elongata* were incubated under natural sunlight from 13:00 to 16:00. In summer, algae
156 were transferred directly to incubation chambers and incubated *in situ* in a large
157 neighbouring pool exposed to direct sunlight. In winter the algae were transported to the
158 *Roscoff Marine Station* in a thermostatic tank within 1 h of collection and incubated outside
159 under natural sunlight in large baths filled with natural seawater pumped at 10 m depth in
160 front of the station at ambient temperature. Temperature during incubation varied from 17
161 to 24°C in summer and from 6 to 9°C in winter (Table 2), being representative of changes in

162 the pool during emersion at these seasons (Table 1). Five sets of individual algae (N = 5)
163 were incubated per irradiance level in 220 and 460 ml clear and black glass bottles.
164 Incubations at all irradiance levels were performed simultaneously in each season for 1 to 3
165 hours in summer and 2 to 4 hours in winter, depending on the expected algal metabolic
166 response. Mean DW of the sets of algae was $0.41 \pm \text{SE } 0.03$ g in summer and $0.81 \pm \text{SE}$
167 0.05 g in winter. Black plastic covers were used to assess respiration and calcification rates
168 in the dark. Four neutral density filters were used to assess net production and calcification
169 in algae simultaneously exposed to 5 different mean irradiance (PAR) levels (I1, I2, I3, I4,
170 and I5; Table 2) of 91, 158, 235, 614, and 1118 $\mu\text{mol photons m}^{-2} \text{ s}^{-1}$ in summer and 30, 75,
171 135, 290, and 500 $\mu\text{mol photons m}^{-2} \text{ s}^{-1}$ in winter, with the highest irradiances
172 corresponding to natural full light. Irradiance was measured with a Li-Cor quantum sensor
173 (LiCor Li-190SA, LI-COR Biosciences, USA). Reference incubations without algae were
174 included for full light and dark conditions to assess and correct the metabolism by any
175 biological activity in the seawater. Oxygen fluxes were estimated by measuring the oxygen
176 concentration at the beginning and the end of each incubation with an oxygen probe
177 (Oxymeter HQ40D, Hach Lange, Ltd portable LDOTM, Loveland USA). pH was measured
178 using a pH meter (HQ40D, Hach Lange, Ltd portable LDOTM, Loveland USA) calibrated on
179 the total scale (pH_T, expressed on the total hydrogen ion concentration scale) using Tris/HCl
180 and 2-aminopyridine/HCl buffer solutions with a salinity of 35 (Dickson *et al.*, 2007). For
181 measurements of total alkalinity (A_T), seawater samples were filtered through 0.7 μm
182 Whatman GF/F filters into 100 mL glass bottles, immediately poisoned with mercuric
183 chloride and stored in a cool dark place pending analyses. A_T was determined
184 potentiometrically using an automatic titrator (Titroline alpha, Schott SI Analytics,
185 Germany) calibrated on the National Bureau of Standards scale and calculated using a Gran
186 function applied to pH values ranging from 3.5 to 3.0 (Dickson *et al.*, 2007). Measurements

187 of A_T were corrected by comparison with standard reference material provided by Andrew
188 G. Dickson (Batch 111). Other parameters of the carbonate chemistry were calculated from
189 pH_T , A_T , temperature and salinity (35) using the software *CO2Sys*, *EXCEL Macro version*
190 *2.1* (originally designed by Lewis & Wallace, 1998). Calculations were based on a set of
191 constants K_1 and K_2 from Mehrbach *et al.* (1973), refit by Dickson and Millero (1987)
192 (Table 1; Table 2).

193

194 *3. Data treatment and statistical analysis*

195

196 Primary production was estimated directly from oxygen fluxes. Net production (P_n) and
197 respiration (R) rates ($\mu\text{mol O}_2 \text{ g}^{-1} \text{ DW h}^{-1}$) were calculated as follows:

$$198 \quad P_n \text{ or } R (\text{O}_2) = \left(\frac{\Delta \text{O}_2}{\Delta t \times V} \right) / DW \quad (1)$$

199 where ΔO_2 is the change in dissolved oxygen ($\mu\text{mol l}^{-1}$), Δt is the incubation time (h), DW is
200 the dry weight of the algae and V is the volume of the incubation chamber (L).

201 The variations of dissolved inorganic carbon (DIC) in the incubations integrate both
202 organic carbon metabolism (photosynthesis and respiration) and inorganic carbon
203 metabolism (calcification and dissolution) (Zeebe & Wolf-Gladrow, 2001). Calcification
204 rates were estimated using the alkalinity anomaly technique (Smith & Key, 1975; Chisholm
205 & Gattuso, 1991) which is based on the relationship: $\text{Ca}^{2+} + 2\text{HCO}_3^- \rightarrow \text{CaCO}_3 + \text{CO}_2 +$
206 H_2O , for which total alkalinity decreases by 2 equivalents for each mol of CaCO_3
207 precipitated.

