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C. M. Armas-Herrera, M.-F Dignac, C. Rumpel, C. D. Arbelo, A Chabbi

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1	Management effects on composition and dynamics of cutin and suberin in topsoil under
2	agricultural use

. CHABBI ^{a,b,d}
. CHABBI ^{a,b}

- 4 ^aUnité Mixte de Recherche Institut d'écologie et des sciences de l'environnement de Paris,
- 5 CNRS-UPMC-UPEC-INRA-IRD, Campus ParisAgroTech, Thiverval-Grignon, France,
- 6 ^bUnité Mixte de Recherche Écologie fonctionnelle et écotoxicologie des agroécosystèmes,
- 7 INRA-AgroParisTech, Campus ParisAgroTech, Thiverval-Grignon, France, ^cDepartamento
- 8 *de Biología Animal, Edafología y Geología, Universidad de La Laguna, Spain, and ^dUnité de*
- 9 Recherche Prairies et Plantes Fourragères, INRA Poitou-Charentes, Lusignan, France
- 10
- 11 Correspondence: C.M. Armas-Herrera: E-mail: cmarmas@unizar.es
- 12 Running title: Cutin and suberin in agricultural soil
- 13 Keywords: shoots, roots, biomarkers, C_3/C_4 chronosequence, ¹³C, bare soil
- 14 Highlights:
- Cutin and suberin monomers should meet certain criteria to be shoot and root markers
 in SOM studies
- Shoot and root markers should be characterised for each plant species to study their
 dynamic in soil
- The concentrations of cutin and suberin in agricultural soil depended on the amount of
 organic inputs
- Roots contributed more to SOM accumulation, but had a shorter residence time in soil
 than shoots
- 23

25

26 Summary

27 We identified and quantified specific biomarkers of shoots and roots (cutin and suberin-28 derived compounds, respectively) of three grassland species (Dactylis glomerata L., Festuca 29 arundinacea Schreb. and Lolium perenne L.) in soil under different land use (grass, crop and 30 bare soil) of the SOERE-ACBB experimental site in Lusignan (France). We also investigated the fate of these markers in soil after conversion from grassland (C₃ plants) to Zea mays L. 31 (maize) (C_4 plant) from natural ¹³C isotope abundances. Our results indicated that 9-hydroxy 32 33 hexadecanedioic acid and 8(9)(10),16-dihydroxy hexadecanoic acid may be used as 34 biomarkers for aboveground biomass, whereas 1,22-docosandioic acid, 22-hydroxy 35 docosanoic acid and 24-hydroxy tetracosanoic acid might be the best belowground 36 biomarkers for the plants investigated under the experimental conditions studied. The 37 presence, concentration and shoot-root allocation pattern of these markers were different 38 from those described for other species, which demonstrates the importance of verifying 39 biomarker specificity for each species. Concentrations of cutin and suberin were largest in 40 soil under maize and smallest under bare soil; this corresponded to the biomass added to the 41 two soils. Suberin decreased by 40-64 % and cutin by 24-40 % during a 6-year bare fallow, 42 which indicates that root markers were more sensitive than shoot markers to degradation. Changes in ¹³C isotopic signatures of specific biomarkers after 6 years of maize showed a 43 44 faster turnover of root than shoot biomarkers, in spite of the much smaller root inputs from 45 maize than from grasses. The sequestration of suberin in soil was more rapid, but less durable than that of cutin. 46

48 Introduction

49 Soil organic matter (SOM) has beneficial effects on soil physical structure, water-retention 50 capacity and plant nutrient availability. It consists of a heterogeneous mixture of substances 51 that have a wide range of decomposability. Organic matter enters the soil from litter fall, root 52 turnover and root exudates. Once incorporated into the mineral soil matrix, a major part of 53 the organic matter is metabolized and mineralized by microorganisms. Another portion 54 (approximately 30 %) remains in soil for longer because of transformation and stabilization 55 processes (von Lützow et al., 2006). The amount of mineralized and stabilized organic matter 56 may differ according to the environment and land use (Marschner et al., 2008). Precise 57 quantification and identification of the origin of labile and more stable SOM pools is 58 necessary to improve the understanding of carbon cycling and the response of SOM to 59 changing environmental conditions (Marschner et al., 2008), in particular climate or land-use 60 change. The conversion from native forest or pasture to arable crops has caused losses that 61 range from -40 to -60 % of original C stocks in soil, whereas an increase in soil C stocks has 62 been reported after a change from crop to pasture and forest (+ 20–60 %) (Guo & Gifford, 63 2002). The preservation of soil C stocks has focused on management practices, such as tillage 64 (conventional or reduced) or no tillage, and soil cover (crop residues, catch crops, intercrops) 65 or none. Many studies have concentrated on bulk C to investigate the effects of management 66 on the formation of SOM (Dungait et al., 2013). It is only recently that the relative 67 contributions of shoots and roots to the SOM pool (Rasse et al., 2005), the root and shoot 68 turnover (Mendez-Millan et al., 2011) or the pattern of biodegradation of different plant 69 tissues (Clemente et al., 2013) have been addressed at the molecular level.

Cutins and suberins are aliphatic plant biopolyesters that occur in vascular plants and could
be among the most recalcitrant plant macromolecules in soil. As a result, they might play an
important role in enriching the slower cycling pool of SOM (Riederer *et al.*, 1993; Nierop *et*

73 al., 2003; Mendez-Millan et al., 2011). Cutins are embedded within intracuticular waxes and 74 covered with epicuticular waxes to form the plant cuticle. The cuticle covers all aerial parts 75 (leaves, fruits, flowers, seeds, and so on) of vascular plants and protects them from desiccation. Cutins are composed mainly of derivatives of saturated C₁₆ (palmitic) acid and 76 C₁₈ acids, such as di- and tri-hydroxy and epoxy fatty acids, interlinked through ester bonds. 77 78 Although cutin from most plants contains both the C_{16} and C_{18} groups, their individual 79 composition varies according to the plant species, specific plant tissue, stage of development 80 and environmental conditions (Kolattukudy, 2001; Kögel-Knabner, 2002; Otto & Simpson, 81 2006). Suberins are wall components of cork cells from which all protective and wound-82 healing layers of bark, woody stems and roots are composed. They are also in the endodermis 83 and in the bundle sheath of grasses. Suberins are composed of an aliphatic polyester and a polyphenolic domain that are spatially segregated (Kolattukudy, 2001; Bernards, 2002; 84 85 Kögel-Knabner, 2002). The most characteristic compounds of the aliphatic domain are a 86 mixture of α, ω -dioic acids, ω -hydroxy acids, very long chain fatty acids, mid-chain-oxidized 87 fatty acids and esterified hydroxycinnamic acids; there is emerging evidence that glycerol is a 88 major component of the aliphatic domain. In turn, the polyphenolic domain has been related 89 to lignin, and more recently has been considered to be composed of a large amount of 90 hydroxycinnamic acids and their derivatives, and monolignols (Bernards, 2002).

The dynamics of cutins and suberins may be studied by analysing their stable carbon isotope composition after extraction from soil under C_3 - C_4 succession (Mendez-Millan *et al.* 2010a). Isotopic analyses have shown that shoot biomarkers in soil under continuous cropping are degraded rapidly, whereas root biomarkers are incorporated into SOM, which suggests their selective preservation (Mendez-Millan *et al.*, 2011). Root biomarkers also contributed considerably to SOM from soil under pasture after forest conversion (Hamer *et al.*, 2012). This suggests that the introduction of ley grasslands into cropping systems could increase SOM content by the accumulation of root-derived C (Rasse *et al.*, 2005). In this research, we addressed the effect of grassland management on the composition and turnover of root and shoot biomarkers in soil. The research took advantage of a long-term agricultural experiment investigating SOM dynamics after the introduction of grassland into the cropping cycle.

