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Organic signatures of fireplaces: experimental references for archaeological interpretations

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ABSTRACT

The use of fire is a well-established human practice, at least from the Late Pleistocene. The variability in fuel type highlights complex practices regarding fire technology throughout this period. This contribution provides the organic signatures from fireplaces and is based on a set of experimental studies using different types of fuel, notably bone and/or wood. Soil layers affected by fire operation were compared with soils unaffected by heating and soils impregnated with unburned bone fat. The carbon content, and lipid and bulk organic matter (OM) composition were determined through

organic carbon measurement, gas chromatography-mass spectrometry (GC-MS), as well as TMAH-assisted pyrolysis coupled to GC-MS.

With the exception of the wood-fueled fireplace, an increase in organic carbon content was caused by fireplace operation. The products of triacylglycerol degradation (diacids, oxo-acids, and glycerol derivatives) reflected the contribution from animal fat affected by the oxidation process with or without thermal alteration. The branched unsaturated C_{9:0} and C_{10:0} acids were detected only in py(TMAH)-GC-MS of soil impregnated with bone fat that was not thermally altered. The branched diacids, ketones, lactones detected in lipid extract, and short chain acids and *n*-alkane/*n*-alkene doublets detected in py(TMAH)-GC-MS, were produced by thermal alteration of animal fat. Finally, in the fireplace fueled with wood only, phenolic compounds, benzoic acids and benzene derivatives were detected and reflected the contribution of charred and uncharred plant OM.

The results from lipid and bulk OM characterization permitted to advance the organic signatures obtained as an experimental reference dataset specifically for the identification of fuel type used in fireplaces.

Highlights

Bone fires lead to a significant increase in soil OC content, whereas no variation was noticed for wood fire

Soil lipid and bulk OM distributions reflected the type of fuel used for experimental fires

Characterization of soil lipids and bulk OM permitted identification of byproducts of thermal alteration

Organic signatures from experimental fireplaces provide a reference for archaeological applications.

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Keywords

56 Fireplace, Prehistory, Experimental References, Organic signature, Soil

57 1. Introduction

58 The use of fire by hominins is a classic field of study within the prehistoric
59 discipline history and has provided a suite of publications, debates, and
60 controversies since the 19th century (see Roebroeks and Villa, 2011 and
61 Sandgathe et al., 2011 for recent examples). Dating the earliest fire use has been
62 a key question since the beginning of the discipline. Moreover, during the last few
63 decades an increase in interest regarding the study of fire related remains has
64 been observed, likely due to the fact that these remains represent both
65 technologically and sociologically significant shifts. The pyrotechnology of ancient
66 societies is most often investigated through several approaches, including
67 anthracology (Théry-Parisot, 2001), ash studies (Canti, 2003), taphonomy (Théry-
68 Parisot et al., 2010), ethnography (Henry and Théry-Parisot, 2014), magnetic
69 susceptibility (Bellomo, 1993), and micromorphology (Mentzer, 2012 and
70 references therein).

71 These studies have significantly improved our understanding of fireplace
72 operation (e.g. type of fuel, re-use and structure management) and proposed
73 several ethnographically-based models of utilization (e.g. specialized vs.
74 multipurpose, expedient vs. curated). However, direct interpretation of
75 archaeological vestiges remains difficult (Coudret et al., 1989) as archaeological
76 data are rarely explicit. The necessity for an experimental dataset, based on the
77 production and utilization of experimental fireplaces, thus seems apparent if we
78 wish to construct an interpretative tool for the deciphering of the archaeological
79 data.

80 Experimental investigations of prehistoric pyrotechnology are a well-established
81 trend in archaeological science; the production of reference databases is regularly
82 used to tackle taphonomic and functional questions (March et al., 2014; Mallol et
83 al., 2013 and references therein). Unfortunately, the numerous parameters
84 inherent in fireplace operation (e.g. type, state and quantity of fuel, type and
85 state of underlying sediment, operator intervention or not, extinction, re-use)
86 combined with the attempt to reconstruct different types of use (e.g. fireplace
87 morphology and layout, different cooking methods, different types of cooked

products) result in complex experimental data, within which it is often difficult to isolate and characterize individual aspects (Aldeias et al. 2016). More specifically these attempts to obtain reference collections rarely took into account the thermal alteration of the organic matter (OM) in archaeological contexts. Ultimately, databases from early experiments were rarely completely published and this lack of documentation hampers our ability to move forward. We are thus faced with the necessary task of systematizing, documenting, and publishing these experimental approaches.

Regarding the investigation of OM in archaeological remains through chemical methods, the exploration of prehistoric fireplaces is less advanced than other domains of archaeological chemistry, the analysis of residues in pottery for example (Evershed et al., 2001; Barnard and Eerkens, 2007; Evershed, 2008; Steele, 2013; Reber et al., 2015). In these studies, the main objectives were to recognize the content and use of the containers, despite the numerous processes that might confound interpretation of the chemical data, such as complex ancient practices, taphonomic modifications, or methodological issues (e.g. Reber and Evershed, 2004; Barnard et al., 2007). Even though they provide methodological and interpretative guidelines, results obtained from pottery analysis cannot be directly transferred to fireplace studies due to the differences that are expected in the archaeological context and then in the processes of organic residue formation and preservation.