208 Calcification rates (G , $\mu\text{mol CaCO}_3 \text{ g}^{-1} \text{ DW h}^{-1}$) was calculated as follows:

$$209 \quad G = - \left(\frac{\Delta A_T}{2 \times \Delta t \times V} \right) / DW \quad (2)$$

210 where ΔA_T is the change in total alkalinity ($\mu\text{eq l}^{-1}$).

211 P_n and R expressed in terms of carbon ($\mu\text{mol CO}_2 \text{ g}^{-1} \text{ DW h}^{-1}$) were calculated by the
 212 difference between the initial and final DIC concentrations (calculated by using the software
 213 *CO2Sys*) as follows:

$$214 \quad P_n \text{ or } R (\text{CO}_2) = \left(\frac{\Delta\text{DIC} - (0.5 \times \Delta\text{AT})}{\Delta t \times V} \right) / \text{DW} \quad (3)$$

215 where ΔDIC is the change in dissolved inorganic carbon (mmol l^{-1}).

216 The relationship between irradiance (E , $\mu\text{mol photon m}^{-2} \text{ s}^{-1}$) and P_n (or G) was
 217 established using P_n (or G) versus E curves following the function of Platt *et al.* (1980)
 218 modified by the addition of a respiration (R) or dark calcification term (G_d , $\mu\text{mol CaCO}_3 \text{ g}^{-1}$
 219 DW h^{-1}):

$$220 \quad P_n = P_s \times (1 - e^{-\alpha \times \frac{E}{P_s}}) \times e^{-\beta \times \frac{E}{P_s}} + R \quad (4)$$

$$221 \quad G = G_s \times (1 - e^{-\alpha \times \frac{E}{G_s}}) \times e^{-\beta \times \frac{E}{G_s}} + G_d \quad (5)$$

222

223 where P_s and G_s are scaling parameters defined as the maximum rate of photosynthesis or
 224 calcification respectively in the absence of photoinhibition (or decrease in calcification rate
 225 under high irradiance; $\mu\text{mol O}_2$, DIC or $\text{CaCO}_3 \text{ g}^{-1} \text{ h}^{-1}$), α is the initial slope of the
 226 relationship ($\mu\text{mol g}^{-1} \text{ h}^{-1} (\mu\text{mol m}^{-2} \text{ s}^{-1})^{-1}$), β is the photoinhibition coefficient ($\mu\text{mol g}^{-1} \text{ h}^{-1}$
 227 ($\mu\text{mol m}^{-2} \text{ s}^{-1})^{-1}$). In the absence of evident photoinhibition, the β value was set to zero.

228 The deduced maximal rate of gross photosynthesis (or calcification) at light saturation, P_g^{max}
 229 (or G_g^{max}), are calculated as:

$$230 \quad P_g^{max} \text{ (or } G_g^{max}) = P_s \text{ (or } G_s) \times \left(\frac{\alpha}{\alpha + \beta} \right) \times \left(\frac{\beta}{\alpha + \beta} \right)^{\frac{\beta}{\alpha}} \quad (6)$$

231 The maximal rate of net production (or calcification), P_{max}^{net} (or G^{max}) can be estimated as:

$$232 \quad P_n^{max} = P_g^{max} + R \quad (7)$$

233 $G^{max} = G_g^{max} + G_d$ (8)

234 The saturating irradiance (E_k , index of light adaptation) is expressed as:

235 $E_k = \frac{P_g^{max}}{\alpha}$ or $\frac{G_g^{max}}{\alpha}$ (9)

236 The compensation irradiance (E_c) where the rate of photosynthesis equals the rate of
237 respiration is expressed as:

238 $E_c = \frac{R}{\alpha}$ (10)

239 P_n - and G - E curves were fitted to the data using the statistical software STATGRAPHICS
240 *Plus, version 5.0 for Windows* (StatPoint Technologies Inc., 2013), establishing 5 individual
241 curves for O₂ flux, CO₂ flux and calcification per season. The statistical software *R 2.15.0*
242 (The R Foundation for Statistical Computing, 2013) was used when examining differences
243 between seasons. Normality of data was estimated via the Shapiro-Wilk normality test. The
244 Welch two sample t-test was used to identify differences between seasons where data was
245 normally distributed or where log or square-root transformation of data yielded normal
246 distribution. When normality was not achieved, a Wilcoxon Rank Sum test was used.
247 Results are expressed as means ± standard error (SE).

248

249 **Results**

250

251 *Respiration and photosynthesis*

252

253 The mean respiration rate (R) was higher in summer than in winter (O₂ fluxes: $t = -6.63$, $p <$
254 0.01 ; CO₂ fluxes: $t = 3.97$, $p < 0.05$; Table 3). In both seasons primary production measured
255 under various irradiance levels showed a strong relationship with irradiance in terms of both

