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1 Melting experiments on the Fe–C binary system up to 255 GPa:

- 2 Constraints on the carbon content in the Earth's core
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ABSTRACT

Phase relations, including the eutectic liquid composition in the Fe-C binary
system, remain unclear under the core pressure range, which makes estimating the
carbon budget in the Earth's core difficult. To explore this issue, we have conducted
melting and subsolidus experiments on Fe-C alloys in a diamond-anvil cell up to 255
GPa. Textural and compositional characterizations of quenched samples show that
carbon concentration in the eutectic liquid slightly decreases with increasing pressure
and is about 3 wt.% at the inner core boundary (ICB) pressure. The solubility of carbon
in solid Fe is found to be almost constant at ~1.0 wt.%. <i>In situ</i> X-ray diffraction data
indicate that Fe forms eutectic melting with Fe ₃ C to 203 GPa and with Fe ₇ C ₃ at 255
GPa.
Previous studies on liquid Fe-C alloys suggested that the density of the outer core is
explained by liquid Fe containing 1.8 to 4.2 wt.% C. If the liquid core includes <3 wt.%
C as a single light element, hexagonal close-packed (hcp) Fe crystallizes at the ICB.
However, the carbon content in such solid Fe is ≤ 1 wt.%, less than that required to
account for the inner core density deficit from pure iron. When the outer core includes
\geq 3 wt.% C, it forms Fe ₇ C ₃ at the ICB, whose density is too small for the inner core.
Carbon is therefore not a primary light element in the core. Nevertheless, the outer core
liquid can be Fe-C-Si, Fe-C-S, or Fe-C-H. Such core liquid crystallizes solid Fe with
light elements including less than 1 wt.% C, which may explain the density and the
sound velocities observed in the inner core.
Keywords: core; iron alloy; carbon; eutectic liquid; high pressure

1. Introduction

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43 Carbon has been believed to be one of feasible light elements contained in planetary 44 cores because of its high abundance in the solar system and strong affinity to molten 45 iron (Benz and Elliott, 1961; Anders and Grevesse, 1989; Wood, 1993). Recent first-46 principles calculations predicted that the incorporation of a small amount of carbon in 47 hcp Fe is compatible with the density and sound velocities observed in the Earth's inner 48 core (Li et al., 2018). 49 In order to reveal the abundance of carbon in the core, it is important to understand 50 the Fe-C phase diagram at high pressure, in particular at 330 GPa, which corresponds to 51 the conditions at the ICB. Wood (1993) first thermodynamically calculated the phase 52 relations in the Fe-C system up to 136 GPa and predicted that the Earth's core could 53 contain ~4 wt.% C. Nakajima et al. (2009) experimentally determined the melting phase 54 relationships in the Fe-rich portion of the Fe-C system up to 29 GPa and demonstrated 55 that Fe₇C₃ appears as a liquidus phase. Using an *in situ* X-radiographic imaging 56 technique, Lord et al. (2009) found that carbon concentration in the eutectic liquid in the 57 Fe-C binary system decreases substantially with increasing pressure to 70 GPa. The 58 thermodynamic modeling by Fei and Brosh (2014) based on multi-anvil experiments up 59 to 25 GPa predicted that the Fe-C eutectic liquid composition includes 2.2 wt.% C at 60 330 GPa. More recently Morard et al. (2017a) reported the change in carbon 61 concentration in the eutectic liquid in the Fe-Fe₃C binary up to 107 GPa and estimated 62 it to be 2.6 wt.% C at the core-mantle boundary (CMB). However, they obtained liquid compositions from X-ray diffraction (XRD) patterns with the Rietveld method, which 63 64 includes great uncertainty. So far, melting experiments on the Fe-C system have not been carried out under the core pressure range (>136 GPa). Moreover, the solubility of 65

carbon in solid iron also remains unknown above 25 GPa (Chabot et al., 2008; Walker et al., 2013; Fei and Brosh, 2014).

It is also important to verify the change in liquidus phase in the Fe–C system. It has been suggested that the liquidus field of Fe₇C₃ expands with respect to that of Fe₃C and eutectic melting occurs between Fe and Fe₇C₃ above 120 GPa (Lord et al., 2009; Fei and Brosh, 2014). Furthermore, the recent experiments by Liu et al. (2016) showed the decomposition of Fe₃C into Fe + Fe₇C₃ above 145 GPa, but it contradicts the formation of Fe₃C at >300 GPa and >5000 K reported by Tateno et al. (2010). Fe₇C₃ has been regarded as a possible constituent in the inner core because it provides a good match for both compressional (V_P) and shear velocities (V_S) (or Poisson's ratio) of the inner core under relevant conditions (Nakajima et al., 2009; Mookherjee et al., 2011; Chen et al., 2014; Prescher et al., 2015). The calculations by Li et al. (2016), however, demonstrated that the density of Fe₇C₃ is too small for the inner core.

Recent progress in melting experiments in a diamond-anvil cell (DAC) combined with a focused ion beam (FIB) technique has enabled us to determine the eutectic liquid composition to the core pressure range (Ozawa et al., 2016; Mori et al., 2017; Hirose et al., 2017). In this study, we have conducted both subsolidus and melting experiments on Fe-C alloys up to 255 GPa in order to determine the change in melting phase relations with increasing pressure, in particular 1) the stability of Fe₇C₃ at liquidus, 2) carbon concentration in eutectic liquid, and 3) the solubility of carbon in solid iron. Based on these results, we discuss the possible carbon content in the Earth's core and the mineralogy of the inner core.

2. Experimental methods

High-pressure and -temperature (*P-T*) conditions were generated in a laser-heated DAC using double-, single-beveled, and flat diamond anvils with 60, 90, 150, and 300 μm culet size. Starting materials were foils of Fe containing 1.9 wt.% C and 4.0 wt.% C, which are homogeneous mixtures of fine-grained Fe and Fe₃C synthesized by an ultrarapid quenching method (Morard et al., 2011, 2017a). Rhenium gaskets were preindented to about 40 μm thickness before laser drilling or spark erosion of sample holes with 30–100 μm diameters. Sample pellets were loaded between two pieces of about 10 μm thick dry Al₂O₃, which served as a pressure-transmitting medium and a thermal insulator from the diamond anvils. After loading, the whole DAC was dried in a vacuum oven at 423 K for more than 12 hrs, and subsequently a sample chamber was flushed with dry argon and squeezed in an argon atmosphere.

After compression to a desired pressure, the sample was heated from both sides with a couple of 100-W single-mode Yb fiber lasers. We used beam-shaping optics, which converts a Gaussian beam to one with a flat energy distribution and thus reduces radial temperature gradient in the sample. A laser-heated spot was up to ~20 μm across. One-dimensional temperature distributions on the surface of the sample were obtained using a spectro-radiometric method (Ohishi et al., 2008). Heating duration was limited to about 5 s in order to avoid fluctuations in temperature, which would otherwise cause complex melting textures. The experimental temperature reported in this study is that at the solid/liquid boundary, which is obtained by combining the measured radial temperature distribution with the melting texture observed in a sample's cross section (Fig. 1) (see Ozawa et al., 2016 for details). We polished the sample using an FIB from both sides little by little (~1 μm for each step) in order to have a cross section at the center of a liquid pool. The sample was thinned to about 3 μm thickness at the end. The

uncertainty in temperature is ± 5 % according to Mori et al. (2017), which includes uncertainties in the width of the liquid pool as well as those derived from the facts that we calculated temperature from integrated light not only from the sample surface but also from the pressure medium and that spatial resolution of the temperature profile was 2 to 3 μ m.

Pressure at room temperature was obtained based on the Raman shift of a diamond anvil (Akahama and Kawamura, 2004). It was then corrected for a contribution of thermal pressure. According to Andrault et al. (1998), when purely isochoric heating is made, the thermal pressure is written as $\Delta P = \alpha K_T T$, in which α is thermal expansivity and K_T is isothermal bulk modulus. Assuming the thermal parameters for liquid Fe-C same as those for pure iron, we employ $\alpha K_T = 9$ MPa/K (Ichikawa et al., 2014) for runs performed above 83 GPa. Andrault et al. (1998) demonstrated that 90% of such isochoric thermal pressure contributes to an experimental pressure increase at the core pressure range. For runs #1, #2, and #5 conducted at relatively low pressures, the lower αK_T value of 4 MPa/K and 60% of the theoretical value were applied (Morard et al., 2011; Andrault et al., 1998). The overall errors in pressure should be less than ± 10 % in the present experiments including the uncertainty in thermal pressure estimate, as argued in Mori et al. (2017). Indeed, we found 191 GPa at 2300 K in run #7 based on the volume and equation of state (EoS) of Fe₃C (Sata et al., 2010), which is consistent with 194 GPa calculated with the thermal pressure correction.