102 The aim of this research was to evaluate the dynamics of specific root and shoot biomarkers 103 after land-use changes from grass to an arable land. For this we (i) determined the 104 composition of cutins and suberins in above- and below-ground biomass of the three 105 dominant grassland species, Dactylis glomerata L., Festuca arundinacea Schreb. and Lolium 106 perenne L., (ii) investigated the composition of cutin and suberin in soil under different land uses (continuous and temporary grassland, arable and bare soil) and (iii) used natural ¹³C 107 108 isotope abundances to follow the fate of specific markers in the cutins and suberins in soil 109 after conversion from grassland (C₃ plants) to arable land (C₄ plants).

110

111 Material and methods

112 Experimental site

113 The fieldwork was conducted at Lusignan (46°25'12.91N"; 0°07"29.35"E) in western France. 114 The site is part of a long-term field experiment initiated in 2005 (SOERE ACBB, Systems of 115 in Environmental Observation and Experimentation Research-Agro-ecosystem, 116 Biogeochemical Cycles and Biodiversity, http://www.soere-acbb.com/), which was designed 117 to increase our understanding of the effects of temporary grassland management on the 118 environmental outputs of mixing arable cropping and grasslands systems. This site had been 119 under agricultural use for at least 200 years and before being cultivated, it supported an oak 120 forest.

121 The experimental treatments (4000 m^2 plot size) were established in 2005 in a randomized 122 block design with four blocks. They consisted of continuous grassland, continuous cropping 123 and temporary grassland including crop rotations of maize (Zea mays (L.)), wheat (Triticum 124 aestivum (L.)) and barley (Hordeum vulgare (L.)) (Figure 1). Continuous cropping treatments 125 were fertilized with N at rates adjusted to achieve the potential yield for each crop in this 126 region. The two treatments with temporary grasslands consisted of rotations of maize, wheat and barley alternating with three or six years of grassland with a large application of N 127 128 adjusted to achieve near maximum forage production. In addition, a continuous grassland 129 composed of a mixture of Festuca arundinacea (Cv Soni), Lolium perenne (Cv Milca) and 130 Dactylis glomerata (Cv Ludac) established with treatments that included large applications of 131 N and no applications of N on a bare soil. Management of the crop rotation followed 132 agricultural practices to achieve a yield close to the potential determined for the region by 133 soil and climate. The rate and timing of N fertilizer application were adjusted every year with 134 PC-AZOTE software (http://www.i-cone.fr/front/viewnode.aspx?typnode=4&idnode=60).

The soil at the site is a Plinthic Cambisol (IUSS Working Group WRB 2014) developed under a temperate climate from loess material over a Mesozoic tropical palaeosoil. It has five soil horizons: a plough layer that overlies two red-brown upper horizons characterized by a loamy texture and two lower red clayey horizons rich in kaolinite, iron nodules and iron oxides. More detailed information about the soil and site characteristics is in Chabbi *et al.* (2009) and Moni *et al.* (2010).

141 Sampling design and preparation of samples

Soil sampling. Samples were taken in November 2011 after the conversion of 3- and 6-year grasslands into arable land (Figure 1). We sampled soil under (i) 6 years of continuous grassland, (ii) converted ley grassland of 6 years, (iii and iv) 6 years of continuous cropping maize, wheat and barley rotation or continuous maize, (v) converted ley grassland of 3 years after 3 years of crop rotation and (vi) 6-years of continuous bare soil. Three cores of soil were sampled at 0–30-cm depth from each plot and were mixed to obtain bulked samples. Samples were air-dried, sieved at 2 mm and ground to pass a 100- μ m sieve. The contents of organic carbon (OC) and nitrogen (N), and carbon isotopic signature (δ^{13} C) of the soil samples are given in Table 1.

151 *Plant sampling.* Samples of the three grass species were collected in triplicate in November 152 2011. Shoots and roots were separated, dried at 60 °C and ground at 100 μ m. Roots were 153 taken in the field by manual separation and subsequently washed with deionized water. The 154 contents of organic carbon (OC) and nitrogen (N), and carbon isotopic signature (δ^{13} C) of the 155 shoots and roots are listed in Table 1. The plant inputs to soil under different land use are 156 summarized in Table 2.

157 Analytical procedures

Extraction of free lipids. To remove free lipids before the depolymerization of cutin and suberin biopolyesters, plant samples of 1 g or soil samples of 3 g were extracted with dichloromethane:methanol (DCM:MeOH) (1:2, volume:volume) at a 1:10 sample:extractant ratio. The suspensions were mixed with a vortex mixer for 30 s, agitated overhead for 2 hours and subsequently centrifuged at 2200 g for 10 minutes. This extraction was repeated. Thereafter, the samples were rinsed with DCM:MeOH (1:2, volume:volume) by means of mixing and centrifugation. The lipid-free samples were air-dried and kept until analysis.

165 Saponification and derivatisation. Saponification was used to release biomarkers of cutins 166 and suberins, as suggested by Mendez-Millan et al. (2010b). A lipid-free plant sample of 100 mg or of 1 g of soil was refluxed for 18 hours in a solution of water:methanol (MeOH) (1:9, 167 168 volume:volume) containing 6% of potassium hydroxide (KOH) (Cardoso & Eglinton, 1975). 169 These conditions lead to the depolymerization of cutin and suberin. Hydroxylated fatty acids 170 are released, and epoxy functions only are transformed into methoxy functions. Then, the 171 solution was filtered (GF/A Whatman glass microfibre filters, 1.6 µm) with a Millipore 172 vacuum filtration system (Darmstadt, Germany) and the residue was washed with

173 water:MeOH (1:9, volume:volume). The pH of the filtrate was adjusted to 2 with 6 N HCl 174 after the addition of 150 ml of distilled water to isolate the acidic products (Naafs & Van Bergen, 2002). The acidified solution was extracted three times with 50 ml of 175 176 dichloromethane (DCM; CH₂Cl₂). The volume of the extracts was reduced with a rotary evaporator and dried completely in a nitrogen atmosphere. All dried extracts were redissolved 177 178 in 2 ml of DCM containing nonadecanoic acid $(C_{19:0})$ as an internal standard and then kept in the freezer until analysis. Prior to analysis, samples were derivatized by silvlation to 179 180 transform hydroxyl and carboxylic acid functions into their trimethylsilyl (TMS) ether and 181 ester derivatives (TMS ether and TMS ester). An aliquot of each sample (0.2–1 ml) was dried 182 in a nitrogen atmosphere, redissolved in 40 µl of pyridine and 10 µl of BSTFA (N,O-183 bis(trimethylsilyl)-trifluoroacetamide) that contained 1% TMCS (trimethylchlorosilane) and 184 heated at 70 °C for 1 hour.

185 Identification and quantification of cutin and suberin monomers

186 The silvlated monomers of cutin and suberin were identified according to their fragmentation 187 pattern after analysis with an Agilent HP6890 gas chromatograph (Santa Clara, CA, USA) 188 coupled to an Agilent HP5973 mass spectrometer (Santa Clara, CA, USA) (GC/MS) and 189 compared with published mass spectra and with a mass spectral library (G1035B Wiley Mass 190 Spectral Database) (Mendez-Millan et al., 2010a). One µl was injected in splitless mode at a 191 temperature of 300 °C. The GC oven temperature was set at 100 °C for 2 minutes, then from 100 to 150 °C at 10 °C minute⁻¹, from 150 to 200 °C at 5 °C minute⁻¹ and finally at a rate of 2 192 °C minute⁻¹ from 200 to 350°C and kept for 5 minutes at 350 °C. Quantification of the 193 194 monomers was done with a flame ionisation detector (FID) using the internal standard $C_{19:0}$ and an external calibration with 16-hydroxyhexadecanoic acid ($\omega OH C_{16:0}$). The 195 chromatographic conditions were the same as for the GC/MS analysis. We obtained a 196

197 response factor for the external standard ω OH-C_{16:0} relative to the internal standard C_{19:0} 198 close to 1.

