

Melting experiments on the Fe–C binary system up to 255 GPa: Constraints on the carbon content in the Earth's core

Izumi Mashino, Francesca Miozzi, Kei Hirose, Guillaume Morard, Ryosuke

Sinmyo

▶ To cite this version:

Izumi Mashino, Francesca Miozzi, Kei Hirose, Guillaume Morard, Ryosuke Sinmyo. Melting experiments on the Fe–C binary system up to 255 GPa: Constraints on the carbon content in the Earth's core. Earth and Planetary Science Letters, 2019, 515, pp.135-144. 10.1016/j.epsl.2019.03.020. hal-02180578

HAL Id: hal-02180578 https://hal.sorbonne-universite.fr/hal-02180578v1

Submitted on 11 Jul 2019

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers. L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1	Melting experiments on the Fe–C binary system up to 255 GPa:
2	Constraints on the carbon content in the Earth's core
3	
4	Izumi Mashino ^{a,*} , Francesca Miozzi ^b , Kei Hirose ^{a,c} , Guillaume Morard ^b , Ryosuke
5	Sinmyo ^a
6	
7	^a Department of Earth and Planetary Science, The University of Tokyo, Hongo, Tokyo,
8	113-0033, Japan
9	^b Sorbonne Université, Muséum National d'Histoire Naturelle, UMR CNRS 7590, IRD,
10	Institut de Minéralogie, de Physique des Matériaux et de Cosmochimie, IMPMC, 75005
11	Paris, France
12	^c Earth-Life Science Institute, Tokyo Institute of Technology, Meguro, Tokyo 152-8550,
13	Japan
14	
15	*Corresponding author.
16	E-mail address: izumi.mashino@erdw.ethz.ch (I. Mashino, Now at ETH)
17	

18 ABSTRACT

19 Phase relations, including the eutectic liquid composition in the Fe–C binary 20 system, remain unclear under the core pressure range, which makes estimating the 21 carbon budget in the Earth's core difficult. To explore this issue, we have conducted 22 melting and subsolidus experiments on Fe-C alloys in a diamond-anvil cell up to 255 23 GPa. Textural and compositional characterizations of quenched samples show that 24 carbon concentration in the eutectic liquid slightly decreases with increasing pressure 25 and is about 3 wt.% at the inner core boundary (ICB) pressure. The solubility of carbon 26 in solid Fe is found to be almost constant at ~1.0 wt.%. In situ X-ray diffraction data 27 indicate that Fe forms eutectic melting with Fe₃C to 203 GPa and with Fe₇C₃ at 255 28 GPa.

29 Previous studies on liquid Fe-C alloys suggested that the density of the outer core is 30 explained by liquid Fe containing 1.8 to 4.2 wt.% C. If the liquid core includes <3 wt.% 31 C as a single light element, hexagonal close-packed (hcp) Fe crystallizes at the ICB. 32 However, the carbon content in such solid Fe is ≤ 1 wt.%, less than that required to 33 account for the inner core density deficit from pure iron. When the outer core includes 34 \geq 3 wt.% C, it forms Fe₇C₃ at the ICB, whose density is too small for the inner core. 35 Carbon is therefore not a primary light element in the core. Nevertheless, the outer core 36 liquid can be Fe-C-Si, Fe-C-S, or Fe-C-H. Such core liquid crystallizes solid Fe with 37 light elements including less than 1 wt.% C, which may explain the density and the 38 sound velocities observed in the inner core.

39

40 Keywords: core; iron alloy; carbon; eutectic liquid; high pressure

42 **1. Introduction**

Carbon has been believed to be one of feasible light elements contained in planetary
cores because of its high abundance in the solar system and strong affinity to molten
iron (Benz and Elliott, 1961; Anders and Grevesse, 1989; Wood, 1993). Recent firstprinciples calculations predicted that the incorporation of a small amount of carbon in
hcp Fe is compatible with the density and sound velocities observed in the Earth's inner
core (Li et al., 2018).

49 In order to reveal the abundance of carbon in the core, it is important to understand 50 the Fe–C phase diagram at high pressure, in particular at 330 GPa, which corresponds to 51 the conditions at the ICB. Wood (1993) first thermodynamically calculated the phase 52 relations in the Fe-C system up to 136 GPa and predicted that the Earth's core could 53 contain ~4 wt.% C. Nakajima et al. (2009) experimentally determined the melting phase 54 relationships in the Fe-rich portion of the Fe-C system up to 29 GPa and demonstrated 55 that Fe₇C₃ appears as a liquidus phase. Using an *in situ* X-radiographic imaging 56 technique, Lord et al. (2009) found that carbon concentration in the eutectic liquid in the 57 Fe–C binary system decreases substantially with increasing pressure to 70 GPa. The 58 thermodynamic modeling by Fei and Brosh (2014) based on multi-anvil experiments up 59 to 25 GPa predicted that the Fe–C eutectic liquid composition includes 2.2 wt.% C at 60 330 GPa. More recently Morard et al. (2017a) reported the change in carbon 61 concentration in the eutectic liquid in the Fe–Fe₃C binary up to 107 GPa and estimated 62 it to be 2.6 wt.% C at the core-mantle boundary (CMB). However, they obtained liquid compositions from X-ray diffraction (XRD) patterns with the Rietveld method, which 63 64 includes great uncertainty. So far, melting experiments on the Fe-C system have not been carried out under the core pressure range (>136 GPa). Moreover, the solubility of 65

carbon in solid iron also remains unknown above 25 GPa (Chabot et al., 2008; Walker
et al., 2013; Fei and Brosh, 2014).

68 It is also important to verify the change in liquidus phase in the Fe–C system. It has 69 been suggested that the liquidus field of Fe_7C_3 expands with respect to that of Fe_3C and 70 eutectic melting occurs between Fe and Fe₇C₃ above 120 GPa (Lord et al., 2009; Fei 71 and Brosh, 2014). Furthermore, the recent experiments by Liu et al. (2016) showed the 72 decomposition of Fe_3C into $Fe + Fe_7C_3$ above 145 GPa, but it contradicts the formation 73 of Fe₃C at >300 GPa and >5000 K reported by Tateno et al. (2010). Fe₇C₃ has been 74 regarded as a possible constituent in the inner core because it provides a good match for 75 both compressional (V_P) and shear velocities (V_S) (or Poisson's ratio) of the inner core 76 under relevant conditions (Nakajima et al., 2009; Mookherjee et al., 2011; Chen et al., 77 2014; Prescher et al., 2015). The calculations by Li et al. (2016), however, demonstrated 78 that the density of Fe_7C_3 is too small for the inner core.

79 Recent progress in melting experiments in a diamond-anvil cell (DAC) combined 80 with a focused ion beam (FIB) technique has enabled us to determine the eutectic liquid 81 composition to the core pressure range (Ozawa et al., 2016; Mori et al., 2017; Hirose et 82 al., 2017). In this study, we have conducted both subsolidus and melting experiments on 83 Fe-C alloys up to 255 GPa in order to determine the change in melting phase relations 84 with increasing pressure, in particular 1) the stability of Fe_7C_3 at liquidus, 2) carbon 85 concentration in eutectic liquid, and 3) the solubility of carbon in solid iron. Based on 86 these results, we discuss the possible carbon content in the Earth's core and the 87 mineralogy of the inner core.