Concerning fireplace investigations, pioneering studies identified lipids in cave fireplaces, using gas chromatography (GC) and thin-layer chromatography (Rottländer, 1989). The latter recognized the presence of reindeer fat and horse fat via comparison with actual reference sets, but he could not determine if their presence was related to fireplace operation (e.g. fuel type) or their utilization (e.g. cooking, animal carcass processing). With the spread of GC-mass spectrometry (GC-MS) in the 1990s, studies in all aspects of archaeological chemistry multiplied (Regert, 2011) and the study of fireplaces became more frequent (e.g. March et al. 1989; March, 1995 ; Buonasera, 2005; Lucquin, 2007; Kedrowski et al., 2009). More recent researches in the field also involved isotopic investigation

including bulk (e.g. Heron et al., 2010) and compound specific $\delta^{13}\text{C}$ analysis (e.g. March, 2013; Buonasera et al., 2015).

These studies highlighted recurrent conservation of organic compounds in hearth sediments, and attempted to associate them with their natural sources (e.g. plant or animal, terrigenous or marine) even sometimes evoking their specific origin. Moreover, recent studies (Regert, 2011; Kaal et al., 2014) mentioned and illustrated the potential interest of pyrolysis-GC-MS, possibly performed with *in situ* tetramethylammonium hydroxide derivatization [py(TMAH)-GC-MS], as a tool for the characterization of archaeological OM in pottery. The technique allows determination of the molecular composition of non-extractable OM, which is often the case for thermally altered fireplace residues. It therefore provides a good opportunity to document an OM fraction usually ignored in GC-MS based studies of solvent extracts.

This study focuses on the type of fuel used to stoke fireplaces in Late Pleistocene French open air sites. Most of the sites suffered post-depositional phenomena that frequently destroyed the upper part of fire structures, and particularly their fuel remains. As the resulting vestiges are limited mainly to mineral elements showing thermal alteration (e.g. underlying sediment and surrounding stones), their interpretation is challenging. In addition, the recurrent presence of burned bones residues (i.e. from slightly charred to completely calcinated) within the sites, but not always directly on fireplaces, raises the question of their frequent use as a common fuel by Paleolithic societies (Perlès, 1977, Villa et al., 2002). Despite experimental data ascertained the practicability of this use (Théry-Parisot, 2001 and 2002; Théry-Parisot and Costamagno, 2005; Costamagno et al., 2009), the interpretation of archaeological artefacts is still equivocal as numerous others anthropogenic and natural processes can produce burned bones (see Costamagno et al., 2009 for a review).

The objectives were thus to establish the geochemical signatures of fireplaces fueled with different materials and to provide a reference dataset for further interpretation of archaeological fireplaces. Experimental fireplaces were realized using wood and/or bone as fuel and compared with the soil layer unaffected by

human practices. The role of the thermal alteration of OM was also tested by analyzing the chemical signature of the soil layer impregnated with unburned animal fat. The chemical composition was determined through organic carbon content (C_{org}) measurement, GC-MS analysis of the lipid extract and py(TMAH)-GC-MS of the bulk OM.

2. Experimental

2.1. Study site and sampling

The study was conducted on the aurignacian site of Régismont-le-Haut (ca. 30 Ky BP, southern France; Pradeau et al., 2014) where an experimental protocol has been in place since 2012. The site is at the base of a strongly eroded Miocene calcareous hill, which provides most of the colluvial and aeolian deposits forming the sediment under study. During the Late Pleistocene, these colluvial and aeolian sediments slowly infilled paleochannels carved into the underlying marlstone. This context and the relatively arid climatic conditions led to the formation of a colluvic and calcaric Regosol (IUSS Working Group WRB, 2015; *Colluviosol calcaire et carbonaté*: AFES, 2008), with few pedogenic features except those linked to the dissolution and precipitation of $CaCO_3$.

Experimental fires were set up to mimic several contexts of hypothetical archaeological operation; the main parameter that was varied was fuel type (wood and/or bone). To emulate the context of archaeological fireplaces and avoid modern agricultural induced bias, the experiment was run in realistic open air conditions on a sterile area of the excavated archaeological level, ca. 50 cm under the plow zone. All experiments and subsequent sampling occurred between August 27th and September 11th, 2012. No significant precipitation was recorded on site during this interval; the local meteorological station of Sètes recorded a mean maximal temperature of 28.6 °C in August and of 23.8 °C in September and a mean minimal temperature of 21.5 °C in August and of 17.6 °C in September. Prior to the experiment a sample of unaffected sediment was collected as a control for the natural organic content and composition in the area (*Control*).

Experiments involved the use of bone and/or pine wood of local origin. The former consisted of up to 15 cm fragments of steer long bones and pelvic bones (*Bos*

taurus). The wood included parasol pine (*Pinus pinea*) dry twigs, small branches and 50 cm long and up to 10 cm in diameter logs. For the first experiment, 10 kg of fragmented steer bones were deposited on the surface of the soil in order to imitate OM impregnation by way of percolation through the sediment. A sample (*Soil with bone OM*) was collected after a week. The second experiment was with a bone-fueled fireplace (*Bone fire*) in which 10 kg of fragmented steer bones were burned. It is noteworthy that a small quantity of pine wood (< 1 kg) was necessary to initiate combustion in this type of hearth (Théry-Parisot and Costamagno, 2005). It operated during 131 min. The third experiment was with a wood-fueled fireplace (*Wood fire*), in which 5 kg of dry pine wood were burned during a total operating time of 137 min. Finally, the last experimental fire involved the use of both 5 kg of pine wood and 3 kg of bone (*Wood+bone fire*), to emulate a mixed-fuel fireplace. The total operating time was of 118 min.

