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Discerning dominant temporal patterns of bio-optical properties in the northwestern Mediterranean Sea (BOUSSOLE site)

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16 Highlights:17

- The Chl-Fluo variability has driven mostly by annual cycle.
- The c_p and b_{bp} coefficients have driven mostly by 6-months cycle.
- During winter, mid-term cycles (greater than 10 days) dominate the intra-seasonal signal for all parameters.
 - During summer, diel cycle has a strong impact on the intra-seasonal variability for all parameters.

23

24 Abstract

25

26 A wavelet analysis has been applied, for the first time, to 3-year high-frequency field observations 27 of bio-optical properties (i.e. chlorophyll-fluorescence, beam attenuation and backscattering 28 coefficients) in the northwestern Mediterranean Sea (BOUSSOLE site), in order to identify their 29 dominant temporal patterns and evolution. A cross-wavelet and coherence analysis has also been 30 applied to paired bio-optical coefficients time-series at the BOUSSOLE site, which allows 31 identifying the temporal relationship between the cycles of the bio-optical properties. Annual, six-32 and four- month, intra-seasonal (i.e., mid- and short-terms) cycles are identified from the time-33 series analysis. The periodicities of chlorophyll-fluorescence, beam attenuation and particulate 34 backscattering coefficients correlate well at different temporal scales and specific seasons. At 35 annual, six- and four- month scales, different bio-optical properties follow rather similar patterns, likely driven by physical forcing. Intra-seasonal variability consists in both mid- and short-term 36 37 variations. The former dominates during the winter and are related to episodic bloom events, while 38 the latter variations (*i.e.*, diel) prevail during summer, in a stratified water column.

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39 1. Introduction

40 Phenology is the study of the timing of periodic life events (Morren, 1849a; Winder et al., 2010). It

was first investigated on terrestrial plants, which are sensitive to climate variability (*Cleland et al.*,
2007; *Winder et al.*, 2010). Phenology of marine ecosystems has been studied for a few decades,

42 2007, *Winder et ut.*, 2010). Filehology of Infamile ecosystems has been studied for a few decades,

- especially by use of bio-optical proxies that are accessible through satellite ocean color remotesensing. In the ocean, phytoplankton are considered as a sentinel of changes in the ecosystems,
- 45 because they respond rapidly to environmental perturbations (Bode et al., 2015). In such a context,
- 46 several studies on phytoplankton phenology (*i.e.* seasonal and annual cycles) have been conducted
- in marine ecosystems, from inland to open ocean waters, using chlorophyll-*a* data (*Behrenfeld et al.*2010, 2016; *Winder et al.* 2010; *Carey et al.* 2016; *Mignot et al.*, 2018). Other cycles (*e.g.* lowest
- 49 temporal cycles) are less well-known.

50 In the ocean, changes in nutrients and light conditions have an impact on the phytoplankton 51 standing stock and on the intracellular chlorophyll concentration (Chl), which is widely used as a 52 proxy for phytoplankton biomass (Volpe et al., 2012; Siegel et al., 2013). In the Mediterranean Sea, 53 under high nutrients and low light, phytoplankton growth rate shows the maximum concentration 54 values. When light intensity increases, there is no longer a need for the cells to produce and sustain 55 large amounts of the energetically expensive chlorophyll pigment. Under these conditions, phytoplankton exploit the nutrients still present in the upper layer. This, together with the light 56 57 conditions provided by the increased stratification, allows phytoplankton to grow, despite 58 phytoplankton chlorophyll concentration decreases. The condition of the strong increase in light 59 determines the decline of phytoplankton pigment demand, while low nutrient concentrations limit phytoplankton population growth and division rates. Under low nutrients (generally the period of 60 61 highest stratification of water column) and high light conditions, phytoplankton shows the minima 62 of abundance. When light starts to decrease and mixing occurs again, phytoplankton assign the 63 energy from the newly available nutrients into the production of chlorophyll, however, limiting their growth rate (Lavigne et al., 2013; Bellacicco et al. 2016; Barbieux et al., 2018). 64

65 In temperate seas like the Mediterranean Sea, a major feature of the annual cycle is the spring bloom, as it occurs, for instance, in the Gulf of Lion (D'Ortenzio et al., 2014; Mayot et al., 2017). 66 This bloom typically persists for a few weeks to months. The variability of phytoplankton annual 67 68 and seasonal patterns has thus been well studied (Behrenfeld et al. 2010; Volpe et al., 2012; 69 Lavigne et al., 2013, Mignot et al., 2014), while no systematic analyses at shorter time scales have 70 been conducted to identify the characteristic periods of biomass variability and recurrence along years at these scales. Winder et al. (2010) defined the mid- and low-term phytoplankton fluctuations 71 72 as irregular blooms that are often responses to short-term weather events which affect sea 73 temperature and vertical mixing dynamics. However, several works have shown that Chl can be 74 complemented with other parameters to describe the ocean ecosystem complexity and 75 phytoplankton dynamics (Behrenfeld et al. 2005, 2006; Ji et al., 2010). These works demonstrated 76 that the physiological processes affect the carbon to chlorophyll ratio (Halsev and Jones, 2015), 77 especially during intermediate periods of light and nutrients variations, as it occurs in the 78 Mediterranean Sea (Bellacicco et al., 2016). Furthermore, biological processes such as 79 phytoplankton photosynthesis and cellular growth and division can be associated with short-term 80 temporal variations like the daily light-dark cycle (Neveux et al., 2003; Poulin et al., 2018). Diel 81 variability is often not taken into account in the current ocean colour algorithms, but many 82 laboratory experiments and studies have demonstrated that, in seawater, there is a diurnal variation

of optical properties, from midday to night, that have implications on biomass measurements 83 84 (Poulin et al., 2018). Thus, the use of other proxies, such as optical coefficients, may be helpful to investigate the different temporal scales of phytoplankton avoiding to take into account the effect of 85 physiological processes on Chl content. In such a context, bio-optical relationships have been 86 87 widely established between inherent optical properties (IOPs) and the main biogeochemical 88 parameters, such as the Chl concentration or fluorescence (Huot et al., 2007, 2008; Dall'Olmo et 89 al., 2009, 2012; Brewin et al., 2012; Antoine et al. 2011; Martinez-Vicente et al., 2013; Barbieux et 90 al., 2018; Bellacicco et al., 2018).

