



HAL
open science

High contribution of Rhizaria (Radiolaria) to vertical export in the California Current Ecosystem revealed by DNA metabarcoding

Andres Gutierrez-Rodriguez, Michael R. Stukel, Adriana Lopes dos Santos, Tristan Biard, Renate Scharek, Daniel Vaultot, Michael Landry, Fabrice Not

► To cite this version:

Andres Gutierrez-Rodriguez, Michael R. Stukel, Adriana Lopes dos Santos, Tristan Biard, Renate Scharek, et al.. High contribution of Rhizaria (Radiolaria) to vertical export in the California Current Ecosystem revealed by DNA metabarcoding. *The International Society of Microbiological Ecology Journal*, 2019, 13 (4), pp.964-976. 10.1038/s41396-018-0322-7 . hal-02323631

HAL Id: hal-02323631

<https://hal.sorbonne-universite.fr/hal-02323631>

Submitted on 21 Oct 2019

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1 **High contribution of Rhizaria (Radiolaria)**
2 **to vertical export in the California Current Ecosystem**
3 **revealed by DNA metabarcoding**

4 **Authors**

5 Andres Gutierrez-Rodriguez^{1,2}, Michael R. Stukel³, Adriana Lopes dos Santos^{1,4}, Tristan
6 Biard^{1,5}, Renate Scharek⁶, Daniel Vaultot¹, Michael R. Landry⁵, and Fabrice Not¹

7
8 **Affiliations**

9 ¹ Sorbonne Université, CNRS, UMR7144, Ecology of Marine Plankton team, Station
10 Biologique de Roscoff, 29680 Roscoff, France

11 ² National Institute of Water and Atmospheric Research, 301 Evans Bay Parade,
12 Wellington 6021, New Zealand

13 ³ Department of Earth, Ocean, and Atmospheric Science, Florida State University,
14 Tallahassee, Florida 32304, USA

15 ⁴ GEMA Center for Genomics, Ecology & Environment, Facultad de Ciencias, Universidad
16 Mayor, Camino La Pirámide 5750, Huechuraba, Santiago, Chile.

17 ⁵ Scripps Institution of Oceanography, University of California, San Diego, La Jolla,
18 California 92093, USA

19 ⁶ Instituto Español de Oceanografía, Centro oceanográfico de Gijón, 33212, Gijón, Spain

20 **Corresponding author**

21 Andres Gutierrez-Rodriguez

22 National Institute of Water and Atmospheric Research, 301 Evans Bay Parade,
23 Wellington 6021, New Zealand

24 Phone: +64 4 386 0815

25 Fax: +64 7 856 0151

26 Email: andres.gutierrez@niwa.co.nz

27 Submitted to ISME Journal, March 6, 2018

28 Resubmitted to ISME Journal, October 19, 2018

29 ABSTRACT

30 Passive sinking of particulate organic matter (POM) is the main mechanism through
31 which the biological pump transports surface primary production to the ocean interior.
32 However, the contribution and variability of different biological sources to vertical export
33 is not fully understood. Here, we use DNA metabarcoding of the 18S rRNA gene and
34 particle interceptor traps (PITs) to characterize the taxonomic composition of particles
35 sinking out of the photic layer in the California Current Ecosystem (CCE), a
36 productive system with high export potential. The PITs included formalin-fixed and
37 'live' traps to investigate eukaryotic communities involved in the export and
38 remineralization of sinking particles. Sequences affiliated with Radiolaria dominated the
39 eukaryotic assemblage in fixed traps (90%), with Dinophyta and Metazoa making minor
40 contributions. The prominence of Radiolaria decreased drastically in live traps, possibly
41 due to selective consumption by copepods, heterotrophic nanoflagellates and phaeodarians
42 that were heavily enriched in these traps. These patterns were consistent across the water
43 masses surveyed extending from the coast to offshore, despite major differences in
44 productivity and trophic structure of the epipelagic plankton community. Our findings
45 identify Radiolaria as major actors in export fluxes in the CCE.

46

47

48

INTRODUCTION

49

50

51

52

53

54

55

56

57

58

59

60

The main mechanisms of the biological pump include the gravitational sinking of particles, the active transport associated with zooplankton, and the mixing and diffusive transport of dissolved and particulate organic matter (DOM and POM) (Turner, 2015). Among them, particle sinking is the main process contributing to carbon export and is responsible for 5-21 PgC y⁻¹ (Eppley and Peterson, 1979; Laws *et al.*, 2000; Henson *et al.*, 2011). Phytoplankton community structure and food-web processes determine the fraction of net primary production exported as well as the size and chemical characteristics of sinking material (Boyd and Newton, 1999; Guidi *et al.*, 2009; Stukel *et al.*, 2011). The fraction of exported particles that reaches the ocean interior is further controlled by biotic (mainly microbial and zooplankton) transformations during their downward transit, which affects remineralization rates and particulate organic carbon (POC) flux attenuation with depth (Lampitt *et al.*, 1990; Steinberg *et al.*, 2008; Giering *et al.*, 2014).

61

62

63

64

65

66

67

68

69

70

71

72

73

Sinking particles are composed of zooplankton fecal pellets (Steinberg and Landry, 2017), organic aggregates of various source ('marine snow') including mucilaginous structures of larger plankton (Caron *et al.*, 1986; Alldredge and Silver, 1988) and intact phytoplankton cells (Martin *et al.*, 2011; Smetacek *et al.*, 2012; Agusti *et al.*, 2015). Drifting particle interceptor traps (PITs) are the most common approach for quantifying vertical fluxes of particles (Knauer *et al.*, 1979). However, the material collected in PITs is often partially degraded and heavily transformed by biological activity, which hampers its taxonomic identification based on morphological attributes and limits the ability to distinguish biological sources and export mechanisms. Moreover, export fluxes from sediment traps do not usually match estimated metabolic demands in the ocean twilight zone (Steinberg *et al.*, 2008; Burd *et al.*, 2010; Herndl and Reinthaler, 2013), suggesting that other organic fluxes besides those typically considered from phytoplankton and fecal pellets contribute significantly to vertical export.

74

75

76

77

78

79

80

Recently developed *in situ* imaging methods have proven useful for identifying larger particles (Guidi *et al.*, 2009; Bochdansky *et al.*, 2013), but they often fall short in resolving the composition of smaller particles and complex aggregates. On the other hand, DNA sequencing analysis coupled with accurate reference databases offer a powerful alternative for extracting detailed taxonomic information from partially degraded material and complex communities. Amacher *et al.* (2009) pioneered this approach in sediment traps, using clone libraries to quantify the relative contributions of protist groups to

81 downward particle fluxes in the eastern subtropical North Atlantic. More recently, a
82 metagenomic approach based on 454 pyrosequencing technology was used to investigate
83 microbial communities associated with sinking particles in the oligotrophic North Pacific
84 Subtropical Gyre (NPSG) (Fontanez *et al.*, 2015). However, analogous information for
85 productive systems with high export potential is lacking.

86 The California Current Ecosystem (CCE) is a coastal upwelling biome characterized
87 by high production and strong advective fields that transport high-nutrient, high-biomass
88 coastal waters to oligotrophic offshore areas (Ohman *et al.*, 2013). This gradient is
89 reflected in phytoplankton composition and productivity (Taylor *et al.*, 2015), food-web
90 interactions (Landry *et al.*, 2009), and export fluxes (Stukel *et al.*, 2011, 2017). The present
91 study aims at characterizing the composition of eukaryotic communities involved in the
92 export and remineralization of sinking particles in the CCE. We capitalize on the
93 capability of DNA metabarcoding of the 18S rRNA gene to retrieve taxonomic
94 information from complex environmental samples such as those collected by PITs. Our
95 specific objectives are to assess 1) the taxonomic compositions of POM sinking out of the
96 euphotic zone and 2) the compositional changes associated with degradation and
97 consumption processes transforming this POM below the euphotic zone. Towards these
98 goals, we deployed formalin-fixed (i.e. fixed) and preservative-free (i.e. live) traps at the
99 base of the euphotic zone, on the premise that microbial activity and degradation of POM
100 would be inhibited in formalin-fixed traps (Knauer *et al.*, 1984) but allowed in live traps
101 (Karl and Knauer, 1984; Lee *et al.*, 1992).

102 103 MATERIALS AND METHODS

104 *Study area and sampling strategy*

105 Hydrographic and biological data were collected during the CCE-P1408 Process
106 Cruise on the R/V Melville (6 Aug - 4 Sept, 2014), as part of the CCE LTER (Long Term
107 Ecosystem Research) program. A quasi-Lagrangian strategy was adopted to sample
108 representative water parcels from coastal to offshore conditions over 3-days sampling
109 periods called 'Cycles'. We used a satellite-tracked drifting array with a drogue at 15 m
110 to follow the water parcels and sampled the water column daily at the array for a suite of
111 physical, chemical and biological measurements (Landry *et al.*, 2009). Cycles 1 to 3
112 were initiated in more productive coastal waters around Point Conception, while Cycles 4
113 and 5 represented typical oligotrophic offshore conditions (Figure 1). Hydrographic data

114 and water samples were acquired at 6-8 depths from a CTD-rosette system with 10-L
115 Niskin bottles with Teflon-coated springs.

116 *Sediment traps and export measurements*

117 A second drifting array with VERTEX-style drifting sediment traps (Knauer *et al.*, 1979)
118 was deployed during each cycle to collect sinking particles and assess export fluxes.
119 The sediment trap consisted of arrays of 8-12 PITs in a cross-like layout attached to the
120 wire at 2-3 depths below the euphotic zone (100 m, 150 m, and base of the euphotic zone if
121 shallower than 100 m). The tubes were filled with a brine solution of 0.1 μm filtered
122 seawater with 50 g L^{-1} NaCl and 80 mg L^{-1} of SrCl_2 added to prevent mixing with *in situ*
123 water and dissolution of acantharian skeletons (Beers and Stewart, 1970), respectively.
124 Most tubes were fixed with formaldehyde (0.4% final concentration, fixed traps), to
125 minimize decomposition and consumption of organic matter (Knauer *et al.*, 1984), while
126 selected tubes were not fixed (live traps) to allow these biotic processes to continue (Karl
127 and Knauer, 1984).

128 Upon recovery, the interface of brine and *in situ* water was visually determined in
129 each tube and the upper layer removed gently by suction. For fixed traps, mesozooplankton
130 swimmers were removed under a dissecting scope, before the sample was mixed, split
131 and subsampled for particulate organic carbon (POC) and nitrogen (PON), chlorophyll *a*
132 (Chl *a*), and phaeopigment analyses as described in Stukel *et al.* (2013). Previous
133 calibration of our PITs using ^{238}U : ^{234}Th disequilibrium methods suggest that they are
134 accurately collecting sinking particles (Stukel *et al.*, 2015). For live traps, swimmers were
135 not removed, and samples were directly split and filtered to minimize processing time and
136 potential degradation. For DNA flux and diversity analysis, a known fraction of fixed- and
137 live- traps was vacuum filtered in parallel through 0.8 and 8 μm Supor membrane filters
138 (Pall Life Sciences, Port Washington, NY, USA) to obtain replicate independent samples
139 for $> 0.8 \mu\text{m}$ and $> 8 \mu\text{m}$ particles. DNA fluxes from fixed traps were estimated by
140 dividing the DNA concentration by an extraction efficiency factor of 0.153 (Amacher *et*
141 *al.*, 2013).

