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1 **Slingram EMI devices for characterizing resistive features using apparent conductivity**
2 **measurements: check of the DualEM-421S instrument and field tests.**

3

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5

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8

9 **Abstract**

10 This paper addresses the characterization of resistive archaeological targets and near surface
11 structures by ElectroMagnetic Induction (EMI). It presents tests achieved with the
12 DualEM421S instrument (Duaem Inc., Canada) in order to be able to quantitatively compare
13 these measurements to the standard technique of direct-current (DC) resistivity. The test was
14 done over the Gallo-roman site of Vieil-Evreux in Normandy, France and 1D and 3D
15 inversions were applied to the data set obtained.

16 We have first investigated the signal/noise ratio of each of the 6 DualEM receiver coils both
17 in a static mode and for a quad-pulled system. The dependence on the roll angle was also
18 measured and it is shown that rotation of DualEM must be taken into account if the roll angle
19 is more than $\pm 10^\circ$. Absolute calibration and in-phase/quadrature (out of phase) component
20 discrimination was checked by measuring the response of a small conductive and non-
21 magnetic sphere. Several EM soundings by measuring the instrument response at different
22 heights were done in order to check the quadrature (out-of-phase) response of the instrument.
23 Inversions of these EM soundings were compared to DC Vertical Electric Soundings (VESs)
24 over 4 locations and found in accordance. Several maps using different coil configurations
25 (HCP, VCP, PERP) and different heights were performed and inverted, both for a wide mesh

26 (5m) and for a finer one (0.5m). The wide mesh allows a global and rapid description of the
27 surface geology context (continuous DC measurements cannot deliver equivalent depth of
28 investigation). The fine mesh conductivity maps clearly show the walls of a *fanum* (temple) as
29 well as other structures and prove that the DualEM-421S is able to map correctly
30 archaeological resistive targets. These maps and their interpretations were compared to
31 previous results obtained by DC technique.

32

33 **Keywords**

34 EMI, multi-receivers, DualEM-421S, calibration, conductivity response, inversion,

35

36

37 **Introduction**

38 In archaeology, magnetometry is the primary technique for the detection of low
39 resistivity earthen targets, like ditches for example, while Direct Current, DC, (earth-
40 resistance) and, more recently, Ground Penetrating Radar, GPR, predominate in the detection
41 of resistive targets, like walls for example. Since both DC and GPR are expensive to acquire
42 and/or process, the detection of resistive features is likely underestimated in most of the
43 assessment of archaeological potential despite the high occurrence of resistive targets in many
44 of the archaeological sites.

45 The use of earth-resistance in the field has two main drawbacks compared to
46 magnetometry: it is more time consuming despite the development of quad-pulled systems
47 such as ARP© (Dabas, 2009), and it might not be practical to use in areas of high contact
48 resistance. The two main advantages of DC resistivity compared to magnetometry are the
49 strong response of both earthen and stone built archaeological features due to a generally high
50 electrical contrast, and also the possibility to achieve more accurate inversions using several

51 depths of investigation. Globally, despite the development of electrostatic multipole arrays
52 (Flageul *et al.* 2013) that fully overcame the galvanic contact limitation, the deployment of
53 such systems remains challenging due to their weight and complexity.

54 EM instrumentation tends to be lighter and thus more convenient than DC
55 instrumentation. These characteristics have made EMI a promising alternative to DC, but only
56 recently have instruments with an accurate phase separation and with multi-receivers become
57 readily available. These developments should open the possibility to have a conductivity
58 inversion and also to improve the use of the in-phase response. A first study aiming at a
59 global comparison of available instruments was done in the area of soil mapping in 2009
60 (Gebbers *et al.*, 2009). In the archaeology field, only a few publications have begun to tackle
61 the use of multi-receivers instruments (Saey *et al.*, 2012; Bonsall *et al.*, 2013) and none to our
62 knowledge have studied the response of multi-receivers systems to resistive targets and
63 compared EMI measurements with resistivity over a whole site.

64 Numerous archaeological sites exhibit stone built remains or compacted horizons that
65 constitute the essence of the site and provide the key to its overall organization and evolution
66 through time. These features, in general, correspond to a marked contrast in electrical
67 resistivity with the surrounding soil and the DC resistivity method has been, and is, logically
68 applied to map apparent resistivity above them. Very good results have been obtained in
69 different climatic and chrono-cultural contexts (Papadopoulos *et al.* 2009, Bossuet *et al.* 2012)
70 and the application of multi-depth arrays allowed exploring the depth and the thickness of the
71 different features (Brinon *et al.* 2012). The sensitivity of EMI devices to resistive targets is
72 indeed worse than that of the DC resistivity technique (Tabbagh 1986(a)) but these devices
73 allow simultaneous measurement of the magnetic susceptibility. Susceptibility can show for
74 example contrasts between stones and their surroundings (Gaffney *et al.* 2000, Saey *et al.*
75 2013), and experiments have shown the possibility of wall detection even with small coil

76 separation instruments (Thiesson *et al.* 2009). Compared to DC deployments for a given depth
77 of exploration, EMI instruments are more compact, which facilitates deeper investigations.

78 In this paper, after a field test in which noise and calibration of DualEM-421S were
79 analyzed, we examine the results obtained with this instrument over some buried Gallo-roman
80 walls at the site of Vieil-Evreux (Eure, France).

81

82 **Site presentation and reference survey**

83 The *Gisacum* religious center is located 7 km east of the capital city of *Aulerques*
84 *Ebuovices* (now Evreux in Normandy). There, inside a great hexagonal sacred polygon
85 covering 230 ha, stands a series of public buildings and temples, the dwelling and craft
86 activities being situated outside the polygon (Guyard and Lepert 1999). This site is located
87 over a great plateau where the superficial layer corresponds to flint clay. Above this clay, of
88 around 15 Ω .m resistivity, the archaeological remains have a variable thickness and may
89 surpass 100 Ω .m in resistivity (Aubry 2003). No remains are visible at the surface, and only
90 the thermal baths building and the main temple have been excavated.

91 The test took place in the area of the *fanum* (outlined blue rectangle in Figure 1) and to
92 the North over a plot named 'Terre noire'. The thickness of the archaeological layer in the
93 *fanum* remains around 90 cm. The test compares the results obtained using a three-depth
94 multipole array called ARP© (Automatic Resistivity Profiling), to those obtained with the
95 DualEM (see description in the next paragraph).

96 The ARP© is used generally for pedological and archaeological applications (Dabas,
97 2009); it weighs 400 kg and has 8 spike-wheels. The system is pulled with a quad-bike over a
98 grass cover and the measurement location is obtained using a combination of dGPS system
99 and a Doppler-radar system that allows checking the distance along the profile. The electrical
100 map obtained with the ARP© system is presented in Figure 1 where all buried structures are

101 clearly shown as resistive anomalies in black: walls of the Gallo-roman temple (*fanum*), walls
102 around the thermal complex (*peribola*), dwelling houses to the West delimited by roads (*via*),
103 aqueducts to the South bringing water to the thermal complex and to the temple of water
104 (*nymphaea*); see Dabas *et al.* (2005) for a complete explanation of archaeological features.

105

106 **Instrument and tests**

107 *Theoretical responses*

108 The DualEM-421S is a multi-receiver EMI instrument (DualEM sensor manual 2010)
109 operating at 9 kHz frequency. Its main characteristic is the provision of one horizontal
110 transmitter loop (TX coil, Figure 2) with three pairs of receivers. In each pair, the first
111 receiver is horizontal, allowing horizontal coplanar, HCP, and, by rotation of all the
112 apparatus, vertical coplanar, VCP, configuration measurements. The second receiver coil is
113 radial to the transmitter allowing perpendicular, PERP, configuration measurements. The first
114 pair is located at 1m and 1.1m from the transmitter, the second at 2 and 2.1 m and the third at
115 4 and 4.1 m. Measurements are triggered by time (up to 10 measurements by second) and
116 other data are available: temperature inside the tube, voltage of the battery and the two
117 rotation angles (roll and pitch).

118 Considering the usual resistivity range of soils and superficial formations, the
119 instrument belongs to the ‘Low Induction Number’ (LIN) EMI group (this number compares
120 the electromagnetic diffusion range characterized by the skin depth, $\sqrt{\frac{2}{\sigma\mu\omega}}$ where σ is the
121 conductivity, ω the angular frequency and μ the magnetic permeability, to the geometrical
122 dimensions of the considered problem: inter-coil spacing, depths of the targets, layer
123 thicknesses). For this group, the conductivity response is mainly in quadrature with the
124 primary field which facilitates the measurement of the magnetic susceptibility in phase. In
125 Figure 3, we illustrate the theoretical variations of the coil responses (defined as the ratio of

126 secondary field to primary field at the receiver coil, measured in parts per million) versus
127 conductivity or susceptibility for a homogeneous soil. In this figure and all along the text, the
128 responses are calculated without approximation using the complete mathematical expressions
129 of the secondary field (Thiesson *et al.* 2014). The left part shows the variation of the response
130 versus susceptibility (for a soil with a fixed 0.01 Sm^{-1} conductivity). The right part shows the
131 variation of the response when changing the conductivity (for a fixed susceptibility of $40 \cdot 10^{-5}$
132 SI in-phase and 0 in quadrature). The coil height is set to 0.1m above ground level and the
133 frequency to 9 kHz. It can be observed that all the dependences are quasi-linear: the
134 sensitivity of the phase responses to the magnetic susceptibility variations are good,
135 respectively, the sensitivity of the quadrature responses to the soil conductivity are also good.
136 But, as frequency, soil conductivity or coil separation increases the quadrature response does
137 not remain linear. For each transmitter-receiver separation, the conductivity limit above which
138 the discrepancy with a linear response surpasses a 10% threshold is illustrated in Table 1 for
139 both HCP and PERP geometries (for a 0.1 m height above a homogeneous ground). It must be
140 underlined that the HCP configuration is far more affected by this limit than the PERP one.
141 Consequently, it could be stated that, (except for 4m HCP pairs) in non-salted soils, the
142 dependences remain linear while this relationship is altered in saline environments (Beamish
143 2011). It must also be noted that for greater separations (4m), the amplitude of the responses
144 generated by the soil conductivity are significantly higher than those generated by the soil
145 susceptibility (Fig. 3): this would necessitate a very accurate phase separation of the
146 instrument responses.