91 One of the most studied IOPs is the particulate beam attenuation coefficient, c_p, used as a proxy for 92 particle concentration and sensitive to a size range that includes phytoplankton cells (Claustre et al., 93 1999; Behrenfeld and Boss, 2003). cp is the sum of particulate scattering and absorption 94 coefficients. Its diel variation is primarily due to planktonic adaptation to the daily light cycle, the 95 dynamics of the upper mixed layer and the variations of particles mass (Siegel et al., 1989; Walsh et 96 al., 1995; Stramski and Reynolds, 1993; Durand and Olson, 1998). The interpretation of diel 97 variability of c_p is difficult because it depends on several factors, such as: i) phytoplankton 98 concentration and composition, ii) physiological status (*i.e.* photoacclimation), and iii) 99 concentrations of detritus and small heterotrophs (e.g. heterotrophic bacteria). The compound effect of variations of these factors on cp remains poorly known (Kheireddine et al., 2014). Recently, 100 101 Gernez et al. (2011) and Kheireddine et al. (2014) have shown that the amplitude and phasing of c_p 102 diel cycles vary seasonally, which may result from seasonal changes in nutrient concentrations, phytoplankton abundance, size distribution, and composition. Unfortunately, c_p is not directly 103 derivable from satellite observations, thus limiting its applications. The IOP that is directly 104 105 proportional to the ocean reflectance and can be retrieved from space is the particulate 106 backscattering coefficient, b_{bp} (Lee et al., 2002). Similarly to c_p, b_{bp} is related to particle concentration to the first order, whereas it also contains information on the particle size distribution, 107 108 refractive index, shape and structure of particles (Twardowski et al., 2001, Neukermans et al., 2012, 109 Slade and Boss, 2015). b_{bp} is more influenced than c_p by submicron non-algal particles (Morel and 110 Ahn, 1991; Stramski and Kiefer, 1991, Stramski et al., 2004), yet their magnitude both covary with 111 phytoplankton concentration, allowing relationships between b_{bp} and Chl to be observed (Bellacicco 112 et al., 2016, 2018; Barbieux et al., 2018). Organelli et al., (2018) suggest that most of b_{bp} signal is due to particles with equivalent diameters between 1 and 10 µm, and thus may be significantly 113 114 influenced by phytoplankton. Several studies have shown good relationships between both the optical coefficients and phytoplankton in terms of Chl, carbon or fluorescence, on both global and 115 116 regional scales (Behrenfeld and Boss, 2003; Behrenfeld et al., 2005; Dall'Olmo et al., 2009, 2012; Antoine et al., 2011; Brewin et al., 2012; Martinez-Vicente et al., 2013; Barbieux et al., 2018; 117 118 Bellacicco et al., 2018). The complexity of temporal relationships between c_p, b_{bp}, and 119 phytoplankton biomass proxies should be, thus, studied and observed at a range from diel to annual 120 cycles.

The objectives of this study are to determine the intra-annual dominant temporal patterns (from the lowest up to annual cycles) of the above-mentioned bio-optical parameters, the changes of these cycles over time, the characteristics and recurrence strength at those periods and to define the temporal relationship between the cycles of the bio-optical properties. To this aim a Wavelet Analysis (WA) was applied, for the first time, to a 3-year time-series (2011–2013) of surface Chlfluorescence, b_{bp} and c_p at the BOUSSOLE site in the northwestern (NW) Mediterranean Sea (*Antoine et al.*, 2006; Figure 1), enabling the detection of all the intra-annual dominant temporalpatterns.

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- 130 2. Area of study, data and methods
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132 **2.1 BOUSSOLE site**

133 The BOUSSOLE (BOUee pour l'acquiSition d'une Série Optique a Long termE) project started in 1999, and its activities are developed on a site located in the northwestern Mediterranean Sea, at 134 135 about 32 nautical miles from the French coast (Figure 1). Essential information about the site 136 characteristics, the measurement platform, and the instrumentation are also provided in Antoine et al. (2006, 2008a,b). The site is protected from coastal inputs by the Ligurian Current, which flows 137 138 along the coast toward the southwest (Millot, 1999). The physical conditions of the area show strong seasonality (D'Ortenzio et al., 2014), with deep (~400 m) mixed layers in winter, and a 139 140 marked stratification in summer (~20 m; Antoine et al., 2011; Mayot et al., 2017). Hydrodynamics 141 drive seasonal changes in phytoplankton abundance, which shows a typical mid-latitude temporal pattern. Oligotrophic conditions prevail in summer where Chl is about 0.1 mg m⁻³ (minima ~0.05 142 mg m⁻³), and concentrations increase up to 3–5 mg m⁻³ during the spring bloom, and stay between 143 0.1-0.3 mg m⁻³ the rest of the year (Gernez et al., 2011, Kheireddine et al., 2014). There is, 144 accordingly, a large range of optical properties (Antoine et al., 2006; Gernez et al., 2011), as 145 observed over the entire northwestern Mediterranean Sea (Bosc et al., 2004). 146

A moored buoy has been permanently deployed at the BOUSSOLE site since September 2003 and operates in a quasi-continuous mode, with data acquisition for one minute every 15 min both night and day. Adequate measures have to be taken to minimize or eliminate bio-fouling, which is unavoidable with moored instrument. All instruments installed on the BOUSSOLE buoy are cleaned by divers about every 2 weeks.