Synchrotron XRD measurements were carried out *in situ* at high *P-T* in runs #7 and #8 at BL10XU, SPring-8 (Ohishi et al., 2008) using an X-ray beam with an energy of ~30 keV. Angle-dispersive XRD spectra were collected on a flat panel detector (Perkin Elmer) with typical exposure time of 200 ms. A monochromatic incident X-ray

beam was collimated to 2 μ m (FWHM). Visible fluorescent light induced by X-rays in a diamond was used to precisely align the laser-heated spot with the X-ray beam. Two-dimensional XRD images were integrated to produce a conventional one-dimensional diffraction profile using the IPAnalyzer software (Seto et al., 2010).

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Recovered samples from our DAC experiments were cross-sectioned using a Ga⁺ focused ion beam (FIB) instrument (FEI VersaTM3D DualBeamTM). During FIB sectioning, sample was monitored by scanning electron microscopy using backscattered electrons to precisely section the center of a heated spot. The sample cross sections were then examined by a field-emission-type scanning electron microscope (FE-SEM) and energy dispersive X-ray spectrometry (EDS) with a silicon drift detector. The section was then mounted on a Cu grid and further thinned to approximately 3 µm thickness for further analysis with a field-emission-type electron probe microanalyzer (FE-EPMA, JXA-8530F, JEOL) in order to determine chemical compositions. Since Fe-C samples are good electric conductors, no coating material was necessary for EPMA analysis with an accelerating voltage of 10 kV and a beam current of 15 nA. A focused electron beam for spot analyses was less than 50 nm in size. We used both Fe₃C and Fe₇C₃ as standards that were synthesized at 5 GPa/1473 K and applied the ZAF correction (Nakajima et al., 2009). Corundum (for Al and O) and pure iron (for Fe) were also used for standards. The X-ray counting time for peak/background was 20s/10s. LDE2 (C), TAP (Al), LIFH (Fe), and LDE1H (O) were analyzing crystals.

Very minor amounts of aluminum and oxygen were found when analyzing the sample, most likely due to a signal from a neighboring/underlying Al_2O_3 insulation layer. We estimated the carbon concentrations in quenched liquid and solid metals after subtracting Al_2O_3 from raw analyses. In addition, we analyzed two samples (runs #2 &

#6) by EPMA with a liquid nitrogen cold finger (JXA-8530FPlus HyperProbe Electron Probe Microanalyzer) at both 10 kV and 15 kV accelerating voltage (Table S1). The results demonstrate that the carbon contents obtained with and without the cold finger are similar to each other. We repeated the analyses of the Fe₃C standard twenty times and found that the carbon concentration changed only by ± 0.1766 atm%. While chemical compositions were calculated by using both the 'standard' ZAF correction and the phi-rho-z type procedure, the results are consistent with each other (see Table S1).

3. Results

3.1. Carbon contents in eutectic liquid and coexisting solid iron

We have conducted eight separate DAC runs for melting experiments in a pressure range from 23 GPa to 255 GPa (Table 1). Figs. 1 and 2 show typical melting textures on recovered samples. There is a chemically homogeneous part at the center of a laserheating spot, which is non-stoichiometric in composition and should represent a quenched liquid. The liquidus phase (solid phase in contact with liquid) is found at both sides of the liquid. The outermost low-temperature area remained subsolidus. We observed iron carbide (Fe₃C or Fe₇C₃) or metallic iron (or both) as liquidus phase(s). In runs #2, #4, #5, and #8, both iron carbide and iron were present, suggesting that the quenched liquid in these runs may represent a eutectic liquid in the Fe–C binary system. Carbon diffusivity in liquid iron was estimated to be about 10^{-8} m²/s under the present *P-T* conditions (Helffrich, 2014), indicating that 1 s is long enough for carbon to diffuse in the present liquid pool that was less than 20 μ m across (Figs. 1 and 2). Since melting/crystallization at the liquid/solid boundary occurs almost instantaneously, chemical homogeneity in the liquid assures that chemical equilibrium was attained in

the present heating duration of about 5 s. The Fe-C liquids obtained in the core pressure range included about 1 wt.% oxygen in addition to carbon (Table 1), which might have affected phase relations in the Fe-C binary system.

We determined the carbon concentration in quenched liquid up to 255 GPa, the conditions corresponding to the middle of the outer core, more than twice the pressure compared to previous experimental studies (e.g., Lord et al., 2009; Fei and Brosh, 2014; Morard et al., 2017a). Fig. 3 demonstrates the carbon concentrations in liquids as a function of pressure, which constrains the change in eutectic liquid composition, considering that they indicate the lower and upper bounds for the carbon content in eutectic liquid at each pressure when coexisting with iron and iron carbides, respectively. These data show a small reduction in carbon content in the eutectic liquid with increasing pressure above 23 GPa. Our results are not consistent with the continuous reduction that was thermodynamically predicted by Fei and Brosh (2014) and experimentally reported by Lord et al. (2009) (Fig. 3). The difference from Lord et al. (2009) may be attributed to the difference in analytical method: Lord and others employed an X-radiographic imaging technique for chemical analysis, while we based our measurements on electron microprobe analyses.

We also obtained the carbon concentrations in solid iron coexisting with Fe-C liquid as a function of pressure (Fig. 4). Liquid Fe-C coexists with face-centered cubic (fcc) iron at relatively low pressure range, in which carbon atoms occupy the octahedral interstitial sites (Fei and Brosh, 2014). Previous studies reported that the maximum solubility of carbon in fcc-Fe is <2.1 wt.% at 1 bar and predicted that it decreases at higher pressures (Walker et al., 2013; Fei and Brosh, 2014). Our results show about 1.0 wt.% in fcc-Fe from 23 to 54 GPa, in good agreement with such predictions. Above 138

GPa in the present experiments, hcp Fe should have crystallized from Fe-C liquid and again included 1.0 wt.% or less C. The effect of pressure on the solubility of carbon in hcp Fe was not observed.

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214 3.2. Eutectic temperature in the Fe–C system

The temperature at the solid/liquid boundary gives the upper bound for eutectic temperature. These temperatures are plotted as a function of pressure in Fig. 5a, which constrains the eutectic melting curve to mid-outer core pressures. We fit the Simon-Glatzel equation to the data (Simon and Glatzel, 1929):

$$T_m = T_{mref} \left(\frac{P - P_{ref}}{a} + 1\right)^{1/c} \tag{1}$$

220 where T_m is eutectic temperature, and both a and c are fitting parameters. $T_{mref} = 1513$ K at $P_{ref} = 5$ GPa is employed as a reference eutectic temperature from previous data 221 222 obtained using a large-volume press (Fei and Brosh, 2014). The best fitted parameters 223 are $a = 8.4 \pm 3.3$ GPa and $c = 4.1 \pm 0.5$. The obtained eutectic melting curve shown in 224 Fig. 5a is consistent with those of previous studies by Liu et al. (2016) and Morard et al. 225 (2017a) within experimental uncertainties. Fig. 5b compares the eutectic melting curves 226 of iron alloys. Under the core pressure range, the Fe–C system exhibits a eutectic 227 temperature lower than those of Fe (Anzellini et al., 2013), Fe–FeSi (Fischer et al., 228 2013), and Fe–FeO (Morard et al., 2017a). And, the Fe–C eutectic melting temperature 229 is higher than that in Fe-Fe₃S below ~250 GPa, but is lower at higher pressures 230 including the ICB pressure (Mori et al., 2017).

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232 3.3. Liquidus field of Fe₇C₃

233	In runs #7 and #8 using the Fe-4.0 wt.%C starting material, the sample was
234	initially compressed at room temperature to 176 GPa and 226 GPa, respectively (Fig.
235	6). Broad XRD peaks from hcp Fe and Fe ₃ C were observed before heating in both runs.
236	The peaks from both phases became sharper and spotty upon heating to 2300 K in run
237	#7, in which more than 10 peaks were indexed to Fe ₃ C cementite (Sata et al., 2010), and
238	only Fe ₃ C phase was observed from the sample in the XRD pattern during melting at
239	203 GPa and 3350 K. In addition, we confirmed with an FE-EPMA that the quenched
240	liquid was surrounded by Fe_3C in the recovered sample (Fig. 7a). These indicate that Fe
241	forms eutectic melting with Fe ₃ C at 203 GPa.
242	In the case of run #8, the diffraction peaks from hcp Fe and Fe ₃ C became sharper
243	and spotty upon heating. We then observed the appearance of the peaks that are
244	assigned to Fe ₇ C ₃ (orthorhombic, <i>Pbca</i>) (Prescher et al., 2015) after we melted the
245	sample by increasing temperature to 3570 K at 255 GPa. After quenching this sample,
246	we confirmed with the FE-EPMA that the liquid was in contact with both Fe and Fe $_7$ C $_3$,
247	whereas the subsolidus part was composed of Fe and Fe $_3$ C (Fig. 7b). These show that
248	eutectic melting occurs between Fe and Fe ₇ C ₃ at 255 GPa. This observation is not
249	consistent with Liu et al. (2016), who argued that Fe_3C dissociates into $Fe + Fe_7C_3$ at

4. Discussion

above eutectic temperature.