142 *Water column chemical and biological analysis*

143 Water-column dissolved inorganic nutrients (DIN) and total and size-fractionated Chl
144 *a* were obtained from <http://oceaninformatics.ucsd.edu/datazoo/catalogs/ccelter/datasets>,
145 where sampling and analytical methods are described in detail. For DIN, seawater was
146 filtered directly from the Niskin bottle using a Suporcap filter capsule (0.1 μm pore size)

147 and major nutrient concentrations ($\text{NO}_3^- + \text{NO}_2^-$, NO_2^- , PO_4^{3-} , NH_4^+ , and SiOH_3) analyzed by
148 autoanalyzer using standard methods (Gordon *et al.*, 1992). For total Chl *a* analysis, the
149 samples were filtered onto 25 mm GF/F filters (Whatman, Maidstone, UK), and pigments
150 were extracted in 90% acetone at -18°C in the dark for 24 h and quantified on a calibrated
151 10AU fluorometer (Turner Designs, Sunnyvale, CA, USA) using the acidification method.
152 For size-fractionated Chl *a*, 0.10-0.25 L samples from the surface mixed layer were filtered
153 through a series of filters with different pore sizes (20 μm Nitex mesh, 8 μm , 3 μm and 1
154 μm Nucleopore and GF/F filters, Whatman, Maidstone, UK) and analyzed
155 fluorometrically as above.

156 Water-column samples for DNA analysis were collected in two ways: 1) seawater
157 was collected from the Niskin spigot into an acid-washed and Milli-Q rinsed
158 polypropylene bottle, with a known volume (1.5-3 L) then filtered on a 0.8 μm Supor filter
159 (i.e. called small-fraction). 2) We also emptied entire Niskin bottles into 10-L carboys by
160 opening the bottom closure and using a large funnel to collect larger organisms
161 (Michaels *et al.*, 1995). These samples were then concentrated to 250 mL using a 5 μm
162 mesh plankton net and subsequently vacuum-filtered onto 8 μm Supor filters (i.e. called
163 large size-fraction).

164 *Genomic DNA extraction, PCR amplification and library preparation*

165 DNA was extracted with the Nucleospin Plant kit (Macherey-Nagel, Düren,
166 Germany), mini version for the water-column 0.8- μm filters and the midi version for 8.0-
167 μm filters, and PIT samples (Supplementary Table S1). The V4 region of the 18S rRNA
168 gene was amplified using the eukaryotic primers V4F_illum (5'
169 CCAGCASCYGCGGTAATTCC 3') and V4R_illum (5'
170 ACTTTCGTTCTTGATYRATGA 3') with Illumina overhang adapters (Forward 5'
171 TCGTCGGCAGCGTCAGATGTGTATAAGAGACAG 3' and Reverse 5'
172 GTCTCGTGGGCTCGGAGATGTGTATAAGAGACAG 3') for attaching Nextera
173 indexes (Piredda *et al.* 2017). PCR amplifications were done in triplicate for each sample.
174 PCR products were visualized on 1.5% agarose gel and pooled together before purification
175 with Agencourt AMPure XP purification system (Beckman Coulter, Brea, CA, USA).
176 Randomly selected 20 purified amplicons were sized and validated using the Agilent High
177 Sensitivity DNA Assay in the 2100 Bioanalyzer (Agilent Technologies, Palo Alto, CA,
178 USA). For 6 distinct samples, the purified PCR product was split into three subsamples and
179 also included in the library to investigate the reproducibility of the sequencing step. A
180 second PCR step to attach index and Illumina adapters was conducted with the Nextera

181 DNA library Preparation Kit (Illumina), followed by additional AMPure purification and
182 library validation. The library was then quantified and prepared for 2 x 250 bp sequencing
183 on a MiSeq platform. For Cycle 1, only fixed-trap samples were available for sequencing,
184 while the > 0.8 μ m live-trap samples from Cycle 3 were lost during processing. We
185 sequenced a total of 54 samples, which included water-column samples from mixed-layer,
186 deep chlorophyll maximum and below (n=22), live-trap samples from 60, 100 and 150 m
187 (n=21), and fixed-trap samples from 100 m (n=11) (Supplementary Table S1). Raw
188 sequence have been deposited to GenBank under Bioproject number PRJNA432581.

189 *Processing and taxonomic assignation of sequencing reads*

190 Fastq files were checked using FastQC on the Galaxy platform (Goecks *et al.*, 2012)
191 for sequence length (L) and Quality Score (Q). Forward reads with L > 200 and Q > 20
192 over 75% of the sequence were retained. For reverse reads, the first 85 bases were of very
193 bad quality. These 85 bases were removed and sequences with L > 150 and Q > 20 in 50%
194 of the sequence were retained. Unpaired sequences were removed. All further processing
195 was performed using mothur v.1.33.0 (Schloss *et al.*, 2009). First, contigs were assembled
196 from forward and reverse reads, keeping only contigs free of ambiguity. Singletons were
197 removed and sequences were aligned to 18S rRNA Silva reference. Sequences with
198 two nucleotide differences were pre-clustered, and chimeras were removed using
199 UCHIME (Edgar *et al.*, 2011) as implemented in mothur. Only pre-clusters with > 10
200 sequences were retained. Pre-clustered sequences were taxonomically annotated using
201 classify.seqs against the PR2 reference database (Guillou *et al.*, 2013) version 4.4
202 available from https://github.com/vaulot/pr2_database/releases. Sequences were clustered
203 using the average neighbour algorithm to determine OTUs at 97% similarity level. OTUs
204 were taxonomically assigned using classify.otu and BLASTed against GenBank to confirm
205 the assignation provided by mothur and to get their percentage of similarity to existing
206 sequences. Since the community composition for triplicated amplicons was not
207 substantially different (Supplementary Figure S1), sequences from replicated amplicons
208 were pooled for subsequent analysis. Similarly, sequences from water-column and live-
209 trap samples obtained at different depths were combined for downstream analysis. The
210 OTU abundance table and OTU sequences are available as supplementary material on
211 Figshare (<https://doi.org/10.6084/m9.figshare.5944291>).

212 *Statistics and diversity analysis of sequence data*

213 Data analyses and statistics were done using R version 3.2.4 (R Core Team, 2016)
214 with the *vegan* 2.4-3 (Oksanen *et al.*, 2017), *ggplot2* (Wickham, 2009), *treemap*
215 (*Tennekes*, 2017) packages, GraphPad 5.0 (GraphPad Software, La Jolla, CA, USA) and
216 the Paleontological Statistics software (*Hammer et al.*, 2001). OTU richness was estimated
217 with *rarefy* function (*vegan* package) on a random subsample of size determined by
218 the minimum number of sequences found among the samples compared (e.g. fixed vs.
219 live traps).

220 RESULTS

221 *Water-column physical, chemical and biological conditions*

222 During the sampling cycles (C1 to C5), we explored 5 water parcels (Figure 1) with
223 physical and chemical properties that reflected primarily their coastal vs. offshore
224 characteristics (Table 1, Supplementary Figure S2a). A hierarchical dendrogram based on
225 both physico-chemical and biological properties clusters coastal C1 and C2 together,
226 separated from offshore C4 and C5, with intermediate conditions for C3 (Supplementary
227 Figure S2b). C1 and C2 exhibited colder and saltier surface waters and shallower
228 nitracline depth indicative of coastal upwelling (Table 1). Surface Chl *a* concentration
229 was also higher with large phytoplankton cells (> 20 μm) accounting for ~40% of Chl
230 *a* for C1 and C2, while the contribution of large cells was minor (~2%) for C4 and C5
231 (Table 1, Supplementary Figure S2c).

232 POC and PON fluxes were higher at coastal cycles and decreased offshore (Figure
233 2). DNA fluxes measured in fixed traps at 100 m followed the same trend (Figure 2a) and
234 were significantly correlated with POC fluxes ($r = 0.88$, $p < 0.05$, $n = 5$). The Chl *a* to
235 phaeopigment ratio of sediment trap material indicated that pigment flux during all cycles
236 was dominated by fecal-derived material (Figure 2b). The higher values observed at the
237 coastal C1 and C2 cycles, however, indicated a higher contribution of ‘fresh’ algae
238 compared to offshore C4 and C5 (Figure 2b).

239 *18S V4 OTU distribution and diversity patterns*

240 Over 18 million paired reads were obtained from the sequencing run, half of which
241 were removed after filtering based on quality and length (Supplementary Table S2). After
242 processing, we obtained 2,662 OTUs (97% similarity) corresponding to 2,802,466
243 sequences (Supplementary Table S2).

244 For protists, rarefaction curves for different stations were not saturated for all
245 samples, although the analysis of pooled sequences from different sample types suggest
246 an adequate recovery of epipelagic protist diversity for cycles C2-C5 (Supplementary
247 Figure S3a). Cycle C1 was excluded from this analysis because only fixed-trap samples
248 were available. Protistan OTU richness increased from coastal to oceanic locations,
249 being higher in water-column compared to fixed- and live-trap samples for all cycles
250 (Figure 3a). More protistan OTUs were recovered in live compared to fixed traps for C2 and
251 C3, while fixed traps had higher or similar OTU numbers compared to live traps for C4
252 and C5 (Figure 3a). OTU richness for water-column protists was higher in the smaller
253 compared to the larger size-fraction samples (Figure 3b). This difference decreased in
254 sediment traps, particularly in fixed traps, where protistan richness was similar or higher
255 in the larger fraction (Figure 3b). Similar diversity patterns were observed for all eukaryotic
256 OTUs, including metazoans (Supplementary Figure S3b).

257 Non-metric multidimensional scaling (NMDS) analysis, based on protistan OTUs
258 abundance and Bray-Curtis distances, ordinated samples into three main clusters
259 corresponding to the different sampling methods (Figure 4). Compositional changes in
260 water-column samples were significantly correlated with surface temperature, salinity,
261 nitracline depth, and Chl *a*, with size fraction (NMDS₁, $r^2 = 0.61$, $P = 0.001$) and cycle
262 (NMDS₂, $r^2 = 0.30$, $P = 0.02$) being the primary ordination factors (Supplementary Figure
263 S4). Sampling cycle emerged as primary ordination factor for the correlations of trap
264 samples with physico-chemical and biogeochemical variables (Supplementary Figure S4).

265 *Taxonomic composition of the water-column eukaryotic community*

266 OTU taxonomy was assigned based on the curated 18S rRNA database PR² (Guillou
267 *et al.*, 2013) which uses 8 different taxonomical levels from Kingdom to Species.
268 Protistan sequences dominated the community across all cycles and size fractions ($58 \pm$
269 27% of eukaryotic sequences), with a substantial contribution of metazoan sequences
270 ($39 \pm 26\%$, Supplementary Figure S5).

271 Among protists, Dinophyta was the most abundant group ($72 \pm 13\%$ of protistan
272 sequences) followed by Radiolaria ($12 \pm 13\%$) and Chlorophyta ($4.7 \pm 7.1\%$) (Figure 5).
273 Ochrophyta ($2.1 \pm 2.0\%$) contributed less on average but represented a substantial
274 percentage of the protistan community at times (Figures 5 and 6). Dinophyta included
275 sequences affiliated with Syndiniales and Dinophyceae in similar relative abundances
276 ($\sim 20\%$, Figure 6), although their distributional pattern and partitioning between size

277 fractions differed markedly (Figure 7). In addition to *Ptychodiscus noctiluca*, uncultured
278 Dinophyceae and Syndiniales OTUs were among the most abundant in water-column
279 samples (Supplementary Table S3 and Figure S6).

280 Most radiolarian sequences were affiliated with Collodaria and Acantharia,
281 although groups such as Spumellaria and RAD-A were occasionally important (Figure 8).
282 Chlorophyta (green algae) were mostly picoplanktonic belonging to Chloropicophyceae
283 (*Chloroparvula*, previously prasinophytes clade VII-B1, Lopes dos Santos *et al.*, 2017a)
284 and Mamiellophyceae (*Bathycoccus*) (Figure 6). Both picoplanktonic groups peaked in
285 coastal waters (e.g. Chloropicophyceae represented ca. 20% of protistan sequences in C2),
286 and decreased offshore (Figures 6 and 7). Ochrophyta (photosynthetic stramenopiles)
287 sequences were also more abundant in the smaller size fraction (Figure 5) and
288 comprised mainly of sequences assigned to pelagophytes and diatoms (Figure 6), with the
289 latter reaching ~5-fold higher relative abundance toward the coast (Figure 7). The
290 contribution of Phaeodaria, a rhizarian group related to Radiolaria, also peaked in coastal
291 waters, but in the larger size fraction (Figure 7).