88

89 2. Experimental methods

90	High-pressure and -temperature $(P-T)$ conditions were generated in a laser-heated
91	DAC using double-, single-beveled, and flat diamond anvils with 60, 90, 150, and 300
92	μ m culet size. Starting materials were foils of Fe containing 1.9 wt.% C and 4.0 wt.% C,
93	which are homogeneous mixtures of fine-grained Fe and Fe ₃ C synthesized by an ultra-
94	rapid quenching method (Morard et al., 2011, 2017a). Rhenium gaskets were pre-
95	indented to about 40 μ m thickness before laser drilling or spark erosion of sample holes
96	with 30–100 μ m diameters. Sample pellets were loaded between two pieces of about 10
97	μ m thick dry Al ₂ O ₃ , which served as a pressure-transmitting medium and a thermal
98	insulator from the diamond anvils. After loading, the whole DAC was dried in a
99	vacuum oven at 423 K for more than 12 hrs, and subsequently a sample chamber was
100	flushed with dry argon and squeezed in an argon atmosphere.
101	After compression to a desired pressure, the sample was heated from both sides

102 with a couple of 100-W single-mode Yb fiber lasers. We used beam-shaping optics, 103 which converts a Gaussian beam to one with a flat energy distribution and thus reduces 104 radial temperature gradient in the sample. A laser-heated spot was up to ~20 µm across. 105 One-dimensional temperature distributions on the surface of the sample were obtained 106 using a spectro-radiometric method (Ohishi et al., 2008). Heating duration was limited 107 to about 5 s in order to avoid fluctuations in temperature, which would otherwise cause 108 complex melting textures. The experimental temperature reported in this study is that at 109 the solid/liquid boundary, which is obtained by combining the measured radial 110 temperature distribution with the melting texture observed in a sample's cross section 111 (Fig. 1) (see Ozawa et al., 2016 for details). We polished the sample using an FIB from 112 both sides little by little (~1 µm for each step) in order to have a cross section at the 113 center of a liquid pool. The sample was thinned to about 3 µm thickness at the end. The

 $\mathbf{5}$

114 uncertainty in temperature is ± 5 % according to Mori et al. (2017), which includes 115 uncertainties in the width of the liquid pool as well as those derived from the facts that 116 we calculated temperature from integrated light not only from the sample surface but 117 also from the pressure medium and that spatial resolution of the temperature profile was 118 2 to 3 μ m.

119 Pressure at room temperature was obtained based on the Raman shift of a diamond 120 anvil (Akahama and Kawamura, 2004). It was then corrected for a contribution of 121 thermal pressure. According to Andrault et al. (1998), when purely isochoric heating is 122 made, the thermal pressure is written as $\Delta P = \alpha K_T T$, in which α is thermal expansivity 123 and $K_{\rm T}$ is isothermal bulk modulus. Assuming the thermal parameters for liquid Fe-C 124 same as those for pure iron, we employ $\alpha K_T = 9$ MPa/K (Ichikawa et al., 2014) for runs 125 performed above 83 GPa. Andrault et al. (1998) demonstrated that 90% of such 126 isochoric thermal pressure contributes to an experimental pressure increase at the core 127 pressure range. For runs #1, #2, and #5 conducted at relatively low pressures, the lower 128 $\alpha K_{\rm T}$ value of 4 MPa/K and 60% of the theoretical value were applied (Morard et al., 129 2011; Andrault et al., 1998). The overall errors in pressure should be less than ± 10 % in 130 the present experiments including the uncertainty in thermal pressure estimate, as 131 argued in Mori et al. (2017). Indeed, we found 191 GPa at 2300 K in run #7 based on 132 the volume and equation of state (EoS) of Fe₃C (Sata et al., 2010), which is consistent 133 with 194 GPa calculated with the thermal pressure correction. 134 Synchrotron XRD measurements were carried out in situ at high P-T in runs #7 135 and #8 at BL10XU, SPring-8 (Ohishi et al., 2008) using an X-ray beam with an energy 136 of ~30 keV. Angle-dispersive XRD spectra were collected on a flat panel detector 137 (Perkin Elmer) with typical exposure time of 200 ms. A monochromatic incident X-ray

beam was collimated to 2 µm (FWHM). Visible fluorescent light induced by X-rays in a
diamond was used to precisely align the laser-heated spot with the X-ray beam. Twodimensional XRD images were integrated to produce a conventional one-dimensional
diffraction profile using the IPAnalyzer software (Seto et al., 2010).

142 Recovered samples from our DAC experiments were cross-sectioned using a Ga⁺ 143 focused ion beam (FIB) instrument (FEI VersaTM3D DualBeamTM). During FIB 144 sectioning, sample was monitored by scanning electron microscopy using back-145 scattered electrons to precisely section the center of a heated spot. The sample cross 146 sections were then examined by a field-emission-type scanning electron microscope 147 (FE-SEM) and energy dispersive X-ray spectrometry (EDS) with a silicon drift detector. 148 The section was then mounted on a Cu grid and further thinned to approximately 3 µm 149 thickness for further analysis with a field-emission-type electron probe microanalyzer 150 (FE-EPMA, JXA-8530F, JEOL) in order to determine chemical compositions. Since Fe-151 C samples are good electric conductors, no coating material was necessary for EPMA 152 analysis with an accelerating voltage of 10 kV and a beam current of 15 nA. A focused 153 electron beam for spot analyses was less than 50 nm in size. We used both Fe₃C and 154 Fe₇C₃ as standards that were synthesized at 5 GPa/1473 K and applied the ZAF 155 correction (Nakajima et al., 2009). Corundum (for Al and O) and pure iron (for Fe) were 156 also used for standards. The X-ray counting time for peak/background was 20s/10s. 157 LDE2 (C), TAP (Al), LIFH (Fe), and LDE1H (O) were analyzing crystals. 158 Very minor amounts of aluminum and oxygen were found when analyzing the 159 sample, most likely due to a signal from a neighboring/underlying Al₂O₃ insulation 160 layer. We estimated the carbon concentrations in guenched liquid and solid metals after subtracting Al₂O₃ from raw analyses. In addition, we analyzed two samples (runs #2 & 161

162#6) by EPMA with a liquid nitrogen cold finger (JXA-8530FPlus HyperProbe Electron163Probe Microanalyzer) at both 10 kV and 15 kV accelerating voltage (Table S1). The164results demonstrate that the carbon contents obtained with and without the cold finger165are similar to each other. We repeated the analyses of the Fe₃C standard twenty times166and found that the carbon concentration changed only by ±0.1766 atm%. While167chemical compositions were calculated by using both the 'standard' ZAF correction and168the phi-rho-z type procedure, the results are consistent with each other (see Table S1).