256 oxygen production ($r^2 = 0.88 - 0.98$ in summer; $r^2 = 0.78 - 0.99$ in winter) and net CO₂
257 uptake ($r^2 = 0.87 - 0.98$ in summer; $r^2 = 0.87 - 0.98$ in winter; Fig. 1). Photoinhibition was
258 recorded for primary production at irradiance levels above about 400 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$
259 in summer and a negligible photoinhibition above about 200 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$ was
260 observed in winter (Table 3; Fig. 1). In summer, maximum gross O₂ production and CO₂
261 uptake at light saturation (P_g^{max}) were about three times those recorded for winter (O₂
262 fluxes: W=25, $p < 0.01$; CO₂ fluxes: W=25, $p < 0.01$; Table 2). In winter the maximum net
263 production (P_n^{max}) was about 40% that of summer (O₂ fluxes: W=25, $p < 0.01$; CO₂ fluxes:
264 W=25, $p < 0.01$; Table 3). The saturating irradiances (E_k) for primary production were about
265 three times higher in summer, relative to winter (O₂ fluxes: $t = 5.42$, $p < 0.05$; CO₂ fluxes: t
266 $= 5.26$, $p < 0.05$; Table 3) and the compensation irradiances (E_c) were higher in summer
267 than in winter (O₂ fluxes: $t = 4.66$, $p < 0.01$; CO₂ fluxes: $t = 3.32$, $p < 0.05$; Table 3).
268 Coefficient α (the initial slope of the curve) did not differ between seasons for both O₂ and
269 CO₂ fluxes. Coefficient β (photoinhibition coefficient) did not differ between seasons for O₂
270 fluxes. No statistical comparison between seasons was conducted for β considering CO₂
271 fluxes since photoinhibition was not indicated by all five $P-E$ curves calculated for each
272 season. The net production rates under 30 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$ and 15 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$
273 (*in situ* midday irradiance in summer and winter respectively) were estimated from the $P-E$
274 curves to $8.4 \pm 3.1 \mu\text{mol O}_2 \text{ g}^{-1} \text{ DW h}^{-1}$ and $-5.6 \pm 4.5 \mu\text{mol C g}^{-1} \text{ DW h}^{-1}$ in summer and 7.4
275 $\pm 1.0 \mu\text{mol O}_2 \text{ g}^{-1} \text{ DW h}^{-1}$ and $-11.3 \pm 1.5 \mu\text{mol C g}^{-1} \text{ DW h}^{-1}$ in winter. Gross production
276 ($P_g = P_n + R$) was thus estimated to $20.4 \pm 3.1 \mu\text{mol O}_2 \text{ g}^{-1} \text{ DW h}^{-1}$ and $-22.5 \pm 1.9 \mu\text{mol C}$
277 $\text{g}^{-1} \text{ DW h}^{-1}$ in summer and $8.4 \pm 1.0 \mu\text{mol O}_2 \text{ g}^{-1} \text{ DW h}^{-1}$ and $-11.8 \pm 1.6 \mu\text{mol C g}^{-1} \text{ DW h}^{-1}$
278 in winter.

279

280 *Calcification*

281
282 Calcification rates in the dark (G_d) differed between seasons ($t = 3.07, p < 0.05$; Table 3). G_d
283 was positive in summer but negative in winter, indicating net dissolution of calcium
284 carbonate. A strong relationship was observed between irradiance and calcification in
285 summer ($r^2 = 0.72 - 1.00$; Fig 2a) and winter ($r^2 = 0.93 - 0.99$; Fig. 2b). At high irradiances
286 in summer the relationship between irradiance and calcification is negligible with a high
287 variance observed for calcification at high irradiances (Fig 2a). Calcification was inhibited
288 at irradiances higher than about $200 \mu\text{mol m}^{-2} \text{s}^{-1}$ in both summer and winter. Maximum net
289 calcification (G_n^{max}) differed between seasons ($t = 3.07, p < 0.05$; Table 3) and was about
290 two times higher in summer relative to winter. E_k was approximately two times higher in
291 summer than in winter ($t = 2.50, p < .005$; Table 3). Since calcification in the dark was
292 positive in summer, an estimate for E_c in summer was unavailable and consequently a
293 statistical comparison between seasons was irrelevant. For calcification, coefficients α and β
294 did not differ between seasons (Table 3). Calcification at midday *in situ* irradiance
295 estimated from the $G-E$ curves was $3.3 \pm 0.3 \mu\text{mol CaCO}_3 \text{ g}^{-1} \text{ DW h}^{-1}$ in summer at $30 \mu\text{mol}$
296 $\text{photons m}^{-2} \text{ s}^{-1}$ and $1.1 \pm \text{SE } 0.1 \mu\text{mol CaCO}_3 \text{ g}^{-1} \text{ DW h}^{-1}$ in winter at $15 \mu\text{mol photons m}^{-2}$
297 s^{-1} .