292 *Taxonomic composition of eukaryotic community in fixed traps*

293 Fixed-trap samples were overwhelmingly dominated by protistan sequences, mainly
294 Radiolaria ($88 \pm 8.6\%$ of eukaryotic sequences) and Dinophyta ($8.6 \pm 6.3\%$) (Figures 5
295 and Supplementary Figure S5). Metazoan sequences accounted for $6.1 \pm 5.0\%$ and were
296 dominated by Crustacea, with Mollusca and Cnidaria contributing less on average
297 (Supplementary Figure S5).

298 The dominance of Radiolaria was a consistent feature across cycles (Figure 6),
299 with contributions from a diverse suite of OTUs from the four major radiolarian groups
300 (Acantharia, Collodaria, Nassellaria and Spumellaria) (Figure 8, Supplementary Table S3).
301 Overall, Spumellaria was the most abundant group ($60 \pm 22\%$ of radiolarian sequences)
302 followed by Acantharia ($22 \pm 28\%$) and Collodaria ($4.8 \pm 11\%$) (Figure 8). Among
303 Dinophyta, Dinophyceae (true dinoflagellates) were relatively more abundant than
304 Syndiniales, and their contribution increased in offshore cycles (Figure 7) with
305 uncultured Dinophyceae and *P. noctiluca* being the most abundant OTUs (Supplementary
306 Table S3 and Figure S6). Diatoms comprised most Ochrophyta sequences in fixed traps
307 from cycles C1 and C2 (Figure 6), belonging mainly to the same genera (*Pseudo-*
308 *nitzschia* and *Chaetoceros*) as those in the water column (Supplementary Figure S6). They
309 virtually disappeared offshore, where sequences affiliated with the heterotrophic

310 nanoflagellate (HNF) *Paraphysomonas imperforata* (Chrysophyceae) became the
311 dominant Ochrophyta group (Figure 6).

312 Taxonomic composition of eukaryotic community in live traps

313 Metazoan sequences dominated live-trap samples ($78 \pm 18\%$ of eukaryotic
314 sequences, Supplementary Figure S5). Among them, Crustaceans belonging to different
315 copepod genera, notably *Metridia* spp., contributed most to the eukaryotic community
316 followed by Cnidaria and Mollusca (Supplementary Table S3).

317 Among protists, Dinophyta ($51 \pm 5.7\%$ of protistan sequences) and Stramenopiles_X
318 (heterotrophic stramenopiles, $24 \pm 7.2\%$) were the dominant groups, with Radiolaria
319 ($9.6 \pm 3.7\%$) and Phaeodaria ($5.4 \pm 5.5\%$) contributing less (Figures 5 and 6). In addition
320 to the dinoflagellate *P. noctiluca*, the HNFs *Caecitellus parvus* and *C. paraparvulus* (order
321 Anoeales) were the most abundant protistan species in live traps (Supplementary Table S3
322 and Figure S6), accounting for the high proportion of Stramenopiles_X. The abundance of
323 these HNFs decreased significantly in the $> 8 \mu\text{m}$ fraction, suggesting that they are free
324 living or feeding on particles and aggregates to which they are loosely attached
325 (Figures 6 and 7). Diatoms and chrysophytes recovered from live traps showed the same
326 spatial and size-fraction distribution as in fixed traps, although the contribution of
327 chrysophytes in offshore cycles was higher in live traps (Figure 7). Most radiolarian
328 sequences recovered from live traps belonged to Acantharia and Spumellaria (Figure 8).
329 RAD-A and Collodaria were detected in all cycles but showed higher contribution in
330 offshore cycles C4 and C5 (Figure 8). Phaeodaria, mostly affiliated with *Aulacantha* spp.
331 (Supplementary Table S3), were more abundant in coastal cycles C1 and C2, where they
332 represented up to 20% of protistan sequences (Figure 7).

334 DISCUSSION

335 Overall, the protistan communities in water-column and trap samples were markedly
336 dissimilar (Figures 5 and 6), showing higher similarity in sinking material across cycles
337 than between traps and the overlying euphotic zone (Figure 4). Such pattern has been
338 observed in previous molecular-based studies (Amacher *et al.*, 2013; Fontanez *et al.*, 2015)
339 and suggests either that a common suite of organisms is responsible for export despite
340 dramatic differences in ambient microbial communities or that these organisms are
341 resistant to degradation of the genetic material. Radiolaria were identified as the most
342 abundant group in POM sinking out of the euphotic zone, in addition to phytoplankton

343 and metazoan zooplankton, usually considered as having a major role in vertical export
344 in the CCE (Figures 5 and 6). The observed dominance of radiolarian sequences in fixed
345 traps (Figures 5 and 6) is also consistent with clone library analysis of trap samples
346 from the eastern subtropical North Atlantic, where radiolarian clones were most abundant
347 (Amacher *et al.*, 2009), and water column samples from the Sargasso sea, where high
348 proportion of radiolarian clones were obtained below the euphotic zone (Not *et al.*, 2007).
349 In contrast to these observations, the relative abundance of radiolarians determined by 18S
350 rRNA metabarcoding in the western Antarctic Peninsula was negatively linked to
351 community export potential (Lin *et al.*, 2017), suggesting that the role of this group in
352 vertical export may differ across systems.

353 Metabarcoding of 18S rRNA genes is largely used for assessing the composition of
354 aquatic microbial communities (De Vargas *et al.*, 2015; Pearman *et al.*, 2016; Piredda *et al.*,
355 2017; Lopes dos Santos *et al.*, 2017b) and the spatio-temporal patterns of specific
356 taxonomic groups (Egge *et al.*, 2015; Ichinomiya *et al.*, 2016; Zouari *et al.*, 2018). Yet, a
357 number of well identified limitations and potential biases are acknowledged with respect to
358 absolute quantification of plankton groups. Along with PCR biases, the presence of
359 multiple copies of 18S rRNA genes and its variation across taxa (Zhu *et al.*, 2005) affects
360 the quantitative interpretation of community compositional changes from read abundance
361 data. In this regard, the prevalence and high relative contributions of Dinophyceae and
362 Syndiniales in this study (Figures. 6 and 7) and other meta-genetic surveys (Guillou *et al.*,
363 2008; Amacher *et al.*, 2009; Lie *et al.*, 2013; De Vargas *et al.*, 2015; Piredda *et al.*, 2017)
364 would be partially due to the high DNA content and number of gene copies in these
365 groups. Similarly, the multi-nuclear nature of Radiolaria (Anderson, 1983; Suzuki *et al.*,
366 2009; Suzuki and Aita, 2011) and the high gene copy number observed in Collodaria
367 (Biard *et al.*, 2017) could explain their high relative abundances in meta-genetic surveys
368 (Decelle *et al.*, 2014; De Vargas *et al.*, 2015, Biard *et al.*, 2017).

369 Nonetheless, a positive relationship between 18S rRNA copy number and cell length
370 has been reported across protists spanning orders of magnitude in cell size (Zhu *et al.*,
371 2005; Godhe *et al.*, 2008; De Vargas *et al.*, 2015; Biard *et al.*, 2017), which encourages the
372 cautious use of read abundances to infer community composition general patterns and
373 dynamics on a quasi-biomass basis (i.e., the larger organisms have proportionately more
374 DNA reads). Further support for an ecological, rather than bias, interpretation of
375 compositional changes between water-column and trap samples (Figures. 4 and 5) and
376 group-specific relative-abundance spatial patterns (Figures 6 and 7) comes from a recent

377 study that reported good agreement between relative picoeukaryotic cell and environmental
378 sequence abundances (Giner *et al.*, 2016). The community composition obtained by
379 parallel sequencing of replicated PCR products was virtually identical (Figure S1),
380 suggesting that potential errors linked to sequencing were not responsible for the
381 differences observed among samples. Moreover, the comparative approach adopted here
382 provides a robust framework for interpreting diversity and compositional changes between
383 fixed and live traps from an ecological and biogeochemical perspective.

384 The prevalence of radiolarian sequences in sinking POM observed in this study
385 (Figure 6) is remarkable, considering the major differences in productivity and trophic
386 structure of the epipelagic communities along the CCE environmental gradient (Table 1
387 and Supplementary Figure S2). Analysis of our preserved samples under epifluorescence
388 microscopy (data not shown) confirmed the presence of Radiolaria, although the
389 dissolution of hard structures observed for this, and other taxonomic groups (e.g. diatoms),
390 precluded reliable estimates of biomass and finer taxonomic assignment. Analysis of
391 DNA sequences, however, shows that different radiolarian groups (Figure 8) and OTUs
392 (Supplementary Table S3 and Figure S6) contribute broadly to this result, highlighting the
393 functional diversity of the group. Previous microscopical and geochemical analyses of
394 sediment trap material have stressed the importance of major radiolarian groups to export
395 (Takahashi, 1983; Gowing and Coale, 1989; Michaels *et al.*, 1995; Decelle *et al.*,
396 2013; Boltovskoy, 2017). The enrichment of Radiolaria in sediment traps may reflect both
397 the high contribution of this group to particle export and the inadequacy of CTD and net
398 tow sampling to capture these fragile and patchy amoeboid organisms (Michaels *et al.*,
399 1995; Dennett *et al.*, 2002; Suzuki and Not, 2015).

400 Several characteristics of radiolarian cell structure and ecology are consistent with
401 their high export potential. The silica or strontium sulfate skeletons of most Polycystines
402 (Collodaria, Nasellaria, and Spumellaria) and Acantharia, respectively, provide substantial
403 mineral ballast (Takahashi, 1983; Suzuki and Not, 2015). In addition, their amoeboid
404 nature and sticky pseudopodia can catalyze the formation of aggregates with high sinking
405 velocities (Takahashi, 1987). In contrast to recent genomic studies pointing to Acantharia
406 and Collodaria as the key ‘export’ taxa in tropical and subtropical oceans (Fontanez *et al.*,
407 2015; Guidi *et al.*, 2016), we found that Spumellaria are the most important export
408 contributors in the CCE (Figure 8). Gowing (1986) also showed Spumellaria to be the
409 dominant radiolarian group in microscopical analyses of sediment trap material from the
410 oligotrophic NPSG VERTEX station. In addition to Acantharia, which represented > 50%

411 of radiolarian sequences in C1 and C4, we also detected a substantial contribution by
412 Collodaria in C2 (Figures 6 and 8). Considering the patchy distributions and large size
413 (μm -to- cm) of single-celled and colonial Collodaria, it is difficult to assess their
414 contributions to export accurately, based only on discrete sediment trap analysis (Michaels
415 *et al.*, 1995). However, the high relative abundance of Collodaria in the CCE region
416 inferred from molecular (Figures 6 and 8) and *in situ* image analysis (Underwater
417 Vision Profiler 5, UVP5) during this and previous cruises (Ohman *et al.*, 2012; Biard *et al.*,
418 2016) supports their important role in export flux.

419 Dinoflagellates were the most abundant non-radiolarian protists in both water-
420 column and fixed-trap samples from cycles C1 and C2 in this region (Figures 6 and 7).
421 The warm anomaly that developed in the NE Pacific during the 2013-2014 winter (Bond
422 *et al.*, 2015), colloquially referred to as ‘the blob’, had already hit the California coast at
423 the time of our cruise (Gentemann *et al.*, 2017) and was responsible for the weak summer
424 upwelling that we encountered. Off the Oregon coast, changes in plankton community
425 composition associated with ‘the blob’ included higher dinoflagellates abundances and
426 penetration of open-ocean copepod species onto the continental shelf region (Peterson *et al.*,
427 2017). Increased abundance of autotrophic dinoflagellates has also been observed in
428 the Point Conception region during years of delayed upwelling (Taylor *et al.*, 2015).
429 Consistent with this, our molecular survey revealed high sequence abundance of *P.*
430 *noctiluca* (Supplementary Table S3 and Figure S6), a dinoflagellate species with distinctive
431 cell covering characteristics and widespread distribution (Gómez *et al.*, 2016) that is
432 generally less common in the CCE. Among green algae, the dominance of prasinophytes
433 clade VII-B1 (Chloropicophyceae), a group typical of offshore waters (Lopes dos Santos
434 *et al.*, 2017a), over Mamiellophyceae, characteristics of coastal waters, is another
435 indication of the oceanic characteristics of the epipelagic zone. Collodaria is typically
436 associated to warm oligotrophic waters (Dennett *et al.*, 2002; Biard *et al.*, 2016), and the
437 observed abundance in mesotrophic coastal waters of C2 (Figures 6 and 8) could have been
438 favored by the warm anomaly. It therefore seems likely that the relatively low contribution
439 of diatoms to the water-column assemblage and export fluxes was a consequence of
440 anomalous conditions. Whether the high contribution of Radiolaria indicated by our
441 molecular analyses was enhanced by these 2014 conditions cannot be directly addressed
442 without comparable data from years with ‘normal’ conditions, although several lines of
443 evidence argue against this idea.