107

```
170 3. Results
```

171 3.1. Carbon contents in eutectic liquid and coexisting solid iron

172 We have conducted eight separate DAC runs for melting experiments in a pressure 173 range from 23 GPa to 255 GPa (Table 1). Figs. 1 and 2 show typical melting textures on 174 recovered samples. There is a chemically homogeneous part at the center of a laser-175 heating spot, which is non-stoichiometric in composition and should represent a 176 quenched liquid. The liquidus phase (solid phase in contact with liquid) is found at both 177 sides of the liquid. The outermost low-temperature area remained subsolidus. We 178 observed iron carbide (Fe_3C or Fe_7C_3) or metallic iron (or both) as liquidus phase(s). In 179 runs #2, #4, #5, and #8, both iron carbide and iron were present, suggesting that the 180 quenched liquid in these runs may represent a eutectic liquid in the Fe–C binary system. Carbon diffusivity in liquid iron was estimated to be about 10^{-8} m²/s under the present 181 182 *P-T* conditions (Helffrich, 2014), indicating that 1 s is long enough for carbon to diffuse 183 in the present liquid pool that was less than 20 µm across (Figs. 1 and 2). Since 184 melting/crystallization at the liquid/solid boundary occurs almost instantaneously, 185 chemical homogeneity in the liquid assures that chemical equilibrium was attained in

the present heating duration of about 5 s. The Fe-C liquids obtained in the core pressure range included about 1 wt.% oxygen in addition to carbon (Table 1), which might have affected phase relations in the Fe–C binary system.

189 We determined the carbon concentration in quenched liquid up to 255 GPa, the

190 conditions corresponding to the middle of the outer core, more than twice the pressure

191 compared to previous experimental studies (e.g., Lord et al., 2009; Fei and Brosh, 2014;

192 Morard et al., 2017a). Fig. 3 demonstrates the carbon concentrations in liquids as a

193 function of pressure, which constrains the change in eutectic liquid composition,

194 considering that they indicate the lower and upper bounds for the carbon content in

195 eutectic liquid at each pressure when coexisting with iron and iron carbides,

196 respectively. These data show a small reduction in carbon content in the eutectic liquid

197 with increasing pressure above 23 GPa. Our results are not consistent with the

198 continuous reduction that was thermodynamically predicted by Fei and Brosh (2014)

and experimentally reported by Lord et al. (2009) (Fig. 3). The difference from Lord et

al. (2009) may be attributed to the difference in analytical method: Lord and others

201 employed an X-radiographic imaging technique for chemical analysis, while we based

202 our measurements on electron microprobe analyses.

We also obtained the carbon concentrations in solid iron coexisting with Fe-C liquid as a function of pressure (Fig. 4). Liquid Fe-C coexists with face-centered cubic (fcc) iron at relatively low pressure range, in which carbon atoms occupy the octahedral interstitial sites (Fei and Brosh, 2014). Previous studies reported that the maximum solubility of carbon in fcc-Fe is <2.1 wt.% at 1 bar and predicted that it decreases at higher pressures (Walker et al., 2013; Fei and Brosh, 2014). Our results show about 1.0 wt.% in fcc-Fe from 23 to 54 GPa, in good agreement with such predictions. Above 138

GPa in the present experiments, hcp Fe should have crystallized from Fe-C liquid and again included 1.0 wt.% or less C. The effect of pressure on the solubility of carbon in hcp Fe was not observed.

213

214 *3.2. Eutectic temperature in the Fe–C system*

The temperature at the solid/liquid boundary gives the upper bound for eutectic temperature. These temperatures are plotted as a function of pressure in Fig. 5a, which constrains the eutectic melting curve to mid-outer core pressures. We fit the Simon-Glatzel equation to the data (Simon and Glatzel, 1929):

219
$$T_m = T_{mref} \left(\frac{P - P_{ref}}{a} + 1\right)^{1/c} \tag{1}$$

where T_m is eutectic temperature, and both a and c are fitting parameters. $T_{mref} = 1513$ 220 221 K at $P_{ref} = 5$ GPa is employed as a reference eutectic temperature from previous data 222 obtained using a large-volume press (Fei and Brosh, 2014). The best fitted parameters 223 are $a = 8.4 \pm 3.3$ GPa and $c = 4.1 \pm 0.5$. The obtained eutectic melting curve shown in 224 Fig. 5a is consistent with those of previous studies by Liu et al. (2016) and Morard et al. 225 (2017a) within experimental uncertainties. Fig. 5b compares the eutectic melting curves 226 of iron alloys. Under the core pressure range, the Fe–C system exhibits a eutectic 227 temperature lower than those of Fe (Anzellini et al., 2013), Fe–FeSi (Fischer et al., 228 2013), and Fe–FeO (Morard et al., 2017a). And, the Fe–C eutectic melting temperature 229 is higher than that in Fe–Fe₃S below ~250 GPa, but is lower at higher pressures 230 including the ICB pressure (Mori et al., 2017). 231

232 3.3. Liquidus field of Fe_7C_3

233	In runs #7 and #8 using the Fe-4.0 wt.%C starting material, the sample was
234	initially compressed at room temperature to 176 GPa and 226 GPa, respectively (Fig.
235	6). Broad XRD peaks from hcp Fe and Fe_3C were observed before heating in both runs.
236	The peaks from both phases became sharper and spotty upon heating to 2300 K in run
237	#7, in which more than 10 peaks were indexed to Fe_3C cementite (Sata et al., 2010), and
238	only Fe_3C phase was observed from the sample in the XRD pattern during melting at
239	203 GPa and 3350 K. In addition, we confirmed with an FE-EPMA that the quenched
240	liquid was surrounded by Fe_3C in the recovered sample (Fig. 7a). These indicate that Fe
241	forms eutectic melting with Fe ₃ C at 203 GPa.
242	In the case of run #8, the diffraction peaks from hcp Fe and Fe_3C became sharper
243	and spotty upon heating. We then observed the appearance of the peaks that are
244	assigned to Fe_7C_3 (orthorhombic, <i>Pbca</i>) (Prescher et al., 2015) after we melted the
245	sample by increasing temperature to 3570 K at 255 GPa. After quenching this sample,
246	we confirmed with the FE-EPMA that the liquid was in contact with both Fe and Fe_7C_3 ,
247	whereas the subsolidus part was composed of Fe and Fe ₃ C (Fig. 7b). These show that
248	eutectic melting occurs between Fe and Fe_7C_3 at 255 GPa. This observation is not
249	consistent with Liu et al. (2016), who argued that Fe_3C dissociates into $Fe + Fe_7C_3$ at
250	subsolidus above 145 GPa when using Fe ₃ C as a sample. Liu and others based their
251	conclusions purely on the change in XRD patterns, and they may have collected them
252	above eutectic temperature.

4. Discussion

255 4.1. Phase diagram of Fe–C system at ICB

256 The eutectic liquid composition (Fig. 3) and eutectic melting curve obtained in this

study (Fig. 5) indicate that eutectic melting occurs in Fe–Fe₃C at Fe + 3.5 wt.% C at

- 258 135 GPa and 3000 K. And, these data determined up to 255 GPa suggest that the
- eutectic liquid includes about 3 wt.% C at 330 GPa and 3750 K.