147 The absolute calibration was controlled by measuring the DualEM response to a small
148 metallic conductive and non-magnetic sphere as a function to the height and distance to the
149 transmitter coil. Comparing the theoretical response of the sphere (not detailed in this paper,
150 see Thiesson *et al.* 2014) with the measured in-phase response (Figure 4a) and-b)), it appears

151 that most of the points exhibits an error less than $\pm 5\%$ except for the highest values where we
152 can hypothesize some non-linearity in the electronics.

153 Using a metallic sphere, no response should be observed in the three quadrature
154 components. Figure 4c and 4d show that the phase separation of the devices is better than
155 10% (except for the points in the dashed area but which are mainly located in the area of very
156 low amplitudes where noise is predominant). These two tests have shown that the device used
157 is reliable and that the measurements are accurate.

158 The 6 receivers correspond to six different depths of investigation that permit a 6
159 points 'geometrical sounding' at each measurement location. Consequently, the DualEM-421
160 can be used in a wide variety of near surface applications: soil, environmental, engineering
161 and archaeological studies. As with other Slingram instruments (Tabbagh 1986(c)), it also
162 allows detecting buried metallic objects. Several studies have already been published in
163 archaeology and soil studies (Simpson *et al.* 2009, Monteiro Santos *et al.* 2010, Saey *et al.*
164 2012, De Smedt *et al.* 2013) but none of them discussed in detail the ability of the instrument
165 for 2D/3D resistive features characterization.

166 The values displayed by the instrument are specific to the sampling volume of each
167 transmitter-receiver configuration, but the volumes are not the same for susceptibility as for
168 conductivity measurements. For the 6 in-phase channels, these values are directly expressed
169 as the ratio of the secondary field to the primary field in ppt (part per thousands). For the 6
170 quadrature channels, they are transformed to pseudo electrical conductivity by multiplying

171 each channel value by: $-\frac{4}{\mu_0 \omega L^2}$, where ω is the angular frequency, μ_0 the vacuum magnetic

172 permeability and L the separation between the coils corresponding to each channel. This
173 formula is valid only for coils at $h=0$, *i.e.* for a center of coils at ground surface. Even if the
174 instrument is laid on the ground, this hypothesis is not valid (the distance between the center

175 of the coils and the bottom of the boom is 4,5cm). When the instrument is towed on a cart, the
176 distance to the ground is 31.5cm.

177 Consequently:

178 (1) We do not use the pseudo electrical conductivities (output of DualEM), but the quadrature
179 magnetic field ratios using Table 2 (multiplication by a series of coefficients deduced from

180 $-\frac{4}{\mu_0 \omega L^2}$; also quoted in DualEM-421S User's manual, 2010),

181 (2) One must keep in mind that the difference between the output of DualEM and the
182 apparent conductivity is more important in HCP than in PERP and that this importance
183 increases with the average conductivity of the soil (clayey and even more salted soils). This is
184 a consequence of: i) greater non-linearity of the responses when the ground conductivity
185 increases, ii) the elevation of the instrument above ground level. For example, when
186 considering a 20 mS/m ground, the instrument at h=0.045 m height will measure in the HCP 4
187 m channel -5.340 ppt in quadrature and thus display 18.79 mS/m while at h=1 m (carried by
188 the operator) it will display 15.70 mS/m (-4.464 ppt). Under the same conditions, the PERP
189 4.1 m channel would deliver a -5.814 ppt at h=0.045m and display 19.5 mS/m while at h=1 m
190 it would measure -3.328 ppt and display 11.15 mS/m.

191 Therefore, it will be necessary to retranslate the displayed pseudo conductivity values
192 into field ratios when attempting to retrieve quantitative results (following Thiesson *et al.*
193 2014).

194

195 *Field stability tests*

196 A cart was designed, Figure 5, to pull the DualEM-421S with a quad bike on the field
197 at two possible clearances, 0.10 and 0.27 m from the bottom of the tube to the ground surface.
198 The distance between the quad-bike and the cart was tested so that no disturbance was
199 measured due to the presence of the quad-bike with engine switched off.

200 In order to measure the potential noise due to the engine of the quad-bike, DualEM
201 measurements were recorded continuously (10 Hz) during a few minutes with and without
202 engine working (Table 3). It can be observed that the engine has no effect and that the
203 standard deviation remains limited. PERP 1.1m configuration shows a much higher deviation
204 for an unknown reason. This experiment shows also that no “short time” drift (defined as a
205 variation within a 5 minutes time lapse, the time that would be needed to acquire for example
206 a 300m profile at 1m/s) was noticed. “Long term” drift (defined as the variations within the
207 the entire time scale of a survey, typically several hours) was not studied here and is the
208 subject of numerous papers ((Delefortrie *et al.* 2014, Dos Santos and Porsani 2011, Beamish,
209 2011). In archeology, we found that only “short time” drifts are really a problem because they
210 create artefacts in the maps that superpose to the short wavelength of archaeological
211 anomalies (typically several meters, that would correspond to a few seconds in the time
212 domain when using a speed of acquisition of the order of 1m/s). Long time drift, if ever
213 noticed, is easily corrected by standard image filtering procedures.

214 The roll is another important aspect when the instrument is pulled on uneven terrain (the pitch
215 angle also acquired with DualEM was not considered in this study). Figure 6 shows the
216 theoretical relative quadrature out of phase signal change versus roll (grey lines). The
217 maximum roll acceptable is 10 degrees (if we consider a maximum 2% deviation in the
218 quadrature out of phase signal) which defines the limit that must be respected in the field, but
219 PERP is more sensitive to the roll effect than HCP and there is a negative shift of the coil
220 orientation by reference to the vertical. This test also allows verifying the exact coil position
221 inside the housing tube. From the theoretical and practical point of view, the HCP
222 configuration is less sensitive to roll because, for an α roll angle, the HCP emission is reduced
223 by $\cos\alpha$ and the received field is also reduced by $\cos\alpha$ thus the HCP response is reduced by
224 $\cos^2\alpha$, but the rotation generated also a VCP response which, following the same reasoning, is

225 multiplied by $\sin^2\alpha$. As for small values of h , HCP and VCP quadrature responses are very
226 close (the lower the instrument, the closer their values), the roll effect is very limited. On the
227 contrary, in PERP configuration only the transmitter coil is tilted and the response is reduced
228 in $\cos\alpha$ (Figure 6). At least, as the spacing increases, it appears that the theoretical HCP
229 sensitivity to roll decreases, but practically appears another effect (probably due to some lack
230 in the stiffness between the boom and the sensor itself). Fortunately, the roll variation of the
231 signal with the PERP appears to be very repeatable and therefore could be corrected from the
232 value of roll measured.

233

234 *Comparison with vertical electrical soundings*

235 Six DC resistivity Vertical Electrical Soundings (VES) have been achieved to
236 recognize the vertical layering of the site (Figure 7). All exhibit the same distribution: a three
237 layer model, where, beneath a topsoil and an 'archaeological' layer of around 1m thickness
238 and [35-40 Ωm] resistivity, exists a conductive flint clay layer of thickness varying between 2
239 and 3 m and of resistivity between 10 and 20 $\Omega\text{.m}$ above a saturated chalk third layer of [30-
240 40 $\Omega\text{.m}$]. At the same locations, the conductivity responses of DualEM-421S were recorded
241 versus the height above ground surface. This test between 0.045m and 2.25m allowed, for
242 each geometrical configuration, both to assess the noise level (by the interquartile distance)
243 and to compare the displayed pseudo-conductivity values (transformed in ppt using the Table
244 1 coefficients) with the theoretical responses calculated from the VES interpreted parameters.
245 In Table 4 the comparison for the VES number 3 is detailed where $\rho_1=41.3 \Omega\text{m}$, $e_1=0.6 \text{ m}$,
246 $\rho_2=16.5 \Omega\text{.m}$, $e_2=2.06 \text{ m}$ and $\rho_3=26.9 \Omega\text{.m}$.

247 It can be observed in Table 4 that firstly the dispersion of the measurement is limited
248 to 1% of the response and independent of both the height above the ground surface and of the
249 receiver orientation. Secondly, the correlation with the theoretical data is very good for the

250 relative variations but exhibits different offsets for the different channels (if one hypothesizes
251 that the gain is the same for the different channels and is accurate). Thirdly, the comparison
252 with the approximate cumulative responses (McNeill 1980) confirms that this approximate
253 model is not relevant for quantitative interpretation.