444 On one hand, Radiolaria have been shown to dominate clone libraries recovered
445 from sediment traps despite the dominance of diatoms in the upper water column in the
446 eastern subtropical North Atlantic (Amacher *et al.*, 2009). On the other hand, although
447 higher fluxes of polycystine radiolarians (living + empty skeletons) have been found in
448 coastal upwelling compared to oligotrophic offshore waters of the CCE, a clear
449 relationship between primary production and these fluxes could not be established
450 (Gowing and Coale, 1989). In the subarctic Pacific, however, a positive relationship
451 between polycystine radiolarian fluxes and primary production was reported at station
452 PAPA (Takahashi, 1987), suggesting differences in the export role of these group among
453 systems. In our 2014 CCE study, microscopical counts of large phaeodarians from fixed
454 traps and in UVP5 profiles, yielded lower abundances compared to previous and
455 following year estimates for this group of rhizarians (Stukel *et al.*, 2018; Biard *et al.*, 2018).
456 Nonetheless, our study showed that phaeodarians represented a significant fraction of the
457 protistan sequences in live traps (Figures 5 and 7). Altogether, these findings argue against
458 the idea that Radiolaria dominance was due to anomalous conditions and support the key
459 role of rhizarians in export fluxes as a general feature of the CCE, and potentially other
460 eastern boundary upwelling systems.

461 One striking result was the markedly lower contribution of Radiolaria taxa in live
462 traps (Figures 5 and 6), indicating rapid remineralization of organic matter associated
463 with this group. Sinking particles and aggregates serve as natural ‘hot spots’ for both
464 microbial and metazoan activity (Karl and Knauer, 1984; Taylor *et al.*, 1986). Live traps
465 were highly enriched with sequences from larger copepods like *Eucalanus* and *Metridia*
466 spp. (Supplementary Table S3 and Figures S6), consistent with the potential of copepods to
467 consume and transform sinking organic particles (Lampitt *et al.*, 1990; Noji, 1991; Iversen
468 and Poulsen, 2007).

469 Rather than ingesting sinking particles directly, detritivorous zooplankton may
470 benefit from the microbial growth enhanced by particle fragmentation (Mayor *et al.*,
471 2014). Such “microbial gardening” would be consistent with the dramatic increase of
472 opportunistic HNFs such as *Caecitellus* spp. and *P. imperforata* (Supplementary Table
473 S3 and Supplementary Figure S6), which dramatically increase in enrichment cultures
474 (Lim *et al.*, 1999). Heterotrophic protists are known to thrive around marine snow particles
475 (Caron *et al.*, 1982, 1986; Silver *et al.*, 1984), but whether this microbial boost is directly
476 exploited by larger metazooplankton or consumed first by small microzooplankton is
477 unclear. The increase of ciliates in both fixed (1.8 % of protistan sequences) and live traps

478 (2.2 %) relative to the water column (0.66 %) in this and previous studies (Amacher *et al.*,
479 2009), indicates that they likely prey on HNF, which function as a trophic link in these
480 rich microenvironments. Further evidence of heterotrophic protistan activity associated
481 with the degradation of sinking organic material is indicated by the enrichment of
482 Phaeodaria in live compared to fixed traps (Figures 5 and 7). This amoeboid group has
483 been shown to feed on HNF and eukaryotic algae and proposed to consume marine snow
484 particles (Gowing, 1986; Gowing and Bentham, 1994). Data from the UVP5 have shown
485 large (> 600 µm) species of phaeodarian forming a high-density layer below the euphotic
486 zone (Ohman *et al.*, 2012), although their abundances were notably lower during this cruise
487 compared to previous ones (Stukel *et al.* 2018; Biard *et al.* 2018). The high relative
488 abundance of Phaeodaria in live traps (Figures 5), together with their trophic biology and
489 vertical distribution in the CCE, further supports the important role that this group may have
490 in the cycling of particulate organic matter sinking below the euphotic zone.

491 CONCLUSION

492 While phytoplankton and crustacean zooplankton are generally viewed as the
493 major biological sources of vertical export in productive systems, our results, showing
494 the abundance and diversity of rhizarian sequences in particulate material sinking below
495 the euphotic zone, indicate an important role for this group in the export and cycling of
496 particulate organic matter in the CCE. Recent studies using advanced molecular and
497 imaging technologies have revealed unprecedented global abundances of Rhizaria
498 (Stemmann *et al.*, 2008; Not *et al.*, 2009; Biard *et al.*, 2016), but the functions and impacts
499 of this group in pelagic ecosystems have yet to be accurately quantified. The high
500 abundance and diversity of Radiolaria reported here, and their prevalence from coastal
501 upwelling to oceanic oligotrophic conditions, stress the need to better characterize the
502 group's functional diversity to improve understanding of biological controls on vertical
503 fluxes.

504 ACKNOWLEDGEMENTS

506 We would like to acknowledge Estelle Bigeard for her help during the DNA
507 extraction and Illumina library preparation, Ralf Goericke for the Chl *a* and nutrient data,
508 and the crew and research technicians of RV Melville for their assistance during the CCE-
509 P1408 Process Cruise. CCE-LTER ship and science support were provided by NSF grants
510 OCE-1026607 and OCE-1637632. We are grateful to the Genomer and ABIMS
511 platforms for access to the Illumina sequencer and bioinformatic cluster, Brittany region

SAD grant SYMBIOX for funding AGR research and Sorbonne Universités Emergence grant to FN, NIWA for its support during data analysis and writing of the manuscript via New Zealand Strategic Science Investment Funding to the National Coasts & Oceans Centre, and Scott Nodder (NIWA) for support and fruitful discussions. Funds for ALS was provided by FONDECYT grant PiSCO South (N1171802).

CONFLICT OF INTEREST

The authors declare no conflict of interests.

BIBLIOGRAPHY

- Anderson OR. (1983). Radiolaria. Springer: New York.
- Agusti S, González-Gordillo JI, Vaqué D, Estrada M, Cerezo MI, Salazar G, *et al.* (2015). Ubiquitous healthy diatoms in the deep sea confirm deep carbon injection by the biological pump. *Nat Commun* **6**: 7608.
- Allredge AL, Silver MW. (1988). Characteristics, dynamics and significance of marine snow. *Prog Oceanogr* **20**: 41–82.
- Amacher J, Neuer S, Anderson I, Massana R. (2009). Molecular approach to determine contributions of the protist community to particle flux. *Deep Sea Res Part I Oceanogr Res Pap* **56**: 2206–2215.
- Amacher J, Neuer S, Lomas M. (2013). DNA-based molecular fingerprinting of eukaryotic protists and cyanobacteria contributing to sinking particle flux at the Bermuda Atlantic time-series study. *Deep Sea Res Part II Top Stud Oceanogr* **93**: 71– 472.
- Beers JR, Stewart GL. (1970). Preservation of acantharians in fixed plankton samples. *Limnol Oceanogr* **15**: 825-827
- Biard T, Stemmann L, Picheral M, Mayot N, Vandromme P, Hauss H, *et al.* (2016). In situ imaging reveals the biomass of giant protists in the global ocean. *Nature* **532**: 504–507.
- Biard, T. *et al.* Biogeography and diversity of Collodaria (Radiolaria) in the global ocean. (2017). *ISME J.* **11**, 1331–1344.
- Biard T, Krause JW, Stukel MR, Ohman MD. (2018). The significance of giant phaeodarians (Rhizaria) to biogenic silica export in the California Current Ecosystem. *Global Biogeochem Cycles* **32**: 987-1004.
- Bochdansky AB, Jericho MH, Herndl GJ. (2013). Development and deployment of a point-source digital inline holographic microscope for the study of plankton and particles to a depth of 6000 m. *Limnol Oceanogr* **11**: 28–40.
- Boltovskoy D. (2017). Vertical distribution patterns of Radiolaria Polycystina (Protista) in the World Ocean: living ranges, isothermal submersion and settling shells. *J Plankton Res* **39**: 330–349.
- Bond NA, Cronin MF, Freeland H, Mantua N. (2015). Causes and impacts of the 2014 warm anomaly in the NE Pacific. *Geophys Res Lett* **42**: 3414–3420.
- Boyd PW, Newton PP. (1999). Does planktonic community structure determine downward particulate organic carbon flux in different oceanic provinces? *Deep Sea Res Part I-Oceanographic Res Pap* **46**: 63–91.
- Burd AB, Hansell DA, Steinberg DK, Anderson TR, Aristegui J, Baltar F, *et al.* (2010). Assessing the apparent imbalance between geochemical and biochemical indicators of meso- and bathypelagic biological activity: What the @#! is wrong with present

557 calculations of carbon budgets? *Deep Sea Res Part II-Topical Stud Oceanogr* **57**: 1557
558 1571.

559 Caron DA, Davis PG, Madin LP, Sieburth JM. (1986). Enrichment of microbial- populations in
560 macroaggregates (marine snow) from surface waters of the North- Atlantic. *J Mar Res* **44**:
561 543–565.

562 Caron DA, Davis PG, Madin LP, Sieburth JM. (1982). Heterotrophic bacteria and bacterivorous
563 protozoa in oceanic macro-aggregates. *Science* **218**: 795–797.

564 Decelle J, Martin P, Paborstava K, Pond DW, Tarling G, Mahé F, *et al.* (2013). Diversity,
565 Ecology and Biogeochemistry of Cyst-Forming Acantharia (Radiolaria) in the Oceans.
566 *PLoS One* **8**: e53598.

567 Decelle J, Romac S, Sasaki E, Not F, Mahé F. (2014). Intracellular diversity of the V4 and V9
568 regions of the 18S rRNA in marine protists (radiolarians) assessed by high-throughput
569 sequencing. *PLoS One* **9**: e104297.

570 Dennett MR, Caron DA, Michaels AF, Gallager SM, Davis CS. (2002). Video plankton
571 recorder reveals high abundances of colonial Radiolaria in surface waters of the central
572 North Pacific. *J Plankton Res* **24**: 797–805.

573 Edgar R, Haas B, Clemente J, Quince C, Knight R. (2011). UCHIME improves sensitivity and
574 speed of chimera detection. *Bioinformatics* **27**:2194-200.

575 Egge ES, Johannessen TV, Andersen T, Eikrem W, Bittner L, Larsen A, *et al.* (2015). Seasonal
576 diversity and dynamics of haptophytes in the Skagerrak, Norway, explored by high-
577 throughput sequencing. *Mol Ecol* **24**:3026-3042

578 Eppley RW, Peterson BJ. (1979). Particulate organic -matter flux and planktonic new
579 production in the deep ocean. *Nature* **282**: 677–680.

580 Fontanez KM, Eppley JM, Samo TJ, Karl DM, DeLong EF. (2015). Microbial community
581 structure and function on sinking particles in the North Pacific Subtropical Gyre. *Front*
582 *Microbiol* **6**: 1–14.

583 Gentemann CL, Fewings MR, García-Reyes M. (2017). Satellite sea surface temperatures along
584 the West Coast of the United States during the 2014-2016 northeast Pacific marine heat
585 wave. *Geophys Res Lett* **44**: 312–319.