260 We model the liquidus curve for the Fe-rich side of the eutectic to 330 GPa (Fig.

8). Considering an ideal solution, a molar ratio of Fe in liquid at *T* is given by:

262
$$\chi_{\text{Fe liq}} = \exp\left\{\frac{\bar{G}_{\text{Fe sol}}^{0} - \bar{G}_{\text{Fe liq}}^{0}}{RT}\right\} = \exp\left\{-\int_{T_{m \text{Fe}}}^{T} \frac{\bar{H}_{\text{Fe sol}}^{0} - \bar{H}_{\text{Fe liq}}^{0}}{RT^{2}} dT\right\}$$
(2)

where $\bar{G}_{Fe\,i}^{0}$ and $\bar{H}_{Fe\,i}^{0}$ are standard Gibbs free energy and the enthalpy of phase *i*, respectively, $T_{m\,Fe}$ is the melting temperature of pure Fe, and *R* is gas constant. When we assume that an enthalpy change upon fusion, $\bar{H}_{Fe\,liq}^{0} - \bar{H}_{Fe\,sol}^{0}$, is independent from temperature, Eq. (2) is written as:

267
$$\chi_{\text{Fe liq}} = \exp\left\{\left(\frac{\overline{H}_{\text{Fe liq}}^{0} - \overline{H}_{\text{Fe sol}}^{0}}{RT_{m \text{Fe}}}\right) \times \left(1 - \frac{T_{m \text{Fe}}}{T}\right)\right\} = \exp\left\{A \times \left(1 - \frac{T_{m \text{Fe}}}{T}\right)\right\}$$
(3)

Here *A* is constant and calculated from $T_{m \text{ Fe}}$ (Anzellini et al., 2013), eutectic temperature, and eutectic composition (Figs. 3 and 5) at a given pressure (see also Mori et al., 2017).

271

272 *4.2. Fe-C core liquid?*

Melting phase relations in the Fe–C system constrain the carbon budget in the Earth's core and the mineralogy of the solid inner core. The maximum carbon content in the outer core has been estimated from the core density deficit. Recent experimental and computational studies on the density of liquid Fe-C alloys indicate 1.8 to 4.2 wt.% C in the liquid core at the ICB (Badro et al., 2014: Nakajima et al., 2015; Morard et al., 2017b), although earlier estimates based on the EoS of solid Fe₃C suggested higher 279 carbon concentration (7.7–9.2 wt.% C) (Sata et al., 2010). Since the eutectic point is 280 located at about 3 wt.% C at 330 GPa (Fig. 8), if the Fe-C outer core includes less than 281 3 wt.% C, it crystallizes hcp Fe with ≤ 1 wt.% C at the ICB (Fig. 4). In this case, 282 however, the difference in carbon concentration between the liquid and solid core is too 283 small to explain the density jump (4.5 to 6.5%) across the ICB (Shearer and Masters, 284 1990; Masters and Gubbins, 2003). Moreover, assuming a linear density change 285 between pure Fe (Sakamaki et al., 2016), Fe₃C (Sata et al., 2010), and Fe₇C₃ (Prescher 286 et al., 2015), hcp Fe with 1.0 wt.% C is found to be only 1.3% lighter than pure iron, 287 which is not enough to account for the inner core density deficit of 4–6% (Brown and 288 McQueen, 1986; Shearer and Masters, 1990; Masters and Gubbins, 2003; Dewaele et 289 al., 2006; Komabayashi and Fei, 2010). This is supported by the recent calculations by 290 Li et al. (2018). Less than 3 wt.% C as a single light element in the core is therefore 291 unlikely.

292 Next we consider the case that the liquid core contains more than 3 wt.% C as the 293 sole light component, from which Fe_7C_3 crystallizes at the ICB (Fig. 8). It has been 294 suggested that both V_P and V_S —in particular the slow V_S (high Poisson's ratio)—of the 295 inner core could be reconciled with non-magnetic Fe₇C₃ (Mookherjee et al., 2011; Chen 296 et al., 2014; Prescher et al., 2015). If the inner core is composed of solid Fe₇C₃, the 297 liquid outer core should include 5.4-6.7 wt.% C according to Badro et al. (2014) or 5.2-298 6.0 wt.% C based on Nakajima et al. (2015), considering that the outer core should be 299 4.5–6.5% less dense than solid Fe₇C₃ (Chen et al., 2014). More than 5 wt.% C in liquid 300 iron, however, leads to a bulk sound velocity and a density much higher and lower than 301 those observed for the outer core, respectively (Badro et al., 2014; Nakajima et al., 302 2015; Morard et al., 2017b) and is thus unlikely. The computational study by Li et al.

303 (2016) also showed that the density of Fe_7C_3 is too low to be a major constituent of the 304 inner core.

These suggest that carbon cannot be a predominant light element in the outer core. It is indeed consistent with Wood et al. (2013), who proposed that the maximum carbon content in the core is 1.0 wt.%, which was deduced from the carbon isotopic signature in silicate Earth and the silicate-metal fractionation factors. It has also been argued that the outer core velocity is explained by the incorporation of a small amount of carbon that is not enough to account for its density deficit from pure iron (Badro et al., 2014; Nakajima et al., 2015).

312

313 4.3. Fe-C-X core liquid?

314 Furthermore, we consider the Fe-C-X (X = Si, S, O, or H) ternary systems to seek 315 possible ranges of outer core composition, from which hcp Fe crystallizes (Figs. 9a-d). 316 The eutectic liquid composition in each Fe-X binary system has been estimated to be Fe 317 with 3 wt.% C (this study), 1.5 wt.% Si (Ozawa et al., 2016), 5 wt.% S (Mori et al., 318 2017), and >10 wt.% O (Morard et al., 2017a) at 330 GPa. The hydrogen concentration 319 in the Fe–H eutectic liquid is unclear, but it would be around 1 wt.% H according to 320 Fukai (1992) and Shibazaki et al. (2014). The ternary eutectic points in these systems 321 are not exactly known, but the recent melting experiments performed by Tateno et al. 322 (2018) have demonstrated that the Fe–Si–S ternary eutectic point is located 1) near the 323 tie line between the Fe-FeSi and Fe-Fe₃S binary eutectic points and 2) close to the Fe-324 Fe₃S binary join because the eutectic temperature in Fe–Fe₃S is much lower than that in 325 Fe–FeSi. The present data show that the eutectic temperature in the Fe–C binary system 326 is low compared to those in other systems (Fe–Si, Fe–S, and Fe–O) at 330 GPa (Fig. 5).

327 On the basis of these observations, Figs. 9a-d illustrate the liquidus phase relations 328 (showing a compositional range from which a specific solid phase first crystallizes) in 329 the Fe–C–X ternary systems at the ICB, with each ternary eutectic point located 330 relatively close to the Fe–C binary eutectic point except for the Fe–C–H system. 331 Previous experiments and calculations have estimated the maximum amount of 332 each light element in the liquid core to explain the outer core density deficit. Estimates 333 based on liquid iron alloys show 4.6-6.4 wt.% Si (Morard et al., 2013; Badro et al 334 2014), 5.8–14.0 wt.% S (Morard et al., 2013; Badro et al., 2014; Umemoto et al., 2014; 335 Kawaguchi et al., 2017), 5.5 wt.% O (Badro et al., 2014), and 1.0 wt.% H (Umemoto 336 and Hirose, 2015) at the ICB, which are again less than those obtained by the earlier 337 study on the EoSs of solid FeSi, FeS, and $Fe_{0.95}O$ (Sata et al., 2010). These give the 338 possible compositional range of Fe–C–X ternary outer core liquid in Figs. 9a–d (green 339 area).