For the 4 experiments, sampling was conducted following visual distinction of layers (Fig. 1a). Sampling first concerned the altered soil (Upper sub sample, Table 1) but also the underlying and seemingly unaltered soil (Lower sub sample, Table 1). The micro-local organic content was measured to test the significance of the visual layer distinction.

Samples were air-dried over two days at 35 °C (Froilabo-AC120 oven). After manual removing of the most fragile coarse elements (calcitic nodules and organic macroscopic remains such as roots), samples were sieved and the > 2 mm fraction was discarded. Samples were then ground with a mortar and pestle, and eventually with a Retsch-PM200 planetary ball mill (6 min, at 450 rpm), to pass a 250 µm sieve.

2.2. *C_{org}* measurement

Due to the naturally high carbonate content of the sediment, acid fumigation was carried out prior to elemental analysis, following Harris et al. (2001). After weighing the samples in Ag foil capsules, they were humidified (15 µl deionized water) and placed in a desiccator under vacuum, with a beaker of HCl (12 M), during 6 h. Lastly, the sample was air-dried for 12 h at 40 °C.

The C_{org} was measured on fumigated samples using an elemental analyzer (Vario PYRO cube, Elementar). To allow C_{org} quantification, two Ag capsules of the same sample were added for a single analysis. The resulting reproducibility was assessed to be between 0.3% (instrumental) and 0.8% (methodological).

2.3. Lipid extraction and GC-MS

Samples for lipid analysis were extracted using an accelerated solvent extractor system (Dionex-ASE100). Between 5 and 7 g of sediment were placed in a 10 ml Inox cell topped with sterile glass fiber. The extraction program ran at 60 °C, with a static time of 20 min, using dichloromethane/methanol (DCM/MeOH) (2:1, v/v). The extract was concentrated using a rotary evaporator and dried under a gentle N₂ flow, dissolved in DCM in a 1,5 ml vial with a PTFE screw cap and stored at 4 °C until analysis (adapted from Quénéa et al., 2012)

Each extract was analyzed with an Agilent 6890 gas chromatograph coupled to an Agilent 5973N mass spectrometer using electron ionization at 70 eV. The GC instrument was equipped with a 30 m Restek 5 Sil MS column (i.d. 0.25 mm, film thickness 0.5 µm). The carrier gas was He, at a constant flow rate of 1 ml/min. Samples were injected splitless with the injector at 300 °C. The oven temperature was programmed from 80 °C (0.5 min) to 100 °C at 10 °C/min then to 320 °C at 4 °C/min. Compound assignment was based on comparison with published data (references thereafter in the text) and the NIST mass spectral library.

The whole device was operated with a Gerstel multipurpose sampler (MPS) and Gerstel Maestro software. The workflow included derivatization of extracts by adding 10% (v/v) BSTFA [N,O-bis(trimethylsilyl)trifluoroacetamide] and heating at 60 °C for 10 min to obtain trimethylsilylated (TMS) derivatives.

2.4. py(TMAH)-GC-MS

Bulk OM from the upper subsample was characterized through Curie-Point pyrolysis (Pilodist) coupled to a GC (Thermo Trace – 30 m RXI 5 Sil MS capillary column, i.d. 0.25 mm, film thickness 0.5 µm with Integra guard column) and a MS (DSQ). Samples mixed with TMAH (25% w/v in MeOH) were loaded in a ferromagnetic tube and pyrolyzed for 9.9 s at 650 °C under a 1 ml/min He flow

(adapted from Quénéa et al., 2005). The pyrolysis products were separated in the GC system, where the temperature of the GC oven was held at 50 °C for 10 min, before an increase at 2 °C/min to 310 °C. The ion source of the mass spectrometer was at 220 °C, which was set to scan from m/z 35 to 800. The total ion current (TIC) trace was recorded and products of pyrolysis identified using comparison with published data (references thereafter in the text) and the mass spectral NIST library.

3. Results and discussion

3.1. Macroscopic features and C_{org} variations

The *Control* displayed general features illustrative of the local soil type: a homogenous sandy loam with numerous calcitic features and few organic macro-elements (Table 1). An expected low C_{org} (0.3%) was measured and was consistent with that of the four lower subsamples ($0.3 \pm 0.1\%$) collected below each experimental structures. This result validated the visual identification of fire-affected soil layers and supported the low soil variability needed to track variables influenced by the experimental parameters.