586 Giering SLC, Sanders R, Lampitt RS, Anderson TR, Tamburini C, Boutrif M, *et al.* (2014).
587 Reconciliation of the carbon budget in the ocean’s twilight zone. *Nature* **507**: 480– 517.

588 Giner CR, Forn I, Romac S, Logares R, De Vargas C, Massana R. (2016). Environmental
589 sequencing provides reasonable estimates of the relative abundance of specific
590 picoeukaryotes. *Appl Environ Microbiol* **82**: 4757–4766.

591 Godhe A, Asplund ME, Harnstrom K, Saravanan V, Tyagi A, Karunasagar I. (2008).
592 Quantification of diatom and dinoflagellate biomasses in coastal marine seawater samples
593 by real-time PCR. *Appl Env Microbiol* **74**: 7174–7182.

594 Goecks J, Nekrutenko A, Taylor J, Galaxy T, Ieee. (2012). Lessons Learned from Galaxy, a
595 Web-based Platform for High-throughput Genomic Analyses. In: *2012 IEEE 8th*
596 *International Conference on E-Science*.

597 Gómez F, Qiu D, Dodge JD, Lopes RM, Lin S. (2016). Morphological and molecular
598 characterization of *Ptychodiscus noctiluca* revealed the polyphyletic nature of the order
599 Ptychodiscales (Dinophyceae). *J Phycol* **52**: 793–805.

600 Gordon LI, Jennings JC, Ross AA, and Krest JM. (1992). A suggested protocol for continuous
601 flow automated analysis of seawater nutrients in the WOCE hydrographic program and the
602 Joint Global Ocean Fluxes Study. *Grp Tech Rpt 92-1, OSU Coll Oceanogr Descr Chem*
603 *Oc*.

604 Gowing MM. (1986). Trophic biology of phaeodarian radiolarians and flux of living
605 radiolarians in the upper 2000-m of the North Pacific Central Gyre. *Deep Sea Res Part A-*
606 *Oceanographic Res Pap* **33**: 655–674.

607 Gowing MM, Bentham WN. (1994). Feeding ecology of phaeodarian radiolarians at the

608 VERTEX North Pacific time-series site. *J Plankton Res* **16**: 707–719.

609 Gowing MM, Coale SL. (1989). Fluxes of living radiolarians and their skeletons along a
610 northeast Pacific transect from coastal upwelling to open ocean waters. *Deep Sea Res Part*
611 *A-Oceanographic Res Pap* **36**: 561–576.

612 Guidi L, Chaffron S, Bittner L, Eveillard D, Larhlimi A, Roux S, *et al.* (2016). Plankton
613 networks driving carbon export in the oligotrophic ocean. *Nature* **532**: 465–470.

614 Guidi L, Stemann L, Jackson GA. (2009). Effects of phytoplankton community on production
615 ,size and export of large aggregates : A world-ocean analysis. *Limnol Oceanogr* **54**: 1951–
616 1963.

617 Guillou L, Viprey M, Chambouvet A, Welsh RM, Kirkham AR, Massana R, *et al.* (2008).
618 Widespread occurrence and genetic diversity of marine parasitoids belonging to
619 Syndiniales (Alveolata). *Env Microbiol* **10**: 3349–3365.

620 Guillou L, Bachar D, Audic S, Bass D, Berney C, Bittner L, *et al.* (2013). The Protist
621 Ribosomal Reference database (PR2): a catalog of unicellular eukaryote small sub- unit
622 rRNA sequences with curated taxonomy. *Nucleic Acids Res* **41**: D597-604.

623 Hammer Ø, Harper DAT, Ryan PD. (2001). PAST: paleontological statistics software package
624 foreducation and data analysis Harper DAT (ed). *Palaeontol Electron* **4**: 1–9.

625 Henson SA, Sanders R, Madsen E, Morris PJ, Le Moigne F, Quartly GD. (2011). A reduced
626 estimate of the strength of the ocean’s biological carbon pump. *Geophys Res Lett* **38**: 1–5.

627 Herndl GJ, Reinthaler T. (2013). Microbial control of the dark end of the biological pump. *Nat*
628 *Geosci* **6**: 718–724.

629 Ichinomiya M, dos Santos AL, Gourvil P, Yoshikawa S, Kamiya M, Ohki K, *et al.* (2016).
630 Diversity and oceanic distribution of the Parmales (Bolidophyceae), a picoplanktonic
631 group closely related to diatoms. *ISME J* **10**: 2419–2434.

632 Iversen MH, Poulsen LK. (2007). Coprorhexy, coprophagy, and coprochaly in the copepods
633 *Calanus helgolandicus*, *Pseudocalanus elongatus*, and *Oithona similis*. *Mar Ecol Prog Ser*
634 **350**: 79–89.

635 Karl DM, Knauer GA. (1984). Detritus-microbe interactions in the marine pelagic environment
636 - selected results from the VERTEX experiment. *Bull Mar Sci* **35**: 550– 556.

637 Knauer GA, Karl DM, Martin JH, Hunter CN. (1984). *In situ* effects of selected preservatives
638 on total carbon, nitrogen and metals collected in sediment traps. *JMar Res* **42**: 445–462.

639 Knauer GA, Martin JH, Bruland KW. (1979). Fluxes of particulate carbon, nitrogen, and
640 phosphorus in the upper water column of the Northeast Pacific. *Deep Sea Res* **26**: 97–108.

641 Lampitt RS, Noji T, Vonbodungen B. (1990). What happens to zooplankton fecal pellets -
642 implications for material flux. *Mar Biol* **104**: 15–23.

643 Landry MR, Ohman MD, Goericke R, Stukel MR, Tsyrklevich K. (2009). Lagrangian studies of
644 phytoplankton growth and grazing relationships in a coastal upwelling ecosystem off
645 Southern California. *Prog Oceanogr* **83**: 208–216.

646 Laws EA, Falkowski PG, Smith WO, Ducklow H, Mccarthy JJ. (2000). Temperature effects on
647 export production in the open ocean. *Global Biogeochem Cycles* **14**: 1231–1246.

648 Lee C, Hedges JI, Wakeham SG, Zhu N. (1992). Effectiveness of various treatments in
649 retarding microbial activity in sediment trap material and their effects on the collection of
650 swimmers. *Limnol Oceanogr* **37**: 117–130.

651 Lie AAY, Kim DY, Schnetzer A, Caron DA. (2013). Small-scale temporal and spatial
652 variations in protistan community composition at the San Pedro Ocean Time-series station
653 off the coast of southern California. *Aquat Microb Ecol* **70**: 93–110.

654 Lim EL, Dennett MR, Caron DA. (1999). The ecology of *Paraphysomonas imperforate* based
655 on studies employing oligonucleotide probe identification in coastal water samples and
656 enrichment cultures. *Limnol Oceanogr* **44**: 37–51.

657 Lin Y, Cassar N, Marchetti A, Moreno C, Ducklow H, Li Z. (2017). Specific eukaryotic
658 plankton are good predictors of net community production in the Western Antarctic

659 Peninsula. *Sci Rep* **7**: 14845.

660 Lopes dos Santos A, Pollina T, Gourvil P, Corre E, Marie D, Garrido JL, *et al.* (2017a).
661 Chloropicophyceae, a new class of picophytoplanktonic prasinophytes. *Sci Rep* **7**: 14019.

662 Lopes dos Santos A, Gourvil P, Tragin M, Noel MH, Decelle J, Romac S, *et al.* (2017b).
663 Diversity and oceanic distribution of prasinophytes clade VII, the dominant group of green
664 algae in oceanic waters. *ISME J* **11**: 512–528.

665 Martin P, Lampitt RS, Perry MJ, Sanders R, Lee C, D’Asaro E. (2011). Export and mesopelagic
666 particle flux during a North Atlantic spring diatom bloom. *Deep Sea Res Part I-
667 Oceanographic Res Pap* **58**: 338–349.

668 Mayor DJ, Sanders R, Giering SL, Anderson TR. (2014). Microbial gardening in the ocean’s
669 twilight zone: detritivorous metazoans benefit from fragmenting, rather than ingesting,
670 sinking detritus: fragmentation of refractory detritus by zooplankton beneath the euphotic
671 zone stimulates the harvestable productio. *Bioessays* **36**: 1132– 1137.

672 Michaels AF, Caron D a., Swanberg NR, Howse F a., Michaels CM. (1995). Planktonic
673 sarcodines (Acantharia, Radiolaria, Foraminifera) in surface waters near Bermuda:
674 abundance, biomass and vertical flux. *J Plankton Res* **17**: 131–163.

675 Noji TT. (1991). The influence of macrozooplankton on vertical particulate flux. *Sarsia* **76**: 1–
676 9.

677 Not F, Gausling R, Azam F, Heidelberg JF, Worden AZ. (2007) Vertical distribution of picoeukaryotic
678 diversity in the Sargasso Sea. *Environ. Microbiol.* **9**, 1233–52.

679 Not F, del Campo J, Balagué V, De Vargas C, Massana R. (2009). New insights into the
680 diversity of marine picoeukaryotes. *PLoS One* **4**: e7143.

681 Ohman M, Barbeau K, Franks P, Goericke R, Landry M, Miller A. (2013). Ecological
682 Transitions in a Coastal Upwelling Ecosystem. *Oceanography* **26**: 210–219.

683 Ohman MD, Powell JR, Picheral M, Jensen DW. (2012). Mesozooplankton and particulate
684 matter responses to a deep-water frontal system in the southern California Current System.
685 *J Plankton Res* **34**: 815–827.

686 Oksanen J, Blanchet FG, Friendly M, Kindt R, Legendre P, McGlenn D, *et al.* (2017). Vegan:
687 Community Ecology Package. R package version 2.4-5 [https://CRAN.R-
688 project.org/package=vegan](https://CRAN.R-project.org/package=vegan).

689 Pearman JK, Casas L, Merle T, Michell C, Irigoien X. (2016). Bacterial and protist community
690 changes during a phytoplankton bloom. *Limnol Oceanogr* **61**: 198–213.

691 Peterson WT, Fisher JL, Strub PT, Du X, Risien C, Peterson J, *et al.* (2017). The pelagic
692 ecosystem in the Northern California Current off Oregon during the 2014-2016 warm
693 anomalies within the context of the past 20 years. *J Geophys Res Ocean* **122**: 7267–7290.

694 Piredda R, Tomasino MP, D’Erchia AM, Manzari C, Pesole G, Montresor M, *et al.* (2017).
695 Diversity and temporal patterns of planktonic protist assemblages at a Mediterranean Long
696 Term Ecological Research site. *FEMS Microbiol Ecol* **93**. e-pub ahead of print, doi:
697 10.1093/femsec/fiw200.

698 Schloss PD, Westcott SL, Ryabin T, Hall JR, Hartmann M, Hollister EB, *et al.* (2009).
699 Introducing mothur: Open-Source, Platform-Independent, Community-Supported Software
700 for Describing and Comparing Microbial Communities. *Appl Env Microbiol* **75**: 7537–
701 7541.

702 Silver MW, Gowing MM, Brownlee DC, Corliss JO. (1984). Ciliated protozoa associated with
703 oceanic sinking detritus. *Nature* **309**: 246–248.

704 Smetacek V, Klaas C, Strass VH, Assmy P, Montresor M, Cisewski B, *et al.* (2012). Deep
705 carbon export from a Southern Ocean iron-fertilized diatom bloom. *Nature* **487**: 313 -319.

706 Steinberg DK, Landry MR. (2017). Zooplankton and the Ocean Carbon Cycle. *Ann Rev*
707 *Mar Sci* **9**: 413–444.

708 Steinberg DK, Van Mooy BAS, Buesseler KO, Boyd PW, Kobari T, Karl DM. (2008).
709 Bacterial vs. zooplankton control of sinking particle flux in the ocean’s twilight zone.

710 *Limnol Oceanogr* **53**: 1327–1338.