340 We find that hcp Fe crystallizes from possible liquid core compositions in all of the 341 Fe-C-Si, Fe-C-S, Fe-C-O, and Fe-C-H systems. The recent calculations by Li et al. 342 (2018) proposed that all of the density, $V_{\rm P}$, and $V_{\rm S}$ of the solid inner core are explained 343 by hcp $Fe_{30}C_1Si_1$ (Fe + 0.7 wt.% C + 1.6 wt.% Si). Such inner core composition is 344 feasible, considering that the maximum solubilities of silicon and sulfur in hcp Fe when 345 coexisting with liquid alloys are ≤ 1.5 wt.% Si and 4 wt.% S at the ICB, respectively 346 (Ozawa et al., 2016; Mori et al., 2017). Hcp Fe + 1 wt.% C + 4.0 wt.% S also could 347 account for the inner core density (Sakai et al., 2012). Hydrogen is known to be soluble 348 in solid Fe (Fukai, 1992; Shibazaki et al., 2012; Terasaki et al., 2012), suggesting that 349 the hcp Fe-C-H inner core is also feasible. On the other hand, solid Fe crystallizing from 350 liquid Fe-C-O contains ≤ 1 wt.% carbon and a negligible amount of oxygen (Ozawa et

al., 2010), which does not account for the inner core density deficit (Li et al., 2018).

352 Therefore, while carbon is not a predominant light element in the core, the outer core

- 353 liquid can be Fe-C-Si, Fe-C-S, or Fe-C-H.
- 354

355 **5.** Conclusion

356 We have conducted melting experiments on the Fe-C binary system up to 255 GPa 357 by using a laser-heated DAC. The carbon concentration in eutectic liquid and the 358 solubility of carbon in solid iron were determined based on textual and chemical 359 characterizations for samples recovered from the DAC. Results show that 1) the carbon 360 content in eutectic liquid slightly decreases with increasing pressure from 4.3 wt.% C at 361 ambient pressure to 3.6 wt.% C at 255 GPa, 2) eutectic melting occurs between Fe and 362 Fe₃C to 203 GPa and between Fe and Fe₇C₃ at 255 GPa, and 3) the maximum solubility 363 of carbon in solid iron is almost constant (≤ 1 wt.%) up to 255 GPa. The eutectic liquid 364 in the Fe–C binary system includes about 3 wt.% C at 330 GPa. 365 If carbon is the sole light element in the core and the outer core includes <3 wt.% 366 C, hcp Fe crystallizes at the ICB but includes only less than 1 wt.% C, which does not 367 explain the inner core density deficit. If the liquid core contains more than 3 wt.% C, it 368 forms Fe_7C_3 at the ICB, whose density is too low for the inner core. These indicate that 369 carbon is not a primary light element in the core. On the other hand, the outer core can 370 be liquid Fe-C-Si, Fe-C-S, or Fe-C-H. It crystallizes hcp Fe including light elements 371 with less than 1 wt.% C, which may explain the density and the sound velocities 372 observed in the inner core. 373

374 Acknowledgments

375 We appreciate Y. Nakajima for his help in quantitative determination of carbon 376 content with an EPMA. We also thank K. Yonemitsu, H. Yoshida, S. Tagawa, S. 377 Tateno, and Y. Kuwayama for their support during experiments. Comments provided 378 anonymous reviewers were helpful, in particular on the EPMA analysis of carbon. This 379 work was supported by the JSPS research grant 16H06285 to K.H. XRD measurements 380 were performed at BL10XU, SPring-8 (proposals no. 2017A0072 and 2017B0072). 381 F.M. and G.M. received funding from the European Research Council (ERC) under the 382 European Union's Horizon 2020 research and innovation Programme (grant agreement 383 no. 670787).

385 **References**

- 386 Akahama, Y., Kawamura, H., 2007. Diamond anvil Raman gauge in multimegabar
- 387 pressure range. High Press. Res. 27, 473–482.
- 388 https://doi.org/10.1080/08957950701659544
- 389 Anders, E., Grevesse, N., 1989. Abundances of the elements: meteoritic and solar.
- 390 Geochim. Cosmochim. Acta 53, 197–214. https://doi.org/10.1016/0016-
- 391 7037(89)90286-X
- 392 Andrault, D., Fiquet, G., Itie, J.P., Richet, P., Gillet, P., Hausermann, D., Hanfland, M.,
- 393 1998. Thermal pressure in the laser-heated diamond-anvil cell: an X-ray diffraction
 394 study. Eur. J. Mineral. 10, 931–940.
- Anzellini, S., Dewaele, A., Mezouar, M., Loubeyre, P., Morard, G., 2013. Melting of
 iron at earth's inner core boundary based on fast X-ray diffraction. Science 340,
- 397 464–466. https://doi.org/10.1126/science.1233514
- 398 Badro, J., Cote, A.S., Brodholt, J.P., 2014. A seismologically consistent compositional
- model of Earth's core. Proc. Natl. Acad. Sci. USA 111, 7542–7545.
- 400 https://doi.org/10.1073/pnas.1316708111
- 401 Benz, M., Elliott, J., 1961. The austenite solidus and revised iron-carbon diagram.
- 402 Trans. Metall. Soc. AIME 221, 323–331.
- 403 Brown, J.M., McQueen, G., 1986. Phase transitions, Grüneisen parameter, and elasticity
- 404 for shocked iron between 77 GPa and 400 GPa. J. Geophys. Res. 91, 7485–7494.
- 405 https://doi.org/10.1029/JB091iB07p07485
- 406 Chabot, N.L., Campbell, A.J., McDonough, W.F., Draper, D.S., Agee, C.B., Humayun,
- 407 M., Watson, H.C., Cottrell, E., Saslow, S.A., 2008. The Fe–C system at 5 GPa and

- 408 implications for Earth's core. Geochim. Cosmochim. Acta 72, 4146–4158.
- 409 https://doi.org/10.1016/j.gca.2008.06.006
- 410 Chen, B., Li, Z., Zhang, D., Liu, J., Hu, M.Y., Zhao, J., Bi, W., Alp, E.E., 2014. Hidden
- 411 carbon in Earth's inner core revealed by shear softening in dense Fe_7C_3 . Proc.
- 412 Natl. Acad. Sci. USA 111, 17755–17758. https://doi.org/10.1073/pnas.1411154111
- 413 Dasgupta, R., Walker, D., 2008. Carbon solubility in core melts in a shallow magma
- 414 ocean environment and distribution of carbon between the Earth's core and the
- 415 mantle. Geochim. Cosmochim. Acta 72, 4627–4641.
- 416 https://doi.org/10.1016/j.gca.2008.06.023
- 417 Dewaele, A., Loubeyre, P., Occelli, F., Mezouar, M., Dorogokupets, P. I., Torrent, M.,
- 418 2006. Quasihydrostatic equation of state of iron above 2 Mbar. Phys. Rev. Lett. 97,
- 419 215504. https://doi.org/10.1103/PhysRevLett.97.215504
- 420 Fei, Y., Brosh, E., 2014. Experimental study and thermodynamic calculations of phase
- 421 relations in the Fe–C system at high pressure. Earth Planet. Sci. Lett. 408, 155–
- 422 162. https://doi.org/10.1016/j.epsl.2014.09.044
- 423 Fischer, R.A., Campbell, A.J., Reaman, D.M., Miller, N.A., Heinz, D.L., Dera, P.,
- 424 Prakapenka, V.B., 2013. Phase relations in the Fe–FeSi system at high pressures
- 425 and temperatures. Earth Planet. Sci. Lett. 373, 54–64.
- 426 https://doi.org/10.1016/j.epsl.2013.04.035
- 427 Fukai, Y., 1992. Some properties of the Fe–H system at high pressures and
- 428 temperatures, and their implications for the Earth's core. In: Syono Y., Manghnani
- 429 M.H. (Eds.), High-Pressure Research: Application to Earth and Planetary Sciences.
- 430 AGU monogr. ser. 67, pp. 373–385. https://doi.org/10.1029/GM067p0373