4.1. Phase diagram of Fe-C system at ICB

conclusions purely on the change in XRD patterns, and they may have collected them

- 256 The eutectic liquid composition (Fig. 3) and eutectic melting curve obtained in this 257 study (Fig. 5) indicate that eutectic melting occurs in Fe–Fe₃C at Fe + 3.5 wt.% C at
- 258 135 GPa and 3000 K. And, these data determined up to 255 GPa suggest that the
- eutectic liquid includes about 3 wt.% C at 330 GPa and 3750 K.
- We model the liquidus curve for the Fe-rich side of the eutectic to 330 GPa (Fig.
- 8). Considering an ideal solution, a molar ratio of Fe in liquid at T is given by:

$$\chi_{\text{Fe liq}} = \exp\left\{\frac{\bar{G}_{\text{Fe sol}}^{0} - \bar{G}_{\text{Fe liq}}^{0}}{RT}\right\} = \exp\left\{-\int_{T_{m \text{ Fe}}}^{T} \frac{\bar{H}_{\text{Fe sol}}^{0} - \bar{H}_{\text{Fe liq}}^{0}}{RT^{2}} dT\right\}$$
(2)

- 263 where $\bar{G}^0_{{\rm Fe}\,i}$ and $\bar{H}^0_{{\rm Fe}\,i}$ are standard Gibbs free energy and the enthalpy of phase i,
- respectively, $T_{m \text{ Fe}}$ is the melting temperature of pure Fe, and R is gas constant. When
- 265 we assume that an enthalpy change upon fusion, $\overline{H}_{\text{Fe liq}}^0 \overline{H}_{\text{Fe sol}}^0$, is independent from
- temperature, Eq. (2) is written as:

$$\chi_{\text{Fe liq}} = \exp\left\{ \left(\frac{\overline{H}_{\text{Fe liq}}^0 - \overline{H}_{\text{Fe sol}}^0}{RT_{m \text{ Fe}}} \right) \times \left(1 - \frac{T_{m \text{ Fe}}}{T} \right) \right\} = \exp\left\{ A \times \left(1 - \frac{T_{m \text{ Fe}}}{T} \right) \right\}$$
(3)

- Here A is constant and calculated from $T_{m \text{ Fe}}$ (Anzellini et al., 2013), eutectic
- temperature, and eutectic composition (Figs. 3 and 5) at a given pressure (see also Mori
- 270 et al., 2017).

- 272 *4.2. Fe-C core liquid?*
- Melting phase relations in the Fe–C system constrain the carbon budget in the
- 274 Earth's core and the mineralogy of the solid inner core. The maximum carbon content in
- 275 the outer core has been estimated from the core density deficit. Recent experimental and
- computational studies on the density of liquid Fe-C alloys indicate 1.8 to 4.2 wt.% C in
- the liquid core at the ICB (Badro et al., 2014: Nakajima et al., 2015; Morard et al.,
- 278 2017b), although earlier estimates based on the EoS of solid Fe₃C suggested higher

carbon concentration (7.7–9.2 wt.% C) (Sata et al., 2010). Since the eutectic point is located at about 3 wt.% C at 330 GPa (Fig. 8), if the Fe-C outer core includes less than 3 wt.% C, it crystallizes hcp Fe with ≤1 wt.% C at the ICB (Fig. 4). In this case, however, the difference in carbon concentration between the liquid and solid core is too small to explain the density jump (4.5 to 6.5%) across the ICB (Shearer and Masters, 1990; Masters and Gubbins, 2003). Moreover, assuming a linear density change between pure Fe (Sakamaki et al., 2016), Fe₃C (Sata et al., 2010), and Fe₇C₃ (Prescher et al., 2015), hcp Fe with 1.0 wt.% C is found to be only 1.3% lighter than pure iron, which is not enough to account for the inner core density deficit of 4–6% (Brown and McQueen, 1986; Shearer and Masters, 1990; Masters and Gubbins, 2003; Dewaele et al., 2006; Komabayashi and Fei, 2010). This is supported by the recent calculations by Li et al. (2018). Less than 3 wt.% C as a single light element in the core is therefore unlikely. Next we consider the case that the liquid core contains more than 3 wt.% C as the sole light component, from which Fe₇C₃ crystallizes at the ICB (Fig. 8). It has been suggested that both V_P and V_S —in particular the slow V_S (high Poisson's ratio)—of the inner core could be reconciled with non-magnetic Fe₇C₃ (Mookherjee et al., 2011; Chen et al., 2014; Prescher et al., 2015). If the inner core is composed of solid Fe₇C₃, the liquid outer core should include 5.4-6.7 wt.% C according to Badro et al. (2014) or 5.2-6.0 wt.% C based on Nakajima et al. (2015), considering that the outer core should be 4.5–6.5% less dense than solid Fe₇C₃ (Chen et al., 2014). More than 5 wt.% C in liquid iron, however, leads to a bulk sound velocity and a density much higher and lower than those observed for the outer core, respectively (Badro et al., 2014; Nakajima et al., 2015; Morard et al., 2017b) and is thus unlikely. The computational study by Li et al.

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(2016) also showed that the density of Fe₇C₃ is too low to be a major constituent of the inner core.

These suggest that carbon cannot be a predominant light element in the outer core. It is indeed consistent with Wood et al. (2013), who proposed that the maximum carbon content in the core is 1.0 wt.%, which was deduced from the carbon isotopic signature in silicate Earth and the silicate-metal fractionation factors. It has also been argued that the outer core velocity is explained by the incorporation of a small amount of carbon that is not enough to account for its density deficit from pure iron (Badro et al., 2014; Nakajima et al., 2015).

4.3. Fe-C-X core liquid?

Furthermore, we consider the Fe-C-X (X = Si, S, O, or H) ternary systems to seek possible ranges of outer core composition, from which hcp Fe crystallizes (Figs. 9a–d). The eutectic liquid composition in each Fe–X binary system has been estimated to be Fe with 3 wt.% C (this study), 1.5 wt.% Si (Ozawa et al., 2016), 5 wt.% S (Mori et al., 2017), and >10 wt.% O (Morard et al., 2017a) at 330 GPa. The hydrogen concentration in the Fe–H eutectic liquid is unclear, but it would be around 1 wt.% H according to Fukai (1992) and Shibazaki et al. (2014). The ternary eutectic points in these systems are not exactly known, but the recent melting experiments performed by Tateno et al. (2018) have demonstrated that the Fe–Si–S ternary eutectic point is located 1) near the tie line between the Fe–FeSi and Fe–Fe₃S binary eutectic points and 2) close to the Fe–Fe₃S binary join because the eutectic temperature in Fe–Fe₃S is much lower than that in Fe–FeSi. The present data show that the eutectic temperature in the Fe–C binary system is low compared to those in other systems (Fe–Si, Fe–S, and Fe–O) at 330 GPa (Fig. 5).