On the contrary, most of the experiments led to an increase in the organic content of upper subsamples (Table 1). In the *Soil with bone OM*, a 2 cm deep impregnation of the sediment was observed on the surface (Fig. 1b) and in section. During the experiment, OM (likely bone marrow) leaked from the bone fragments and was responsible for the impregnation, as well as the increase in C_{org} , which reached 4.7%. The macroscopic visual aspects of the *Bone fire* consisted of an irregular shaped darkened and hardened hearth covered with carbonized to calcinated bones. Vertically, the sediment was affected by the combustion down to 4 cm depth, displaying darkening and reddening. During the operation, a significant leakage of organic fluid was observed around the structure (Fig. 1c) and was also noted vertically on the structure's periphery. The leakage resulted in a slightly lower increase of C_{org} (2.3%) than for the *Soil with bone OM*. The phenomenon can likely be explained by a shorter interaction time between the soil and the bone OM, as well as the combustion process that likely degraded part of the bone marrow, thereby limiting its influence on C_{org} . The

operation of the *Wood fire* resulted in a circular (ca. 50 cm diam.) collection of ash and charcoal overlying a reddened sediment (Fig. 1e). This reddening was 0.5 to 1.5 cm deep. No significant difference was detected in the C_{org} of this reddened sediment compared with the *Control*, likely because the coarse fraction, which may have included some combustion residue, was discarded during sample collection and preparation. The results of the *Wood+bone fire* experiment were, unsurprisingly, intermediate between those from other experimental fireplaces. A superficial covering, consisting logically of a mix of wood and bone combustion residue, capped the structure (Fig. 1d). Vertically, the structure displayed a 1 to 2.5 cm deep reddening, whose first 1 cm was heavily influenced by the combustion residues, leaving a darkened layer whose reddened nature was not visible to the naked eye (Table 1). Despite the C_{org} (0.9%) did not prove to be significantly higher than the C_{org} of the *Wood fire* (or *Control*), the value was intermediate between the *Wood Fire* and the *Bone fire*, what seemed coherent with the field observation of a lower input of bone OM due to 1) a lower quantity of bone used as fuel and 2) the formation of wood embers that might have prevented OM leakage into the soil by increasing the efficiency of combustion.

3.2. Local organic background

In accord with its low C_{org} , the *Control* did not present any detectable component from py(TMAH)-GC-MS analysis. No phenol, lignin-derived components, or long chain alkanes/alkenes and acids, traditionally seen in the pyrolysis products of natural soil OM (Hempfling and Schulten, 1990), were visible in the archaeological soil layer. Lipids detected from GC-MS were diverse, yet present in such low concentration as to render them negligible. The identification of such a diversified lipid fraction was made possible through a higher concentration of sample during preparation. The absence of a significant organic signal confirmed its suitability as a control for the experiments that followed.

3.3. Sources of organic compounds

3.3.1. Non-diagnostic

The saturated *n*-acids from $C_{6:0}$ to $C_{18:0}$, dominated by $C_{16:0}$ and $C_{18:0}$, and monomethylated acids, were ubiquitous in both lipid extracts and

thermochemolysis products (Figs. 2 and 3, Tables 2 and 3). Their presence in all the samples might reflect animal, plant and microbial biomass origins. Nevertheless, the C chain distribution appeared more specific, especially the high relative proportion of short-chain fatty acids in *Soil with bone OM*, and is discussed later.

Due to its widespread occurrence in animal tissue (Christie, 2014) cholesterol can be considered as an indicator of an animal contribution to lipid extracts, but a fungal or even bacterial origin cannot be completely excluded (Huang and Meinschein, 1976; Volkman, 2005; Christie, 2014). These last two elements were likely the source of cholesterol [25] in the *Control* (Fig. 2a) and the *Wood fire* (Fig. 2e).

The significant contribution of Me-C₁₅ and Me-C₁₇ acids (iso and anteiso) to the lipids and py(TMAH)-GC-MS trace of the *Soil with bone OM* (Figs. 2b and 3a), the *Bone fire* (Figs. 2c and 3b) and the *Wood+bone fire* (Figs. 2d and 3c), might point to a fungal and bacterial OM signature (Saiz-Jimenez and De Leeuw, 1986; Dudd et al., 1998; Kanthilatha et al., 2014; Spangenberg et al., 2014). However those branched acids can also be produced by bacteria in ruminant intestines (Dudd et al., 1999; Christie, 2014) and an animal origin can not be ruled out in the context of these experiments.

The presence of squalene [24], a precursor of numerous sterols, might have a natural origin in the extract of the *Control* (Fig. 2a) and the *Wood fire* (Fig. 2e). However, it is also a broadly-used plasticizer, so in-situ or handling contamination could not be totally dismissed (Evershed, 1993; Heron et al., 2010).

3.3.2. *Plant OM*

The presence of stigmasterol [26] and sitosterol [28], both common phytosterols (Huang and Meinschein, 1976), and of taraxerone [27], α -amyrone [29], and lupenone [30] in the extract of the *Control* (Fig. 2a) point to a significant contribution from plant OM (Vilegas et al., 1997; Mathe et al., 2004; Simoneit, 2005; Hernández-chávez et al., 2012). The contribution of aliphatic compounds

with relatively long chains is commonly considered to reflect a plant-derived OM origin. In the extract of the *Control* (Fig. 2a) and *Wood fire* (Fig. 2e), the *n*-alkanes (C₂₁, C₂₉, C₃₁ and C₃₃) might be attributed to an plant epicuticular wax origin (Spangenberg et al., 2014), whereas *n*-alcohols longer than C₁₇, and saturated *n*-acids longer than C₁₈, support the assignment of slightly altered plant derived compounds (Evershed et al., 1999; Poirier et al., 2005). Finally, phenol derivatives [4 and 8] and benzoic acids [1 and 3] in the *Wood+bone fire* (Fig. 2d) and the *Wood fire* lipids (Fig. 2e) highlight byproducts of lignin degradation (Simoneit et al., 1993; Quenea et al., 2004; Simoneit, 2005; Regert et al., 2006; Shadkami and Helleur, 2010).