199 Compound specific isotopic analysis

We measured the δ^{13} C values (expressed in % relative to Vienna PeeDee Belemnite) of 200 201 individual compounds in the plant and soil samples from the 6-year grassland, 6-year maize 202 and 6-year bare soil treatments. We did the analysis with an isotopic ratio mass spectrometer (Micromass-GVI Optima, Manchester, UK) coupled with a combustion interface to a GC 203 204 (GC-C-IRMS), and used the same chromatographic conditions as for the identification and 205 quantification of the monomers. The carbon atoms of BSTFA were assumed to have the same 206 isotopic ratio as that reported by Dignac et al. (2005) in a previous study of lignin-derived phenols. We corrected the δ^{13} C of the C introduced by the derivatisation process (C atoms 207 from the trimethylsilvl groups) by measuring the δ^{13} C off-line with an Elemental Analyser 208 (NA 1500, Carlo Erba) coupled to the IRMS. The value obtained was used to correct the ¹³C 209 210 concentrations of the cutin and suberin monomers in the samples, according to a mass 211 balance equation, following the procedure of Dignac et al. (2005).

212 Calculations

Suberin: cutin ratios. We selected monomers that were specific to cutin (C), to suberin (S) or specific to both molecules (SC), and calculated the respective sums of their contents ($\sum C$, $\sum S$, $\sum SC$). We calculated the following suberin: cutin ratio adapted from Otto & Simpson (2006):

217 Suberin: cutin ratio =
$$(\sum S + \sum SC/2) / (\sum C + \sum SC/2)$$
 (1)

and the sum of suberin and cutin

219 $\sum SC = \sum S + \sum C + \sum SC$ (2)

221 *Carbon isotopic signature* ($\delta^{13}C$)

We computed the δ^{13} C of each compound class (δ_{class}) for both shoots and roots following Mendez-Millan *et al.* (2011). This δ^{13} C of individual compounds with similar chemical structure were weighted by their concentrations with the following equation:

225
$$\delta_{\text{class}} = \sum_{i=1}^{N} (\delta_{i\text{comp}} \times C_{i\text{comp}}) / \sum_{i=1}^{N} C_{i\text{comp}} , \qquad (3)$$

where $C_{i\text{comp}}$ is the concentration of the *i*th compound of the chemical class, $\delta_{i\text{comp}}$ is the isotopic ratio of this compound and *N* is the number of individual compounds within each class.

The proportion, *F*, of maize-derived C_4 -C in soil after 6-years of conversion was calculated for each compound class according to the following equation from Balesdent & Mariotti (1996):

232
$$F = (\delta^{13}C_{\text{maize-soil}} - \delta^{13}C_{\text{grassland-soil}})/(\delta^{13}C_{\text{maize}} - \delta^{13}C_{\text{grasses}}).$$
(4)

233 The $\delta^{13}C_{\text{grasses}}$ is the average value of $\delta^{13}C$ of the three grasses studied.

234

235 **Results and discussion**

236 Identification of characteristic compounds for shoots and roots

The monomers obtained after saponification of the cutin and suberin biopolyesters were grouped into six chemical classes: *n*-carboxylic acids, *n*-alcohols, α -hydroxy carboxylic acids, ω -hydroxy carboxylic acids, mid-chain hydroxy acids and α,ω -alkanedioicacids. Because the *n*-carboxylic acids, *n*-alcohols and α -hydroxy carboxylic acids (Appendix 1) can appear in plant polymers other than cutins and suberins (Otto *et al.*, 2005; Amelung *et al.*, 2008), they were not considered appropriate plant biomarkers. 243 We detected seven ω -hydroxy carboxylic acids in the range C₁₆ to C₂₆ (Figures 2, 3). This 244 chemical class was most abundant in grass roots, with a contribution of 46-49% to the total 245 released monomers. Their abundance and contribution (13-17%) in the plant shoots were 246 much smaller. The ω -hydroxy carboxylic acids showed differences in their shoot-root 247 allocation patterns. For example, the $\omega C_{16:0}$ was almost twice as abundant in roots than in 248 shoots; the $\omega C_{18:0}$ was detected only in the shoots at very small concentrations, and the ω -249 hydroxy carboxylic acids from $C_{20:0}$ to $C_{26:0}$, dominated by the $\omega C_{22:0}$ and $\omega C_{24:0}$, were about 250 three to ten times more abundant in roots than in shoots. The ω -hydroxy carboxylic acids 251 with more than 20 atoms of C are generally considered to be more frequent in the suberins 252 than in the cutins (Kolattukudy, 2001; Kögel-Knabner, 2002; Otto & Simpson, 2006), which 253 is consistent with our results. Some authors also found a larger abundance of these 254 compounds in the roots than in the shoots of other species (Spielvogel et al., 2014; Bull et al., 255 2000). However, irregular or non-significant plant allocation patterns have been reported 256 (Hamer et al., 2012; Andreetta et al., 2013) or an even larger content of ω -hydroxy 257 carboxylic acids in shoots than in roots (Mendez-Millan et al., 2011, for wheat and maize, 258 Figure 2).

259 We detected three α, ω -alkanedioic acids (C_{18:1}, C_{20:0} and C_{22:0}) in grass roots, but none in 260 shoots (Figures 2,4). Their concentrations and contribution to the total of monomers (5–7 %) 261 were small. In maize, α, ω -alkanedioic acids were also exclusively present in roots and their concentrations were smaller than in the grasses (Mendez-Millan *et al.*, 2010a). The $C_{18:1}$ 262 263 diacid was by far the most abundant α,ω -alkanedioic acid, which is the case for several plants (Otto & Simpson, 2006; Mendez-Millan et al., 2010a, 2011; Hamer et al., 2012; Spielvogel et 264 al., 2014), whereas the $C_{22:0}$ diacid was only slightly more abundant than $C_{20:0}$. For this 265 266 chemical class, there is general agreement about the exclusive presence of the α,ω -

alkanedioic acids in the suberins of plants (Kolattukudy, 2001; Kögel-Knabner, 2002; Otto &
Simpson, 2006; Amelung *et al.*, 2008; Mendez-Millan *et al.*, 2010a).

269 We identified six mid-chain hydroxy acids, which were about five to seven times more abundant in the shoots than in the roots (relative abundance of 32-39 % and 9-11 %, 270 271 respectively) (Figures 2,5). The 9,10-epoxy, 18-OH $C_{18:0}$ was predominant and was one of the major monomers released in the shoots of the three grasses. These compounds were 272 273 detected in the shoots of maize only (Mendez-Millan et al., 2010a), and they were less 274 abundant than in the grasses. The 9,10-epoxy, 18-OH C_{18:0}, the 9(10),16-diOH C_{16:0} and the 9,10,18-triOH C_{18:0} are considered the most common monomers in cutins (Kolattukudy, 275 276 2001). In our study, only the 9,10,18-triOH $C_{18:0}$ showed no differences in its concentration 277 between shoots and roots in *Dactylis glomerata*. This is not an exception in the literature, other authors have also found an irregular pattern of distribution in certain mid-chain hydroxy 278 279 acids, including the 9,10,18-triOH C_{18:0}, which has been found to be more abundant in the 280 roots than in the shoots of several plant species (Hamer et al., 2012; Spielvogel et al., 2014). 281 Moreover, our results showed that *Dactylis glomerata* and *Lolium perenne* released more 282 aliphatic monomers and had larger suberin/cutin ratios than Festuca arundinacea (Table 3). 283 The amount of monomers in these grasses contrasted with those found for different species 284 (Mendez-Millan et al., 2010a; Hamer et al., 2012; Andreetta et al., 2013; Spielvogel et al., 285 2014). These findings emphasize the need to identify specific cutin and suberin biomarkers 286 for each plant species in order to use them as indicators of relative fluctuations in above- and 287 below-ground biomass in soil.