327 On the basis of these observations, Figs. 9a–d illustrate the liquidus phase relations 328 (showing a compositional range from which a specific solid phase first crystallizes) in 329 the Fe-C-X ternary systems at the ICB, with each ternary eutectic point located 330 relatively close to the Fe-C binary eutectic point except for the Fe-C-H system. 331 Previous experiments and calculations have estimated the maximum amount of 332 each light element in the liquid core to explain the outer core density deficit. Estimates 333 based on liquid iron alloys show 4.6-6.4 wt.% Si (Morard et al., 2013; Badro et al 334 2014), 5.8–14.0 wt.% S (Morard et al., 2013; Badro et al., 2014; Umemoto et al., 2014; 335 Kawaguchi et al., 2017), 5.5 wt.% O (Badro et al., 2014), and 1.0 wt.% H (Umemoto 336 and Hirose, 2015) at the ICB, which are again less than those obtained by the earlier 337 study on the EoSs of solid FeSi, FeS, and Fe_{0.95}O (Sata et al., 2010). These give the 338 possible compositional range of Fe-C-X ternary outer core liquid in Figs. 9a-d (green 339 area). 340 We find that hcp Fe crystallizes from possible liquid core compositions in all of the 341 Fe-C-Si, Fe-C-S, Fe-C-O, and Fe-C-H systems. The recent calculations by Li et al. 342 (2018) proposed that all of the density, V_P , and V_S of the solid inner core are explained 343 by hcp $Fe_{30}C_1Si_1$ (Fe + 0.7 wt.% C + 1.6 wt.% Si). Such inner core composition is 344 feasible, considering that the maximum solubilities of silicon and sulfur in hcp Fe when 345 coexisting with liquid alloys are ≤1.5 wt.% Si and 4 wt.% S at the ICB, respectively 346 (Ozawa et al., 2016; Mori et al., 2017). Hcp Fe + 1 wt.% C + 4.0 wt.% S also could 347 account for the inner core density (Sakai et al., 2012). Hydrogen is known to be soluble 348 in solid Fe (Fukai, 1992; Shibazaki et al., 2012; Terasaki et al., 2012), suggesting that 349 the hcp Fe-C-H inner core is also feasible. On the other hand, solid Fe crystallizing from 350 liquid Fe-C-O contains ≤1 wt.% carbon and a negligible amount of oxygen (Ozawa et

al., 2010), which does not account for the inner core density deficit (Li et al., 2018). Therefore, while carbon is not a predominant light element in the core, the outer core liquid can be Fe-C-Si, Fe-C-S, or Fe-C-H.

5. Conclusion

We have conducted melting experiments on the Fe–C binary system up to 255 GPa by using a laser-heated DAC. The carbon concentration in eutectic liquid and the solubility of carbon in solid iron were determined based on textual and chemical characterizations for samples recovered from the DAC. Results show that 1) the carbon content in eutectic liquid slightly decreases with increasing pressure from 4.3 wt.% C at ambient pressure to 3.6 wt.% C at 255 GPa, 2) eutectic melting occurs between Fe and Fe₃C to 203 GPa and between Fe and Fe₇C₃ at 255 GPa, and 3) the maximum solubility of carbon in solid iron is almost constant (\leq 1 wt.%) up to 255 GPa. The eutectic liquid in the Fe–C binary system includes about 3 wt.% C at 330 GPa.

If carbon is the sole light element in the core and the outer core includes <3 wt.% C, hcp Fe crystallizes at the ICB but includes only less than 1 wt.% C, which does not explain the inner core density deficit. If the liquid core contains more than 3 wt.% C, it forms Fe₇C₃ at the ICB, whose density is too low for the inner core. These indicate that carbon is not a primary light element in the core. On the other hand, the outer core can be liquid Fe-C-Si, Fe-C-S, or Fe-C-H. It crystallizes hcp Fe including light elements with less than 1 wt.% C, which may explain the density and the sound velocities observed in the inner core.

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385 **References**

- 386 Akahama, Y., Kawamura, H., 2007. Diamond anvil Raman gauge in multimegabar
- pressure range. High Press. Res. 27, 473–482.
- 388 https://doi.org/10.1080/08957950701659544
- Anders, E., Grevesse, N., 1989. Abundances of the elements: meteoritic and solar.
- 390 Geochim. Cosmochim. Acta 53, 197–214. https://doi.org/10.1016/0016-
- 391 7037(89)90286-X
- 392 Andrault, D., Fiquet, G., Itie, J.P., Richet, P., Gillet, P., Hausermann, D., Hanfland, M.,
- 393 1998. Thermal pressure in the laser-heated diamond-anvil cell: an X-ray diffraction
- 394 study. Eur. J. Mineral. 10, 931–940.
- 395 Anzellini, S., Dewaele, A., Mezouar, M., Loubeyre, P., Morard, G., 2013. Melting of
- iron at earth's inner core boundary based on fast X-ray diffraction. Science 340,
- 397 464–466. https://doi.org/10.1126/science.1233514
- 398 Badro, J., Cote, A.S., Brodholt, J.P., 2014. A seismologically consistent compositional
- model of Earth's core. Proc. Natl. Acad. Sci. USA 111, 7542–7545.
- 400 https://doi.org/10.1073/pnas.1316708111
- 401 Benz, M., Elliott, J., 1961. The austenite solidus and revised iron-carbon diagram.
- 402 Trans. Metall. Soc. AIME 221, 323–331.
- Brown, J.M., McQueen, G., 1986. Phase transitions, Grüneisen parameter, and elasticity
- for shocked iron between 77 GPa and 400 GPa. J. Geophys. Res. 91, 7485–7494.
- 405 https://doi.org/10.1029/JB091iB07p07485
- 406 Chabot, N.L., Campbell, A.J., McDonough, W.F., Draper, D.S., Agee, C.B., Humayun,
- M., Watson, H.C., Cottrell, E., Saslow, S.A., 2008. The Fe–C system at 5 GPa and

- implications for Earth's core. Geochim. Cosmochim. Acta 72, 4146–4158.
- 409 https://doi.org/10.1016/j.gca.2008.06.006
- 410 Chen, B., Li, Z., Zhang, D., Liu, J., Hu, M.Y., Zhao, J., Bi, W., Alp, E.E., 2014. Hidden
- carbon in Earth's inner core revealed by shear softening in dense Fe₇C₃. Proc.
- 412 Natl. Acad. Sci. USA 111, 17755–17758. https://doi.org/10.1073/pnas.1411154111
- Dasgupta, R., Walker, D., 2008. Carbon solubility in core melts in a shallow magma
- ocean environment and distribution of carbon between the Earth's core and the
- 415 mantle. Geochim. Cosmochim. Acta 72, 4627–4641.
- 416 https://doi.org/10.1016/j.gca.2008.06.023
- Dewaele, A., Loubeyre, P., Occelli, F., Mezouar, M., Dorogokupets, P. I., Torrent, M.,
- 418 2006. Quasihydrostatic equation of state of iron above 2 Mbar. Phys. Rev. Lett. 97,
- 419 215504. https://doi.org/10.1103/PhysRevLett.97.215504
- 420 Fei, Y., Brosh, E., 2014. Experimental study and thermodynamic calculations of phase
- relations in the Fe–C system at high pressure. Earth Planet. Sci. Lett. 408, 155–
- 422 162. https://doi.org/10.1016/j.epsl.2014.09.044
- 423 Fischer, R.A., Campbell, A.J., Reaman, D.M., Miller, N.A., Heinz, D.L., Dera, P.,
- 424 Prakapenka, V.B., 2013. Phase relations in the Fe–FeSi system at high pressures
- and temperatures. Earth Planet. Sci. Lett. 373, 54–64.
- 426 https://doi.org/10.1016/j.epsl.2013.04.035
- 427 Fukai, Y., 1992. Some properties of the Fe–H system at high pressures and
- temperatures, and their implications for the Earth's core. In: Syono Y., Manghnani
- M.H. (Eds.), High-Pressure Research: Application to Earth and Planetary Sciences.
- 430 AGU monogr. ser. 67, pp. 373–385. https://doi.org/10.1029/GM067p0373

- Helffrich, G., 2014. Outer core compositional layering and constraints on core liquid
- transport properties. Earth Planet. Sci. Lett. 391, 256–262.
- 433 https://doi.org/10.1016/j.epsl.2014.01.039
- Hillgren, V.J., Gessmann, C.K., Li, J., 2000. An experimental perspective on the light
- element in Earth's core. In: Canup, R.M., Righter, K. (Eds.), Origin of the Earth
- and Moon. Univ. Arizona Press, pp. 245–263.
- Hirose, K., Morard, G., Sinmyo, R., Umemoto, K., Hernlund, J.W., Labrosse, S., 2017.
- 438 Crystallization of silicon dioxide and compositional evolution of the Earth's core.
- 439 Nature 543, 99–102. https://doi.org/10.1038/nature21367
- Ichikawa, H., Tsuchiya, T., Tange, Y., 2014. The P-V-T equation of state and
- thermodynamic properties of liquid iron. J. Geophys. Res. Solid Earth 119, 240–
- 442 252. https://doi.org/10.1002/2013JB010732
- Kawaguchi, S. I., Nakajima, Y., Hirose, K., Komabayashi, T., Ozawa, H., Tateno, S.,
- Kuwayama, Y., Tsutsui, S., Baron, A.Q., 2017. Sound velocity of liquid Fe–Ni–S
- at high pressure. J. Geophys. Res. Solid Earth 122, 3624–3634.
- 446 https://doi.org/10.1002/2016JB013609
- Komabayashi, T., Fei, Y., 2010. Internally consistent thermodynamic database for iron
- to the Earth's core conditions. J. Geophys. Res. Solid Earth 115, 1–12.
- 449 https://doi.org/10.1029/2009JB006442
- 450 Li, Y., Vočadlo, L., Brodholt, J., Wood, I.G., 2016. Thermoelasticity of Fe₇C₃ under
- inner core conditions. J. Geophys. Res. Solid Earth 121, 5828–5837.
- 452 https://doi.org/10.1002/2016JB013155