- 154 Figure 1: The area of the north-western Mediterranean Sea showing the main current branches (grey arrows),
- and the location of the BOUSSOLE site in the Ligurian Sea (black square) (*Antoine et al.*, 2011).

156 2.2 Optical measurements

The volume scattering function at 140°, $\beta(140)$, is measured using in alternation two HOBI Labs 157 Hydroscat-4 backscattering meters installed at the lower measurements depth of the buoy (ca. 9 m). 158 159 Instruments are calibrated before deployments (which last about 6 to 12 months) and are equipped with filters at 442, 488, 550, and 620 nm, here only the green band is used. The instruments operate 160 161 at 1 Hz, so that about 60 measurements are recorded during 1 minute, from which the median is taken as representative for $\beta(140)$. Dark current measurements are performed on site with a 162 neoprene cap covering the instrument windows, average dark readings are subtracted to the time-163 series for each deployment. The $\beta(140)$ values are also corrected for attenuation along the 164 measurement path (the $\sigma(\lambda)$ correction of Maffione and Dana, 1997) using c_p (see below) and the 165 total absorption coefficient derived from inversion of the diffuse attenuation coefficient for 166 167 downward irradiance (K_d) and the irradiance reflectance (R). K_d and R are retrieved from parallel measurements performed with a set of Satlantic OCR-200 series radiometers. b_{bp} is derived from 168 169 the corrected $\beta(140)$ as follows (Maffione and Dana 1997; Boss and Pegau, 2001):

170
$$b_{bp}(550) = 2\pi \chi_p(\beta(140,550) - \beta_w(140,550))$$
 (m⁻¹) [1]

171 where $\chi_p = 1.13$ (*D. R. Dana and R. A. Maffione*, unpublished manuscript, 2014) and β_w (140), the 172 pure seawater scattering at 140°, is computed following *Zhang et al.*, (2009); *Zhang and Hu* (2009) 173 using the temperature and salinity measured at the same depth with a Sea-Bird Scientific SBE-37SI 174 CTD sensor.

The particulate transmittance $(Tr_p, \%)$ at 650 nm is measured at 4 and 9 m with 25 cm path length WETLabs C-Star transmissometers (acceptance angle is 1.2°). Instruments are factory calibrated with deionized, ultra-filtered, UV-screened water. The corresponding particulate beam attenuation coefficient, c_p , is then calculated as:

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$$c_{p}(650) = -\frac{1}{0.25} \ln \left(\frac{Tr_{p}(650)}{100} \right) \qquad (m^{-1})$$
[2]

181

This assumes that absorption by colored dissolved organic matter (CDOM) is negligible at 650 nm
(*Bricaud et al.*, 1981).

184 Chlorophyll-a fluorescence (Chl-Fluo) is measured in RFU with an ECOFLNTUS WET Labs (now 185

185 Sea-Bird Scientific) fluorimeter (470 nm EX/695 nm EM; see
 186 https://www.seabird.com/combination-sensors/eco-flntu/family?productCategoryId=54758054352)

- 187 at 4 m and 9 m depth.
 - 188 In this study, we selected b_{bp} , c_p and Chl Fluo at the depth of 9 m in order to compare them together.
 - Here, b_{bp} is referred to b_{bp} at 550 nm (m⁻¹), c_p to c_p at 650 nm (m⁻¹).

190 2.3 Multi-Channel Spectral Analysis (M-SSA)

191 For each parameter, the period from to 2011 to 2013 was used for a total of 105216 measurements

192 after quality control (Figure 2). It consists in removing outliers using three standard deviations $(\pm \sigma)$ 193 confidence limit.

194 An important prerequisite for applying WA is that the time-series has to be continuous at the 195 minimum considered frequency. Missing data, including those not passing quality control, represented 19.57%, 13.40% and 22.38% of Chl-Fluo, b_{bp} and c_p time-series, respectively. Gaps 196 197 were filled using a Multi-Channel Singular Spectral Analysis (M-SSA) technique which is a non-198 parametric spectral estimation method relying on data only (Ghil et al., 2002; Kondrashov et al. 199 2006, 2010). This technique is not based on *a priori* parametrized family of probability distribution. 200 The method uses temporal correlation to fill in the missing data and represents a generalization of 201 the Beckers et al., (2003) spatial empirical orthogonal functions-(EOFs) based reconstruction. 202 Kondrashov and Ghil (2006) demonstrated that an increased number of gaps yields the same effect 203 as an increase of the noise in the measurements. Two different inputs are required to apply M-SSA 204 for field reconstruction: window-length (W) and components (M). Both depend on the 205 characteristics of the time-series, and need to be accurately defined to avoid any bias in the 206 reconstructed fields. The W represents the length of the sliding window (expressed in number of observation) used in the M-SSA in order to identify the leading components of the time-series (Ghil 207 208 et al., 2002; Kondrashov et al. 2006, 2010). Diversely, M is the number of eigen-functions used for 209 signal reconstruction. Here, we applied the M-SSA to the three time-series using specific W(W=5000) and M components (*i.e.*, M = 1 up to 20) following the recommendations listed in *Ghil et* 210 al., (2002) and Kondrashov et al. (2005, 2006, 2010). These settings are compatible with the 211 properties of the time-series hereby analyzed, taking into account long, mid- and short-term 212 213 variations. Figure 2 (a, b, c) shows the time series of each parameter with missing data reconstructed after application of M-SSA technique. Figure 2d is the M-SSA spectrum of filled 214 215 time-series with W=5000. The optimum number M=20 corresponds to the number of modes that 216 explain more than 95% of the variance the M-SSA spectrum.