254

255 *Reminder about the investigation depths*

256 We do not consider here the magnetic susceptibility (in-phase responses) changes
257 which are known to correspond to clearly lower depths of investigation than conductivity
258 changes (Scollar *et al.* 1990).

259 The quadrature measurement sensitivity of the HCP configuration to deep horizontal
260 layering is usually recognized as higher than that of the PERP configuration. However, the
261 comparison of their respective abilities is more complex for a 2D/3D structure: for example
262 the ability of the PERP configuration is generally higher than that of the HCP in the 3D case
263 (Tabbagh 1986(b)). Consequently we have decided not to use a definition which is linked to a
264 1D case and to approximate sensitivity functions using arbitrary levels (70% of the total
265 sensitivity for DOE, ‘depth of exploration’, and 50% for the DOI50, DualEM manual) even if
266 these latter permits very simple calculations. We think that the notion of investigation depth
267 we have chosen is of better use for surveyors as it is linked to the 3D case, to the full
268 calculation of the sensitivity functions, and to a standard level of ‘noise’ of 10%. Depth of
269 investigation is defined as to the depth where a change of more than 10% in apparent
270 conductivity can be measured when changing the target is true electrical conductivity.

271

272 **DualEM results over Terre noire (wide mesh): spatial variations observed in the**
273 **superficial geologic context (1D inversion)**

274 In the DualEM–421S, the largest inter-coil separations allow a better characterization
275 of the superficial geological context (the first 5 m approximately) than presently available
276 mobile multi-pole resistivity systems (limited to approximately the first 2m). The 1D
277 inversion process has been applied with good results in a wide variety of near-surface studies
278 (Bendjoudi *et al.* 1998, Bobée *et al.* 2010). Even if it blurs the exact location of lateral limits
279 of the features (Guérin *et al.* 1996), it is very relevant for determining the thicknesses of the
280 different layers and their low frequency spatial variations, and in archaeological sites to
281 evaluate the bulk thickness of the remains.

282 The 1D inversion (program QwInv1D, UMR METIS, using an iterative least square
283 inversion (Guérin *et al.* 1996)) is here applied to the wide mesh (5m) EM data acquired in the
284 ‘Terre noire’ plot (Fig. 1) with nine different data series corresponding to HCP 1m, VCP 1m,
285 PERP 1.1m, HCP 2m, VCP 2m, PERP 2.1m, HCP 4m, VCP 4m and PERP 4.1m (all recorded
286 at 0.27m elevation and acquired continuously in two successive surveys). The original
287 pseudo-conductivity data were pre-processed to transform them into magnetic field ratio at a
288 height of 0.27m. Then, at each point of the surveyed area, the 9 experimental ratio values
289 were inverted to obtain the 5 following parameters: resistivity and thickness of the first layer,
290 resistivity and thickness of the second layer and resistivity of the third layer. This process was
291 done for the whole plot (3906 points for the 5 x 5 m² mesh). With a modal value of 6% for the
292 relative RMS error, the modal values of the first layer parameters are 41 Ω m and 1.0 m, those
293 of the second layer are 13 Ω m and 2.3 m and the resistivity of the third layer 39 Ω m. The
294 spatial variations of these parameters are presented in Figure 8. The first layer has a limited
295 resistivity which can be interpreted by an absence of stone remains except in the northern part
296 where such remains can be expected. A marked deep resistivity variation is present in the
297 second and third layer; it probably corresponds to a huge change in the chalk hardness that
298 outcrops with a reduction of the second layer thickness.

299 To take into account the fact that, with the DualEM 421S, VCP configuration would
300 necessitate a second survey, we repeat the inversion with 6 data only corresponding to HCP
301 and PERP configurations. The results obtained are quasi –identical to those of the previous 9
302 data inversion with a modal value of 5.5% for the relative RMS error, 41 Ωm and 1.02 m for
303 the first layer 12.3 Ωm and 2.3 m for the second layer and 40 Ωm for the third.

304

305 **DualEM results over the Fanum (fine mesh): detection of 3D resistive targets (3D**
306 **inversion and comparison with DC maps)**

307 An extensive survey was done by pulling the DualEM on its cart with a 0,1m clearance (only
308 in HCP configuration for the TX coils) over the *fanum*. Acquisition is in a continuous mode
309 (10Hz) following parallel lines separated by 0.5m and in zig-zag mode (10m between the
310 blocks of lines). Average velocity being 6.7 km.h⁻¹, the distance between acquisition points is
311 0.19 m. Positioning of data is done by a GNSS system using broadcasted differential
312 corrections from a base station situated less than 200m from the *fanum* area. The horizontal
313 relative accuracy is estimated around 1cm. The survey was done in 2 hours and 1 minute
314 (distance covered: 12780m; area 0,52ha). The statistics for the roll angle are such that no
315 correction was applied (mean=-3.8°, interquartile distance=7.7°).The change of temperature
316 was less than 2°C during the two hours survey (median value =24.1°C). Positioning of the
317 data was done knowing the distance between the GPS and the different coils (Figure 2) and
318 also the time lag between DualEM and GPS. Data, converted to apparent electrical resistivity
319 for comparison with DC maps, were re-interpolated over a square mesh of 0.2 x 0.2m using a
320 bi-cubic spline function. We have decided not to use any 1D or 2D filters for these data in
321 order to show the raw data.

322

323 *Qualitative comparison of EM and DC maps*

324 In Figure 9 are presented the 6 apparent resistivity maps obtained with the DualEM,
325 together with the two apparent resistivity maps obtained with ARP©. As stated before, no
326 filtering was undertaken to generate these maps. No outlier points exist for the 4m coils. The
327 first outliers appear for the HCP2m coil and for both 1m coils. HCP configuration seems to be
328 more sensitive to these shallow outliers. We think that most of these outliers come from small
329 metallic objects spread all over the surface (the in-phase maps, not shown in this article, show
330 high amplitude anomalies at the same position; this area is very close to an air-force base that
331 was bombed during last world war). These outliers can be easily filtered by a 2D median
332 filtering. Some “long term” drift does appear (see the ‘bands’ in PRP 4.1m for example) in
333 these maps but does not affect the visibility of archaeological features in this case.

334 It can be observed that:

- 335 (1) Except for the larger coil separation in EMI, all the main resistive features are clearly
336 delineated even if the apparent resistivity anomaly magnitudes are more limited in
337 EMI than in DC.
- 338 (2) The apparent resistivity values are in good coherence with the depth of investigation
339 expected from sensor geometries. The 1D point by point inversion applied in that area
340 with a three layer model delivers the following mode values: 70 Ωm and 0.2 m for the
341 first topsoil layer, 26 Ωm and 1.3 m for the ‘archaeological’ layer and 20 Ωm for the
342 flint clay.
- 343 (3) As expected from both theoretical (Tabbagh 1986(b)) and physical (Frischknecht *et al.*
344 1991) models, the HCP responses are oscillating and exhibit two parallel resistivity
345 maxima beside each wall and a stronger minimum just above it. This ‘triple arch’
346 response would be confusing when direct reading of the map in absence of 2D/3D
347 modelling reference, but it does not impede the interpretation (see below).

348 (4) In agreement with the expected depths of the walls (tops at around 0.3m), the ‘best’
349 results are given by PERP 1m configuration.

350

351 *3D interpretation of the wall of the Cella: quantitative comparison of inverted 3D DC data*
352 *and DualEM data*

353 To go further in the comparison, one can apply over a limited area, the two-step
354 1D/3D rapid inversion described in (Brinon *et al.* 2012). This process is based for 1D on an
355 analytical forward modelling and for 3D on the moment method. The latter is relevant to
356 quantify the responses of simple-shape features in both EM modelling, for resistivity and
357 magnetic susceptibility contrasts (Tabbagh 1985) and in DC resistivity modelling (Dabas *et*
358 *al.* 1991). This approach allows a direct comparison of the resistivity contrasts of the detected
359 features.

360 We consider, for example, a small area limited by X= 519018.56 and 519023.36 in
361 easting and Y= 145232.66 and 145236.26 in northing, delineated in Figure 9, where can be
362 observed the external *cella* wall. Using the data of the three channels of the ARP© one
363 obtains for the surrounding terrain 1D model $\rho_1=70 \Omega\text{m}$, $e_1=0.2 \text{ m}$, $\rho_2=32 \Omega\text{m}$, $e_2=1.3 \text{ m}$ and
364 $\rho_3=19 \Omega\text{m}$. For the wall, the inversion gives an orientation at 100° from X axis, a section of
365 $1.0 \times .88 \text{ m}^2$ for a center at X=519020.96, Y=145234.46 and Z=0.7m and a resistivity of 70.5
366 Ωm . The contrast ratio between the wall and the surrounding is thus 2.2.

367 Adopting these values, it is possible to calculate using the moment method the
368 responses that would be obtained for the different DualEM coil configurations. Due to the
369 depth of the wall, the calculations are limited to HCP 1m, HCP 2m, PERP 1.1 m and PERP
370 2.1m. The results of the comparison with the DualEM measured profile are presented in
371 Figure 10 (Hs/Hp values in ppt) for the four coil configurations. The agreement is good both

372 for magnitude and width. It also confirms the anomaly shapes with a single maximum in
373 PERP configuration and the three arch responses for HCP configuration.