As the phenolic and benzoic compounds were only recovered in *Wood+bone fire* and *Wood fire* (Table 4), they may be illustrative of the contribution of the wood OM used as fuel to the organic signatures of these experiments (March et al., 2014).

3.3.3. Thermally-altered plant OM

Aromatic compounds such as benzene derivatives ([5] to [10]) in the py(TMAH)-GC-MS trace of the *Wood+bone fire* (Fig. 3c) and the *Wood fire* (Fig. 3d), as well as naphthalene [10] in the extract of the *Wood fire* (Fig. 2e), reflect a high degree of OM carbonization (Simoneit et al., 1993; Knicker et al., 2005; Kaal et al., 2009; Alexis et al., 2012; De la Rosa et al., 2012). This type of compounds could also result from artefact production by Curie point pyrolysis from carbohydrates, as observed by Pastorova et al. (1994). However it is noteworthy that in the present study, the benzene and naphthalene derivatives were only detected in the experiments involving wood as dominant fuel and that the numerous furan and pyranone derivatives expected in cellulose pyrolysis were not detected (Table 4). This result suggests that they likely come from carbonized wood residues. Moreover the absence of odd C chain predominance in the alkane series in the *Wood fire* extract (Fig. 2e) could also be indicative of the thermal alteration process of wood fuel OM (Eglinton and Calvin, 1967; Wiesenberg et al., 2009; De la Rosa et al., 2012; March et al., 2014).

Many of these compounds were readily documented in GC-MS and py-GC-MS studies of forest fire sediments (e.g. Knicker et al., 2005). Their observation in fireplace context validates the crossing of results between these close fields of research. Finally, their association in wood fueled fires with compounds specific of uncharred wood OM (Table 4) provides a good overall signature for this type of fire operation.

3.3.4. Bone OM

In the *Soil with bone OM* and the *Bone fire*, the dimethoxypropanol [1], trimethoxypropane [2], saturated and unsaturated *n*-acids, diacids, epoxy- [13], and (di)hydroxyacids ([12] and [14] to [17]) were detected from py(TMAH)-GC-MS (Figs. 3a and b). Monoacylglycerol [15, 17, 19 to 23] and diacylglycerol [31 and 32] were observed in the extract (Figs. 2b and c). These compounds likely resulted from triacylglycerol degradation (Dudd et al., 1998; Evershed et al., 2002; Van Den Berg et al., 2002). They could originate from plant (Van Den Berg et al., 2002; Maher and Bressler, 2007) or animal fats (Evershed et al., 2002; Ben Hassen-Trabelsi et al., 2014). However, the presence of *n*-acid C_{18:1}, one of the most important *n*-acid in animal fat (Malainey et al., 1999; Nieuwenhuyse et al., 2015), points to the latter origin. Moreover, several compounds likely reflect the contribution of fat degraded by hydrolysis or oxidation. The short chain *n*-acids (C₈ and C₉) and monomethyl branched unsaturated acids (Me-C₉ and Me-C₁₀) in the py(TMAH)-GC-MS products (Figs. 3a and b), as well as the short chain diacids (C₄ to C₁₁) and *n*-acid (C₈ and C₉) in the lipid extract (Figs. 2b and c), could originate from drying and/or oxidation of unsaturated *n*-acids (Wexler, 1964; Mills, 1966; Erhardt, 1998; Van Den Berg et al., 2002; Rontani and Aubert, 2008).

These compounds were all present in the *Soil with bone OM* and occasionally present in *Bone fire* and *Wood+bone fire* (Table 4). Those results were in accordance with existing references on animal fat (e.g. Evershed et al., 2002), despite short-chain methyl branched acids have been rarely described in such context. Moreover, the functioning of the fire seems to induce an important reduction of the number of recovered diagnostic compounds.

3.3.5. Thermally-altered bone OM

In the py(TMAH)-GC-MS TIC traces of the *Bone fire* (Fig. 3b) and of the *Wood+bone fire* (Fig. 3c), the short chain *n*-alkane/*n*-alkene doublets could be related to thermal degradation processes (Almendros et al., 1988; Nawar, 1989; Simoneit, 2002; Eckmeier and Wiesenberg, 2009; Wiesenberg et al., 2009). In the extract of the same samples (Figs. 2c and d), the series of short chain monomethyl branched diacids (C₄ to C₁₀), short chain *n*-alcohols (< C₁₇), oxoacids, as well as ketones (C₁₄ to C₁₇ and C₂₉, ₃₁, ₃₃ and ₃₅) and lactones [13 and 14] were also common markers of, more specifically, animal OM thermal degradation (Nawar, 1989; Evershed et al., 2002). Ketones and γ -lactones are frequently reported in thermally altered OM and are usually attributed to animal-fat and meat degradation by heating (Nawar, 1969; Evershed et al., 1995). In this context, the triacylglycerols and *n*-acids are both sources of the long chain ketones, whereas short chain methyl ketones are considered to be secondary products formed during more intense heating. Besides, the γ -lactones are formed by cyclisation of hydroxyacids following dehydration upon thermal stress (Nawar, 1969). The disappearance of 10, ω -dihydroxyoctanoic acid [16], in parallel with the formation of γ -lactones [13 and 14] in *Bone fire* (Fig. 2c) compared with *Soil with bone OM* (Fig. 2b) might be illustrative of this process.