Figure 2: Time-series of Chl-Fluo (a), b_{bp} (b), c_p (c) for 2011 – 2013 at the BOUSSOLE site. Original data are displayed with blue lines, while the gaps-filled by the M-SSA technique are highlighted with red lines. The letters W stands for winter, S for spring, S for summer and F for fall. Panel d is the M-SSA spectrum of filled time-series with the break of the slope spectrum at 20.

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- 223

224 **2.4 Wavelet analysis**

225 Since it was first introduced by Morlet (1982; Part I and II), WA has been widely applied to 226 different fields of science. The main characteristic of the WA lays on the decomposition time-227 series, and its time-scale localization and amplitude. Usually, a signal of the series can be 228 decomposed into different harmonic components using, for example, the Fourier method. This can be defined as a partition of the variance of the series into its different oscillating components with 229 230 different frequencies (i.e. the periods). The spectral frequency analysis based on the widely-used 231 Fourier method makes the assumption that the statistical properties of the time-series do not vary with time, being stationary and constant. In such a context, the oceanographic processes do not 232 233 respect the stationary assumption, and there are evidences of the non-stationary nature of biooptical properties (e.g. bbp or cp) along a single year (Antoine et al., 2011; Gernez et al., 2011; 234 Dall'Olmo et al., 2012; Barnes et al., 2014; Kheireddine et al., 2014). But, the WA overcomes this 235 236 problem of non-stationary conditions by performing a local time-scale decomposition of the signal. Thus, WA provides time-dependent spectra (Lau and Weng 1995; Torrence and Campo, 1998, 237 238 Percival and Walden, 2000; Ampe et al., 2014). This approach helps to track how the different 239 scales are related to the periodic components of the signal. WA is applicable to stationary or non-240 stationary time-series and quantifies correlation between two signals (Daubechies, 1992; Lau and Weng 1995; Cazalles et al., 2008; Garcia-Reyes et al., 2013). Figure 3 gives an example of 241 application of WA to a stationary synthetic signal and provides elements to interpret the results 242 243 obtained by this analysis as well as definition of the main WA indexes.

Additionally, in order to understand which are the main dominant cycles that explain the variability on a three-year's time-series, the coefficient AWP* is computed as follows:

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$$AWP_{k}^{*} = \frac{AWP_{period}}{AWP_{maximum}}$$

248

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AWP $_{k}^{*}$ is, thus, the ratio between the average wavelet power (AWP_{period}; Table 1) of each local maximum with respect to the absolute AWP maximum for each parameter k (AWP_{maximum}; Table 1). The ratios allow the rank of cycles for each parameter evaluating which cycles are the most important (values approximately 1) and which are not (values close to 0).

253 A cross-wavelet analysis (CWA) has also been applied to paired bio-optical coefficients time-series 254 at BOUSSOLE site, which allows identification of the temporal relationship between the cycles of 255 the bio-optical properties. The CWA is fundamentally a comparison between the spectra of two time-series, x(t) and y(t) (*Chatfield*, 1989), sampled with the same time step. It results in a quantity, 256 257 the cross-wavelet coherence, which can assume values between 0 and 1, indicating the cross-258 correlation between the spectra of two time-series, as a function of the period. The most important 259 information obtained by CWA is the identification of the portions of x(t) that covaries with v(t) at 260 specific periods. An output from CWA, coupled to the wavelet coherence spectra (WCS), is its 261 time-average (the average coherence; AC). The AC is equal to 1 when there is a perfect linear relationship at particular periods between the two time-series spectra. In this work, the CWA has 262 263 been used to investigate the strength of the relationship between bio-optical properties at different temporal scales (e.g. Figs. 6, 9, 10, 11) and to understand the relative phases between the time-264 265 series. For more information about the theoretical background and applications of WA and CWA 266 see Torrence and Combo, (1998) and Cazalles et al., (2008).



268 Figure 3: Theoretical example of a three-year time-series of data at 15 minutes' interval, built by overlapping 6 continuous sinusoids at 1, 32, 64, 128, 256 and a two-year 365 days' cycles (a). The wavelet power spectrum (WPS) resulting from the WA is shown in (b) as a function of time. The y axis is the period and colors indicate the power levels of the time-series (high levels in red and low levels in blue). The time-series has a strong (cyclical) signal for the periods and duration of time in correspondence of the black lines in the WPS. The shaded area has not to be considered as it might provide false periodic events (Torrence and Compo, 1998). The thin white contours surrounding regions of stronger variance in the spectra indicate coherent time-frequency regions that are significant (*i.e.* > 95% significance). Panel c shows the average of WPS which is the average of the power levels for each period over the whole time-series. Local maxima in the AWP (vellow dots) indicate periods contributing significantly to the variance of the time-series, providing an efficient method to detect and identify periodicities, if present. In red are highlighted periods, and correspondent AWP, with low statistical significance (less than 95%; Roesch and Schmidbauer, 2014).

Acronym	Definition
Chl	Chlorophyll Concentration (mg m ⁻³)
$b_{bp}(\lambda)$	Particulate backscattering coefficient (m ⁻¹)
$c_p(\lambda)$	Beam attenuation coefficient (m ⁻¹)
Chl-Fluo	Chlorophyll-Fluorescence (RFU)
M-SSA	Multi-Channel Singular Spectral Analysis
W	Window Length (N° of observations)
М	Components
WA	Wavelet Analysis
CWA	Cross-Wavelet Analysis
WPS	Wavelet Power Spectra
AWP	Average Wavelet Power
AWP*	Average Wavelet Power Ratio
WCS	Wavelet Coherence Spectra
AC	Average Coherence

289

Table 1: Symbol and acronym definitions.