374 In this example, it is clear that the walls are detected by the four coil configurations as
375 resistive features, and would be correctly interpreted using only one of each. This result
376 confirms what has been already established (Thiesson *et al.* 2009) for a smaller instrument in
377 VCP configuration. In fact, the relative variation of amplitude of the responses above the wall
378 is rather limited, 10% for the PERP and about 7% for HCP compared to 25% observed in the
379 ARP© channels. However, here again, the relevant criterion is the signal to noise ratio, i.e. the
380 relative sensitivity of instruments to the searched-for features as opposed to the unwanted
381 features ('geophysical noise'). In the Vieil-Evreux site where the global resistivity is of the
382 order of 100Ωm and lower, the signal to noise ratio can be comparable to that of the DC
383 resistivity method.

384

385 **Conclusions**

386 We presented here a test of a single frequency domain EMI instrument over an
387 archaeological site in order to better grasp the capabilities of this exploration technique. Even
388 though this technique has been used for fifty years in archaeological survey and is sensitive to
389 both magnetic susceptibility and electrical conductivity, it has been, and still is, less applied
390 than magnetic and earth-resistance prospection.

391 The Vieil-Evreux roman site remains (mainly walls) are situated in a sedimentary (clay)
392 geological context where the conductivity response is significantly high; consequently the two
393 major limitations of EMI cannot be considered in the present study but they merit to be
394 recalled in this conclusion because they, at least for a part, explain the restricted use of the
395 technique. They are: (1) metallic objects and features disturb measurements and practically
396 disqualify its application specially in urban contexts and (2) in highly resistive soil contexts

397 the conductivity response is too small and difficult to separate from the quadrature magnetic
398 susceptibility in the quadrature response (Tabbagh 1986(c)) in spite of recent research works
399 aiming to overcome this problem (Simon *et al.* 2015).

400 The series of tests conducted with the DualEM in Vieil-Evreux have allowed verifying that
401 the electronic and external electromagnetic noises are fairly rejected, the phase of the received
402 signals are well defined, the absolute calibration is fine, the drifts limited and the roll effect
403 characterized. Measured quadrature data are in full accordance with the theoretical complete
404 calculations taking the true resistivities and layer thicknesses as input data (from separate
405 VES). Using this instrument, it is thus possible to assess the advantage and drawbacks of the
406 method itself within the particular objective of the detection of stone built resistive features.

407 The first point to be emphasized is that frequency domain multi-receiver EMI instruments,
408 lighter and easier to move than multi-pole DC, can be used to characterize the superficial
409 geological context with a sufficient depth of investigation. This aspect is important for
410 extended archaeological sites as it allows a first rough estimation of the volume of the
411 archaeological remains. The second point is the ability to detect and characterize walls. They
412 are detected with 1m and 2m coil separations and the corresponding apparent resistivity maps
413 clearly show the plan of the fanum at Vieil-Evreux. The corresponding signal variations are
414 lower than with DC resistivity but as the geophysical noise is also lower, the observed signal
415 to noise ratios are roughly equivalent. Nevertheless, referring to 3D theoretical modelling is
416 required to correctly interpret the HCP raw data.

417 Finally the Vieil-Evreux test verifies that with sufficient fine meshing and a close referencing
418 to modelling, stone built resistive features can be surveyed with frequency domain multi-
419 receiver EMI.

420

421

422

423 **Acknowledgements**

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426 *and Scott Holladay for their corrections of this manuscript and to the reviewers for their*
427 *helpful contributions.*

428

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527

528 **Figure captions**

529 Figure 1: Composite apparent resistivity map of Vieil-Evreux Gallo-Roman site (RM15-
530 MuCEP and ARP®channel2); ARP® system in the upper right with array configuration;
531 studied area outlined in blue.

532

533 Figure 2: General scheme of coil configuration in the DualEM-421S instrument (cm units).

534

535 Figure 3: Theoretical in-phase (continuous lines) and quadrature responses (dotted lines)
536 versus magnetic susceptibility (left side) and electrical conductivity (right side) for the
537 different HCP and PERP distances from the transmitter in the DualEM-421S.

538

539 Figure 4: a) Results of the calibration with the sphere at $z=60$ cm above the device, b) Results
540 of the calibration with the sphere at $z=23$ cm above the device
541 c) Phase separation for the Perp geometries d) Phase separation for the HCP geometries

542

543 Figure 5: DUALEM-421S on its cart pulled by a quad.

544

545 Figure 6: Relative quadrature out of phase signal changes (theoretical: grey line, practical
546 dark diamonds) as function of roll angle (in degree) a) HCP 4m, b) PERP 4.1m, c) HCP 2m,
547 d) PERP 2.1m, e) HCP 1m, f) PERP 1.1m.

548

549 Figure 7: Position of the 6 VESs and example of inversion for VES N°3

550

551 Figure 8: Terre noire plot: spatial variations of the 5 parameters inverted from the 9
552 quadrature data.

553

554 Figure 9: Apparent resistivity maps over the *fanum* for the 6 different DualEM-421S
555 quadrature channels and two of the ARP© channels (3D inversion area outlined in blue).

556

557 Figure 10: Comparison between the experimental (bold line) values obtained over the small
558 test area and the responses calculated using the result of the ARP© data inversion (thin line).

559

560

561 **Table captions**

562 Table 1: Conductivity values above which the quadrature response differs with more than
563 10% from linear response.

564

565 Table 2: Coefficients allowing the computation of magnetic field ratio.

566

567 Table 3: Measurement dispersion (resistivity in Ωm) at a fixed point with (a) and without (b)
568 engine-turned on.

569

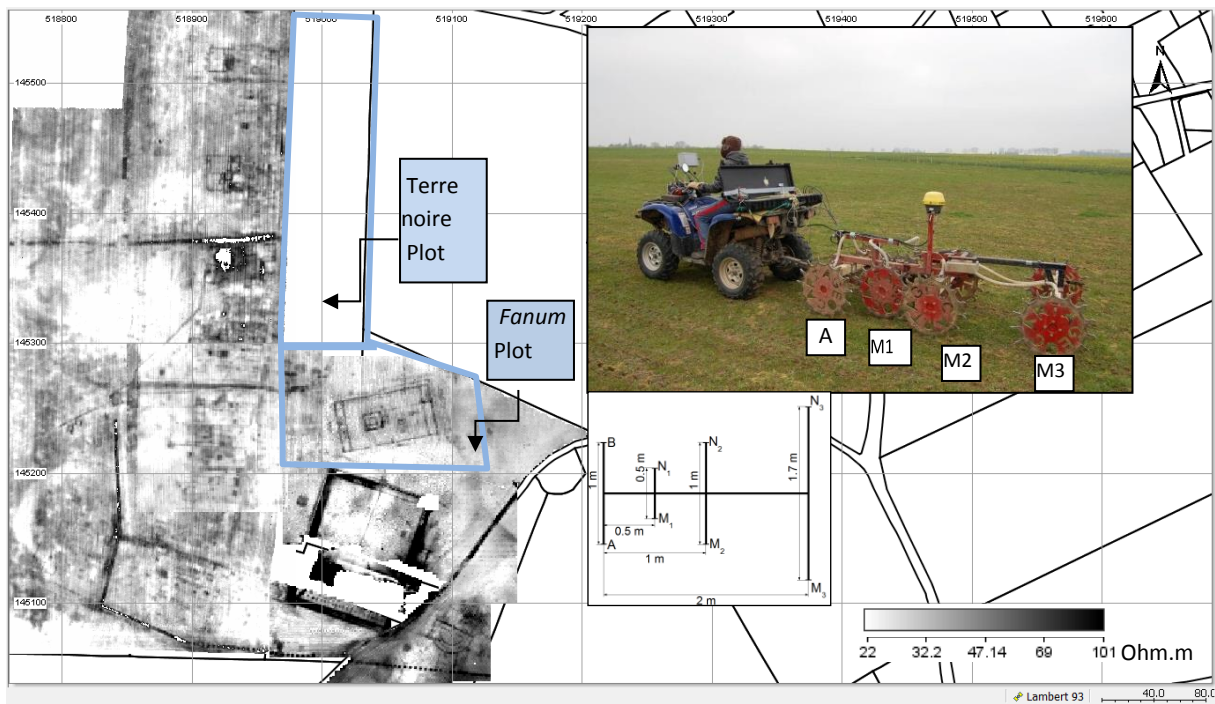
570 Table 4: Experimental dispersion and median values observed versus instrument height above
571 the ground surface at VES 3 location and comparison with the values deduced from the VES
572 three layer interpretation where $\rho_1=41.3 \Omega\text{m}$, $e_1=0.6 \text{ m}$, $\rho_2=16.5 \Omega\text{m}$, $e_2=2.06 \text{ m}$ and $\rho_3=26.9$
573 Ωm using the complete expression of the secondary field (Thiesson *et al.* 2014) and using the
574 approximate cumulative sensitivity response (Wait 1962, Mc Neill 1980). The order of the
575 coil geometries is the following: HCP 1 m, PERP 1.1 m, HCP 2 m, PERP 2.1 m, HCP 4 m
576 and PERP 4.1 m.