- 431 Helffrich, G., 2014. Outer core compositional layering and constraints on core liquid
- 432 transport properties. Earth Planet. Sci. Lett. 391, 256–262.
- 433 https://doi.org/10.1016/j.epsl.2014.01.039
- 434 Hillgren, V.J., Gessmann, C.K., Li, J., 2000. An experimental perspective on the light
- element in Earth's core. In: Canup, R.M., Righter, K. (Eds.), Origin of the Earth
 and Moon. Univ. Arizona Press, pp. 245–263.
- 437 Hirose, K., Morard, G., Sinmyo, R., Umemoto, K., Hernlund, J.W., Labrosse, S., 2017.
- 438 Crystallization of silicon dioxide and compositional evolution of the Earth's core.
- 439 Nature 543, 99–102. https://doi.org/10.1038/nature21367
- 440 Ichikawa, H., Tsuchiya, T., Tange, Y., 2014. The P-V-T equation of state and
- 441 thermodynamic properties of liquid iron. J. Geophys. Res. Solid Earth 119, 240–
- 442 252. https://doi.org/10.1002/2013JB010732
- 443 Kawaguchi, S. I., Nakajima, Y., Hirose, K., Komabayashi, T., Ozawa, H., Tateno, S.,
- 444 Kuwayama, Y., Tsutsui, S., Baron, A.Q., 2017. Sound velocity of liquid Fe–Ni–S
- 445 at high pressure. J. Geophys. Res. Solid Earth 122, 3624–3634.
- 446 https://doi.org/10.1002/2016JB013609
- 447 Komabayashi, T., Fei, Y., 2010. Internally consistent thermodynamic database for iron
- to the Earth's core conditions. J. Geophys. Res. Solid Earth 115, 1–12.
- 449 https://doi.org/10.1029/2009JB006442
- 450 Li, Y., Vočadlo, L., Brodholt, J., Wood, I.G., 2016. Thermoelasticity of Fe₇C₃ under
- 451 inner core conditions. J. Geophys. Res. Solid Earth 121, 5828–5837.
- 452 https://doi.org/10.1002/2016JB013155

- 453 Li, Y., Vočadlo, L., Brodholt, J.P., 2018. The elastic properties of hcp-Fe alloys under
- 454 the conditions of the Earth's inner core. Earth Planet. Sci. Lett. 493, 118–127.
- 455 https://doi.org/10.1016/j.epsl.2018.04.013
- 456 Liu, J., Lin, J., Prakapenka, V.B., Prescher, C., Yoshino, T., 2016. Phase relations of
- 457 Fe₃C and Fe₇C₃ up to 185 GPa and 5200 K: implication for the stability of iron
- 458 carbide in the Earth's core. Geophys. Res. Lett. 43, 12415–12422.
- 459 https://doi.org/10.1002/2016GL071353.
- 460 Lord, O.T., Walter, M.J., Dasgupta, R., Walker, D., Clark, S.M., 2009. Melting in the
- 461 Fe–C system to 70 GPa. Earth Planet. Sci. Lett. 284, 157–167.
- 462 https://doi.org/10.1016/j.epsl.2009.04.017
- 463 Masters, G., Gubbins, D., 2003. On the resolution of density within the Earth. Phys.
- 464 Earth Planet. Inter. 140, 159–167. https://doi.org/10.1016/j.pepi.2003.07.008
- 465 McDonough, W.F., 2003. Compositional model for the Earth's core. In: Holland, H.,
- 466 Turekian, K. (Eds.), Treatise on Geochemistry: 2nd Edition. Elsevier Ltd., pp.
- 467 559–577. https://doi.org/10.1016/B978-0-08-095975-7.00215-1
- 468 Mookherjee, M., Nakajima, Y., Neumann, G.S., Glazyrin, K., Wu, X., Dubrovinsky, L.,
- 469 Mccammon, C., Chumakov, A., 2011. High-pressure behavior of iron carbide
- 470 (Fe₇C₃) at inner core conditions. J. Geophys. Res. Solid Earth 116, 1–13.
- 471 https://doi.org/10.1029/2010JB007819
- 472 Morard, G., Andrault, D., Guignot, N., Siebert, J., Garbarino, G., Antonangeli, D., 2011.
- 473 Melting of Fe–Ni–Si and Fe–Ni–S alloys at megabar pressures: Implications for
- 474 the core-mantle boundary temperature. Phys. Chem. Miner. 38, 767–776.
- 475 https://doi.org/10.1007/s00269-011-0449-9

- 476 Morard, G., Siebert, J., Andrault, D., Guignot, N., Garbarino, G., Guyot, F.,
- 477 Antonangeli, D., 2013. The Earth's core composition from high pressure density
- 478 measurements of liquid iron alloys. Earth Planet. Sci. Lett. 373, 169–178.
- 479 https://doi.org/10.1016/j.epsl.2013.04.040
- 480 Morard, G., Andrault, D., Antonangeli, D., Nakajima, Y., Auzende, A.L., Boulard, E.,
- 481 Cervera, S., Clark, A., Lord, O.T., Siebert, J., Svitlyk, V., Garbarino, G., Mezouar,
- 482 M., 2017a. Fe–FeO and Fe–Fe₃C melting relations at Earth's core-mantle
- 483 boundary conditions : Implications for a volatile-rich or oxygen-rich core. Earth
- 484 Planet. Sci. Lett. 473, 94–103. https://doi.org/10.1016/j.epsl.2017.05.024
- 485 Morard, G., Nakajima, Y., Andrault, D., Antonangeli, D., Auzende, A.L., Boulard, E.,
- 486 Cervera, S., Clark, A.N., Lord, O.T., Siebert, J., Svitlyk, V., Garbarino, G.,
- 487 Mezouar, M., 2017b. Structure and density of Fe–C liquid alloys under high
- 488 pressure. J. Geophys. Res. Solid Earth 122, 7813–7823.
- 489 https://doi.org/10.1002/2017JB014779
- 490 Mori, Y., Ozawa, H., Hirose, K., Sinmyo, R., Tateno, S., Morard, G., Ohishi, Y., 2017.
- 491 Melting experiments on Fe–Fe₃S system to 254 GPa. Earth Planet. Sc. Lett. 464,
- 492 135–141. https://doi.org/10.1016/j.epsl.2017.02.021
- 493 Nakajima, Y., Takahashi, E., Suzuki, T., Funakoshi, K., 2009. "Carbon in the core"
- 494 revisited. Phys. Earth Planet. Inter. 174, 202–211.
- 495 https://doi.org/10.1016/j.pepi.2008.05.014
- 496 Nakajima, Y., Imada, S., Hirose, K., Komabayashi, T., Ozawa, H., Tateno, S., Tsutsui,
- 497 S., Kuwayama, Y., Baron, A.Q.R., 2015. Carbon-depleted outer core revealed by
- 498 sound velocity measurements of liquid iron-carbon alloy. Nat. Commun. 6, 1–7.
- 499 https://doi.org/10.1038/ncomms9942