290 **3. Results**

This section starts from the description of each time-series, throughout the detection of the intraannual dominant temporal patterns, and ends with the description of the changes of the main cycles over time and the definition of the temporal relationships between the different cycles of the biooptical properties here studied.

3.1 Description of the time-series

The time-series of each parameter whose gaps are filled by using the M-SSA technique are shown in Figure 2. The b_{bp} and c_p range of variation is respectively between 10^{-4} m⁻¹ and 10^{-3} m⁻¹ and between 0.1 m⁻¹ and 1.0 m⁻¹ with minima occurring in the summer/fall (*i.e.* oligotrophic season) and maxima occurring in the winter/spring (*i.e.* mesotrophic season), respectively (Figure 2). The Chl-Fluo annual signal ranges between 0.012 RFU, in summer/fall, up to 7.5 RFU in winter/spring.

301 3.2 Dominant temporal patterns

302 Figure 4 and Table 2 display the AWP after the wavelet analysis on the three time-series, and the correspondent AWP*. For the Chl-Fluo nine relative maxima (yellow circles in Figure 4a) emerge 303 304 at periods of approximately 351, 200, 124, 46, 31, 21, 10, 3 and 1 days, indicating nine major patterns of the Chl-Fluo variability. The AWP spectrum for b_{bp} shows seven relative maxima at 305 306 370, 193, 126, 58, 16, 11, 2 days. Finally, for the c_p relative maxima at 372, 192, 130, 57, 27, 15, 11, 4, 2, 1 days are detected. In both cases of b_{bp} and c_p , two additional maxima are found at 979 307 308 and 911 days, respectively. These relative maxima have low AWP (<1) and are related to periods 309 exceeding the upper limit that can be significantly retrieved in this study (*i.e.* 1 year).

For brevity, periods close to 365 days are interchanged with the term "annual", close to 180-days with "6-months", close to 128-days with "4-months".

312 **3.3 Wavelet Power Spectra**

313 The WPS of Chl-Fluo (Figure 5a) reveals a persistent annual periodicity, thereby explaining the

314 greatest amount of variability. Accordingly, the AWP value shows its maximum (30.8) at this

- period (Table 2) with two secondary maxima at 4-months and at 6-months, both accounting for half of the variability with respect to the annual periodicity (14.6 and 15.6 respectively; Table 2).
- 317 In the case of b_{bp}, the annual cycle has an AWP value of 24.3. The most dominant patterns are
- associated to the 6-month cycle with an AWP of 29.7, and 4-month cycle with AWP of 24.3 (Figure
- 4). A significant cycle is also present with a period of 58 days (AWP value of 11.0). The WPS
- 320 (Figure 5b) highlights a range between 11 and 16 days, particularly evident at the winter-to-spring
- 321 transitions, with a clear inter-annual variability and the AWP value for this cycle is less than 5
- 322 (Figure 4). For cycles between 0.5 and 2 days, the AWP is generally lower, yet not zero, indicating
- that some periodicity for b_{bp} at these scales exists though with limited impact on the overall variability (Table 2). Cycles less than 2 days also have less statistical significance along the three
- 325 years of data.
- 326 The c_p cycles at 4-months, 6-months, and 1-year periods are observed in the WPS (Figure 5c),
- resulting in AWP values of 20.3, 33.0 and 31.8, respectively (Figure 4 and Table 2). Other relative
- maxima have average values less than 6. The WPS (Figure 5c) also highlights 1, 2, 4, 11-15, 27 and
- 329 57 day cycles, evident at the transition from winter to spring of 2012. For the period at 1 day, the
- AWP is low, though greater than zero, and still statistically significant, indicating that some c_p variability at this scale exists (Table 2).
- 332 The signal of the annual cycle covers all seasons for all parameters (Figure 5), which have maxima
- at a particular moment of the year: always in winter and spring (Figure 2 and Figure 5). The 6-
- month period is the dominant pattern in the case of b_{bp} and c_{p} , and is the second dominant pattern in
- the Chl-Fluo time-series. The 4-month cycle shows a considerable magnitude for b_{bp} and c_p , but is limited for Chl-Fluo (Table 2). Moving to the lowest temporal cycles, Figure 5 shows how diel
- cycles are detectable for most of the three time-series, confirming the importance of this temporal
 scale, especially for Chl-Fluo and c_p time-series.
- 339 Summarizing, Chl-Fluo shows a strong annual cycle signal along the entire time-series, but 340 particularly from spring 2011 to winter 2013 considering only the lightened areas (Figure 5a). The 341 year 2012 (more specifically from winter 2011 to spring 2013) shows a strong cyclic nature at 342 periods greater than 1 week.
- WPS of b_{bp} shows high power levels from 128 to 365 day periods (Figure 5b). However, for periods
 lower than 64-days, the strength of the WPS is high only during spring and winter.
- 345 Contrarily, WPS of c_p (Figure 5c), shows evident inter-annual variability yet not all the temporal
- footprints (*i.e.* periods) have high values in 2011 and 2013. Dominant cycles are persistent at higher temporal scales from 2011 to 2013, whereas periods lower than 64 days are dominant only in the winter and spring in agreement with the other bio-optical properties.
- 349 The Chl-Fluo AWP is also characterized by two relative maxima in correspondence with 46 and 350 124 days (Table 2), occurring only in the mesotrophic periods of the year (winter – spring of 2012; Figure 5a). The AWP of b_{bp} shows relevant cycles also at the 11-16, 58 and 126 day periods 351 352 occurring in the most productive periods of the year of 2011 - 2013 (winter and spring). 353 Furthermore, the 126-day cycle has a strong influence on the entire time-average spectrum (Figure 4 and 5b). In case of c_p, there is a relevant cycle at 130 days. All these cycles have strong signals 354 355 during the winter and spring seasons (Figure 5). In such a context, Figures 6 illustrates the AC 356 obtained with the CWA between the different bio-optical coefficients along the entire 3-year timeseries of observations. A high correlation (> 0.8) between the periodicities of bio-optical 357 358 coefficients greater than 10 days is observed (Figure 6). Periodicity at these scales has been poorly
- 359 represented in the literature, and with the use of WA all of these cycles can be retrieved and