298

299 **Discussion**

300

301 *Respiration*

302

303 Respiration rates observed here varied seasonally in terms of both O_2 and CO_2 fluxes, being
304 more than 10-fold higher in summer in comparison to winter. The temperature dependence
305 of respiration is well known for seaweeds (Hurd *et al.*, 2014) and has already been reported

306 for several species of coralline algae with a trend of increasing respiration with increasing
307 temperature until temperature reaches beyond the thermal tolerance limit of the alga (see
308 Martin *et al.*, 2013b for a review). For example, respiration rate in the temperate crustose
309 coralline alga, *Lithophyllum yessoense* increased by ten-fold, from 3 to 29 $\mu\text{g O}_2 \text{ cm}^{-2} \text{ h}^{-1}$
310 between 5 and 25°C (Ichiki *et al.*, 2001). Changes in respiration rates observed between
311 winter and summer in *E. elongata* are high in comparison with the threefold increase
312 measured in the temperate unattached coralline alga (maerl) *Lithothamnion corallioides*
313 between 10°C (winter) and 16.5°C (summer) (Martin *et al.*, 2006) or in the temperate
314 crustose coralline alga *Lithophyllum cabiochae* between 13°C (winter) and 22°C (summer)
315 (Martin *et al.*, 2013b). The large differences observed between winter and summer in *E.*
316 *elongata* is mainly due to the high respiration rates recorded in summer. A three- to four-
317 fold lower respiration rate was recorded for this species at a temperature range of 13 to
318 17°C in a Mediterranean subtidal environment (3 $\mu\text{mol O}_2 \text{ g}^{-1} \text{ DW h}^{-1}$; El Haïkali *et al.*,
319 2004) and at a constant temperature of 18°C in a laboratory (4 $\mu\text{mol O}_2 \text{ g}^{-1} \text{ DW h}^{-1}$;
320 Egilsdottir *et al.*, 2013). Temperature was considerably higher in summer in the rock pool at
321 the time of measurements, reaching 23°C at the end of the incubation explaining the high
322 summer respiration rates in *E. elongata* observed in this study. Higher respiration rates in
323 summer in comparison to winter may also reflect high rates of primary production shortly
324 before dark measurements took place as light-enhanced dark respiration has been observed
325 in plants and algae (Raghavendra *et al.*, 1994; Martin *et al.*, 2013a). The large seasonal
326 difference observed for respiration could partly be attributed to the low temperature during
327 winter measurements (reaching below 6°C). In winter, the algae were carefully transported
328 to a different location and incubated in water-baths. Therefore, it is conceivable that
329 different handling and transportation of algae prior to incubations contributed seasonal
330 differences in respiration and other metabolic activity.

331 Although changes between day and night respiration rates occur in subtidal coralline algae
332 in a more stable environment in terms of temperature (Martin *et al.*, 2013a), our results
333 suggest that respiration in coralline algae may vary strongly with temperature changes in
334 tidal pools, even on a very short time scale. Such variations may in turn have consequences
335 for the whole metabolism of algae and in particular for the balance between respiration and
336 photosynthesis.

337

338 *Photosynthesis*

339

340 The photosynthesis of *E. elongata* was significantly influenced by the season. The gross
341 production, P_g , under ambient irradiance was two- to three-fold higher in summer than in
342 winter. These seasonal fluctuations of P_g are related to the changes in both temperature and
343 irradiance which are the fundamental parameters in the controlling algal photosynthesis
344 (Hurd *et al.*, 2014). Photosynthetic characteristics of the P - E curves reflect this with values
345 of P_g^{max} threefold higher in summer than in winter. Similar seasonal influence of
346 temperature and irradiance on photosynthesis has previously been reported in other
347 temperate coralline algal species such as *L. corallioides* or *L. cabiochae*, which exhibited
348 values of P_g^{max} two- to three-fold higher in summer compared to winter (Martin *et al.*, 2006;
349 Martin *et al.*, 2013b). A strong seasonal difference was observed in the photosynthetic
350 parameters E_k and E_c with both being considerably lower in the winter season, reflecting a
351 greater degree of low-light acclimation in winter than in summer. E_k was 2-3 fold higher in
352 summer compared to winter and is in agreement with E_k previously reported for low light
353 acclimated coralline algae, in the range of 7-32 $\mu\text{mol m}^{-2} \text{s}^{-1}$ (Kühl *et al.*, 2001; Payri *et al.*,
354 2001; Roberts *et al.*, 2002; Martin *et al.*, 2013a). E_c was more than ten fold higher in
355 summer than in winter. The value of E_c in summer is consistent with that previously

356 reported for *E. elongata* acclimated to low irradiance ($E_c = 20 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$; Algarra
357 et al., 1991), while the value E_c in winter is similar to that reported for other coralline algae
358 living in dim light condition such as *L. cabiochae* in winter under very low irradiance (< 10
359 $\mu\text{mol photons m}^{-2} \text{ s}^{-1}$; Martin et al. 2013b). Such low E_c reflects the ability of *E. elongata* to
360 make a net gain from photosynthesis at very low light levels.

361 Photoinhibition was observed at high irradiances as has previously been reported for low
362 light acclimated algae (Kain, 1987; Sagert et al., 1997; Kühl et al., 2001; Roberts et al.,
363 2002; Martin et al., 2013b). Recently, photoinhibition was also observed in *E. elongata*
364 under rapid light curves (RLCs) (Korbee et al., 2014) giving further evidence to *E. elongata*
365 being a “shade plant” which Häder et al. (2003) has suggested characterized all geniculate
366 coralline algal species. Coralline algae have been reported to exhibit dynamic
367 photoinhibition strategies which likely enable the algae to tolerate high irradiance levels
368 rather than undergoing photodamage (Burdett et al., 2014). For example, Burdett et al.
369 (2014) reported varying reflectance of the thallus at different times of the day, and varying
370 intracellular concentrations of the antioxidant compound DMSP
371 (dimethylsulphoniopropionate) in crustose coralline algae.