711 Stemann L, Youngbluth M, Robert K, Hosia A, Picheral M, Paterson H, *et al.* (2008).

712 Global zoogeography of fragile macrozooplankton in the upper 100 – 1000 m inferred

713 from the underwater video profiler. *ICES J Mar Sci* **65**: 433–442.

714 Stukel MR, Biard T, Krause J, Ohman MD. (2018) Large Phaeodaria in the twilight zone: Their

715 role in the carbon cycle. *Limnol Oceanogr in press*

716 Stukel MR, Aluwihare LI, Barbeau KA, Chekalyuk AM, Goericke R, Miller AJ, *et al.* (2017).

717 Mesoscale ocean fronts enhance carbon export due to gravitational sinking and subduction.

718 *Proc Natl Acad Sci U S A* **114**: 1252–1257.

719 Stukel MR, Kahru M, Benitez-Nelson CR, Décima M, Goericke R, Landry MR, *et al.* (2015).

720 Using Lagrangian-based process studies to test satellite algorithms of vertical carbon flux

721 in the eastern North Pacific Ocean. *J Geophys Res Ocean* **120**: 7208–7222.

722 Stukel MR, Landry MR, Benitez-Nelson CR, Goericke R. (2011). Trophic cycling and carbon

723 export relationships in the California Current Ecosystem. *Limnol Oceanogr* **56**: 1866–

724 1878.

725 Suzuki N, Aita Y. (2011). Radiolaria: achievements and unresolved issues: taxonomy and

726 cytology. *Plankt Benthos Res* **6**: 69–91.

727 Suzuki N, Not F. (2015). Biology and Ecology of Radiolaria. In: Ohtsuka S, Suzaki T,

728 Horiguchi T, Suzuki N, Not F (eds). Springer: Japan, pp 179–222.

729 Takahashi K. (1983). Radiolaria - sinking population, standing stock, and production-rate. *Mar*

730 *Micropaleontol* **8**: 171–181.

731 Takahashi K. (1987). Radiolarian flux and seasonality: climatic and El Nino response in the

732 subarctic Pacific, 1982–1984. *Global Biogeochem Cycles* **1**: 213–231.

733 Taylor AG, Landry MR, Selph KE, Wokuluk JJ. (2015). Temporal and spatial patterns of

734 microbial community biomass and composition in the Southern California Current

735 Ecosystem. *Deep Sea Res Part II Top Stud Oceanogr* **112**: 117–128.

736 Taylor GT, Karl DM, Pace ML. (1986). Impact of bacteria and zooflagellates on the

737 composition of sinking particles: an in situ experiment. *Mar Ecol Prog Ser* 141–155

738 Tennekes M. (2017). Treemap: Treemap Visualization. R package version 2.4-2.

739 Turner JT. (2015). Zooplankton fecal pellets, marine snow, phytodetritus and the ocean's

740 biological pump. *Prog Oceanogr* **130**: 205–248.

741 Wickham H. (2009). ggplot2: Elegant Graphics for Data Analysis. Springer-Verlag: New York.

742 Zouari AB, Hassen MB, Balague V, Sahli E, Ben Kacem MY, Akrouf F, *et al.* (2018).

743 Picoeukaryotic diversity in the Gulf of Gabes: variability patterns and relationships to

744 nutrients and water masses. *Aquat Microb Ecol* **81**: 37–53.

745

746

747

TABLES

748

Table 1. Water column physical, chemical and biological properties averaged from 7 CTD casts

749

conducted throughout the 3-days duration of each Cycle.

	C1	C2	C3	C4	C5
Salinity	33.47 ± 0.0	33.42 ± 0.0	33.33 ± 0.0	33.02 ± 0.0	33.09 ± 0.0
Temperature (°C)	16.5 ± 0.0	16.9 ± 0.0	18.6 ± 0.1	19.2 ± 0.0	19.8 ± 0.0
Mixed layer depth (m)	17.6 ± 4.3	29.3 ± 5.7	17.3 ± 5.9	24.3 ± 4.8	28.6 ± 4.4
Nitracline depth (m)	30.9 ± 5.8	40.4 ± 14.0	33.1 ± 6.0	70.6 ± 2.1	91.6 ± 5.7
Surface Chl <i>a</i> (µg/L)	0.59 ± 0.15	0.69 ± 0.08	0.21 ± 0.03	0.10 ± 0.02	0.08 ± 0.01
Surface Chl <i>a</i> > 20 µm (%)	42.6 ± 5.0	39.0 ± 5.7	3.4 ± 1.2	1.9 ± 0.3	1.8 ± 0.1

750

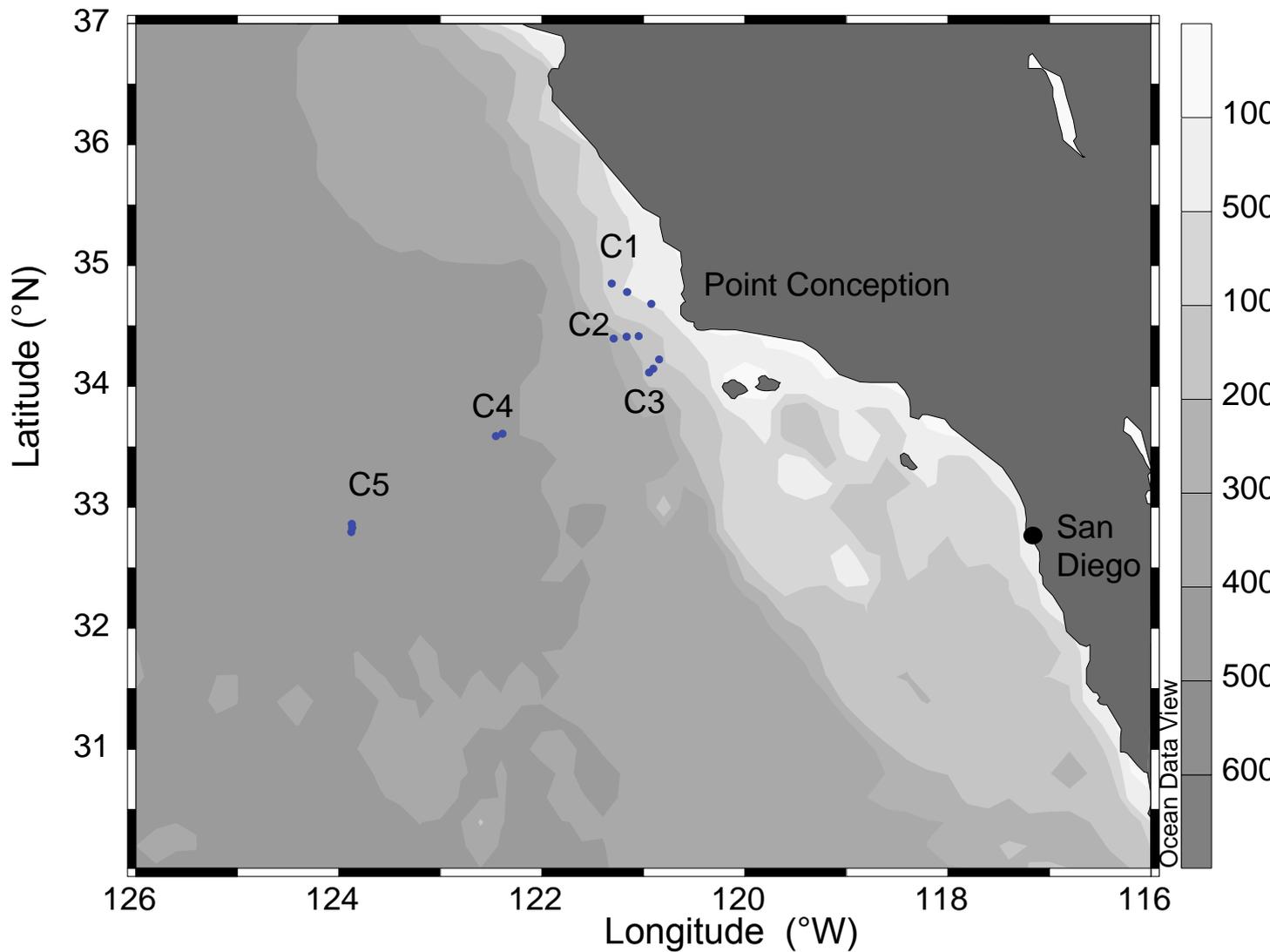


Figure 1. Map showing the location of the different water masses surveyed. Each waterparcel was tracked for three consecutive days called Cycles (C1-C5). Blue dots represent the position of the water parcel during the predawn CTD carried out daily during the cycles.

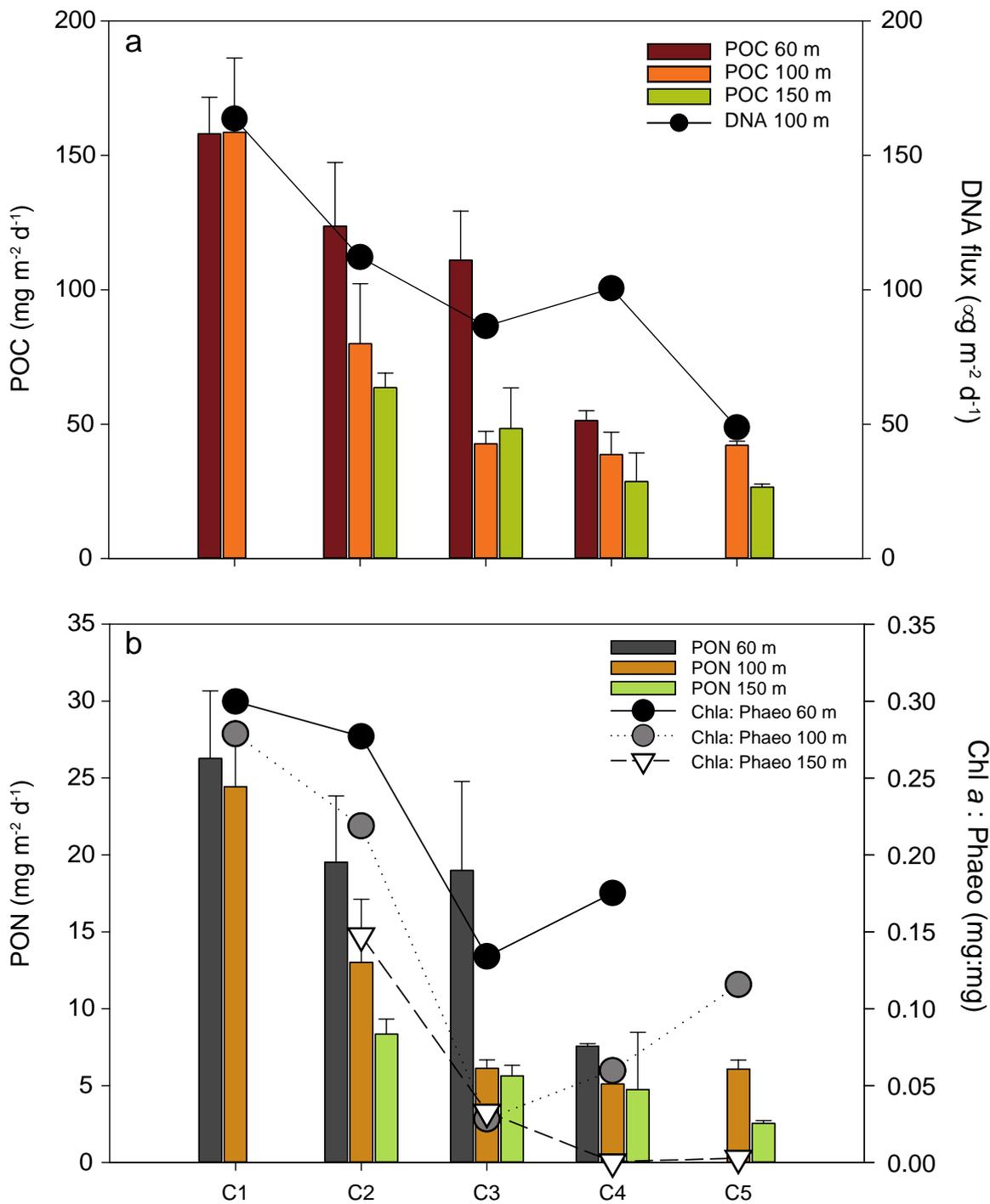


Figure 2. (a) POC and DNA fluxes in standard fixed sediment traps. (b) PON fluxes and Chlorophyll to Phaeopigments ratio in the traps. Error bars refer represent the standard deviation of POC and PON fluxes ($n = 3$) in traps deployed at 60, 100, and 150 m.