- described. Figure 5 displays evident cycles in winter 2012 for all parameters (black lines; *i.e.* a 360 productive period), and, therefore, it is an optimum case of study to highlight those lesser-known 361 cycles, such as the intra-seasonal cycles (i.e. diel, weekly, monthly). In order to complement the 362 analysis, an opposite case of study has been analyzed by applying a specific WA on summer 2012, 363 *i.e.* the most oligotrophic period. Moreover, in these seasons, the diurnal variance of parameters is, 364 365 in general, the highest in winter and the lowest in summer (not shown). Finally, the selected data 366 sets have a limited number of missing observations (less than 11 % in both winter and summer seasons), restricting the use of reconstructed measurements. 367
- In order to highlight the mid and short-term variability, a specific WA has thus been applied to two 368 selected seasons of the time-series: winter and summer 2012 (Figures 7 and 8). Figure 7 shows the 369 370 AWP obtained by WA on winter 2012 and Table 3 contains the AWP* values correspondent to the 371 maxima. For the b_{bp}, four significant maxima are retrieved (yellow circles in Figure7a) at periods of 17, 10, 5 and 1 days indicating four major patterns of b_{bp} variability in this season. There is also a 372 373 12 hour cycle that emerges from the analysis, but with reduced statistical significance in respect to 374 other cycles (less than 95%). The c_p AWP shows five relative maxima at 29, 12, 3, 2, and 1 days. A 375 diel cycle is evident and represents the third maximum even if with a low AWP. Ultimately, for the 376 Chl-Fluo, detected relative maxima are 29, 18, 11, 3 and 1 days. In winter, the dominant temporal 377 patterns are of cycles greater than 10 days for all three parameters.
- 378 The opposite season, summer, is the counterpart case of study. This is the period of lowest 379 productivity along the entire year at BOUSSOLE site. Figure 8 shows a time-series for each parameter and the corresponding AWP and AWP* (Table 4). For the b_{bp} different maxima emerged 380 (yellow circles in Figure7a) in correspondence to 18, 10, 7, 4, 2, 1 days and 12 hours indicating 381 382 these cycles as the main driven temporal patterns of variability during the season. The diel cycle is 383 the second important temporal pattern for this coefficient. The c_p AWP shows five relative maxima at 23, 10, 4, 3, 1 days. The diel cycle is the most dominant for c_p in summer (AWP of 22.0). Lastly, 384 the Chl-Fluo has eight relative maxima at 22, 12, 8, 4, 3, 1 days and 12 hours and the diel cycle is 385 386 the most significant local pattern. One could argue that, during summer, the quenching effect can 387 dominate and drive the diel cycle of Chl-Fluo. Therefore, this has to be taken into account in the interpretation of the results for the Chl-Fluo diel cycle (Xing et al., 2017). 388

389 3.4 Cross-Wavelet Analysis

390 The Chl-Fluo vs b_{bp} time-series have high AC values for periods ranging from 1 day ($\cong 0.7$) to 1 year (\cong 1). At lower scales, there is a high correlation only during the winter and spring (Figure 9). 391 This is also true for the Chl-Fluo vs cp and the cp vs bbp time-series which have coherence values 0.7 392 and 0.8 for the 1 day period, respectively, and up to \cong 1 at the annual scale. At periods lower than 1 393 day, the correlation between the three bio-optical parameters is low (<0.5). The Chl-Fluo to c_p 394 395 coherence analysis reveals that the correlation is higher at daily scales (> 0.7) with respect to Chl- $Fluo/b_{bp.}$ The AC value is higher with respect to Chl-Fluo and $b_{bp.}$ This is confirmed by what is 396 397 expected about the relationship between b_{bp} and phytoplankton cells (Loisel et al., 2001; Stramski et 398 al., 2004, Dall'Olmo et al., 2009; 2012).



400 Figure 4: AWP of the Chl-Fluo (a), b_{bp} (b) and c_p (c). The statistical significance is 95% in case of the 401 highest temporal scales, from day to annual scales, for each parameter (blue line). At the lowest temporal 402 scales, the significance is lower than 95% (red line) and the results have to be interpreted with caution. The 403 statistical significance is computed following *Roesch and Schmidbauer* (2014). Yellow dots are the relative 404 local maxima for the single parameter with also information of the exactly days.