288 Differences in the ${}^{13}C$ isotopic content of aliphatic monomers in plants

There were small differences in the ¹³C content between the classes of aliphatic monomers in the three grasses (Table 4, Figure 6): ¹³C content of the α,ω -alkanedioic acids in *Dactylis glomerata* and *Festuca arundinacea* was slightly smaller than for the mid-chain hydroxy 292 acids and ω -hydroxy carboxylic acids, whereas *Lolium perenne* showed no clear differences. Most of the monomers in Lolium perenne were more ¹³C depleted than in the other two 293 grasses. We found the biopolyesters of the grasses investigated to be ¹³C-depleted compared 294 to the bulk plant tissues (Figure 6). Similar results were obtained by Mendez-Millan et al. 295 (2011) and Hamer *et al.* (2012) for several plant species. This ¹³C depletion is common for 296 297 lignins and lipids because of isotopic fractionation along transport pathways (Hobbie & Werner, 2004). The same explanation might be applied here to the monomers of cutin and 298 suberin; they have similar biosynthetic pathways to extractable lipids. The ¹³C of the bulk 299 300 samples of the grasses agree with those for some C₃ plants (Wiesenberg & Schwark, 2006; 301 Mendez-Millan et al., 2011), such as ryegrass, oats, barley or wheat (from -28.1 to -32.3 ‰ 302 on average). The isotopic signature of maize (C₄ plant) reported by the same authors was much larger (12.5 ‰ on average) than for C_3 plants. To the best of our knowledge, the ¹³C 303 304 isotopic signature of different cutin and suberin biopolyesters has been determined for a few 305 plant species only: wheat and maize (Mendez-Millan et al., 2010a, 2011), the grass Setaria 306 sphacelata (Schumach.) and the bracken fern Pteridium arachnoideum (Kaulf.) (Hamer et *al.*, 2012) and in the grasses investigated here. The differences in 13 C content of the aliphatic 307 308 monomers for different C₃ plants also demonstrates the need to study each plant species 309 independently.

310 Criteria for selecting cutin and suberin biomarkers in soil

The fact that these aliphatic monomers are considered to be specific for cutin or suberin is not enough to consider them as adequate above- or below-ground biomass markers in soil. The $\omega C_{22:0}$ and $\omega C_{24:0}$ were the predominant ω -hydroxy carboxylic acids (Figure 3) in the soils studied here. Their largest contents were in the 6-year maize crop and the smallest in the 6year bare soil, with no clear pattern in the other soils. The same was true for the $\omega C_{16:0}$, $\omega C_{18:0}$ and $\omega C_{20:0}$. The $\omega C_{18:1}$ was not detected in the 6-year maize crop. Four alkanedioic 317 acids (C_{16:0}, C_{18:1}, C_{20:0}, C_{22:0}) were identified in soil under grass and crops (Figure 4). In the 318 6-year bare soil C_{22:0} only was detected. The largest values for the C_{16:0} diacid were in the 6-319 year maize crop. This monomer was not detected in the grasses but it is present in the roots of 320 maize and wheat (Mendez-Millan et al., 2010a, 2011), which are the likely sources in our study. The C_{18:1} diacid, which was predominant among the alkanedioic acids of the grasses, 321 322 had a much smaller relative abundance in soil. The total concentration of mid-chain hydroxy 323 acids was much smaller in the 6-year bare soil than in the other soils. The 8(9)(10),16-diOH 324 $C_{16:0}$ diacid and the 9-OH $C_{16:0}$ diacid were the most abundant compounds (Figure 5).

325 Moreover, in our study the relative contribution of the different compounds to the total of 326 released monomers often varied considerably from plant to soil. These changes were almost 327 exclusively the result of a drastic decrease in the relative abundance of a few monomers that 328 contained either epoxy functions or double bonds: $\omega C_{18:1}$, $C_{18:1}$ diacid, 11,18-diOH $C_{18:1}$, 329 9(10),18-diOH C_{18:1} and 9,10-epoxy, 18-OH C_{18:0}. In contrast, the 8(9)(10),16-diOH C_{16:0} and 330 the 9-OH C_{16:0} diacid showed a larger relative distribution in soil (4.9–10.5 % and 2.7–4.2 %, 331 respectively) than in the plant tissues (5.0–6.9 % and 0.7–1.1 %). The decrease in the double 332 bond functions from plant to soil might be explained by their preferential degradation with 333 respect to the saturated building blocks (Nierop, 2001; Nierop et al., 2003) because the epoxy 334 groups are considered as first intermediates in the oxidation of double bonds (Watkinson & 335 Morgan, 1990). There is more debate about the preservation of cutin and suberin monomers 336 in soil. Some authors have reported a similar rate of decomposition irrespective of their 337 chemical composition (e.g. Riederer et al. (1993) in an in vitro decomposition experiment of 338 cutin in Fagus sylvatica (L.) leaves; Nierop et al. (2003), for suberin monomers in an oak 339 forest), whereas other authors have proposed different mechanisms for preservation of these 340 monomers in soil. Mendez-Millan et al. (2011) suggested chemical recalcitrance for some 341 cutin monomers and soil physical protection for suberin in soils cultivated with wheat and

maize, whereas Hamer *et al.* (2012) found that organo-mineral interactions were responsible for the long-term survival of several cutin and suberin monomers in soil under different land use, and chemical recalcitrance for the x,16-diOH C_{16:0}.

345 We suggest that for cutin or suberin to be considered as a useful marker to study SOM 346 dynamic, each compound should meet the following criteria: (i) significant differences in 347 their amounts between shoots and roots and (ii) adequate concentrations in the soil. Our results suggested that the 9-OH $C_{16:0}$ diacid and 8(9)(10),16-diOH $C_{16:0}$ were the most 348 349 appropriate aboveground biomarkers for grassland species. On the other hand, the C_{22:0} 350 diacid, the $\omega C_{22:0}$, the $\omega C_{24:0}$ and to a lesser extent the $\omega C_{20:0}$ and $\omega C_{26:0}$, were the most useful 351 belowground biomass biomarkers for Dactylis glomerata, Festuca arundinacea and Lolium 352 perenne.

353 Dynamics of cutin and suberin biomarkers in soil under different land use

354 The concentrations of cutin and suberin in soil (Table 3) were related to the amount of organic inputs into the soil (Table 2); their largest concentrations were in soil cultivated with 355 356 maize and their smallest were in bare soil. The largest amounts of cutin and suberin in the soil 357 under 6-year maize cultivation is easily explained by the larger maize biomass inputs (mostly 358 from shoots) than for wheat, barley and the grasses. The large suberin/cutin ratio in the 6-year 359 maize crop probably results from the larger amounts of ω -hydroxy carboxylic acids in the 360 shoots than in the roots of maize (Mendez-Millan et al., 2010a). In general, we found no 361 marked differences in the concentrations of cutin and suberin monomers between soil under 362 continuous and temporary grasslands. Grass species are characterized by a dense root system 363 in the topsoil, therefore most of their residues that enter the soil come from roots (Table 2). Because the grasses are harvested, the addition of leaf litter to soil is limited to around 20 % 364 365 (Sanaullah et al., 2010). At the same experimental site, Rumpel et al. (2009) and Rumpel & 366 Chabbi (2010) found a rapid decrease in SOM and a change in its composition three months