The *Bone fire* and *Wood+bone fire* provide a signature specific to the use of bone as fuel (Table 4). It is characterized by markers of animal fat heating, resulting from the melting and the burning of marrow initially contained in the bones. Yet, it is noteworthy that, to our knowledge, the formation of short monomethyl branched diacids (C₄ to C₁₀), short chain *n*-alcohols (< C₁₇) and oxoacids is little known in this context.

3.4. Archaeological implications

In order to apply this reference dataset to archaeological contexts, the recovery of previously identified organic signatures has to be questioned. As stated by March et al. (2014), post-depositional processes may occur, namely the potential pollution through input of more recent OM and the degradation processes that likely affect the archaeological organic residues.

First, the pollution issue regards both natural and anthropogenic agents. Natural ones are mostly linked to sedimentary covering of archaeological site and pedogenic processes such as bioturbation and are unavoidable. Additionally, this type of pollution will most likely correspond to vegetal OM lacking thermal alteration byproducts and displaying a weak maturity allowing its recognition (March et al. 2014). Anthropogenic pollution can be related either to agricultural practices or to contamination related to archaeological excavation and sampling. As for agricultural input, markers are generally well known and easily identifiable. They may correspond to modern practices (chemical fertilized), which will not interfere with our compounds of interest, and to manuring (like coprostanol; Evershed et al., 2001). However, conjoint collection of control samples both in stratigraphy and planimetry allows definition of the local soil organic background (e.g. Heron et al., 1991).

Secondly, the ability of the signatures to be preserved over time is uncertain. Except from the work of March et al. (2014), this central matter has been rarely explored in the specific case of fireplaces, despite substantial documentation exists in the field of the aging of fire affected soil OM (e.g. Knicker, 2011) and organic residues in pottery (e.g. Malainey, 2007). Labile compounds such as unsaturated *n*-acids or short-chain saturated *n*-acids are unlikely to survive long period of time (Eglinton et al., 1991). In contrary, due to their relative recalcitrance, aromatic compounds originating from lignin (benzoic acids, phenols) and from charring processes (benzene and naphthalene derivatives) are expected to remain in soil at long timescale (Baldock et al., 2004; Lützow et al., 2006; Marschner et al., 2008). Lipids were also reported to present decreasing stability following the order: alkanes, alkanones, sterols, *n*-alkanoic acids (especially > C₂₀), *n*-alkanols and *n*-alkenoic acids (Cranwell, 1981). Based on this chemical property, among previously identified signatures the best candidates for recovery after millennia are 1) ketones and *n*-alkane/*n*-alkene doublets for *Bone fire*, 2) saturated *n*-acids longer than C₁₈, benzoic acids, ketones, benzene derivatives and *n*-alkane/*n*-alkene doublets for *Wood and bone fire* and 3) saturated *n*-acids longer than C₁₈, *n*-alkanes, benzoic acids, phenolic compounds

and benzene derivatives for *Wood fire*, whereas no compound is expected to remain in *Soil with bone OM* (Table 4).

Nevertheless, the relevance of recalcitrance alone to predict the residence time of compounds has been widely discussed (Marschner et al., 2008; Bol et al., 2009; Kleber et al., 2011). Rather, the spatial inaccessibility or the organo-mineral interactions seem determinant for preservation at long timescale or in subsoil horizon (Lützow et al., 2006). More specifically, in calcareous soils like at Régismont-le-Haut, the high polyvalent cation content may induce the formation of cation bridges, leading to stabilization of organic residues (Lützow et al., 2006; Grünewald et al., 2006). However, no effect of CaCO_3 and pH on mineralization could be observed by Bertrand et al. (2007), questioning the significance of this stabilization mechanism. In this fireplace context, the association with charcoal might also limit the decomposition (Hernandez-Soriano et al., 2016), in this case, mainly through physical protection. Finally, as the substrate limitation can be considered as the main contributing factor of organic matter stabilization in deep soil horizons (Marschner et al., 2008), burial underwent by archaeological level could then have favored various compound preservation. For example, in prehistoric contexts, diacids were documented through GC-MS, in addition to long chain ketones and γ -lactones in Pincevent site fireplace sediments and heated stones (ca. 14 Ky cal. BP, Lucquin, 2007). Short chain diacids were also mentioned in the GC-MS study of sediments related to thermal fat processing in Alaska (ca. 5 Ky BP, Buonasera et al. 2015). In the same geographical context, the study of sediments from fireplaces (ca. 14 Ky cal. BP, Kedrowsky et al. 2009) illustrated the preservation of several branched *n*-acids interpreted as resulting from lipid degradation by bacteria or from ruminant fat. The recovery of compounds similar to those detected in this work, but with various intrinsic recalcitrance, could illustrate the importance of physical conditions for the record of organic signatures.