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583 Fig. 1

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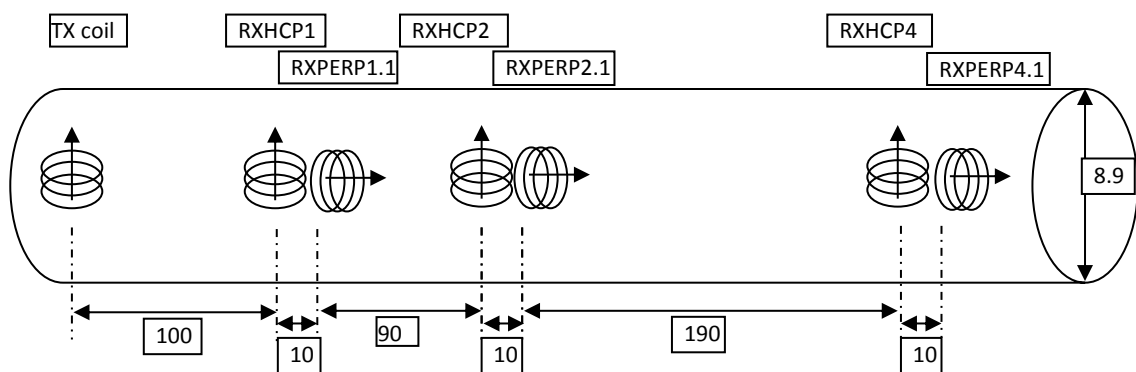
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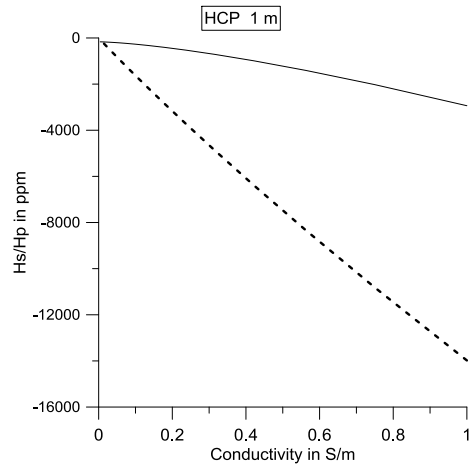
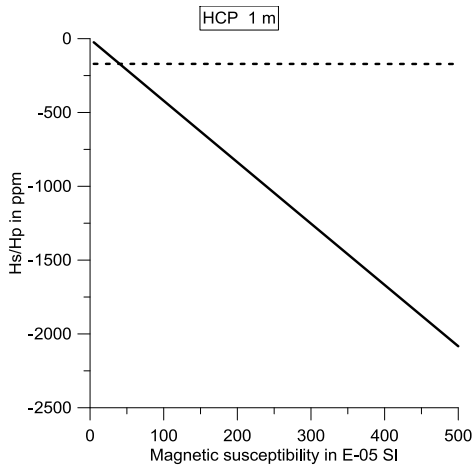
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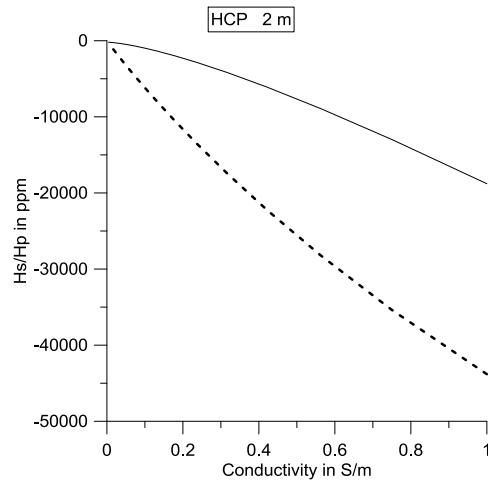
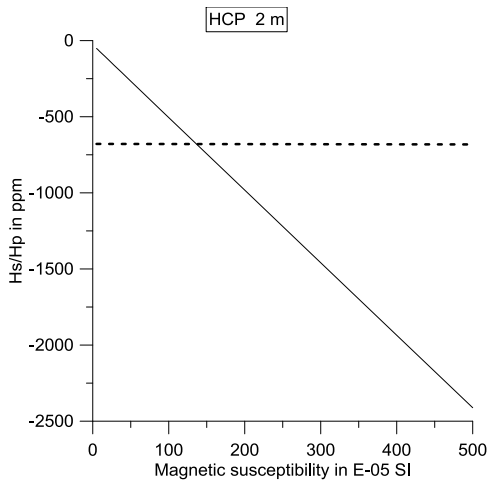
592 Fig. 2