367 after conversion from grassland to cropland. However, this disturbance of SOM composition 368 was transitory and one year only after conversion the chemical characteristics of SOM 369 returned to their initial status. Our results corroborate this finding for cutin and suberin 370 because we found no difference in their concentrations between the continuous and 371 temporary crop rotation and grasslands. The most prominent change in the concentration of 372 monomers was in the 6-year bare soil where they decreased from 40 to 64 % for suberin and 373 from 24 to 40 % for cutin. There was also a smaller suberin/cutin ratio (1.01) than for the 374 other soil uses (1.12-1.54). Suberin is considered to be more resistant to degradation than 375 cutin because of the larger concentration of aromatic compounds from its polyphenolic 376 domain (Riederer et al., 1993; Nierop et al., 2003; Otto & Simpson, 2006). The residence 377 time of organic compounds in soil, however, depends more on their susceptibility to 378 physicochemical stabilization through incorporation into soil aggregates or chemical 379 interactions with the mineral phase or both than on their chemical composition (Marschner et 380 al., 2008). Cutin and suberin compounds sorb strongly to clay mineral surfaces (Feng et al., 381 2005; Simpson et al., 2006). In a long-term bare-fallow experiment (Closeaux experiment, 382 from 1928, in Versailles, France), the clay-SOM association in macroaggregates was the 383 most important sink for stabilized organic C (Balabane & Plante, 2004). In our study, cutins 384 seem either to be more protected from biodegradation than suberins in the 6-year bare soil, or 385 to have more effective mechanisms of stabilization than those for suberins. Thus, root 386 markers seem to be more sensitive than shoot markers to microbial degradation.

387 Turnover rate of shoot and root-derived organic matter in soils after land-use change

388 The ¹³C concentrations of the monomers of cutins and suberins in soil was measured only in 389 the 6-year continuous grassland, 6-year maize crop and 6-year bare soil. In the rotation with 390 crops, the ¹³C signal of biomass input changes every year because both C_3 and C_4 plants are 391 present, which prevents the fate of these plant biomarkers in soil from being followed. 392 The average isotopic signature of the hydroxy carboxylic acids was largest in the 6-year 393 maize crop and tended to be slightly smaller in the 6-year bare soil and the 6-year continuous grassland soil (Table 4, Figure 6). The ¹³C content varied according to the type of monomer. 394 For the 6-year maize crop, the largest enrichment in ¹³C compared to the grassland soil was in 395 ω -C_{22:0}, followed by ω -C_{24:0}. The ¹³C was more depleted in ω -C_{20:0}, ω -C_{24:0} and ω -C_{18:1} for 396 the 6-year bare soil than for the cropland soil. The largest average content of ¹³C in the α, ω -397 alkanedioic acids was for the 6-year bare soil (Table 4, Figure 6). The 6-year maize had 398 399 intermediate values and the smallest average content was for the 6-year continuous grassland soil. The $C_{16:0}$ and $C_{22:0}$ diacids were ¹³C-enriched more in the 6-year maize crop than in the 400 401 grassland soil. The C_{22:0} diacid was the only α, ω -alkanedioic acid detected in the 6-year bare soil and it had the largest ${}^{13}C$ content. Lastly, the $C_{20:0}$ diacid had a slightly larger isotopic 402 signature in the 6-year continuous grassland than for the 6-year maize crop. The average ${}^{13}C$ 403 404 content of the mid-chain hydroxy acids was similar under the 6-year maize crop and 6-year 405 continuous grassland soil, with a slightly smaller value in the 6-year bare soil (Table 4, 406 Figure 6). The isotopic signature of each monomer was somewhat irregular: the 8(9)(10), 16diOH C_{16:0}, 9-OH C_{16:0} diacid and 9,10-epoxy, 18-OH C_{18:0} were ¹³C-enriched only slightly 407 408 more in the soil of the 6-year maize crop than for the 6-year continuous grassland, whereas 9,10,18-triOH C_{18:0} was strongly ¹³C-enriched (Table 4, Figure 6). These results should be 409 410 treated with caution because we do not know the effect of the methylated hydroxy group on the epoxy functions. In all cases, the smallest ¹³C contents were for the 6-year bare soil. 411

We calculated the proportion of cutin and suberin markers with Equation (4) that are maizederived 6 years after maize was introduced on land that had been cultivated previously with C_3 plants. This enabled us to estimate the rate of turnover of shoot- and root-derived OM in soil. The rate was large, about 29 %, for the α, ω -alkanedioic acids, whereas it was much less for the ω -hydroxy carboxylic acids (8 %) and the mid-chain hydroxy acids (5 %). The small 417 rate of incorporation of both ω -hydroxy carboxylic acids and mid-chain hydroxy acids into 418 SOM suggests that they are degraded faster than α,ω -alkanedioic acids, in spite of the 419 increase in shoot inputs after conversion from grassland to maize crop (Table 2). Hamer et al. 420 (2012) found that about 75 % of newly introduced C_4 - ω -hydroxy carboxylic acids was 421 decomposed 15 years after the pasture was abandoned. This relatively small rate of 422 biodegradability was explained by the binding of these compounds to soil minerals (Hamer et 423 al., 2012). The mid-chain hydroxy acids, however, decompose rapidly in soil (Mendez-424 Millan et al., 2011). The small variation in the isotopic signature after 6 years of maize 425 suggest that at least a fraction of these compounds was stabilized previously in the soil that 426 might have come from the ancient forest vegetation (Mendez-Millan et al., 2011; Hamer et 427 al., 2012). In summary, our finding suggests that root C contributed more to the SOM 428 accumulation. However, root C showed a shorter residence time in soil compared to shoot C. 429 Thus, the sequestration of suberin in soils was less durable than that of cutin.

430

431 Conclusions

There were marked differences in monomer composition, abundance and patterns of shootroot allocation of cutin and suberin in the plant species analysed (*Dactylis glomerata*, *Festuca arundinacea* and *Lolium perenne*) than for other plant species. These results emphasize the need to identify specific cutin and suberin biomarkers for each plant species to study the incorporation of their biomass into SOM.

To be shoot and root markers cutin and suberin monomers should have strong differences in their shoot and root concentrations and measurable concentrations in soil. According to these criteria, 9-hydroxy hexadecanedioic and 8(9)(10),16-dihydroxy hexadecanoic acids can be used as aboveground biomarkers, and 1,22-docosandioic, 22-hydroxy docosanoic and 24hydroxy tetracosanoic acids for belowground biomarkers for the plants investigated. The concentrations of cutin and suberin in soil were related to the amount of organic inputs. We found no differences in the amounts of cutin and suberin in soil under continuous and temporary grassland, which might indicate that the disturbance caused by conversion from grassland to cropland was transitory only. For bare soil, suberin decreased by 40–64 % and cutin by 24–40 % during a 6-year fallow, which indicated that root markers were more sensitive to degradation than shoot markers.

The changes detected in the ¹³C isotopic signatures of specific biomarkers after 6 years of maize cropping showed that incorporation into SOM was greater for roots than for shoot markers, in spite of the much smaller root inputs from maize than from grasses. Roots contributed more to SOM accumulation, but had a shorter residence time in soil than shoots. The sequestration of suberin in soil was more rapid, but less durable than that of cutin.

To specify the mechanisms and processes that lead to the turnover of cutin and suberin monomers with land-use changes, future studies should focus on the organo-mineral associations to protect these macromolecules in soil.

456

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464

465 **Figure captions**

466 **Figure 1.** Sequence of soil cover for the various treatments.