4. Conclusion

This study provides the organic signature of different experimental fireplaces and demonstrates that they allow distinguishing if bone and/or wood were used as

fuel. It confirms that the thermal alteration of OM significantly modifies the initial signature of the unburned material, limiting the direct use of markers from unaltered sources to reconstruct fuel nature in archaeological fireplaces. In contrary, several byproducts of thermal degradation can serve as indicators of fuel type and, in this respect, the use of Py(TMAH)-GC-MS for the identification of non-extractable compounds proved particularly adapted. However, because of the variable organic components that constituted the organic signatures of the fireplaces, their differential preservation has to be expected. Their intrinsic recalcitrance, as well as the local environment and depositional context, might influence the recovery of the chemical signatures under investigation. Consequently, when applying this type of reference dataset to archaeological samples, which likely underwent substantial post-depositional processes, caution is needed. Nevertheless, as many identified compounds were found within non-extractable or recalcitrant OM fraction in our study, and subject to the collection of control samples, the reconstruction of archaeological interpretable signatures based on this referential dataset seems possible.

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Table 1

Sample description, experimental parameters and organic carbon content (C_{org}) of upper subsamples (visually-identified as affected by the experiment) and of lower subsamples (visually unaffected by the experiment).

Name	Subsample	Description	C _{org}
Control	-	Yellow sandy loam with frequent calcitic features (e.g. hypocotings, nodules)	0.3 %
Soil with bone OM	Upper	Yellow to brown sandy loam with frequent calcitic features and organic fluids impregnation ; 2 cm thick	4.7 %
	Lower	Similar to control	0.2 %
Bone fire	Upper	Reddened and carbonized sandy loam with frequent calcitic features and fine carbon coatings/infillings ; 4 cm thick	2.3 %
	Lower	Similar to control	0.3 %
Wood+bone fire	Upper	Reddened and carbonized sandy loam with frequent calcitic features and fine carbon coatings/infillings ; 2.5 cm thick	0.9 %
	Lower	Similar to control	0.3 %
Wood fire	Upper	Reddened sandy loam with frequent calcitic features and rare combustion residues on surface ; 1.5 cm thick	0.2 %
	Lower	Similar to control	0.2 %

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Table 2

Compounds in extract of the samples. The molecular weight (MW) and the main mass fragments of TMS derivatives are reported (with decreasing contributions). Compound numbers are reported in Fig. 2. The same peak number may correspond to several co-eluting compounds.

Peak #	Main fragments (<i>m/z</i>)	MW	Name
1	179, 105, 77, 135, 194	194	Benzoic acid, TMS
2	73, 147, 205, 293	308	Glycerol, TMS
3	193, 119, 149, 91, 208	208	2-Methylbenzoic acid, TMS
4	151, 73, 210, 195, 117, 167	210	2-Phenoxy-ethanol, TMS
5	71, 89, 56, 43, 173, 143	216	Propanoic acid, 2-methyl-, 2-ethyl-3-hydroxyhexyl ester
6	43, 71, 145, 83, 161, 98	304	1,3-Diisobutyryl, TMS
7	227, 75, 143, 129	242	3-Methyldecen-1-ol, TMS
8	194, 193, 73, 209, 224	224	Vanillin, TMS
9	57, 71, 83	242	Octyl ether
	257	257	Dimethylbenzoacridine
10	195, 210, 180, 165	210	Naphthalene, 1,2,3-trimethyl-4-propenyl-, (E)-
11	195, 119, 73, 103	252	2-Acetoxy-benzoic acid (TMS)
12	88, 101, 157, 284, 239	284	Hexadecanoic acid, ethyl ester
13	85, 236, 192, 97	254	γ -Palmitolactone
14	85, 97, 220, 264	282	γ -Stearolactone
15	343, 73, 147, 227, 431, 371, 205	446	1-Monomyristin, TMS
16	317, 73, 215, 147	532	10, ω -Dihydroxy-octadecanoic acid, TMS
17	129, 218, 73, 147, 103, 313, 459, 403	474	2-Monopalmitin, TMS
18	369, 73, 129, 147, 55, 370, 457	472	1-Monopalmitolein, TMS
19	371, 73, 147, 459	474	1-Monopalmitin, TMS
20	103, 129, 73, 218, 147	500	2-Monoolein, TMS
21	129, 218, 147, 341, 103, 487	502	2-Monostearin, TMS
22	397, 73, 129, 147, 103	500	1-Monoolein, TMS
23	399, 73, 147, 129, 267, 487	502	1-Monostearin, TMS
24	69, 81, 137, 95, 121	410	Squalene
25	329, 129, 368, 458	458	Cholesterol, TMS
26	83, 129, 484, 394, 255	484	Stigmasterol, TMS
27	204, 300, 133, 285, 424	424	Taraxerone
28	129, 357, 396, 486, 255	486	Sitosterol, TMS
29	218, 424, 203	424	α -Amyrone
30	424, 205, 189, 409	424	Lupenone
31	129, 145, 313, 329, 385, 625	640	1,2-Dipalmitin, TMS
32	371, 372, 73, 385, 625	640	1,3-Dipalmitin, TMS

Table 3

Compounds from py(TMAH)-GC-MS analysis of bulk OM. The MW and main mass fragments of original compounds or methylated derivatives are reported (with decreasing contribution). Peak numbers are reported in Fig. 3. The same peak number may correspond to several co-eluting compounds.