372 Intertidal rock pools typically exhibit large diel fluctuations in temperature and irradiance
373 levels, both of which affect photosynthesis in algae.

374 Temperature directly affects the biochemistry occurring at the algal cellular level. For
375 example, temperature may influence the activity of the enzyme carbonic anhydrase, active
376 transport of CO_2 and HCO_3^- and diffusion rates by which photosynthesis may be affected
377 through the supply of substrate to carbon fixation pathways (Hurd et al., 2014). In this
378 study, temperature exhibited similar variability at all irradiance levels for each season but is
379 likely to contribute to the difference in *P-E* curves between seasons.

380 Irradiance affects seawater in tidal pools through regulation of photosynthesis. As observed
381 in this study, and is well known for photosynthetic organisms, a positive relationship exists
382 between photosynthetic rates and irradiance up to a certain irradiance level, i.e. near the
383 saturating irradiance E_k (Lüning, 1990). In this way, irradiance influences parameters of the
384 seawater chemistry in tidal pools, such as carbonate chemistry and nutrient concentrations
385 (Morris & Taylor, 1983; Hurd *et al.*, 2014). Contrary to respiration, algal photosynthesis
386 removes CO_2 from seawater and subsequently leads to an elevation in pH and a decrease in
387 total inorganic carbon concentrations (Hurd *et al.*, 2014). High pH is known to negatively
388 affect photosynthesis in some macroalgal species (Borowitzka, 1981; Middelboe & Hansen,
389 2007). For example, Björk *et al.* (2004) found certain algal species to be absent in tidal
390 pools where high pH values and low inorganic carbon concentration occurred temporarily.
391 Effect of pH on photosynthesis in *E. elongata* has only been investigated for a pH range of
392 7.7 to 8.1 with no significant effects observed (Egilsdottir *et al.*, 2013; Noisette *et al.*, 2013)
393 but the effects of higher pH levels have not yet been investigated for this species. Like most
394 macroalgae, coralline algae have CO_2 concentrating mechanisms (CCMs) that allow the
395 algae to utilize low concentrations of inorganic carbon (Borowitzka, 1981; Gao *et al.*,
396 1993). These mechanisms are diverse in their exact functionality but for example promote
397 active transport of inorganic carbon across membranes and/or may facilitate the conversion
398 of HCO_3^- to CO_2 for photosynthetic metabolism (Raven *et al.*, 2012). Effective CCMs
399 explain why some algae, including many species of coralline algae, are commonly found in
400 rock pools where pH is often elevated and inorganic carbon concentrations low (Raven,
401 2011).

402

403 *Calcification*

404

405 Although calcification in articulated coralline algal communities may have the potential to
406 significantly influence regional carbon balance (Bensoussan & Gattuso, 2007), very few
407 data are available on calcification rates in articulated coralline algae *in situ* (El Haïkali *et*
408 *al.*, 2004; Bensoussan & Gattuso, 2007; Fisher & Martone, 2014) and none considered the
409 seasonal response of this process.

410 Irradiance had a positive effect on calcification both in winter and in summer. These
411 findings are in agreement with previous studies revealing calcification in coralline algae to
412 be greater in light than in dark (Borowitzka, 1981; Gao *et al.*, 1993).

413 In winter, G_d was below zero indicating a net dissolution of calcium carbonate.
414 Dissolution of calcium carbonate at night in winter has previously been observed within a
415 shallow infralittoral ecosystem dominated by *E. elongata* in the Mediterranean (Bensoussan
416 & Gattuso, 2007). Borowitzka (1981) suggested a positive relationship between Ω_{CaCO_3} of
417 the surrounding seawater and calcification rates in dark in coralline algae. In this study, the
418 seawater within the incubation chambers was supersaturated with respect to aragonite
419 throughout all incubations and, based on calculations by Ries (2011), is unlikely to have
420 become undersaturated with respect to the high-Mg calcite produced by the algae (Mg/Ca
421 mol % \approx 17-19; Egilsdottir *et al.* 2013). However, chemical properties of seawater in micro-
422 environments within the diffusion boundary layer differ from the surrounding seawater as a
423 result of algal metabolism (De Beer & Larkum, 2001; Roberts *et al.*, 2002; Cornwall *et al.*,
424 2014). Thus, it is conceivable that the dissolution observed in dark in winter is indicative of
425 undersaturation of the high-Mg calcite within the diffusion boundary layer as a result of
426 respiration. G_d in summer was significantly different from G_d in winter but was about two-
427 times lower than previously described for the *E. elongata*, $2.5 \mu\text{mol CaCO}_3 \text{ g}^{-1} \text{ DW h}^{-1}$ (El
428 Haïkali *et al.*, 2004). However, El Haïkali *et al.* (2004) attributed this relatively high
429 calcification rate in dark to “residual light metabolism” which has also been suggested for

430 high calcification rates in dark measured in other coralline algae (Chisholm, 2000; Martin *et*
431 *al.*, 2006). Similarly, the seasonal difference in G_d observed in this study may result from
432 different rates of “residual light metabolism”, affected by temperature and irradiance prior
433 to and during incubations and/or reflect different Ω_{CaCO_3} between seasons.