- 467 Figure 2. Concentrations (C / μ g g⁻¹) of plant-specific monomers in shoots and roots of
- 468 Dactylis glomerata L., Festuca arundinacea Schreb., Lolium perenne L., and Zea mays L.
- 469 The standard errors were calculated from three replicates
- 470 Figure 3. Concentrations (C / μ g g⁻¹) of ω -hydroxy carboxylic acids in shoots and roots of
- 471 Dactylis glomerata L., Festuca arundinacea Schreb., Lolium perenne L., and Zea mays L.,
- 472 and in soil at the SOERE-ACBB site in Lusignan (France)
- 473 The standard errors were calculated from three replicates
- 474 **Figure 4.** Concentrations (C /µg g⁻¹) of α , ω -alkanedioic acids in shoots and roots of *Dactylis*
- 475 glomerata L., Festuca arundinacea Schreb., Lolium perenne L., and Zea mays L., and in soil
- 476 at the SOERE-ACBB site in Lusignan (France)
- 477 The standard errors were calculated from three replicates
- 478 Figure 5. Concentrations (C / μ g g⁻¹) of mid-chain hydroxy acids in shoots and roots of
- 479 Dactylis glomerata L., Festuca arundinacea Schreb., Lolium perenne L., and Zea mays L.,
- 480 and in soil at the SOERE-ACBB site in Lusignan (France)
- 481 The standard errors were calculated from three replicates
- 482 Figure 6. Bulk and molecular isotopic signatures in the grasses, maize and soil. The values
- 483 for the compound classes are calculated as the average of the signatures of the monomers
- 484 weighted by their concentration.
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486 **References**

- Amelung, W., Brodowski, S., Sandhage-Hofmann, A. & Bol, R. 2008. Combining
 biomarker with stable isotope analyses for assessing the transformation and turnover of
 soil organic matter. *Advances in Agronomy*, **100**, 155–251.
- 490 Andreetta, A., Dignac, M.-F. & Carnicelli, S. 2013. Biological and physico-chemical
- 491 processes influence cutin and suberin biomarker distribution in two Mediterranean forest
- 492 soil profiles. *Biogeochemistry*, **112**, 41–58.
- Balabane, M. & Plante, A.F. 2004. Aggregation and carbon storage in silty soil using
 physical fractionation techniques. *European Journal of Soil Science*, 55, 415–427.
- 495 Balesdent, J. & Mariotti, A. 1996. Measurement of soil organic matter turnover using ¹³C
- 496 natural abundance. In: *Mass Spectrometry of Soils* (eds T.W. Boutton & S. Yamasaki),

497 pp. 83–111. Marcel Dekker, NY.

- 498 Bernards, M.A. 2002. Demystifying suberin. *Canadian Journal of Botany*, **80**, 227–240.
- Bull I.D, Nott, C.J., van Bergen, P.F., Poulton, P.R. & Evershed, R.P. 2000. Organic
 geochemical studies of soils from the Rothamsted classical experiments—VI. The
 occurrence and source of organic acids in an experimental grassland soil. *Soil Biology & Biochemistry*, **32**, 1367–1376.
- 503 Cardoso, J.N. & Eglinton, G. 1975. The use of cutin acids in the recognition of higher
- 504 plant contribution to recent sediments. In: Advances in Organic Geochemistry (eds R.
- 505 Campos & J. Goni), pp. 273–287. Enadimsa, Madrid.
- 506 Chabbi, A., Kögel-Knabner, I. & Rumpel, C. 2009. Stabilised carbon in subsoil horizons
- is located in spatially distinct parts of the soil profile. *Soil Biology & Biochemistry*, **41**,
 256–261.

- 509 Clemente, J.S., Simpson, M.J., Simpson, A.J., Yanni, S.F. & Whalen, J.K. 2013.
- 510 Comparison of soil organic matter composition after incubation with maize leaves, roots,
- 511 and stems. *Geoderma*, **192**, 86–96.
- 512 Dignac, M.-F., Bahri, H., Rumpel, C., Rasse, D.P., Bardoux, G. & Balesdent, J. 2005.
- 513 Carbon-13 natural abundance as a tool to study the dynamics of lignin monomers in soil:
- 514 an appraisal at the Closeaux experimental field (France). *Geoderma*, **128**, 3–17.
- 515 Dungait, J.A.J., Ghee, C., Rowan, J.S., McKenzie, B.M., Hawes, C., Dixon, E.R., et al.
- 516 2013. Microbial responses to the erosional redistribution of soil organic carbon in arable
- 517 fields. Soil Biology & Biochemistry, **60**, 195-201.
- 518 Feng, X., Simpson, A.J., & Simpson, M.J. 2005. Chemical and mineralogical controls on
- 519 humic acid sorption to clay mineral surfaces. *Organic Geochemistry*, **36**, 1553–1566.
- 520 Guo, L.B. & Gifford, R.M. 2002. Soil carbon stocks and land use change: a meta 521 analysis. *Global Change Biology*, **8**, 345–360.
- Hamer, U., Rumpel. C. & Dignac, M.-F. 2012. Cutin and suberin biomarkers as tracers
 for the turnover of shoot and root derived organic matter along a chronosequence of
- 524 Ecuadorian pasture soils. *European Journal of Soil Science*, **63**, 808–819.
- Hobbie, A.E. & Werner, R.A. 2004. Intramolecular, compound-specific, and bulk carbon
 isotope patterns in C3 and C4 plants: a review and synthesis. *New Phytologist*, 161, 371–
- 527 385.
- 528 IUSS Working Group WRB 2014. World Reference Base for Soil Resources 2014.
- 529 International soil classification system for naming soils and creating legends for soil
- 530 maps. World Soil Resources Reports No 106. FAO, Rome.
- 531 Kögel-Knabner, I. 2002. The macromolecular organic composition of plant and microbial
- residues as inputs to soil organic matter. *Soil Biology & Biochemistry*, **34**, 139–162.

533	Kolattukudy, P.E. 2001. Polyesters in higher plants. In: Advances in Biochemical
534	Engineering/Biotechnology, Volume 71 (ed. Th. Scheper), pp. 1-49. Springer-Verlag,
535	Berlin Heidelberg.
536	Marschner, B., Brodowski, S., Dreves, A., Gleixner, G., Gude, A., Grootes, P.M. et al.

- 537 2008. How relevant is recalcitrance for the stabilization of organic matter in soils?
- 538 *Journal of Plant Nutrition and Soil Science*, **171**, 91–110.
- Mendez-Millan, M., Dignac, M.-F., Rumpel, C. & Derenne, S. 2010a. Can cutin and
 suberin biomarkers be used to trace shoot and root-derived organic matter? A molecular
 and isotopic approach. *Biogeochemistry*, **106**, 26–38.
- Mendez-Millan, M., Dignac, M.-F., Rumpel, C. & Derenne, S. 2010b. Quantitative and qualitative analysis of cutin in maize and a maize-cropped soil: Comparison of CuO oxidation, transmethylation and saponification methods. *Organic Geochemistry*, **41**, 187– 191.
- Mendez-Millan, M., Dignac, M.-F., Rumpel, C., Rasse, D.P. & Derenne, S. 2011.
 Molecular dynamics of shoot vs. root biomarkers in an agricultural soil estimated by
 natural abundance ¹³C labelling. *Soil Biology & Biochemistry*, 42, 169–177.
- 549 Moni, C., Rumpel, C., Virto, I., Chabbi, A. & Chenu, C. 2010. Relative importance of 550 sorption versus aggregation for organic matter storage in subsoil horizons of two 551 contrasting soils. *European Journal of Soil Science*, **61**, 958–969.
- 552 Naafs, D.F.W. & van Bergen, P.F. 2002. Effect of pH adjustments after base hydrolysis:
- implications for understanding organic matter in soils. *Geoderma*, **106**, 191–217.
- Nierop, K.G.J. 2001. Temporal and vertical organic matter differentiation along a vegetation succession as revealed by pyrolysis and thermally assisted hydrolysis and methylation. *Journal of Analytical and Applied Pyrolysis*, **61**, 111-132.