Period (days)	351	200	124	-	46	31	21	-	10	3	-	1
AWP _{Chl} [*] _{-Fluo}	1.0	0.50	0.47	-	0.21	0.27	0.27	-	0.16	0.05	-	0.15
Period (days)	370	193	126	58	-	-	-	16	11	-	2	-
$AWP_{b_{bp}}^{*}$	0.82	1.0	0.82	0.37	-	-	-	0.13	0.09	-	0.02	-
Period (days)	372	192	130	57	-	-	25	15	11	4	2	1
$AWP_{c_p}^*$	0.96	1.0	0.62	0.17	-	-	0.09	0.15	0.13	0.02	0.03	0.05

405 Table 2: AWP* from the AWP of each parameter. Numbers in bold indicate the dominant cycles.



407

Figure 5: WPS for Chl-Fluo (a), b_{bp} (b) and c_p (c). The letters W stands for winter (in blue), S for spring (in green), S for summer (in orange) and F for fall (in brown). The time-series has a strong (cyclical) signal for the periods and duration of time in correspondence of the black lines in the WPS. The shaded area has not to be considered as it might provide false periodic events (*Torrence and Compo*, 1998). The thin white contours surrounding regions of stronger variance in the spectra indicate coherent time-frequency regions that are significant (*i.e.* 95% significance). The significance test is computed following *Roesch and Schmidbauer* (2014).





417 Figure 6: AC obtained from the CWA between paired bio-optical properties: Chl-Fluo vs. b_{bp} (green line), Chl-Fluo vs. c_p (black line) and b_{bp} vs. c_p (blue line). In red are highlighted periods where the AC is with low

418 419 statistical significance (less than 95%).





421 422 Figure 7: Time-series of b_{bp} (a), c_p (b) and Chl-Fluo (c) in Winter 2012. AWP of b_{bp}, c_p, Chl-Fluo are in 423 panels d, e, f. Red lines are located where the statistical significance is less than 95% and the results have to 424 be interpreted with caution. The significance test is computed following Roesch and Schmidbauer (2014).

- 425
- 426

Period (days)	-	17	10	5	-	1
$AWP_{b_{bp}}^{*}$	-	1.0	0.52	0.23	-	0.09
Period (days)	29	-	12	3	2	1
$AWP_{c_p}^*$	0.70	-	1.0	0.18	0.31	0.33
Period (days)	29	18	11	3	-	1
AWP _{Chl} - [*] _{Fluo}	0.76	1.0	0.68	0.21	-	0.33

427 Table 3 AWP* from the AWP of each parameter in Winter 2012. Numbers in bold indicate the dominant 428 cycles.



430 431 Figure 8: Time-series of b_{bp} (a), c_p (b) and Chl-Fluo (c) in Summer 2012. AWP of b_{bp}, c_p, Chl-Fluo are 432 reported in panels d, e, f. Red lines are located where the statistical significance is less than 95% and the 433 results have to be interpreted with caution. The significance test is computed following Roesch and 434 Schmidbauer (2014).

-	_	
4	3	5

Period (days)	18	10	7	4	2	1	0.5
$AWP_{b_{bp}}^{*}$	0.61	1.0	0.34	0.43	0.31	0.63	0.52
Period (days)	23	10	-	4	3	1	-
$AWP_{c_p}^*$	0.32	0.30	-	0.11	0.11	1.0	-
Period (days)	22	12	8	4	3	1	0.5
AWP _{Chl} + Fluo	0.44	0.36	0.31	0.22	0.21	1.0	0.30

436 Table 4: AWP* from the AWP of each parameter for Summer 2012. Numbers in bold indicate the dominant 437 cycles.

438

440 **4. Discussion**

441 **4.1 Annual cycle**

The most generally observed pattern was a phytoplankton maximum in the winter/spring season, minimum in summer and a successive increase in fall season, that corresponds to a typical annual

444 cycle of a temperate ocean, as already reported for the north-western Mediterranean Sea (*Antoine et al.*, 2011; *D'Ortenzio et al.*, 2014).

The main pattern revealed by WA applied to BOUSSOLE data is the periodicity of the seasonal
evolution: from late fall to early spring (essentially the winter) where there is always a convergence
of Chl-Fluo, b_{bp} and c_p signals (Figures 5 and 9).

The spring bloom in the Ligurian Sea is regulated by the increase of light availability after the winter mixing of the water column that redistributes nutrients from deep to surface waters (*Antoine et al.*, 2011; *D'Ortenzio et al.*, 2014). The phasing, duration and intensity of the annual bloom varies from year to year, with a stronger bloom in 2012 in respect to 2011 and 2013 (Figure 2; *Mayot et al.*, 2016). This variability arises from the range of the processes controlling bloom dynamics, including physical forcing such as meteorological extreme events, and/or interaction between different species of phytoplankton organisms (*Winder et al.*, 2010).

456 During summer, the increase of light availability causes the decrease of intra-cellular photosynthetic 457 pigments concentration need and, concurrently, the low nutrient availability limits the 458 phytoplankton population growth and its abundance stays low (*Bellacicco et al.*, 2016). The b_{bp} and 459 c_p are characterized by a similar temporal pattern in summer (Figure 2 and 5). During fall, as light 460 availability decreases and mixed layer deepens, phytoplankton concentration increases again.

The b_{bp} and c_p time-series are coherent with the Chl-Fluo time-series at annual scale because both 461 462 b_{bp} and c_p are sensitive to particle size ranges that include phytoplankton (Stramski et al., 2004; 463 Sosik et al., 2008; Organelli et al., 2018), thus phytoplankton cells abundance strongly impacts the 464 variability of b_{bp} and c_p. Figures 6 and 9 reveal how the bio-optical properties have a WCS and AC 465 close to 1 at annual scale along the entire time-series. In detail, Chl-Fluo is in advance in respect to the annual signal of b_{bp} (Figure 9a), while with c_p they are in phase (Figure 9b). The c_p leads in 466 467 respect to the b_{bp} annual signal as highlighted by the arrows in the Figure 9c. In such a context, the 468 maximum of Chl-Fluo is in shift in advance of approximately 20 days with respect to cp and bbp 469 (Figure 4 and Table 2). This could be attributed to the inter-annual variability (low bloom maxima 470 in 2011 and two maxima in 2013 for Chl-Fluo) which determined a dephasing in the AWP. Another 471 reasonable explanation can be that the Chl-Fluo annual maximum (independently of its strength) is 472 not occurring always at the same time (i.e. indeed there are four maxima in the Chl-Fluo timeseries: March in 2011 and 2012, April and December in 2013; Figure 2). Nevertheless, there is a 473 474 good coherence between the parameters at annual temporal scales (AC \cong 1; Figure 6).