434 The net calcification rate, G_n estimated for ambient irradiance around midday was three-
435 fold higher in summer relative to winter and G^{max} was twofold higher in summer relative to
436 winter. G^{max} in summer was consistent with the mean calcification of *E. elongata* ($9 \mu\text{mol}$
437 $\text{CaCO}_3 \text{ g}^{-1} \text{ DW h}^{-1}$) measured under higher irradiances in the Mediterranean (El Haïkali *et*
438 *al.*, 2004).

439 In agreement with previous studies (see Borowitzka, 1989 and references therein), a
440 relationship between irradiance and calcification was observed both in summer and in
441 winter. Borowitzka and Larkum (1976) suggested that photosynthetic rates and calcification
442 rates were directly related in calcifying algae and an association there between has also been
443 reported by other studies (Pentecost, 1978; Martin *et al.*, 2006; Martin *et al.*, 2013a; Martin
444 *et al.*, 2013b).

445 Above approximately $200 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$ in summer, calcification rates indicate a
446 stronger inhibition than observed for photosynthetic rates. As with photosynthetic rates,
447 inhibition of calcification processes under high irradiances likely reflects a local acclimation
448 of *E. elongata* to a shaded environment. Calcification inhibition under high irradiances (up
449 to $800\text{-}1000 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$) was not observed for ‘light-acclimated’ tropical coralline
450 algae such as *Hydrolithon onkodes*, *Neogoniolithon brassica-florida*, *Hydrolithon reinboldii*
451 and *Neogoniolithon conicum* (Chisholm, 2000) and other species such as the temperate one
452 *L. corallioides* living in shallow waters (1-8 m depth) (Martin *et al.*, 2006). Conversely,
453 inhibition of calcification was reported under high irradiance levels for *L. cabiochae* living
454 in dim-light conditions ($\leq 35 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$; Martin *et al.*, 2013b).

455 Photosynthesis promotes calcification through two major ways: by elevation of Ω_{CaCO_3}
456 and by production of organic tissue or compounds. Photosynthesis (and respiration) regulate
457 pH and Ω_{CaCO_3} within the diffusion boundary layer and at the calcification site via uptake
458 (and production) of CO_2 (Smith & Roth, 1979; Gao *et al.*, 1993; Hurd *et al.*, 2011). As
459 reviewed by Borowitzka (1989), a certain degree of Ω_{CaCO_3} is required to induce nucleation
460 and for the growth of the algal calcium carbonate crystals. Ω_{CaCO_3} may also be relevant for
461 structure, morphology and size of the skeleton. Nucleation and calcification can also be
462 induced by photosynthesis via production of a fibrous organic matrix in the cell walls that is
463 required for forming calcite crystals (Borowitzka, 1989). In coralline algae the organic
464 matrix produced also affects the crystal polymorph, i.e. induces the formation of calcite
465 instead of aragonite (Borowitzka, 1989). The enzyme carbonic anhydrase may be involved
466 in the interplay between photosynthesis and calcification as originally suggested by Digby
467 (1977) and later by Hofmann *et al.* (2012) and Martin *et al.* (2013b), by catalyzing the
468 conversion of CO_2 into HCO_3^- and then CO_3^{2-} . Indeed, the presence of carbonic anhydrase
469 has been reported in *E. elongata* (Mercado *et al.*, 1997). However, carbonic anhydrase is
470 also used in photosynthesis to convert HCO_3^- to CO_2 which could then lower Ω_{CaCO_3} .
471 Therefore, it is conceivable that calcification may be reduced at the same time as
472 photosynthetic rates are high.

473 Results presented here for *E. elongata* suggest a direct relationship between
474 photosynthetic production and calcification. However this relationship was found to be
475 limited to low irradiance levels since calcification and photosynthetic processes seem to
476 evolve differently under high irradiances. In summer, calcification rates varied greatly at
477 high irradiances. This could reflect different acclimations to irradiance in the selected algae,
478 resulting from algae inhabiting distinct localities in the rock pool in terms of irradiance
479 regimes (Häder *et al.*, 1997). Anatomical differences between algae may also explain

480 variable responses. For example, Borowitzka (1979) demonstrated that the uppermost
481 segment calcified at a much higher rate than the segments below and that they also had a
482 much higher carbon fixation rate into organic material than the lower segments. In addition,
483 Borowitzka and Vesk (1979) observed more chloroplasts in young segments compared to
484 older segments. Furthermore, it is conceivable that the equation used to convert wet weight
485 of algae into dry weight may have contributed to the variability observed in the data.