- Nierop, K.G.J., Naafs, D.F.W. & Verstraten, J.M. 2003. Occurrence and distribution of
 ester-bound lipids in Dutch coastal dune soils along a pH gradient. *Organic Geochemistry*, 34, 719–729.
- 560 Otto, A. & Simpson, M.J. 2006. Sources and composition of hydrolysable aliphatic lipids 561 and phenols in soils from western Canada. *Organic Geochemistry*, **37**, 385–407.
- 562 Otto, A., Shunthirasingham, C. & Simpson, M.J. 2005. A comparison of plant and
- 563 microbial biomarkers in grassland soils from the Prairie Ecozone of Canada. *Organic*564 *Geochemistry*, **36**, 425-448.
- Rasse, D.P., Rumpel, C. & Dignac, M.-F. 2005. Is soil carbon mostly root carbon?
 Mechanisms for a specific stabilisation. *Plant and Soil*, 269, 341–56.
- 567 Riederer, M., Matzke, K., Ziegler, F. & Kögel-Knabner, I. 1993. Occurrence, distribution
- and fate of the lipid plant biopolymers cutin and suberin in temperate forest soils. *Organic Geochemistry*, 20, 1063–1076.
- Rumpel, C. & Chabbi, A. 2010. Response of bulk chemical composition, lignin and
 carbohydrate signature to grassland conversion in a ley-arable cropping system. *Nutrient Cycling in Agroecosystems*, 88, 173–182.
- 573 Rumpel, C., Chabbi, A., Nunan, N. & Dignac, M.-F. 2009. Impact of land use change on
- 574 the molecular composition of soil organic matter. *Journal of Analytical and Applied* 575 *Pyrolysis*, **85**, 431–434.
- 576 Sanaullah, M., Chabbi, A., Lemaire, G., Charrier, X. & Rumpel, C. 2010. How does plant
- 577 leaf senescence of grassland species influence decomposition kinetics and litter 578 compounds dynamics? *Nutrient Cycling in Agroecosystems*, **88**, 159–171.
- 579 Simpson, A.J., Simpson, M.J., Kingery, W.L., Lefebvre, B.A., Moser, A., Williams, A.J.,
- 580 et al. 2006. The application of 1H high-resolution magic-angle spinning NMR for the

- study of clay-organic associations in natural and synthetic complexes. *Langmuir*, 22,
 4498-4503.
- Spielvogel, S., Prietzel, J., Leide, J., Riedel, M., Zemke, J. & Kögel-Knabner, I. 2014.
 Distribution of cutin and suberin biomarkers under forest trees with different root
 systems. *Plant and Soil*, **381**, 95–110.
- 586 Von Lützow, M., Kögel-Knabner, I., Ekschmitt, K., Matzner, E., Guggenberger, G.,
- 587 Marschner, B. *et al.* 2006. Stabilization of organic matter in temperate soils: Mechanisms
- 588 and their relevance under different soil conditions a review. *European Journal of Soil*
- 589 *Science*, **57**, 426–445.
- Watkinson, R.J., & Morgan, P. 1990. Physiology of aliphatic hydrocarbon-degrading
 microorganisms. *Biodegradation* 1, 79–92.
- 592 Wiesenberg, G.L.B. & Schwark, L. 2006. Carboxylic acid distribution patterns of
 593 temperate C₃ and C₄ crops. *Organic Geochemistry*, **37**, 1973–1982.

		1	1		12
		$C/mg g^{-1}$	N /mg g^{-1}	C/N	δ^{13} C /‰
Plants					
Dactylis glomerata	Shoot	384 (29)	22.8 (1.4)	16.9 (0.8)	-26.7 (0.1)
	Root	350 (34)	11.5 (1.5)	27.8 (0.8)	-26.2 (0.3)
Festuca arundinacea	Shoot	389 (17)	16.3 (0.3)	23.7 (1.8)	-26.1 (0.2)
	Root	375 (8)	10.6 (0.7)	35.5 (2.4)	-26.3 (0.4)
Lolium perenne	Shoot	385 (27)	27.9 (6.4)	15.5 (0.5)	-28.5 (0.4)
	Root	332 (18)	13.2 (0.8)	25.2 (0.2)	-28.4 (0.1)
Zea mays*	Bulk plant	425	na	na	-12.5
Soil (0–30 cm)					
6-year continuous gras	ssland	11.2 (0.6)	1.20 (0.0)	9.35 (0.2)	-25.7 (0.2)
6-year ley grassland		11.3 (1.3)	1.24 (0.1)	9.10 (0.5)	-25.8 (0.2)
6-year crop rotation		10.1 (1.3)	1.11 (0.1)	9.06 (0.2)	-24.9 (0.2)
6-year maize crop		11.0 (0.4)	1.23 (0.1)	9.59 (0.1)	-25.4 (0.3)
3-year ley grassland		11.0 (0.7)	1.08 (0.1)	9.47 (0.4)	-25.2 (0.0)
6-year bare soil		8.29 (0.5)	0.96 (0.1)	8.59 (0.2)	-25.1 (0.2)

Table 1 Carbon (C) and nitrogen (N) contents and carbon isotopic signature (δ^{13} C) of grassland plants and soil at the SOERE-ACBB site in Lusignan (France)

*Data from Mendez-Millan et al. (2011)

n = 3; SD is in parentheses, na = not analysed

Table 2 Mean annual sources of shoot and root into soil under different management for the study period (2005–2011) at the SOERE-ACBB site in Lusignan (France) (SD in parentheses, n = 4)

	6-year continuous grassland	6-year ley grassland	6-year crop rotation	6-year maize crop	3-year ley grassland	6-year bare soil
Shoot /t ha ⁻¹						
Grasses	1.93*	1.94*	0	0	2.29 (0.66)*	0
Maize	0	0	14.8 (2.74)	14.8 (2.7)	8.02 (0.35)	0
Wheat	0	0	1.62 (0.12)	0	1.86 (0.21)	0
Barley	0	0	1.19 (0.08)	0	1.56 (0.11)	0
Shoot weighted average	1.93	1.94	5.87 (0.98)	14.8 (2.7)	1.14 (0.11)	0
Root /t ha ⁻¹						
Grasses	7.63 (1.47)	8.46 (2.88)	0	0	9.21 (1.04)	0
Maize	0	0	1.09 (0.70)	1.09 (0.70)	4.32 (2.54)	0
Wheat	0	0	1.06 (0.32)	0	0.47 (0.05)	0
Barley	0	0	0.99 (0.23)	0	0.66 (0.17)	0
Root weighted average	7.63 (1.47)	8.46 (2.88)	1.04 (0.41)	1.09 (0.70)	4.72 (1.43)	0

*These values correspond to 20% of total shoot biomass because of removal by mowing the grasses. Source: Sanaullah et al. (2010), at the same experimental site

	$^{a}\Sigma C$ Cutin	^b ∑S	°∑CS	^e Suberin +	^d Suberin /
		Suberin	Cutin+suberin	cutin ∑SC	cutin
Plants	C /µg g ⁻¹	C /µg g ⁻¹	$C/\mu g g^{-1}$	$C/\mu g g^{-1}$	
Dactylis glomerata					
Shoots	9574 (446)	911 (42)	3877 (159)	14 361 (551)	0.25 (0.01)
Roots	1239 (150)	6394 279)	3950 (375)	11 583 (765)	2.61 (0.13)
Festuca arundinacea					
Shoots	6940 (536)	721 (101)	3950 (375)	11 620 (662)	0.30 (0.03)
Roots	1301 (73)	5411 (763)	5282 (394)	11 994 (862)	2.05 (0.22)
Lolium perenne					
Shoots	9341 (632)	891 (113)	3796 (374)	14 028 (930)	0.25 (0.01)
Roots	1186 (173)	7305 (763)	3835 (90)	12 325 (624)	2.99 (0.38)
Soil					
6-year continuous grassland	2258 (279)	3209 (276)	1128 (200)	6595 (259)	1.27 (0.00)
6-year ley grassland	2152 (257)	2451 (294)	1093 (39)	5696 (77)	1.12 (0.21)
6-year crop rotation	1920 (72)	2463 (198)	925 (61)	5308 (209)	1.23 (0.05)
6-year maize crop	2409 (214)	4061 (168)	1338 (110)	7808 (260)	1.54 (0.15)
3-year ley grassland	2340 (77)	3108 (65)	1111 (163)	6560 (22)	1.27 (0.00)
6-year bare soil	1457 (129)	1461 (71)	707 (79)	3625 (137)	1.01 (0.11)