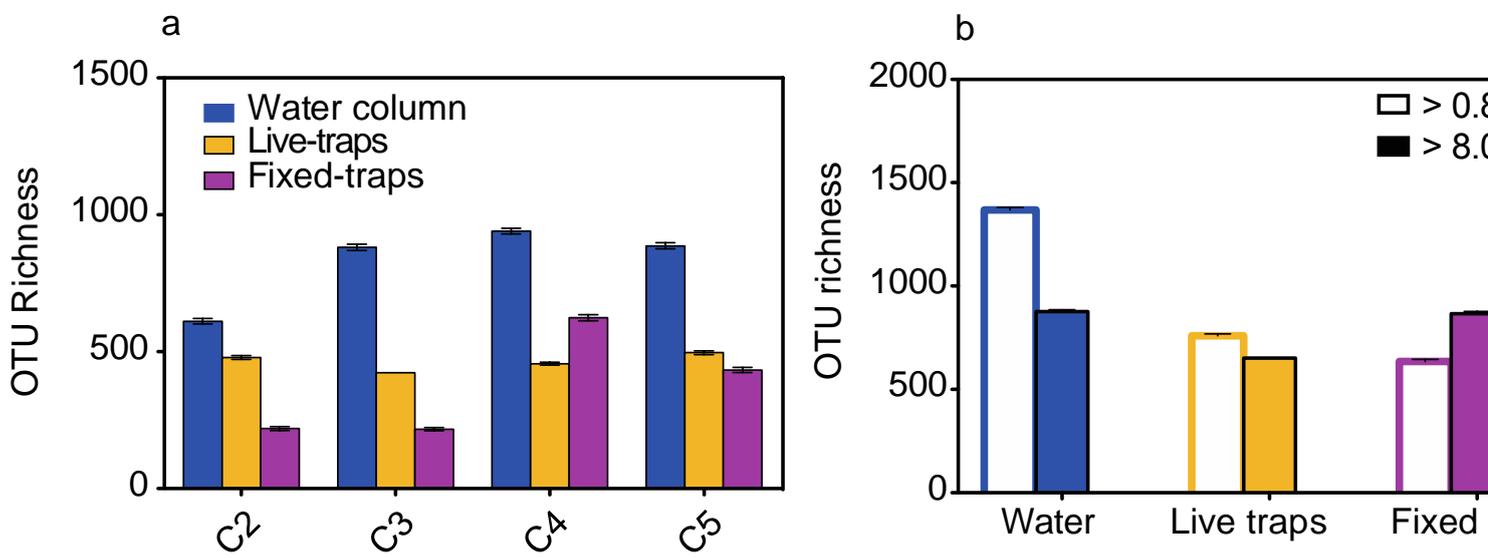


Figure 3. Protistan OTU richness calculated for the minimum sampling effort common to all samples compared. (a) Water column, live- and fixed-sediment trap samples comparison across different cycles. (b) Small and large size fraction comparison in different sample types. Colours refer to samples collected from the water column (blue), live (yellow), and fixed traps (purple).

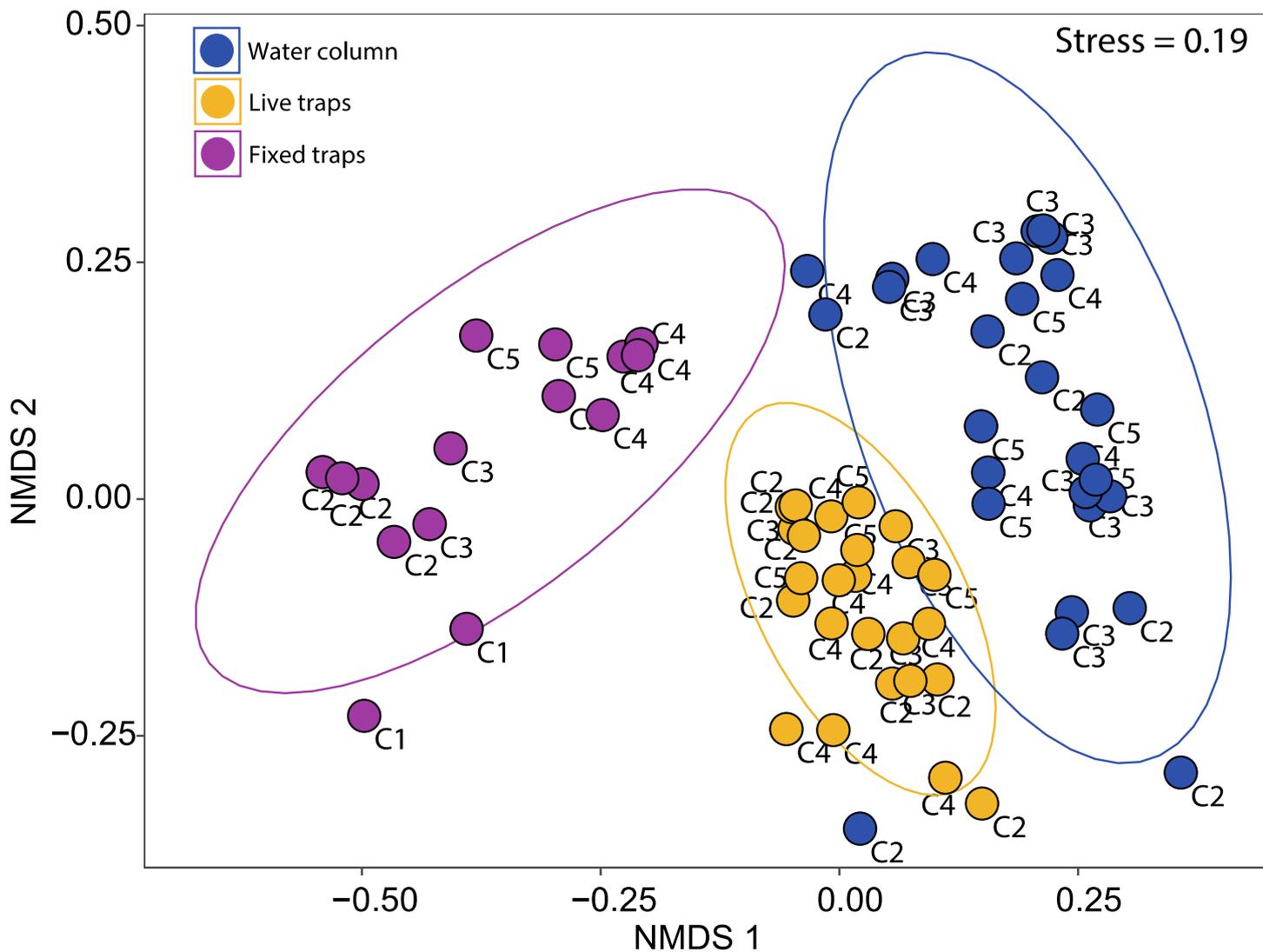


Figure 4. Non-parametric multidimensional scaling ordination (nMDS) plot in 2-dimension configuration (K=2) based on Bray-curtis dissimilarity between protist OTU community composition of all samples. 95% confidence ellipses for each method type are represented.

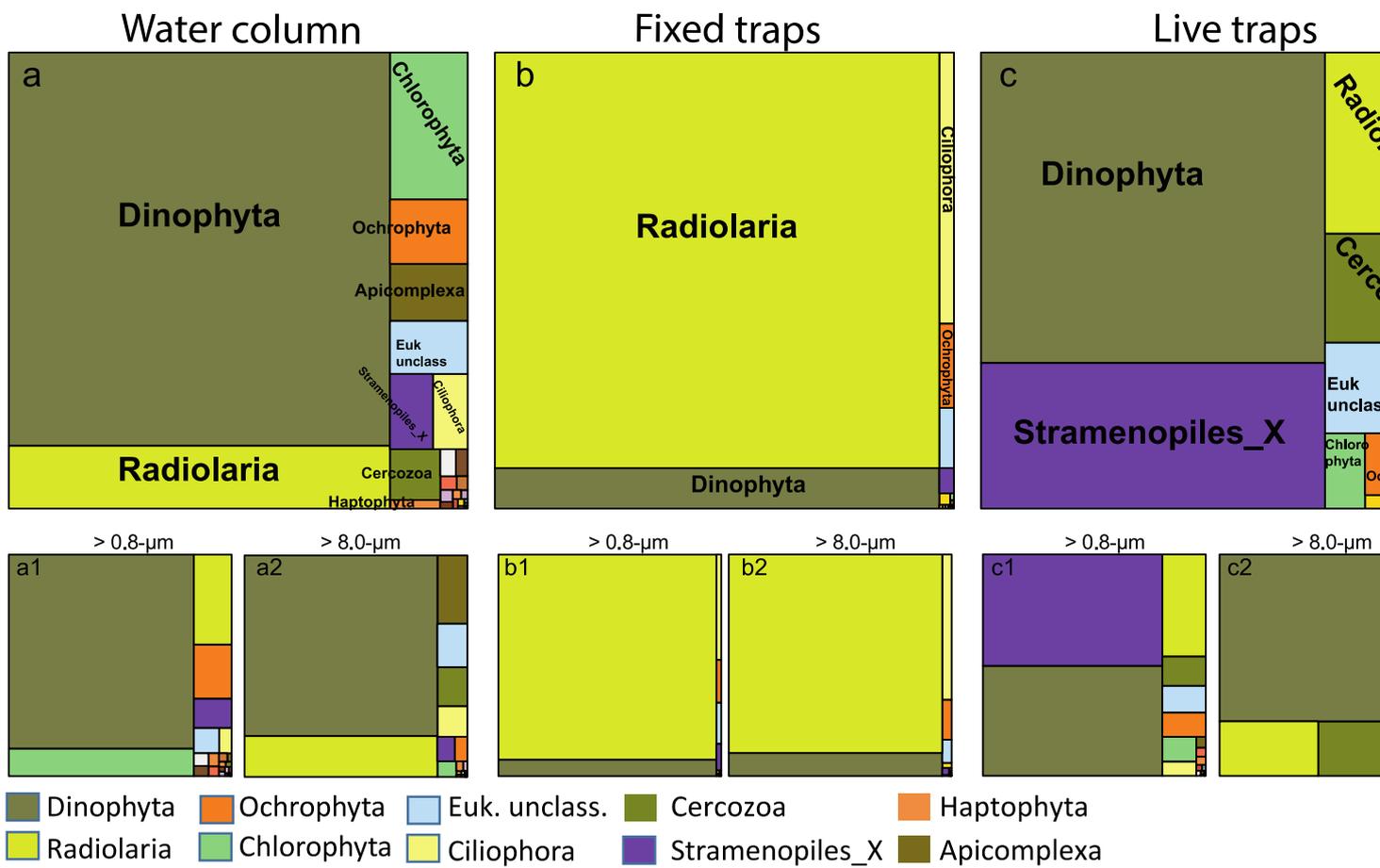


Figure 5. Mean percentage of 18S rDNA reads affiliated to protistan divisions and major taxonomic groups in samples from the water column (a), fixed (b) and live traps (c). Total (> 0.8 µm) and large (> 8 µm) size-fractions are presented at the bottom left (1) and right (2) of each sample type, respectively.