486 Calcification rates in coralline algae are affected by temperature and irradiance, which
487 both exhibit strong diurnal fluctuations in intertidal rock pools. Temperature is positively
488 correlated with Ω_{CaCO_3} , potentially explaining how elevated temperature may promote
489 calcification. This is perhaps especially relevant for dark calcification in the absence of
490 photosynthetic metabolism when the relative importance of abiotic environmental factors
491 increases. Temperature also directly affects diffusion rates and metabolic processes that may
492 play a role in calcification (Hurd *et al.*, 2014). For example, temperature affects carbonic
493 anhydrase activity and active transport of HCO_3^- and CO_2 , both of which are also important
494 for photosynthesis (Hurd *et al.*, 2014). As discussed above, much of the observed response
495 in calcification may be directly associated with photosynthesis and respiration conducted by
496 the algae themselves. In addition, the biological metabolism of all the species residing in a
497 tidal pool can influence the chemical properties of the seawater and subsequently affect
498 physiological processes in coralline algae inhabiting the pool.

499 Since calcifying algae in intertidal rock pools experience large fluctuations in the physico-
500 chemical environment, it has been suggested that they have physiological advantages as
501 ocean acidification and warming advances (Egilsdottir *et al.*, 2013). However, ocean
502 acidification and global warming may affect organisms synergistically in the future. For
503 example, the species investigated in this study inhabited a low-lit tidal pool and respiration
504 increased significantly over the temperature range observed in the study. Thus, the

505 minimum pH and Ω_{CaCO_3} in rock pools might be lowered in the future as a result of
506 increased temperatures leading to elevated respiration rates, but this effect may be
507 somewhat offset due to the positive correlation between temperature and Ω_{CaCO_3} . Coralline
508 algae inhabiting low light environments may be especially at risk by a rise in temperature
509 since higher irradiance levels could be required to accomplish a sustainable net
510 photosynthetic production in the algae.

511

512 *Conclusions*

513

514 Data presented in this study provides valuable information on respiration, photosynthesis
515 and calcification in a coralline alga from an intertidal rock pool and their seasonal
516 differences. This information is relevant in developing an understanding of how climate
517 change will affect coralline algae in the future. This study also highlights the importance of
518 considering different light regimes when studying the response of coralline algae to
519 environmental change. Certainly, the relationship between photosynthesis and calcification
520 metabolism is dynamic and care must be taken when assuming a direct relationship there
521 between.

522

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524

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543 **(Figure captions)**

544

545 **Fig. 1.** Relationship between irradiance and *Ellisolandia elongata* net primary production

546 (in O₂ and CO₂ fluxes) in (a) summer and (b) winter

547

548 **Fig. 2.** Relationship between irradiance and calcification rates in *Ellisolandia elongata* in

549 (a) summer and (b) winter

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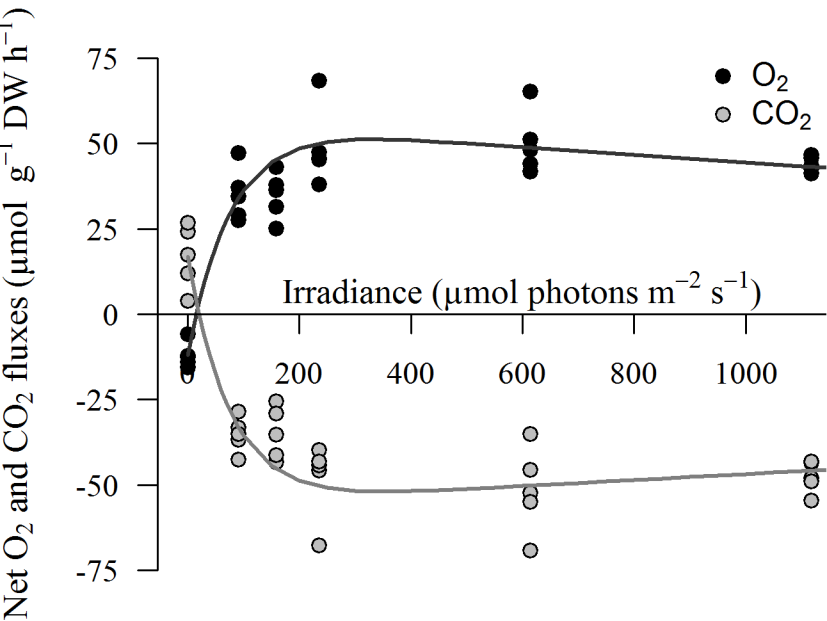