Peak #	Main mass fragments (m/z)	MW	Name
1	45, 75, 88, 102	120	1,3-Dimethoxy-2-propanol
2	59, 45, 89, 102	134	1,2,3-Trimethoxypropane
3	55, 41, 68, 84, 112	112	Cycloheptanone
4	41, 57, 70, 82, 98, 114, 124	142	Nonanal
5	119, 134, 91, 115	134	1,2,3,5-Tetramethylbenzene
6	119, 134, 91, 77, 65	134	1,2,4,5-Tetramethylbenzene
7	133, 148, 105, 91, 115	148	Pentamethylbenzene
8	147, 162, 91, 119, 105, 77	162	1,2,4-Trimethyl-5-(1-methylethyl)benzene
9	147, 162, 133, 91, 105	162	1,2-Diethyl-3,4-dimethylbenzene
10	147, 162, 91, 105, 119, 131	162	Hexamethylbenzene
11	101, 88	298	Heptadecanoic acid, ethyl ester
12	71, 183 95, 294	326	8-Methoxy-9-octadecenoic acid, methyl ester
13	71, 227, 163, 81, 294	340	12,13-Epoxy-11-methoxy-9-octadecanoic acid, methyl ester
14	71, 213, 169, 294	326	10-Methoxy-8-octadecenoic acid, methyl ester
	71, 169, 227, 294	326	9-Methoxy-10-octadecenoic acid, methyl ester
15	131, 71, 241	328	4-Methoxy-octadecanoic acid, methyl ester
	83, 145, 201	302	6,10-Dihydroxyhexadecanoic acid, methyl ester
	229, 143	328	11-Methoxy-octadecanoic acid, methyl ester
16	55, 74, 155	330	9,10-Dihydroxyoctadecanoic acid, methyl ester
17	69, 201, 83, 137,	358	9,18-Dimethoxyoctadecenoic acid, methyl ester

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Table 4

Summary of specific components in the extracts (GC-MS) and in the bulk soil OM composition [py(TMAH)-GC-MS], and possible origins. In grey, compounds with an expected long-term preservation due to their recalcitrance (according to literature, see §3.4).

	<i>Control</i>	<i>Soil with bone OM</i>	<i>Bone fire</i>	<i>Wood + bone fire</i>	<i>Wood fire</i>	Possible origin
Lipid extract (GC-MS)	Saturated <i>n</i> -acids > C ₁₈	X		X	X	Plant OM
	<i>n</i> -Alcohols > C ₁₇	X			X	Plant OM
	<i>n</i> -Alkanes	X			X	Plant OM
	Phytosterols	X			X	Plant OM
	Terpenes	X				Plant OM
	Benzoic acids			X	X	Plant OM
	Phenolic compounds				X	Plant OM
	Other aromatic compounds				X	Thermal alteration of plant OM
	Diacids	X	X	X		Animal fat
	(di)Hydroxyacids	X				Animal fat
	Monomethyl branched dicarboxylic acids < C ₁₀		X			Slight thermal alteration of animal fat
	Oxoacids		X	X		Thermal alteration of animal fat
	Monomethyl branched acids < Me-C ₁₂		X	X		Thermal alteration of animal fat
	<i>n</i> -Alcohols < C ₁₇		X	X		Thermal alteration of animal fat
	Ketones		X	X		Thermal alteration of animal fat
	Lactones		X	X		Thermal alteration of animal fat
Bulk OM [py(TMAH)-GC-MS]	Benzene derivatives			X	X	Thermal alteration of plant OM
	Monomethyl branched unsaturated C ₉ and C ₁₀ acids, high C ₈ and C ₉ saturated <i>n</i> -acids	X				Animal fat
	Diacids	X	X			Animal fat
	(di)Hydroxy-, epoxy-saturated and unsaturated acids	X	X			Animal fat
	Glycerol derivatives	X	X			Animal fat
	<i>n</i> -Alkane/ <i>n</i> -alkene doublets		X	X		Thermal alteration of animal fat

864 **List of figures:**

865 **Fig. 1.** Pictures of Regismont-le-Haut (France) experimental structures during
866 summer 2012 experimental session; (a) section of an experimental fireplace with
867 the localization of upper and lower subsamples, (b) *Soil with bone OM*
868 experiment after removal of bone fragments, (c) *Bone fire* experiment after
869 removal of combustion residues, (d) *Bone+wood fire* experiment after removal of
870 combustion residues, (e) *Wood fire* experiment after removal of combustion
871 residues. The scale is 10 cm long in (a), 24 cm long in (b), (c), (d) and (e).

872 **Fig. 2.** Total ion Chromatogram (TIC) of silylated extract of the (a) *Control*, (b)
873 *Soil with bone OM*, (c) *Bone fire*, (d) *Wood+bone fire* and (e) *Wood fire*. Numbers
874 refer to Table 2; p, plastic derivative; c, column peak. C_x indicates number of
875 carbons in the linear chain.

876 **Fig. 3.** TIC trace from py(TMAH)-GC-MS analysis for (a) *Soil with bone OM*, (b)
877 *Bone fire*, (c) *Wood+bone fire* and (d) *Wood fire*. Numbers refer to Table 3. C_x
878 indicates number of carbons in the linear chain.

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- Saturated n-acids
- Unsaturated n-acids
- Diacids
- m- Monomethyl branched
- ox- Oxo-
- ▼ n-Alcohols
- n-Alkanes
- ◀ Ketones

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