Table 3 Cutin and suberin concentrations and calculated ratios for plants and soil

^aCutin $\sum C (\mu g g^{-1} C) = 8(9)(10), 16$ -diOH C_{16:0} + 9-OH C_{16:0} diacid + 11, 18-diOH C_{18:1} + 9, 10-epoxy, 18-OH C_{18:0}

^bSuberin $\sum S (\mu g g^{-1} C) = \omega C_{20:0} + \omega C_{22:0} + \omega C_{24:0} + \omega C_{26:0} + C_{16:0} \text{ diacid} + C_{18:1} \text{ diacid} + C_{20:0} \text{ diacid} + C_{22:0} \text{ diacid}$

^cSuberin or cutin \sum SC (µg g⁻¹ C) = ω C_{16:0} + ω C_{18:0} + ω C_{18:1} + 9(10),18-diOH C_{18:1} + 9,10,18-triOH C_{18:0}

^dSuberin/cutin ratio = $(\sum S + \sum SC/2) / (\sum C + \sum SC/2)$

^eSum of suberin and cutin \sum SC (µg g⁻¹ C)= \sum S + \sum C + \sum SC

	Plants $/\% \delta^{13}$ C				Soil /‱ δ ¹³ C					
	Dactylis glomerata		Festuca an	rundinacea	Lolium perenne		Zea mays ^a	6-year	6-year	6-year
	Shoot	Root	Shoot	Root	Shoot	Root	Bulk plant	grassland	maize crop	bare soil
ω -Hydroxy carboxylic acids										
16-Hydroxy hexadecanoic acid ($\omega C_{16:0}$)	na	na	na	na	na	na	-17.1 (0.9)	na	na	na
18-Hydroxy octadecanoic acid ($\omega C_{18:0}$)	na	na	na	na	na	na	na	na	na	na
18-Hydroxy octadecenoic acid ($\omega C_{18:1}$)	-35.5	-35.2	-34.3	-34.7	-36.0	-36.2	-17.1 (0.3)	-38.3 (0.5)	na	-39.3 (0.7)
20-Hydroxy eicosanoic acid ($\omega C_{20:0}$)		-36.2		-34.8		-37.8	na	-35.5 (0.6)	-35.3 (1.1)	-37.0 (0.1)
22-Hydroxy docosanoic acid ($\omega C_{22:0}$)	-36.9	-36.8	-32.6	-38.2	-37.8	-37.9	-19.8 (0.7)	-36.8 (0.8)	-35.4 (0.2)	-36.2 (2.7)
24-Hydroxy tetracosanoic acid ($\omega C_{24:0}$)	-36.5	-29.5	-30.3	-36.6	-37.0	-37.8	-21.2 (0.4)	-35.2 (0.2)	-34.6 (0.4)	-36.6 (0.1)
26-Hydroxy hexacosanoic acid ($\omega C_{26:0}$)		-28.0		-33.0		-34.6	-19.1 (0.8)	na	na	na
Weighted average for ω -hydroxy carboxylic acids	-35.7	-34.2	-33.7	-35.4	-36.3	-37.1	na	-36.4 (0.5)	-35.1 (0.3)	-36.8 (1.5)
α, ω -Alkanedioic acids										
1,16-Hexadecadioic acid (C _{16:0} diacid)	na	na	na		na	na	-15.1 (0.5)	-32.4 (0.1)	-29.0 (0.8)	na
1,18-Octadecendioic acid (C18:1 diacid)	na	-35.4	na	-36.3	na	-36.2	-17.0 (2.2)	-36.8 (4.6)	na	na
1,20-Neodecandioic acid (C _{20:0} diacid)	na	-36.6	na	-33.1	na	-31.3	na	-30.6 (2.2)	-31.2 (1.7)	na
1,22-Docosandioic acid (C22:0 diacid)	na	-34.9	na	-37.5	na	-38.7	na	-37.4 (0.4)	-34.8 (0.9)	-30.5 (1.1)
Weighted average for α, ω -alkanedioic acids		-35.4		-36.2		-36.1	na	-33.6 (1.2)	-31.3 (1.0)	-30.5 (1.1)
Mid-chain hydroxy acids										
8(9)(10),16-Dihydroxy hexadecanoic acids (8(9)(10),16-diOH C _{16:0})	-34.6	-32.4	-34.7	-30.8	-32.4	-35.8	-16.4 (0.5)	-34.2 (0.3)	-33.8 (0.1)	-35.0 (0.9)
9-Hydroxy hexadecanedioic acid (9-OH $C_{16:0}$ diacid)	na	na	-34.7	na	-33.9	na	-16.1 (0.5)	-34.5 (0.3)	-34.1 (0.3)	-36.6 (0.4)
11,18-Dihydroxyoctadecenoic acid (11,18-diOH $C_{18:1}$) ^b	-35.9	-37.4	-35.8	-37.5	-37.3	-39.6	-17.1 (0.5)	-37.5 (0.4)	-38.4 (0.1)	-36.9 (0.2)
9(10),18-Dihydroxyoctadecenoic acid (9(10),18- diOH $C_{18:1}$) ^b	-33.6	-29.6	-34.7	-36.7	-36.1	-38.0	-17.7 (0.5)	-31.6 (0.3)	-32.9 (0.5)	na
9,10-Epoxy, 18-hydroxyoctadecanoic acid (9,10- epoxy, 18-OH $C_{18:0}$) ^c	-34.2	-30.4	-34.9	-34.7	-33.6	-37.4	-19.1 (0.5)	-33.8 (0.9)	-33.5 (0.6)	na
9,10,18-Trihydroxyoctadecanoic acid (9,10,18- triOH C _{18:0})	-34.6	-36.8	-39.4	-37.3	-41.8	na	-17.9 (0.5)	-35.3 (0.5)	-31.9 (0.3)	na
Weighted average for mid-chain hydroxy acids	-34.4	-33.1	-35.4	-34.8	-34.3	-37.4	na	-34.4(0.4)	-34.1 (0.2)	-35.7 (0.6)

Table 4 Carbon isotopic signature (δ^{13} C) values of the aliphatic monomers of cutins and suberins present in grassland, crop plants and soil at the SOERE-ACBB site in Lusignan (France)

^aData from Mendez-Millan *et al.* (2011); ^bPosition of the double bond not confirmed; ^cIsomeric mixture of 9-methoxy, 10,18-dihydroxyoctadecanoic acid and 9-hydroxy,10-methoxy, 18-hydroxyoctadecanoic acids

(SD in parentheses, n = 2, na = not analysed)

2005	2006	2007	2008	2009	2010	2011

Continuous cropping	Μ	W	В	Μ	W	В	Μ
Ley grassland	Μ	WBGrass 3 years			Μ		
Ley grassland	Grass 6 years					Μ	
Bare soil							
Continuous grassland	L, F,	D					

M= maize, W= wheat, B= barley

L= Lolium, F= Festuca, D= Dactylis











 \Box bulk sample $\triangle \omega$ -OH carboxylic acids \bigcirc mid-chain OH acids $\Diamond \alpha, \omega$ -alkanedioic acids