- Li, Y., Vočadlo, L., Brodholt, J.P., 2018. The elastic properties of hcp-Fe alloys under
- the conditions of the Earth's inner core. Earth Planet. Sci. Lett. 493, 118–127.
- 455 https://doi.org/10.1016/j.epsl.2018.04.013
- Liu, J., Lin, J., Prakapenka, V.B., Prescher, C., Yoshino, T., 2016. Phase relations of
- 457 Fe₃C and Fe₇C₃ up to 185 GPa and 5200 K: implication for the stability of iron
- carbide in the Earth's core. Geophys. Res. Lett. 43, 12415–12422.
- 459 https://doi.org/10.1002/2016GL071353.
- Lord, O.T., Walter, M.J., Dasgupta, R., Walker, D., Clark, S.M., 2009. Melting in the
- 461 Fe–C system to 70 GPa. Earth Planet. Sci. Lett. 284, 157–167.
- 462 https://doi.org/10.1016/j.epsl.2009.04.017
- Masters, G., Gubbins, D., 2003. On the resolution of density within the Earth. Phys.
- Earth Planet. Inter. 140, 159–167. https://doi.org/10.1016/j.pepi.2003.07.008
- McDonough, W.F., 2003. Compositional model for the Earth's core. In: Holland, H.,
- Turekian, K. (Eds.), Treatise on Geochemistry: 2nd Edition. Elsevier Ltd., pp.
- 467 559–577. https://doi.org/10.1016/B978-0-08-095975-7.00215-1
- 468 Mookherjee, M., Nakajima, Y., Neumann, G.S., Glazyrin, K., Wu, X., Dubrovinsky, L.,
- Mccammon, C., Chumakov, A., 2011. High-pressure behavior of iron carbide
- 470 (Fe₇C₃) at inner core conditions. J. Geophys. Res. Solid Earth 116, 1–13.
- 471 https://doi.org/10.1029/2010JB007819
- 472 Morard, G., Andrault, D., Guignot, N., Siebert, J., Garbarino, G., Antonangeli, D., 2011.
- 473 Melting of Fe–Ni–Si and Fe–Ni–S alloys at megabar pressures: Implications for
- the core-mantle boundary temperature. Phys. Chem. Miner. 38, 767–776.
- 475 https://doi.org/10.1007/s00269-011-0449-9

- 476 Morard, G., Siebert, J., Andrault, D., Guignot, N., Garbarino, G., Guyot, F.,
- Antonangeli, D., 2013. The Earth's core composition from high pressure density
- 478 measurements of liquid iron alloys. Earth Planet. Sci. Lett. 373, 169–178.
- 479 https://doi.org/10.1016/j.epsl.2013.04.040
- 480 Morard, G., Andrault, D., Antonangeli, D., Nakajima, Y., Auzende, A.L., Boulard, E.,
- Cervera, S., Clark, A., Lord, O.T., Siebert, J., Svitlyk, V., Garbarino, G., Mezouar,
- 482 M., 2017a. Fe–FeO and Fe–Fe₃C melting relations at Earth's core-mantle
- boundary conditions: Implications for a volatile-rich or oxygen-rich core. Earth
- 484 Planet. Sci. Lett. 473, 94–103. https://doi.org/10.1016/j.epsl.2017.05.024
- 485 Morard, G., Nakajima, Y., Andrault, D., Antonangeli, D., Auzende, A.L., Boulard, E.,
- 486 Cervera, S., Clark, A.N., Lord, O.T., Siebert, J., Svitlyk, V., Garbarino, G.,
- Mezouar, M., 2017b. Structure and density of Fe–C liquid alloys under high
- 488 pressure. J. Geophys. Res. Solid Earth 122, 7813–7823.
- 489 https://doi.org/10.1002/2017JB014779
- 490 Mori, Y., Ozawa, H., Hirose, K., Sinmyo, R., Tateno, S., Morard, G., Ohishi, Y., 2017.
- Melting experiments on Fe–Fe₃S system to 254 GPa. Earth Planet. Sc. Lett. 464,
- 492 135–141. https://doi.org/10.1016/j.epsl.2017.02.021
- Nakajima, Y., Takahashi, E., Suzuki, T., Funakoshi, K., 2009. "Carbon in the core"
- revisited. Phys. Earth Planet. Inter. 174, 202–211.
- 495 https://doi.org/10.1016/j.pepi.2008.05.014
- Nakajima, Y., Imada, S., Hirose, K., Komabayashi, T., Ozawa, H., Tateno, S., Tsutsui,
- S., Kuwayama, Y., Baron, A.Q.R., 2015. Carbon-depleted outer core revealed by
- sound velocity measurements of liquid iron-carbon alloy. Nat. Commun. 6, 1–7.
- 499 https://doi.org/10.1038/ncomms9942

500 Ohishi, Y., Hirao, N., Sata, N., Hirose, K., Takata, M., 2008. Highly intense 501 monochromatic X-ray diffraction facility for high-pressure research at SPring-8. 502 High Press. Res. 28, 163–173. https://doi.org/10.1080/08957950802208910 503 Ozawa, H., Hirose, K., Tateno, S., Sata, N., Ohishi, Y. 2010. Phase transition boundary 504 between B1 and B8 structures of FeO up to 210 GPa. Phys. Earth Planet. Inter. 505 179, 157–163. https://doi:10.1016/j.pepi.2009.11.005 506 Ozawa, H., Hirose, K., Yonemitsu, K., Ohishi, Y., 2016. High-pressure melting 507 experiments on Fe-Si alloys and implications for silicon as a light element in the 508 core. Earth Planet. Sci. Lett. 456, 47–54. https://doi.org/10.1016/j.epsl.2016.08.042 509 Prescher, C., Dubrovinsky, L., Bykova, E., Kupenko, I., Glazyrin, K., Kantor, A., 510 McCammon, C., Mookherjee, M., Nakajima, Y., Miyajima, N., Sinmyo, R., 511 Cerantola, V., Dubrovinskaia, N., Prakapenka, V., Rüffer, R., Chumakov, A., 512 Hanfland, M., 2015. High Poisson's ratio of Earth's inner core explained by carbon 513 alloying. Nat. Geosci. 8, 220–223. https://doi.org/10.1038/ngeo2370 514 Sakai T., Ohtani, E., Kamada, S., Terasaki, H., Hirao, N. 2012. Compression of 515 Fe_{88.1}Ni_{9.1}S_{2.8} alloy up to the pressure of Earth's inner core. J. Geophys. Res. Solid 516 Earth 117, B02210. https://doi:10.1029/2011JB008745 517 Sakamaki, T., Ohtani, E., Fukui, H., Kamada, S., Takahashi, S., Sakairi, T., Takahata, 518 A., Sakai, T., Tsutsui, S., Ishikawa, D., Shiraishi, R., Seto, Y., Tsuchiya, T., Baron, 519 A.Q.R., 2016. Constraints on Earths inner core composition inferred from 520 measurements of the sound velocity of hcp-iron in extreme conditions. Sci. Adv. 2, 521 e1500802-e1500802. https://doi.org/10.1126/sciadv.1500802 522 Sata, N., Hirose, K., Shen, G., Nakajima, Y., Ohishi, Y., Hirao, N., 2010. Compression 523 of FeSi, Fe₃C, Fe_{0.95}O, and FeS under the core pressures and implication for light