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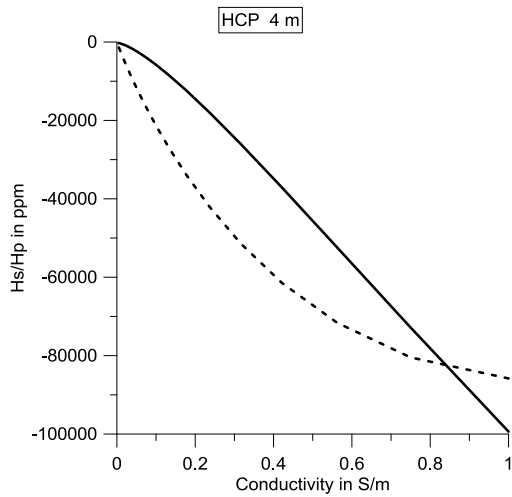
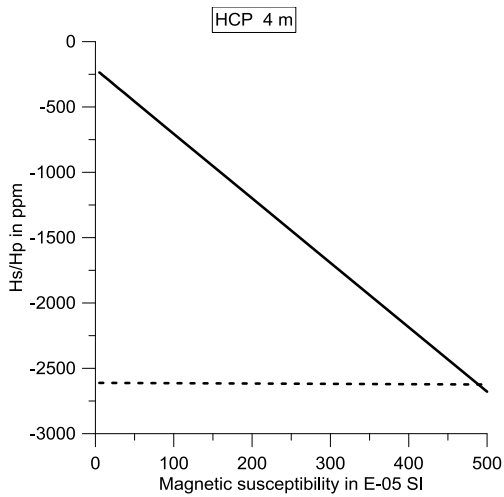
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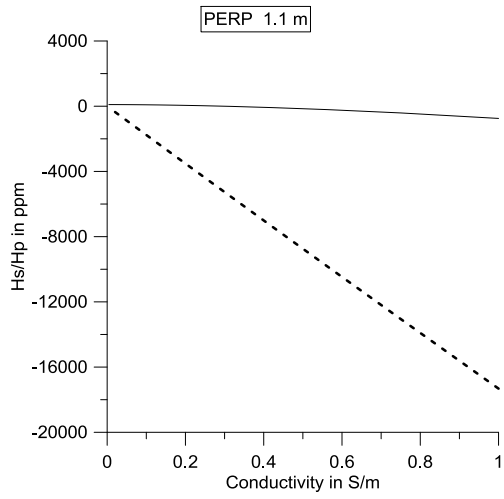
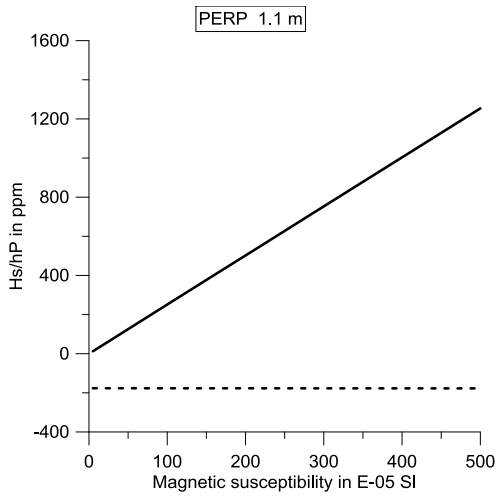
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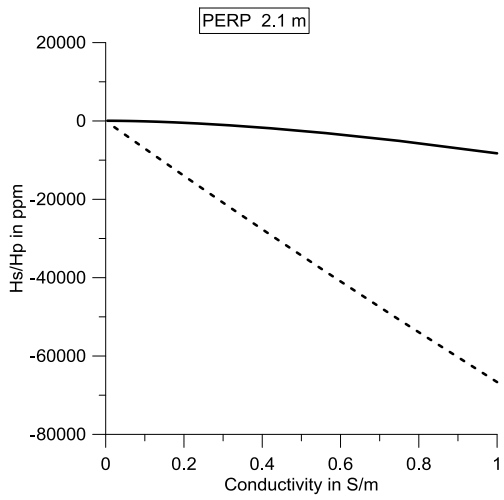
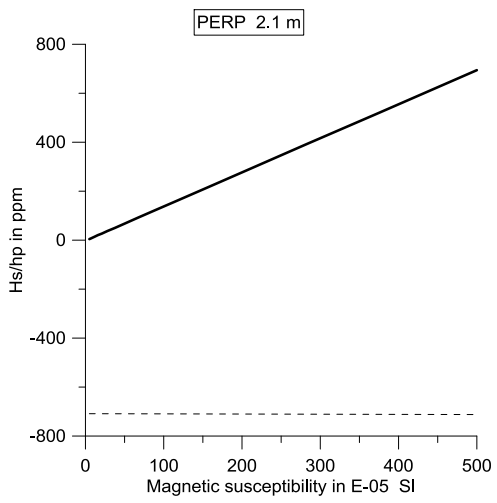
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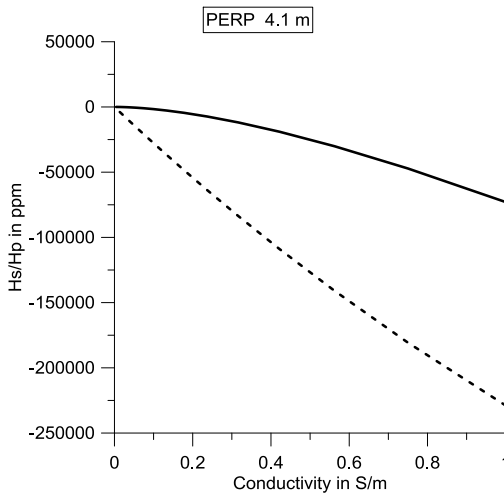
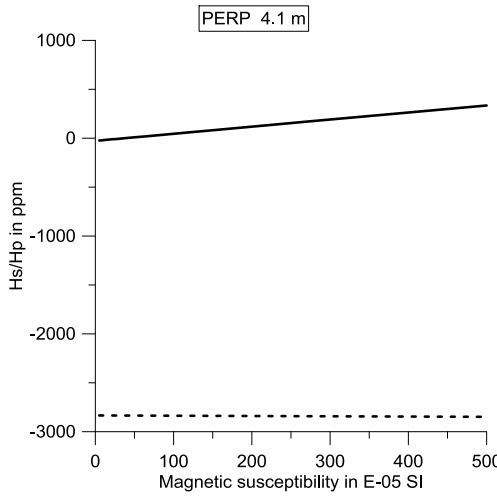
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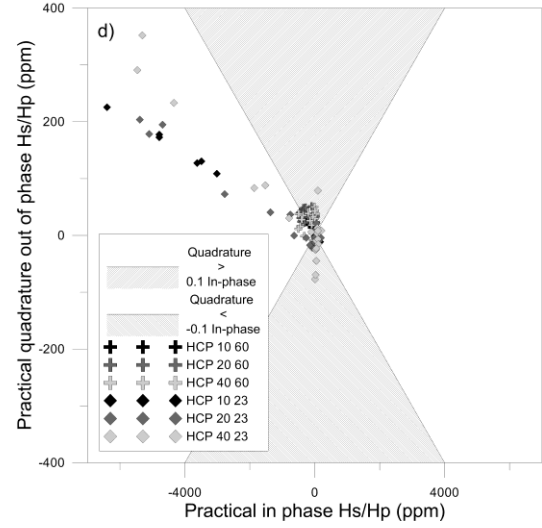
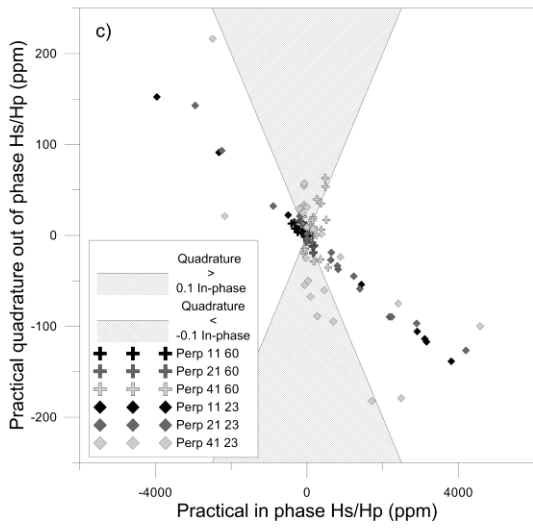
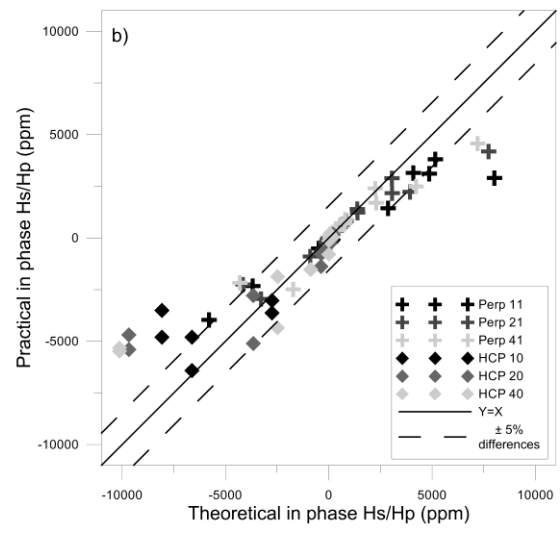
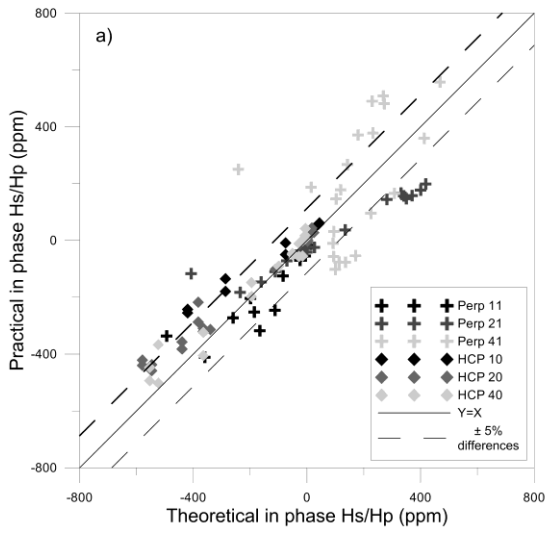


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601 Fig. 3

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606 Fig. 4

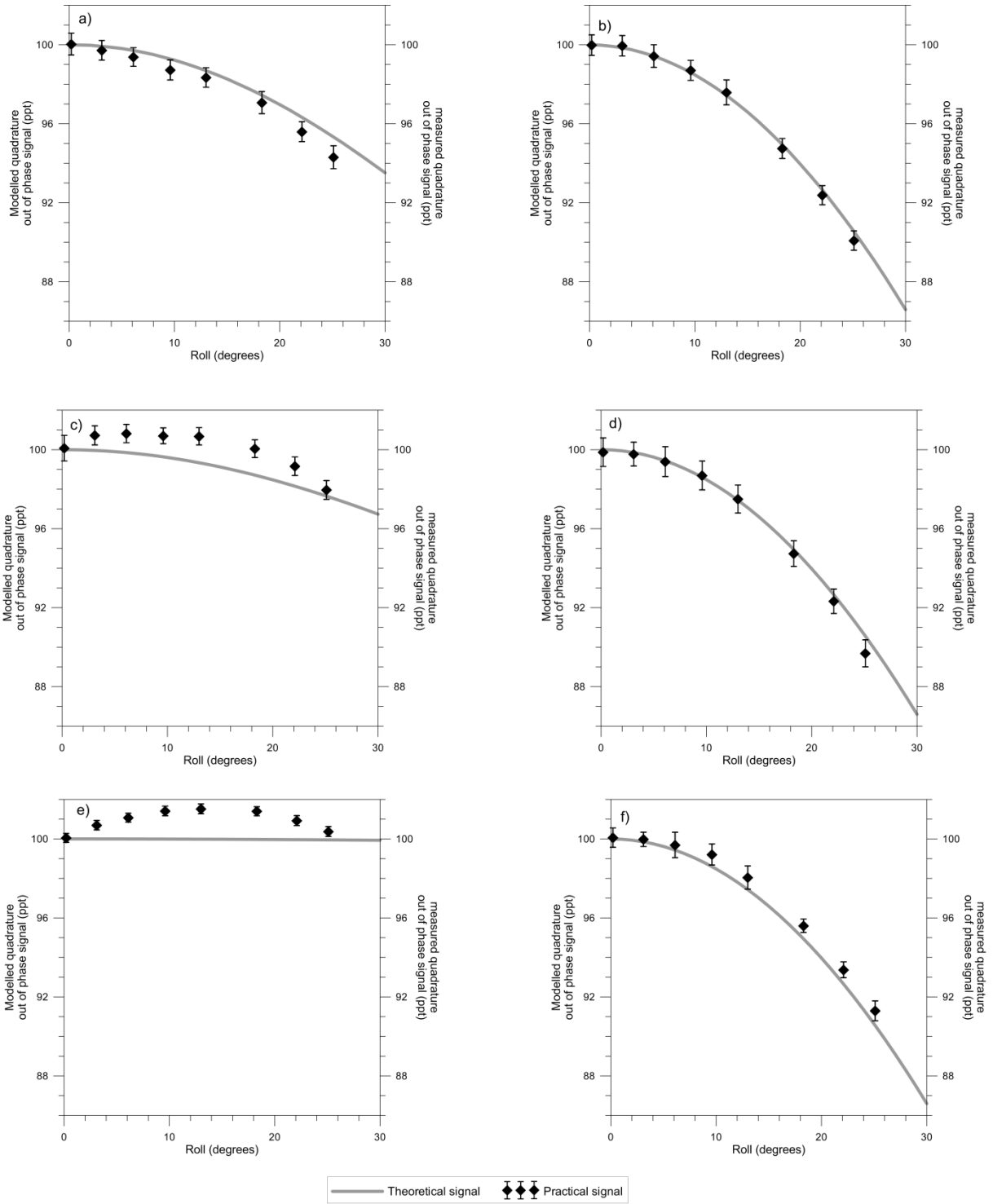
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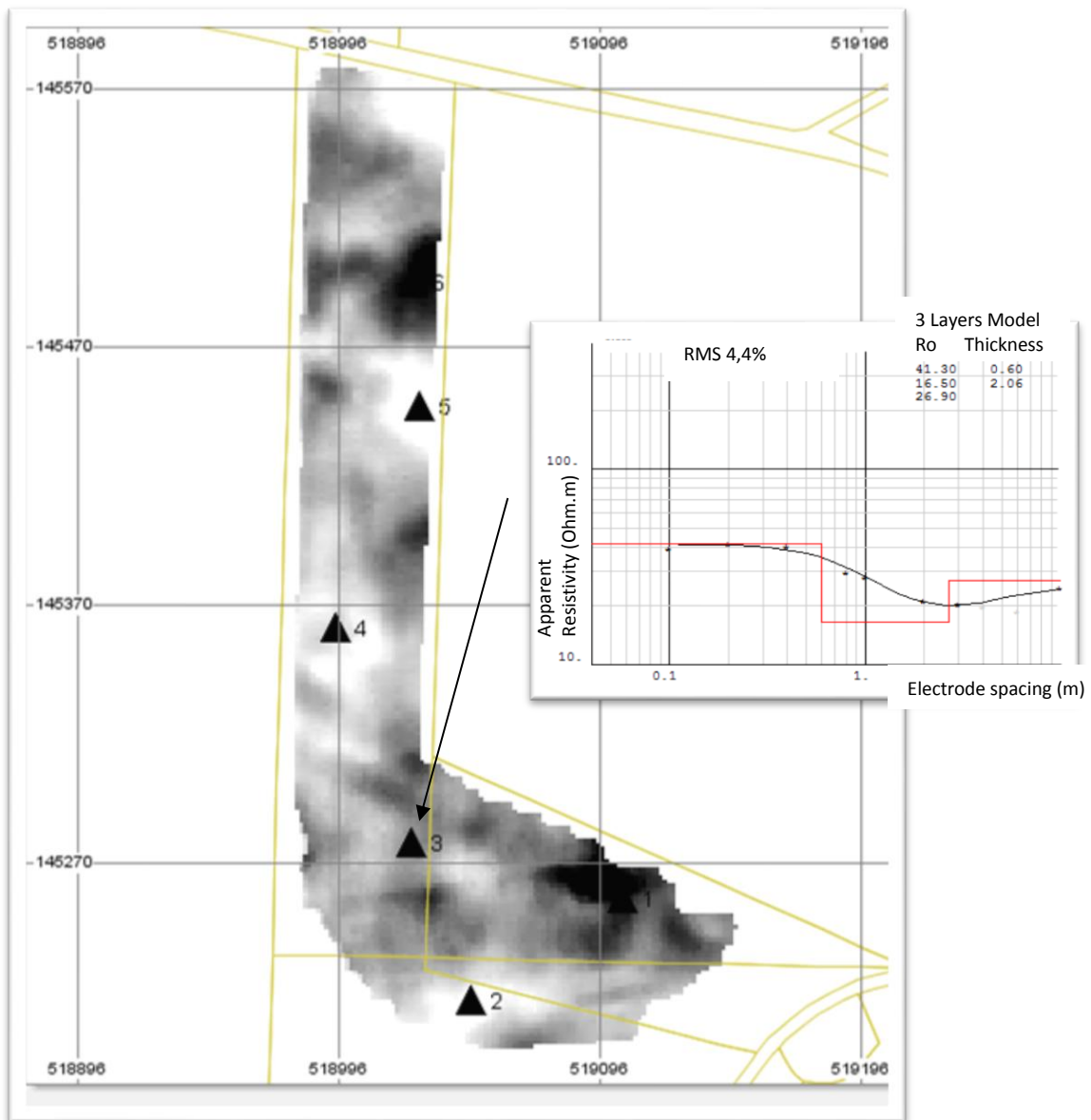
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609 Fig. 5