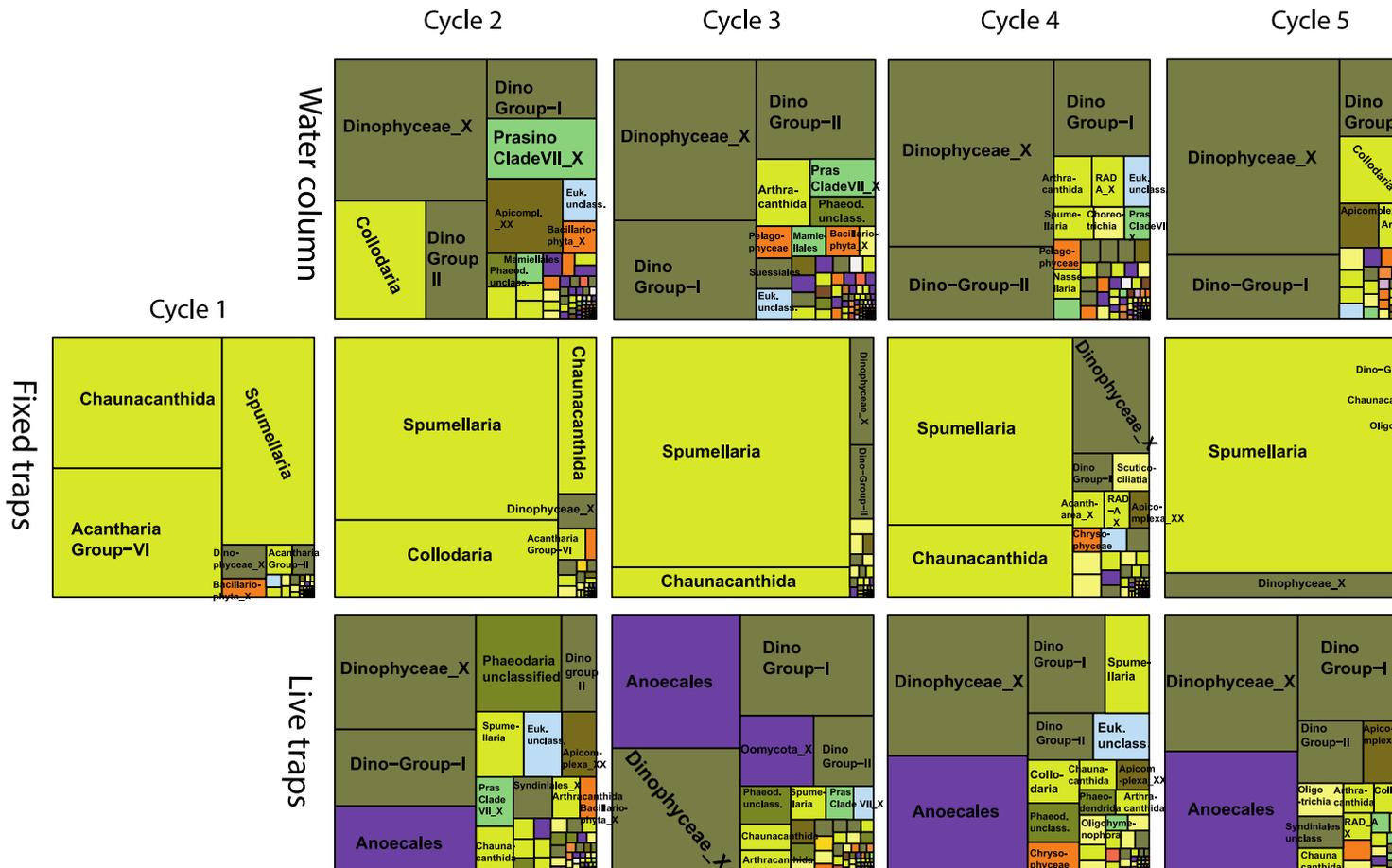


Figure 6. Mean percentage of 18S rDNA reads affiliated to protistan taxonomic groups (classes and orders mainly) in water-column, fixed- and live-trap samples across different sampling cycles. The area represents the mean percentage of reads affiliated to each protistan group (only groups with > 1% mean contribution are labelled). Color codes, as in Figure 5, represent the taxonomic affiliation, mainly at the division level.

● > 8.0- μ m ▲ > 0.8- μ m

Fraction of protistan sequences

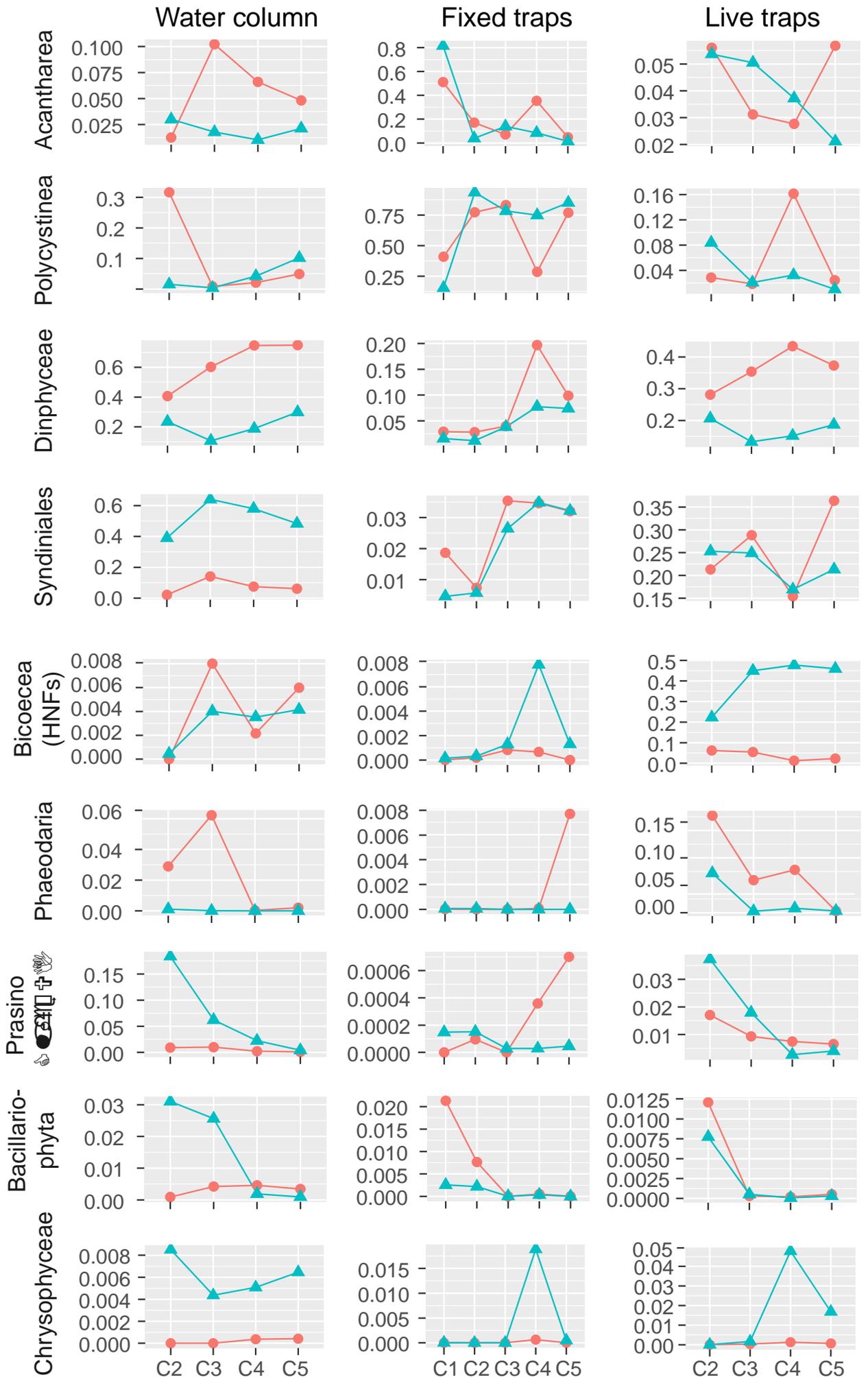


Figure 7. Mean relative abundance of 18S rRNA sequences of most abundant protistan groups in the total (> 0.8 μm) and large (> 8 μm) size-fraction samples from the upper water column (0-100 m), fixed- and live-traps, across different cycles.

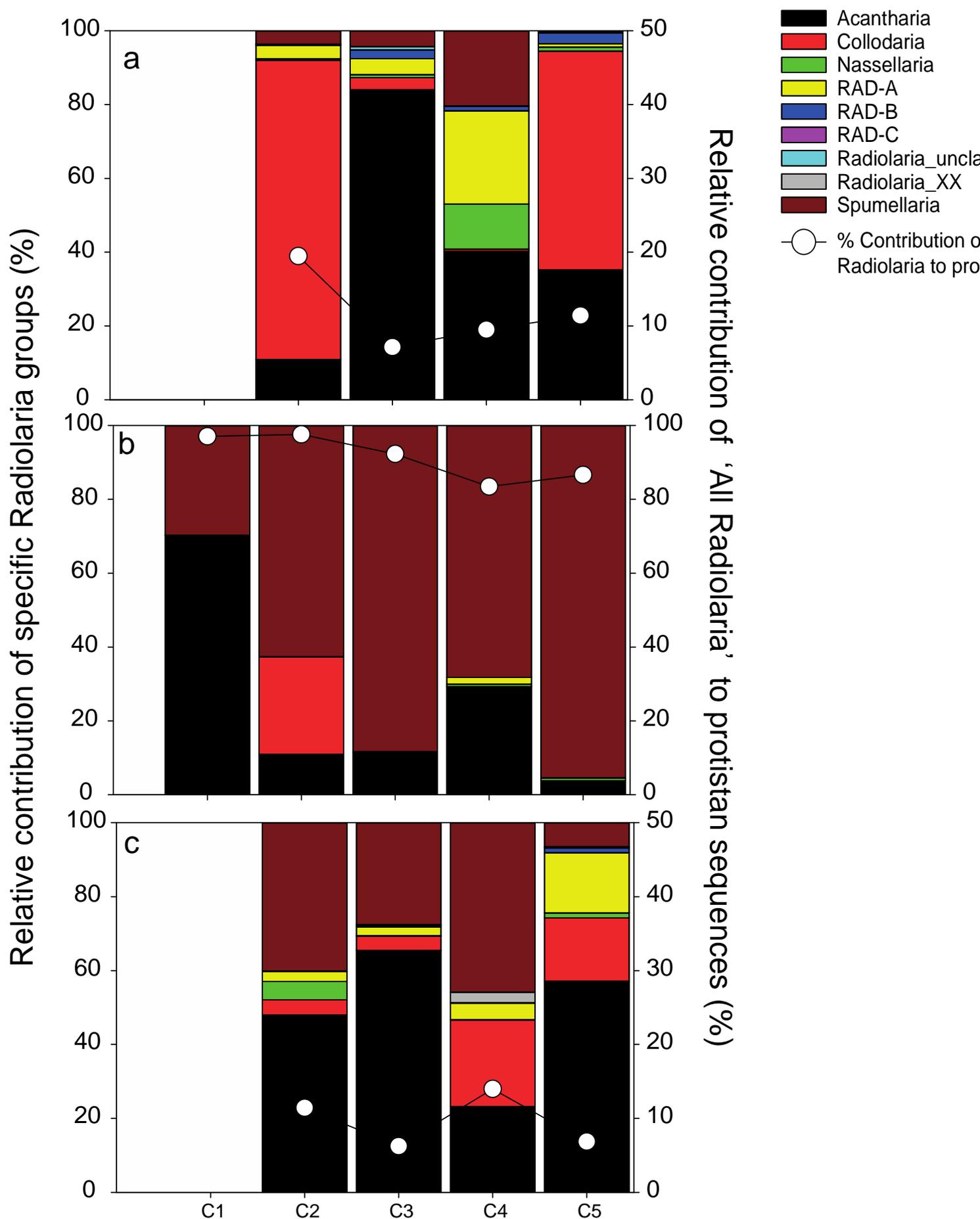


Figure 8. Relative contribution of Radiolaria to total protistan sequences (solid points) and percentage contribution of specific radiolarian groups in the water column (a), fixed (b) and live (c) trap samples across different cycles.