- 500 Ohishi, Y., Hirao, N., Sata, N., Hirose, K., Takata, M., 2008. Highly intense
- 501 monochromatic X-ray diffraction facility for high-pressure research at SPring-8.
- 502 High Press. Res. 28, 163–173. https://doi.org/10.1080/08957950802208910
- 503 Ozawa, H., Hirose, K., Tateno, S., Sata, N., Ohishi, Y. 2010. Phase transition boundary
- between B1 and B8 structures of FeO up to 210 GPa. Phys. Earth Planet. Inter.
- 505 179, 157–163. https:// doi:10.1016/j.pepi.2009.11.005
- 506 Ozawa, H., Hirose, K., Yonemitsu, K., Ohishi, Y., 2016. High-pressure melting
- 507 experiments on Fe–Si alloys and implications for silicon as a light element in the
- 508 core. Earth Planet. Sci. Lett. 456, 47–54. https://doi.org/10.1016/j.epsl.2016.08.042
- 509 Prescher, C., Dubrovinsky, L., Bykova, E., Kupenko, I., Glazyrin, K., Kantor, A.,
- 510 McCammon, C., Mookherjee, M., Nakajima, Y., Miyajima, N., Sinmyo, R.,
- 511 Cerantola, V., Dubrovinskaia, N., Prakapenka, V., Rüffer, R., Chumakov, A.,
- 512 Hanfland, M., 2015. High Poisson's ratio of Earth's inner core explained by carbon

513 alloying. Nat. Geosci. 8, 220–223. https://doi.org/10.1038/ngeo2370

- 514 Sakai T., Ohtani, E., Kamada, S., Terasaki, H., Hirao, N. 2012. Compression of
- 515 Fe_{88.1}Ni_{9.1}S_{2.8} alloy up to the pressure of Earth's inner core. J. Geophys. Res. Solid
- 516 Earth 117, B02210. https://doi:10.1029/2011JB008745
- 517 Sakamaki, T., Ohtani, E., Fukui, H., Kamada, S., Takahashi, S., Sakairi, T., Takahata,
- 518 A., Sakai, T., Tsutsui, S., Ishikawa, D., Shiraishi, R., Seto, Y., Tsuchiya, T., Baron,
- 519 A.Q.R., 2016. Constraints on Earths inner core composition inferred from
- 520 measurements of the sound velocity of hcp-iron in extreme conditions. Sci. Adv. 2,
- 521 e1500802–e1500802. https://doi.org/10.1126/sciadv.1500802
- 522 Sata, N., Hirose, K., Shen, G., Nakajima, Y., Ohishi, Y., Hirao, N., 2010. Compression
- 523 of FeSi, Fe₃C, Fe_{0.95}O, and FeS under the core pressures and implication for light

- element in the Earth's core. J. Geophys. Res. Solid Earth 115, 1–13.
- 525 https://doi.org/10.1029/2009JB006975
- 526 Seto, Y., Nishio-Hamane, D., Nagai, T., Sata, N., 2010. Development of a software
- 527 suite on X-ray diffraction experiments. Rev. High Press. Sci. Technol. 20, 269–
- 528 276. https://doi.org/10.4131/jshpreview.20.269
- Shearer, P., Masters, G., 1990. The density and shear velocity contrasts at the inner core
 boundary. Geohys. J. Int. 102, 408–491. https://doi.org/10.1111/j.1365-
- 531 246X.1990.tb04481.x
- 532 Shibazaki, Y., Ohtani, E., Fukui, H., Sakai, T., Kamada, S., Ishikawa, D., Tsutsui, S.,
- 533 Baron, A.Q.R., Nishitani, N., Hirao, N., Takemura, K., 2012. Sound velocity
- 534 measurements in dhcp-FeH up to 70 GPa with inelastic X-ray scattering:
- 535 implications for the composition of the Earth's core. Earth Planet. Sci. Lett. 313–
- 536 314, 79–85. https://doi.org/10.1016/j.epsl.2011.11.002
- 537 Shibazaki, Y., Terasaki, H., Ohtani, E., Tateyama, R., Nishida, K., Funakoshi, K., Higo,
- 538 Y., 2014. High-pressure and high-temperature phase diagram for Fe_{0.9}Ni_{0.1}–H
- alloy. Phys. Earth Planet. Inter. 228, 192–201.
- 540 https://doi.org/10.1016/j.pepi.2013.12.013
- 541 Simon, F., Glatzel, G., 1929. Remarks on fusion pressure curve. Z. Anorg. Allg. Chem.
- 542 178, 309–316. https://doi.org/10.1002/zaac.19291780123
- 543 Tateno, S., Hirose, K., Ohishi, Y., Tatsumi, Y., 2010. The structure of iron in Earth's
- 544 inner core. Science 330, 359–361. https://doi.org/10.1126/science.1194662
- 545 Tateno, S., Kuwayama, Y., Hirose, K., Ohishi, Y. 2015. The structure of Fe–Si alloy in
- 546 Earth's inner core. Earth Planet. Sci. Lett. 418, 11–19.
- 547 http://dx.doi.org/10.1016/j.epsl.2015.02.008

- 548 Tateno, S., Hirose, K., Sinmyo, R., Morard, G., Hirao, N., Ohishi, Y., 2018. Melting
- 549 experiments on Fe–Si–S alloys to core pressures: silicon in the core? Am. Miner.

550 103, 742–748. http://doi.org/10.2138/am-2018-6299

- 551 Terasaki, H., Ohtani, E., Sakai, T., Kamada, S., Asanuma, H., Shibazaki, Y., Hirao, N.,
- 552 Sata, N., Ohishi, Y., Sakamaki, T., Suzuki, A., Funakoshi, K., 2012. Stability of
- 553 Fe–Ni hydride after the reaction between Fe–Ni alloy and hydrous phase (δ-
- AlOOH) up to 1.2Mbar: possibility of H contribution to the core density deficit.
- 555 Phys. Earth Planet. Inter. 194–195, 18–24.
- 556 https://doi.org/10.1016/j.pepi.2012.01.002
- 557 Umemoto, K., Hirose, K., 2015. Liquid iron–hydrogen alloys at outer core conditions
- 558 by first-principles calculations. Geophys. Res. Lett. 42, 7513–7520.
- 559 https://doi.org/10.1002/2015GL065899
- 560 Umemoto, K., Hirose, K., Imada, S., Nakajima, Y., Komabayashi, T., Tsutsui, S.,
- 561 Baron, A.Q.R., 2014. Liquid iron–sulfur alloys at outer core conditions by first-
- 562 principles calculations. Geophys. Res. Lett. 41, 6712–6717.
- 563 https://doi.org/10.1002/2014GL061233
- 564 Walker, D., Dasgupta, R., Li, J., Buono, A., 2013. Nonstoichiometry and growth of
- some Fe carbides. Contrib. Mineral. Petrol. 166, 935–957.
- 566 https://doi.org/10.1007/s00410-013-0900-7
- 567 Wood, B.J., 1993. Carbon in the core. Earth Planet. Sci. Lett. 117, 593–607.
- 568 Wood, B.J., Li, J., Shahar, A., 2013. Carbon in the core : its influence on the properties
- of core and mantle. In: Hazen, R.M., Jones, A.P., Baross, J.A. (Eds.), Carbon in
- 570 Earth. Rev. Mineral. Geochem. 75, pp. 231–250.
- 571 https://doi.org/10.2138/rmg.2013.75.8