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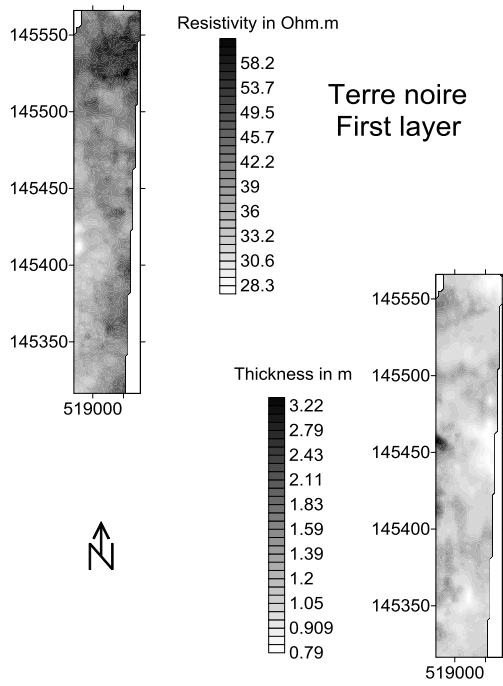


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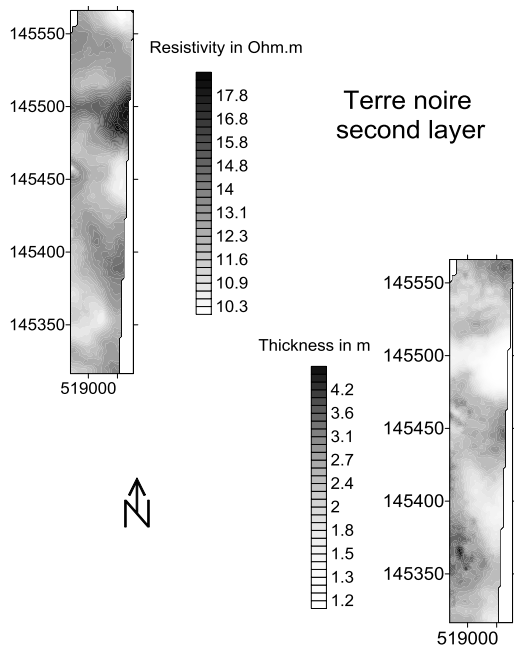
617 Fig.7