- element in the Earth's core. J. Geophys. Res. Solid Earth 115, 1–13.
- 525 https://doi.org/10.1029/2009JB006975
- Seto, Y., Nishio-Hamane, D., Nagai, T., Sata, N., 2010. Development of a software
- suite on X-ray diffraction experiments. Rev. High Press. Sci. Technol. 20, 269–
- 528 276. https://doi.org/10.4131/jshpreview.20.269
- 529 Shearer, P., Masters, G., 1990. The density and shear velocity contrasts at the inner core
- boundary. Geohys. J. Int. 102, 408–491. https://doi.org/10.1111/j.1365-
- 531 246X.1990.tb04481.x
- 532 Shibazaki, Y., Ohtani, E., Fukui, H., Sakai, T., Kamada, S., Ishikawa, D., Tsutsui, S.,
- Baron, A.Q.R., Nishitani, N., Hirao, N., Takemura, K., 2012. Sound velocity
- measurements in dhcp-FeH up to 70 GPa with inelastic X-ray scattering:
- implications for the composition of the Earth's core. Earth Planet. Sci. Lett. 313–
- 536 314, 79–85. https://doi.org/10.1016/j.epsl.2011.11.002
- 537 Shibazaki, Y., Terasaki, H., Ohtani, E., Tateyama, R., Nishida, K., Funakoshi, K., Higo,
- 538 Y., 2014. High-pressure and high-temperature phase diagram for Fe_{0.9}Ni_{0.1}–H
- 539 alloy. Phys. Earth Planet. Inter. 228, 192–201.
- 540 https://doi.org/10.1016/j.pepi.2013.12.013
- 541 Simon, F., Glatzel, G., 1929. Remarks on fusion pressure curve. Z. Anorg. Allg. Chem.
- 542 178, 309–316. https://doi.org/10.1002/zaac.19291780123
- Tateno, S., Hirose, K., Ohishi, Y., Tatsumi, Y., 2010. The structure of iron in Earth's
- inner core. Science 330, 359–361. https://doi.org/10.1126/science.1194662
- Tateno, S., Kuwayama, Y., Hirose, K., Ohishi, Y. 2015. The structure of Fe–Si alloy in
- Earth's inner core. Earth Planet. Sci. Lett. 418, 11–19.
- 547 http://dx.doi.org/10.1016/j.epsl.2015.02.008

- Tateno, S., Hirose, K., Sinmyo, R., Morard, G., Hirao, N., Ohishi, Y., 2018. Melting
- experiments on Fe–Si–S alloys to core pressures: silicon in the core? Am. Miner.
- 550 103, 742–748. http://doi.org/10.2138/am-2018-6299
- Terasaki, H., Ohtani, E., Sakai, T., Kamada, S., Asanuma, H., Shibazaki, Y., Hirao, N.,
- Sata, N., Ohishi, Y., Sakamaki, T., Suzuki, A., Funakoshi, K., 2012. Stability of
- Fe–Ni hydride after the reaction between Fe–Ni alloy and hydrous phase (δ-
- AlOOH) up to 1.2Mbar: possibility of H contribution to the core density deficit.
- 555 Phys. Earth Planet. Inter. 194–195, 18–24.
- 556 https://doi.org/10.1016/j.pepi.2012.01.002
- 557 Umemoto, K., Hirose, K., 2015. Liquid iron–hydrogen alloys at outer core conditions
- by first-principles calculations. Geophys. Res. Lett. 42, 7513–7520.
- 559 https://doi.org/10.1002/2015GL065899
- Umemoto, K., Hirose, K., Imada, S., Nakajima, Y., Komabayashi, T., Tsutsui, S.,
- Baron, A.Q.R., 2014. Liquid iron–sulfur alloys at outer core conditions by first-
- principles calculations. Geophys. Res. Lett. 41, 6712–6717.
- 563 https://doi.org/10.1002/2014GL061233
- Walker, D., Dasgupta, R., Li, J., Buono, A., 2013. Nonstoichiometry and growth of
- some Fe carbides. Contrib. Mineral. Petrol. 166, 935–957.
- 566 https://doi.org/10.1007/s00410-013-0900-7
- 567 Wood, B.J., 1993. Carbon in the core. Earth Planet. Sci. Lett. 117, 593–607.
- Wood, B.J., Li, J., Shahar, A., 2013. Carbon in the core: its influence on the properties
- of core and mantle. In: Hazen, R.M., Jones, A.P., Baross, J.A. (Eds.), Carbon in
- 570 Earth. Rev. Mineral. Geochem. 75, pp. 231–250.
- 571 https://doi.org/10.2138/rmg.2013.75.8

572 573 **Figure captions** 574 575 Fig. 1. (a) Temperature profile and (b) composite X-ray maps for oxygen (red), carbon 576 (purple), iron (green), and aluminum (light blue) for run #3 performed at 138 GPa (see 577 Fig. S1 for raw radiation spectra). Homogeneous quenched liquid is found at the center 578 (hot part) surrounded by solid layers of iron. A measured temperature profile is 579 combined in (a), which provides temperature at the liquid/solid boundary. 580 581 Fig. 2. Typical melting texture of a quenched sample and composite X-ray maps for 582 oxygen, carbon, aluminum, and iron for run #5 at 40 GPa. SIM, scanning ion 583 microscope image; SL, secondary electron image; Comp, backscattered image. 584 Homogeneous quenched liquid is found at the center surrounded by solid layers of Fe 585 and Fe₃C. 586 587 Fig. 3. Carbon concentrations in liquids obtained in this study. Liquid compositions are 588 shown by triangles (blue) when coexisting with solid Fe, inverse triangles (purple) when 589 in contact with Fe₃C, squares (red) when both solid Fe and Fe₃C were found, and a 590 diamond (orange) when solid Fe and Fe₇C₃ were formed. The triangles and inverse 591 triangles give lower and upper bounds for carbon concentration in eutectic liquid, 592 respectively. Open and cross symbols represent eutectic compositions reported by 593 previous studies (Chabot et al., 2008; Lord et al., 2009; Fei and Brosh, 2014; Morard et 594 al., 2017a). Green curve shows the change in eutectic liquid composition with 595 increasing pressure.

596

- **Fig. 4.** Carbon contents in solid iron found in this study. Closed symbols are the same as
- 598 in Fig. 3. Previous reports on the maximum solubility of carbon in solid iron are shown
- by open symbols (Walker et al., 2013; Fei and Brosh, 2014).

600

- **Fig. 5.** (a) Fe–Fe₃C (or Fe–Fe₇C₃ above 255 GPa) eutectic melting curve (orange)
- obtained in this study (closed symbols). It agrees well with the previously determined
- Fe-C melting curves by Morard et al. (2017a) (blue) and Liu et al. (2015) (green). Black
- dotted curve represents the solidus of Fe₃C composition reported by Liu et al. (2016).
- Purple dashed curve shows the melting temperature of pure Fe (Anzellini et al., 2013).
- 606 (b) Eutectic melting curve of the Fe–C system obtained in this study compared with
- those of the Fe-Fe₃S (Mori et al., 2017), Fe-FeO (Morard et al., 2017a), and Fe-FeSi
- 608 (Fischer et al., 2013) binary systems.

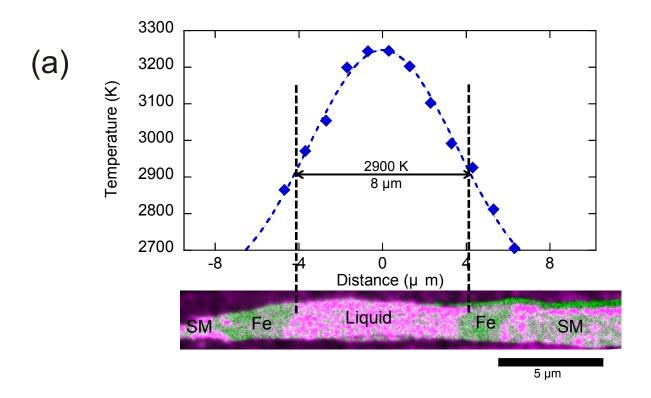
609

- Fig. 6. XRD patterns of Fe + 4.0 wt.% C obtained before and after heating (no melting),
- and during and after melting (a) at 203 GPa in run #7 and (b) 255 GPa in run #8. The
- unit-cell parameters and volumes of Fe₃C and Fe₇C₃ are a = 3.95(1) Å, b = 4.54(1) Å, c = 4.54(1) Å, c = 4.54(1) Å
- 613 = 5.95(1) Å, and V = 106.9(4) Å³ for Fe₃C, and a = 10.40(1) Å, b = 3.96(1) Å, c = 10.40(1) Å
- 614 12.06(1) Å, and V = 496.4(3) Å³ for Fe₇C₃.

- **Fig. 7.** Quenched textures and composite X-ray maps for carbon, oxygen, and iron at (a)
- 617 203 GPa (run #7) and (b) 255 GPa (run #8). While liquid was surrounded by Fe₃C in
- the sample quenched at 203 GPa, it was in contact with Fe₇C₃ and Fe at 255 GPa (the
- subsolidus part was composed of Fe₃C and Fe).