573 Figure captions

574

575 Fig. 1. (a) Temperature profile and (b) composite X-ray maps for oxygen (red), carbon 576 (purple), iron (green), and aluminum (light blue) for run #3 performed at 138 GPa (see 577 Fig. S1 for raw radiation spectra). Homogeneous quenched liquid is found at the center 578 (hot part) surrounded by solid layers of iron. A measured temperature profile is 579 combined in (a), which provides temperature at the liquid/solid boundary. 580 581 Fig. 2. Typical melting texture of a quenched sample and composite X-ray maps for 582 oxygen, carbon, aluminum, and iron for run #5 at 40 GPa. SIM, scanning ion 583 microscope image; SL, secondary electron image; Comp, backscattered image. 584 Homogeneous quenched liquid is found at the center surrounded by solid layers of Fe 585 and Fe₃C. 586 587 Fig. 3. Carbon concentrations in liquids obtained in this study. Liquid compositions are 588 shown by triangles (blue) when coexisting with solid Fe, inverse triangles (purple) when 589 in contact with Fe₃C, squares (red) when both solid Fe and Fe₃C were found, and a 590 diamond (orange) when solid Fe and Fe_7C_3 were formed. The triangles and inverse 591 triangles give lower and upper bounds for carbon concentration in eutectic liquid, 592 respectively. Open and cross symbols represent eutectic compositions reported by 593 previous studies (Chabot et al., 2008; Lord et al., 2009; Fei and Brosh, 2014; Morard et 594 al., 2017a). Green curve shows the change in eutectic liquid composition with 595 increasing pressure.

597 Fig. 4. Carbon contents in solid iron found in this study. Closed symbols are the same as598 in Fig. 3. Previous reports on the maximum solubility of carbon in solid iron are shown

599 by open symbols (Walker et al., 2013; Fei and Brosh, 2014).

600

601 **Fig. 5.** (a) Fe–Fe₃C (or Fe–Fe₇C₃ above 255 GPa) eutectic melting curve (orange)

602 obtained in this study (closed symbols). It agrees well with the previously determined

603 Fe–C melting curves by Morard et al. (2017a) (blue) and Liu et al. (2015) (green). Black

dotted curve represents the solidus of Fe₃C composition reported by Liu et al. (2016).

605 Purple dashed curve shows the melting temperature of pure Fe (Anzellini et al., 2013).

606 (b) Eutectic melting curve of the Fe–C system obtained in this study compared with

607 those of the Fe–Fe₃S (Mori et al., 2017), Fe–FeO (Morard et al., 2017a), and Fe–FeSi

608 (Fischer et al., 2013) binary systems.

609

Fig. 6. XRD patterns of Fe + 4.0 wt.% C obtained before and after heating (no melting),

and during and after melting (a) at 203 GPa in run #7 and (b) 255 GPa in run #8. The

612 unit-cell parameters and volumes of Fe₃C and Fe₇C₃ are a = 3.95(1) Å, b = 4.54(1) Å, c

613 = 5.95(1) Å, and V = 106.9(4) Å³ for Fe₃C, and a = 10.40(1) Å, b = 3.96(1) Å, c =

614 12.06(1) Å, and V = 496.4(3) Å³ for Fe₇C₃.

615

616 **Fig. 7.** Quenched textures and composite X-ray maps for carbon, oxygen, and iron at (a)

617 203 GPa (run #7) and (b) 255 GPa (run #8). While liquid was surrounded by Fe₃C in

618 the sample quenched at 203 GPa, it was in contact with Fe_7C_3 and Fe at 255 GPa (the

619 subsolidus part was composed of Fe_3C and Fe).

Fig. 8. Phase diagrams for the iron-rich portion of the Fe–C binary system at 54, 138,
203, and 255 GPa from the present experiments and at ICB conditions (330 GPa) by
extrapolation. Red and blue (and gray) curves represent Fe–Fe₃C and Fe–Fe₇C₃
eutectic, respectively.
Fig. 9. Likely liquidus phase relations in the Fe–C–X ternary systems at 330 GPa. Red
circles represent the eutectic compositions in the Fe–C (this study), Fe–Si (Ozawa et al.,
2016), Fe–S (Mori et al., 2017), Fe–O (Morard et al., 2017a), and Fe–H binary systems

629 (Fukai, 1992). Green regions show the range of possible composition of the outer core

that explains its density based on studies on liquid iron alloys (Morard et al., 2013;

Badro et al., 2014; Umemoto et al., 2014; Umemoto and Hirose, 2015; Nakajima et al.,

632 2015; Kawaguchi et al., 2017).

Table	1									
Exper	imental results in th	le Fe-C bina	rry system							
Run#	Starting material	P (GPa)	eutectic T (K)	Phases present	Liquid			Solid iron		
					C (wt%)	Fe (wt%)	O (wt%)	C (wt%)	Fe (wt%)	O (wt%)
	Fe-1.9 wt%C	23(2)	1950(100)	Liq + Fe	4.04(30)	96.30(30)	0.16(5)	0.85(24)	99.08(16)	0.07(11)
7	Fe-1.9 wt%C	54(5)	2550(130)	$Liq + Fe + Fe_3C$	4.45(32)	95.70(30)	0.34(30)	1.00(45)	98.89(51)	0.11(7)
3	Fe-1.9 wt%C	138(14)	2900(150)	Liq + Fe	3.95(22)	94.84(21)	1.70(17)	0.68(48)	99.38(29)	0.05(11)
4	Fe-1.9 wt%C	211(21)	3230(160)	$Liq + Fe + Fe_3C$	4.19(20)	95.23(30)	0.95(30)	1.04(22)	98.96(24)	not detected
5	Fe-4.0 wt%C	40(4)	2300(120)	$Liq + Fe + Fe_3C$	3.83(24)	96.50(23)	0.17(22)	0.64(4)	99.49(20)	not detected
9	Fe-4.0 wt%C	83(8)	2640(130)	$Liq + Fe_3C$	3.90(23)	95.72(12)	0.88(8)		,	I
7	Fe-4.0 wt%C	203(20)	3350(170)	$Liq + Fe_3C$	3.92(9)	95.39(22)	1.19(21)			I
8	Fe-4.0 wt%C	255(25)	3570(180)	$Liq + Fe + Fe_7C_3$	3.60(21)	96.84(49)	1.06(50)	0.61(34)	99.43(31)	not detected
Numb	ers in parentheses a	tre standard	deviation uncert	ainties in the last dig	it(s).					





Fig. 1





Fig. 3







Pressure (GPa)

Fig. 5 (a)



Fig. 5 (b)







Fig. 7



Fig. 8