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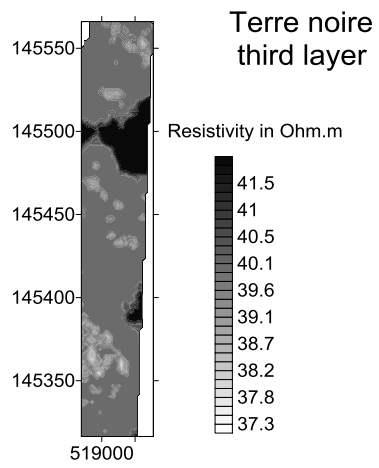
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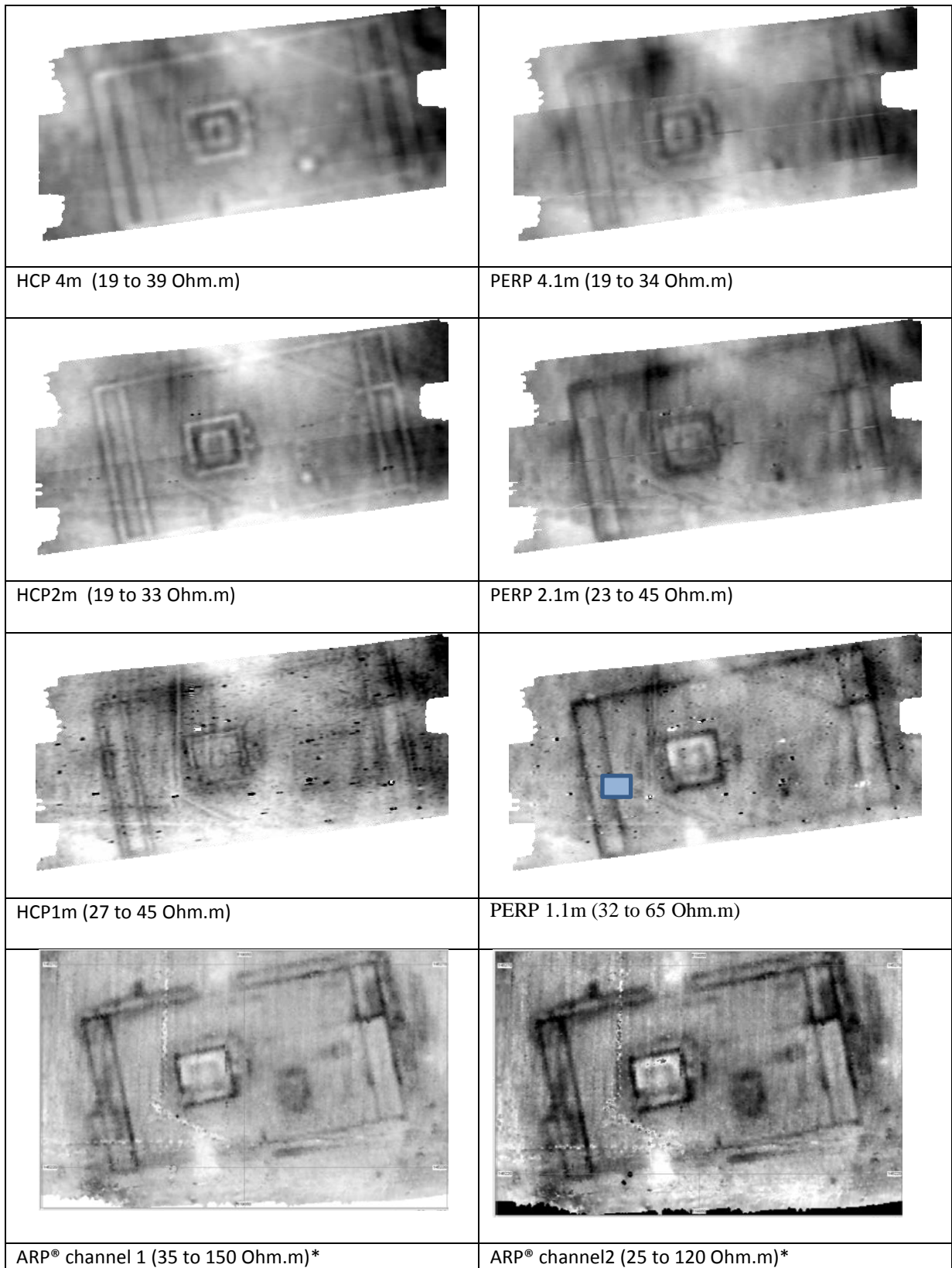
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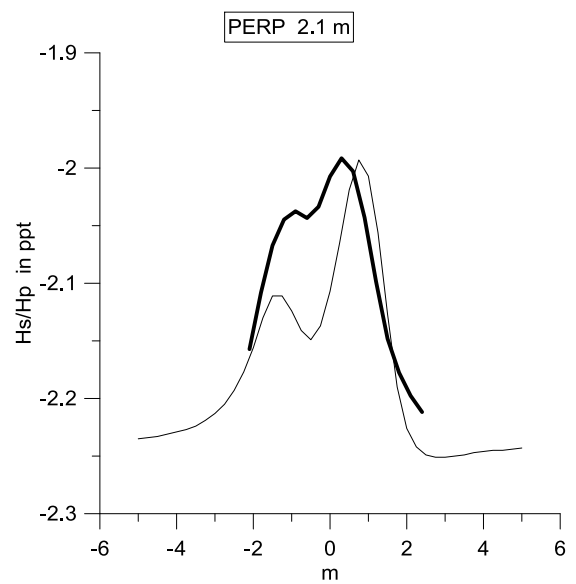
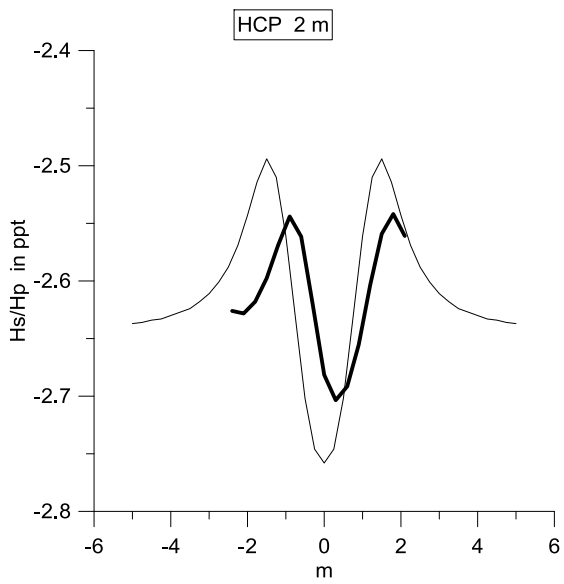
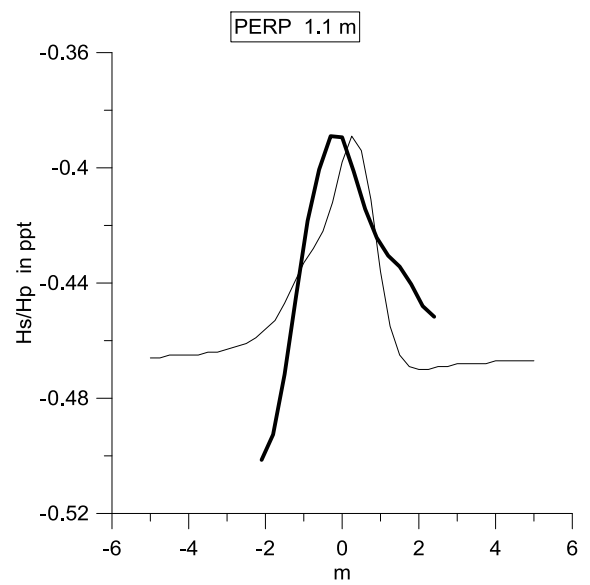
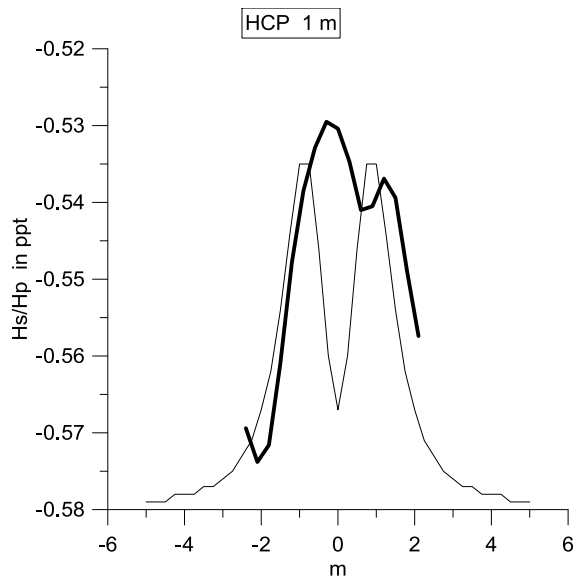


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623 Fig. 8

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628 Fig. 10

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Max	29.1	27.2	28.7	82.6	43.1	333.3
Min	27.6	26	26.2	74.1	39.5	243.9
Average	28.32	26.60	27.53	78.02	41.28	279.7
Variance	0.029	0.033	0.023	1.842	0.150	95.9
Standard deviation	0.17	0.18	0.15	1.36	0.39	9.75

636 (b)

637

638 Table 4

Height above ground surface (m)	Experimental Dispersion Third quartile – first quartile (ppt)	Median experimental value (ppt)	theoretical value deduced from VES data (ppt)	Approximate cumulative responses from VES data (ppt)
0.045	0.005	0.552	0.722	1.169
	0.004	0.516	0.654	0.796
	0.028	2.936	3.098	4.125
	0.024	2.670	3.043	3.667
	0.057	11.246	11.043	10.769
	0.119	12.661	13.151	17.062
0.112	0.005	0.531	0.691	1.143
	0.004	0.449	0.570	0.757
	0.028	2.922	3.036	4.100
	0.031	2.466	2.820	3.577
	0.057	11.161	11.003	10.921
	0.119	12.243	12.662	16.849
0.179	0.005	0.518	0.655	1.113
	0.004	0.400	0.493	0.714
	0.028	2.887	2.966	4.067
	0.031	2.286	2.609	3.480
	0.057	11.076	10.948	11.054
	0.090	11.705	12.180	16.628
0.246	0.005	0.493	0.617	1.081
	0.004	0.355	0.425	0.668
	0.028	2.823	2.889	4.026
	0.023	2.122	2.410	3.378
	0.057	10.962	10.878	11.168
	0.090	11.227	11.707	16.397
0.347	0.004	0.452	0.560	1.031
	0.002	0.284	0.340	0.598
	0.021	2.709	2.765	3.952
	0.031	1.864	2.134	3.214
	0.057	10.792	10.747	11.305
	0.119	10.541	11.016	16.034
0.449	0.005	0.409	0.506	0.983
	0.004	0.224	0.273	0.531
	0.009	2.567	2.633	3.868

	0.031	1.644	1.884	3.045
	0.057	10.565	10.587	11.406
	0.090	9.824	10.343	15.651
0.75	0.007	0.304	0.377	0.872
	0.004	0.120	0.152	0.365
	0.028	2.211	2.239	3.602
	0.031	1.167	1.307	2.546
	0.057	9.883	9.990	11.524
	0.090	8.062	8.538	14.456
1.05	0.005	0.231	0.291	0.801
	0.004	0.073	0.094	0.249
	0.028	1.962	1.889	3.360
	0.023	0.830	0.925	2.099
	0.085	9.088	9.280	11.451
	0.060	6.659	7.020	13.225
1.35	0.005	0.182	0.232	0.756
	0.004	0.049	0.063	0.167
	0.028	1.585	1.600	3.165
	0.031	0.564	0.674	1.723
	0.057	8.293	8.531	11.273
	0.090	5.449	5.772	12.015
1.65	0.004	0.146	0.192	0.725
	0.004	0.032	.044	0.107
	0.028	1.322	1.367	3.013
	0.031	0.446	0.506	1.415
	0.057	7.526	7.795	11.049
	0.090	4.539	4.963	10.869
1.95	0.005	0.117	0.162	0.704
	0.004	0.024	0.033	0.062
	0.028	1.145	1.181	2.898
	0.023	0.353	0.390	1.163
	0.057	6.816	7.103	10.817
	0.090	3.792	3.954	9.810
2.25	0.005	0.098	0.139	0.690
	0.004	0.17	0.026	0.026
	0.028	1.003	1.030	2.810
	0.031	0.290	0.307	0.955
	0.057	7.696	6.467	10.597
	0.090	3.195	3.506	8.845

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