620 621 Fig. 8. Phase diagrams for the iron-rich portion of the Fe–C binary system at 54, 138, 622 203, and 255 GPa from the present experiments and at ICB conditions (330 GPa) by 623 extrapolation. Red and blue (and gray) curves represent Fe–Fe₃C and Fe–Fe₇C₃ 624 eutectic, respectively. 625 626 Fig. 9. Likely liquidus phase relations in the Fe–C–X ternary systems at 330 GPa. Red 627 circles represent the eutectic compositions in the Fe-C (this study), Fe-Si (Ozawa et al., 628 2016), Fe-S (Mori et al., 2017), Fe-O (Morard et al., 2017a), and Fe-H binary systems 629 (Fukai, 1992). Green regions show the range of possible composition of the outer core 630 that explains its density based on studies on liquid iron alloys (Morard et al., 2013; 631 Badro et al., 2014; Umemoto et al., 2014; Umemoto and Hirose, 2015; Nakajima et al., 632 2015; Kawaguchi et al., 2017).

Table 1 Experim	Fable 1 Experimental results in the Fe-C binary system	he Fe-C bina	ary system							
Run	Run# Starting material P (GPa)	P (GPa)	eutectic T (K)	eutectic T (K) Phases present	Liquid			Solid iron		
)	,	,	•	C (wt%)	C (wt%) Fe (wt%) O (wt%)	O (wt%)	C (wt%)	C (wt%) Fe $(wt%)$ O $(wt%)$	O (wt%)
_	Fe-1.9 wt%C	23(2)	1950(100)	Liq + Fe	4.04(30)	96.30(30)	0.16(5)	0.85(24)	99.08(16)	0.07(11)
7	Fe-1.9 wt%C	54(5)	2550(130)	$Liq + Fe + Fe_3C$	4.45(32)	95.70(30)	0.34(30)	1.00(45)	98.89(51)	0.11(7)
\mathcal{E}	Fe-1.9 wt%C	138(14)	2900(150)	Liq + Fe	3.95(22)	94.84(21)	1.70(17)	0.68(48)	99.38(29)	0.05(11)
4	Fe-1.9 wt%C	211(21)	3230(160)	$Liq + Fe + Fe_3C$	4.19(20)	95.23(30)	0.95(30)	1.04(22)	98.96(24)	not detected
2	Fe-4.0 wt%C	40(4)	2300(120)	$Liq + Fe + Fe_3C$	3.83(24)	96.50(23)	0.17(22)	0.64(4)	99.49(20)	not detected
9	Fe-4.0 wt%C	83(8)	2640(130)	$Liq + Fe_3C$	3.90(23)	95.72(12)	0.88(8)	•		1
7	Fe-4.0 wt%C	203(20)	3350(170)	$Liq + Fe_3C$	3.92(9)	95.39(22)	1.19(21)			1
∞	Fe-4.0 wt%C	255(25)	3570(180)	$Liq + Fe + Fe_7C_3$	3.60(21)	96.84(49)	1.06(50)	0.61(34)	99.43(31)	99.43(31) not detected
Nun	bers in parentheses	are standarc	deviation uncert	Numbers in parentheses are standard deviation uncertainties in the last digit(s)	;it(s).					



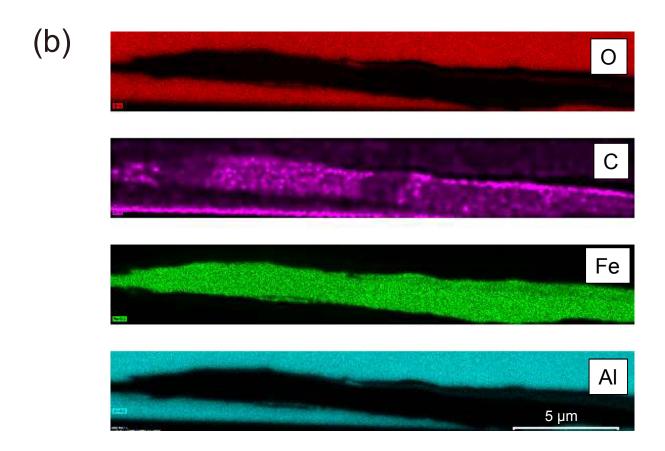


Fig. 1

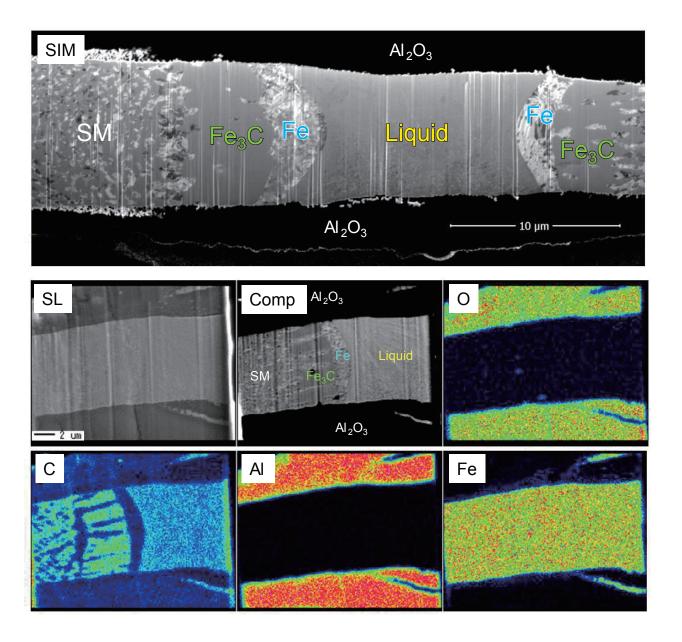


Fig. 2

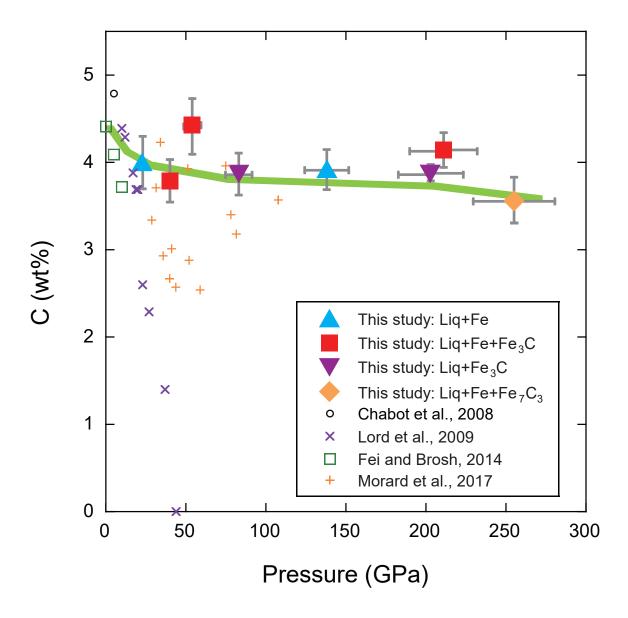


Fig. 3

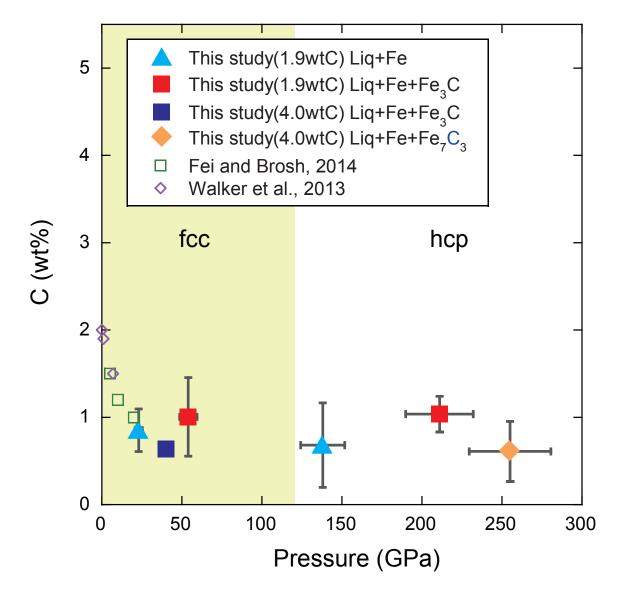


Fig. 4

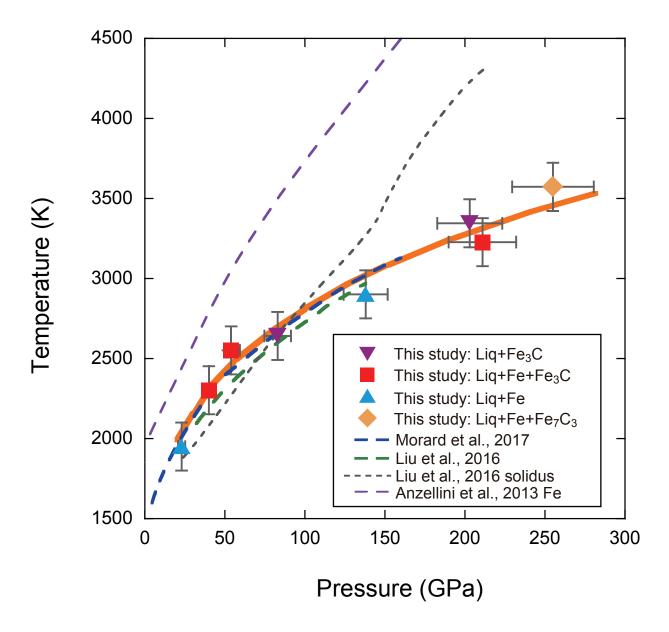


Fig. 5 (a)