Figure 1a

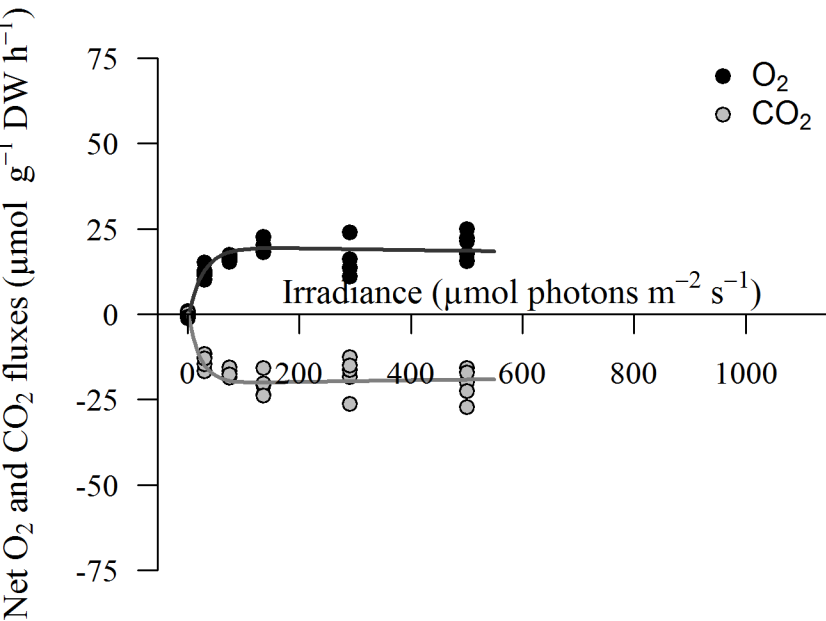


Figure 1b

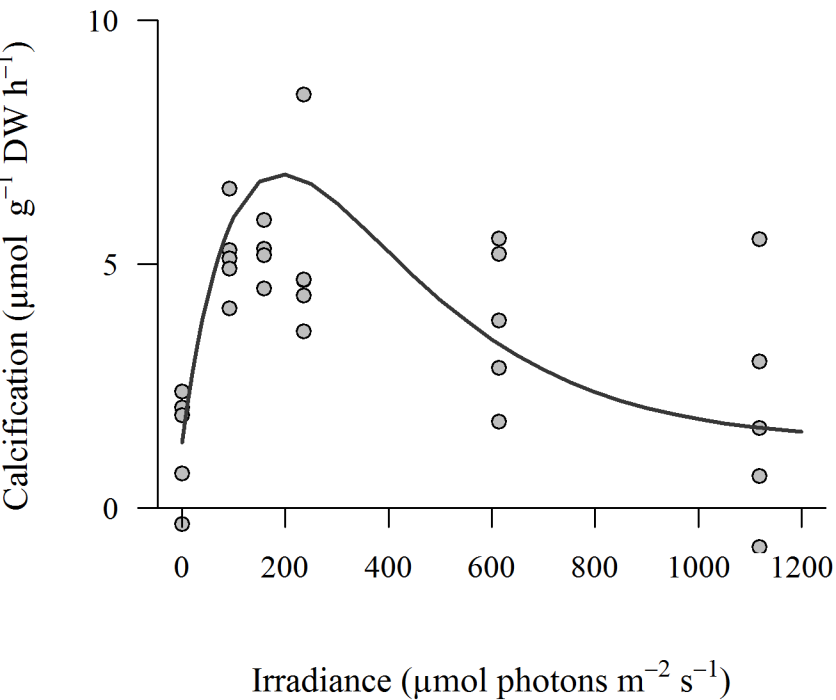


Figure 2a

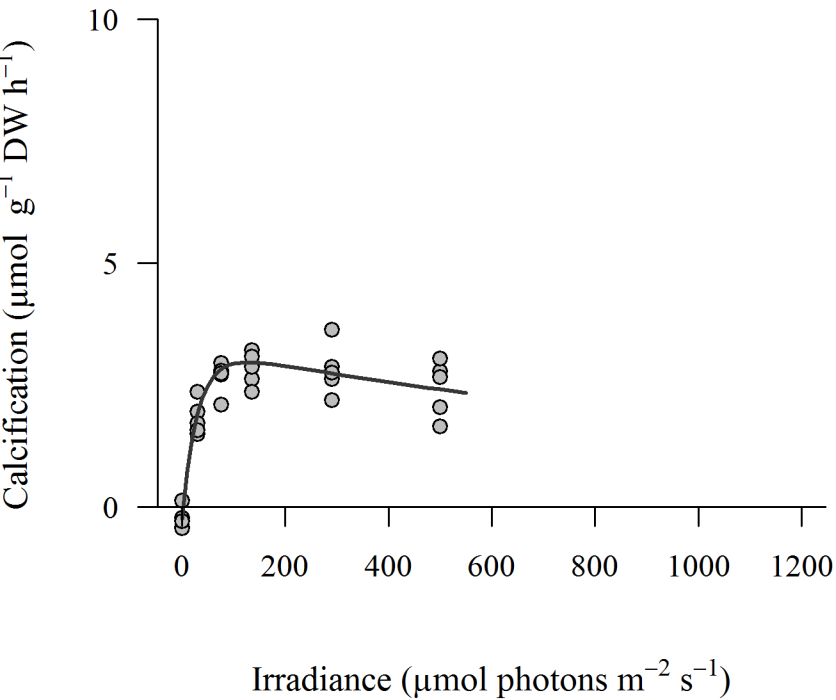


Figure 2b