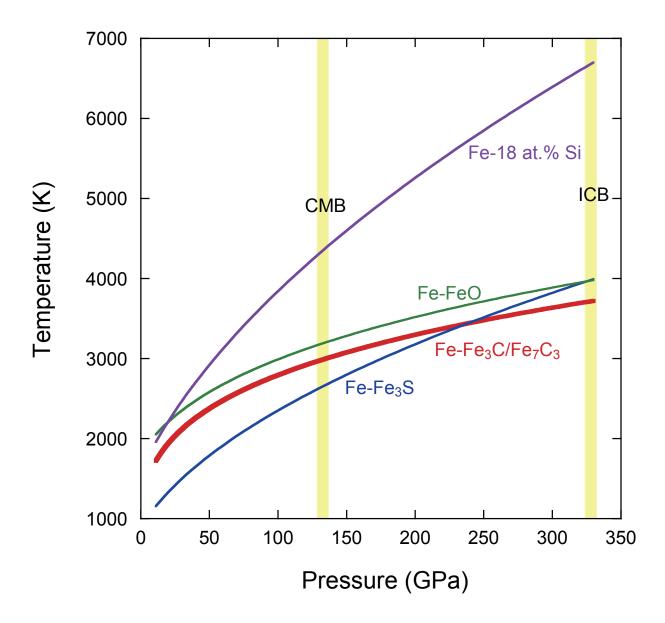
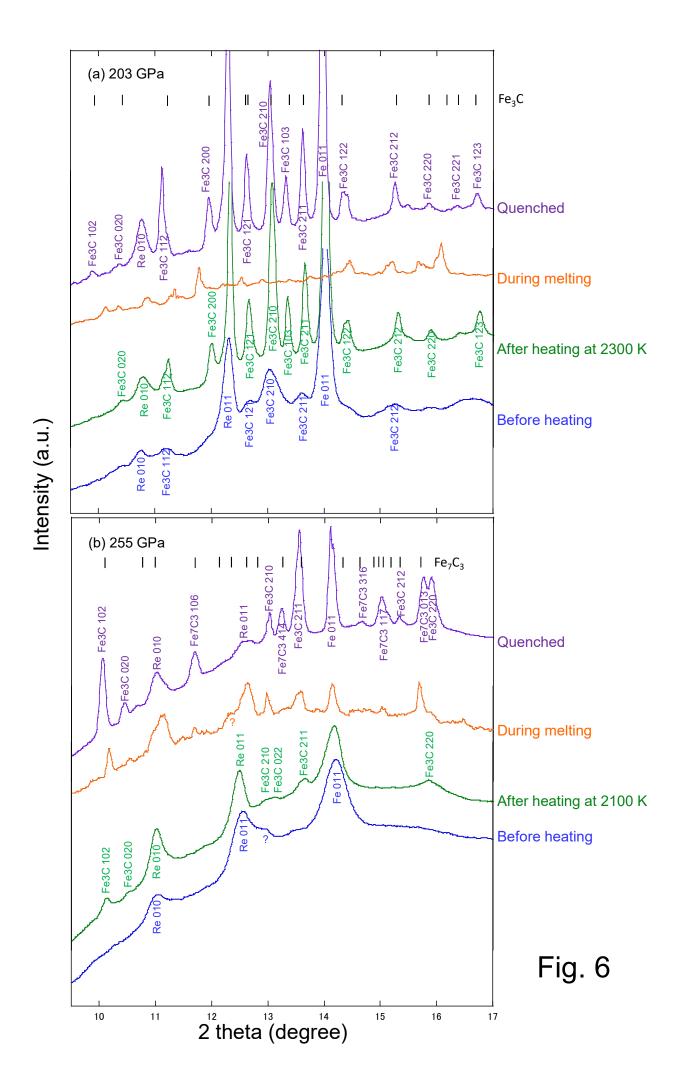


Fig. 5 (b)



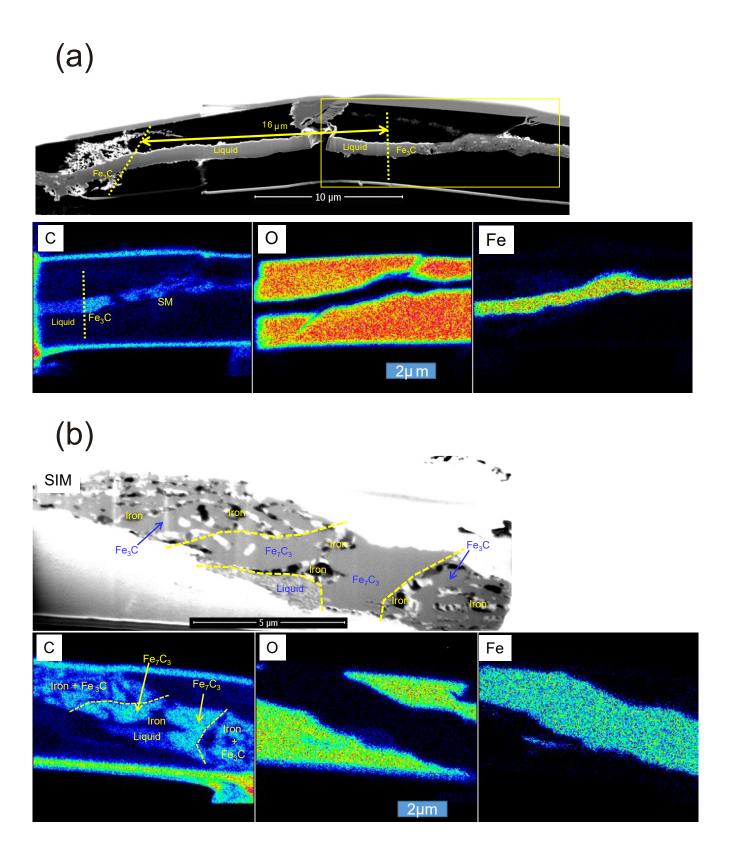


Fig. 7

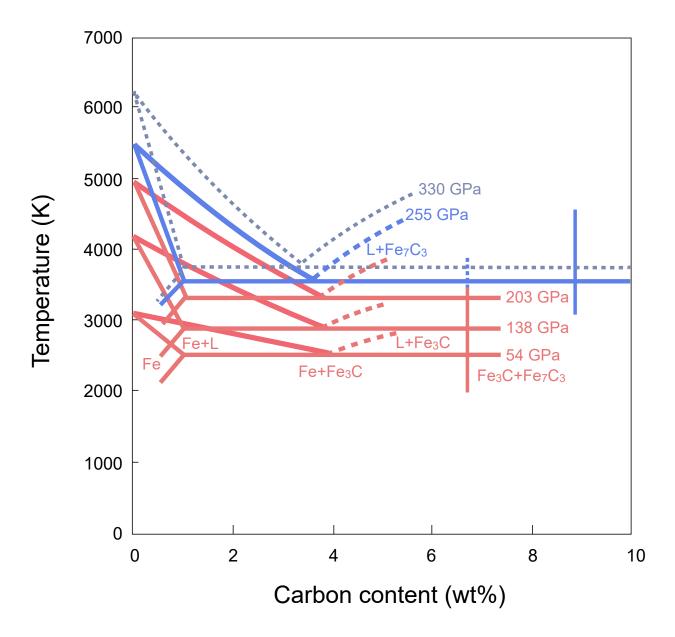


Fig. 8