475 **4.2 Six- and four- month cycles**

476 Other fundamental cycles retrieved by WA are at 6 months and 4 months. The 6-month cycle is 477 unexpectedly the most important cycle in cases of b_{bp} and c_p , while for Chl-Fluo it is second in 478 terms of dominance. The 4-month cycle is the third important cycle for all of the parameters, 479 especially for b_{bp} (Table 2). AC shows a strong correlation of the WPS for all combinations of bio-480 optical coefficients at these periods (\cong 1; Figures 6 and 9). The 6-month cycle is dominant for b_{bp} 481 and c_p , whereas it has a lower strength for the Chl-Fluo, a consequence of the inter-annual 482 variability of its WPS (Figure 5a). The 6- and 4-month cycles are interpreted here as mainly due to

483 the winter-to-spring modification of the mixing intensity (D'Ortenzio et al., 2005), nutrient and light availability, grazing and shift in phytoplankton community structures (Mignot et al., 2014; 484 Sammartino et al., 2015). Bellacicco et al., (2016) highlight that the use of Chl as a proxy of 485 phytoplankton biomass, as well as Chl-Fluo, is strictly influenced by intracellular processes, 486 especially in intermediate seasons, such as late spring and early fall. Alternately, b_{bp} and c_p are 487 488 sensitive to the abundance of phytoplankton cells and non-algal particles, as well as particle size 489 distribution, refractive index, and the shape and structure of particles in the seawater. Therefore, in 490 these intermediate periods, where the Chl-Fluo signal is low while cp and bbp are relatively high, the 6-month cycle signal detection could be reduced impacting total seasonal and annual cycles. For 491 492 example, in early fall, there are low nutrients and low light conditions and the result is an increase 493 of phytoplankton cells, as indicated by the increase of c_p and b_{bp}, however phytoplankton does not 494 have a high concentration of photosynthetic pigment, and consequently a low Chl-Fluo. In late 495 spring, the photoacclimation process is particularly relevant and impacts on Chl-Fluo due to high 496 nutrients and high light availability. The b_{bp} and c_p proxies are not affected by the phytoplankton 497 physiological state (Bellacicco et al., 2016; Barbieux et al., 2018) and the effect of the carbon 498 accumulation in phytoplankton cells dominates the cp and bbp WPS (Figures 4 and 5b, c), while the 499 physiological signal (i.e. photoacclimation) has an effect on the strength of the signal in WPS of the Chl-Fluo signal (Figures 4 and 5a). Figure 9 displays how cp signal is in advance with respect to bbp 500 501 at 6 month cycles along 2012 and 2013. On the other hand, the b_{bp} signal is in delay in respect to 502 Chl-Fluo in 2012, while with c_p they are in phase from summer 2011 to spring 2012. Figure 9 also 503 shows the relationship between parameters at the period of 4-months. In this period, Chl-Fluo signal 504 is delayed in respect to c_p as in the case of 6-months, while c_p signal seems to be in advance 505 compared to the b_{bp} signal. From the analysis, the intensity of a 4-month cycle differs in cases of 506 Chl-Fluo in respect to b_{bp} and c_p (Table 2).



507

Figure 9: Cross-wavelet coherence spectra between (a) Chl-Fluo and b_{bp} , (b) Chl-Fluo and c_p , and (c) b_{bp} and c_p at BOUSSOLE site. Color indicates the level of covariability between the time-series (values between 0 and 1); arrows denote the relative phase between the time-series (right: in phase; left: antiphase; up or down: one series leads the other by 90°) with significance covariability (> 95%). Shaded area has not to be considered as it might provide false periodic events (*Torrence and Compo*, 1998). The letters W stands for winter (in blue), S for spring (in green), S for summer (in orange) and F for fall (in brown), as for Figure 5.

515 **4.3 Intra-seasonal variability: mid- and short-term cycles**

516 WA detected relevant cycles at scales other than annual, 6- and 4-months are interpreted here as 517 caused by intra-seasonal (*i.e.* from diel to monthly) variability of the examined bio-optical 518 parameters: the highest cycles are defined as mid-term cycles, while the lowest as short-term 519 counterparts. In such a context, the specific WA applied to these particular periods of winter and summer 2012 enables the detection and power of these additional sources of variability (from 12

hours to 29 days; Tables 3 and 4).

522 At the beginning of winter and over the 16 day period, c_p and Chl-Fluo signals are in phase and then

- 523 c_p is in advance compared to the Chl-Fluo footprint. Contrarily, b_{bp} and Chl-Fluo indicators are in
- 524 phase along the entire season. From 2 day up to 16 day periods, Chl-Fluo is partially in phase with
- b_{bp} and c_{p} , however, some hotspots occur in which both b_{bp} and c_{p} have phase differences in respect
- to the Chl-Fluo signal. In cases of b_{bp} to c_p , at these highest periods, c_p leads the b_{bp} mark (Figure 10).
- 528 In summer, at the range of 8-16 days, Chl-Fluo is in advance in respect to the b_{bp} signal, while on
- 529 the reverse, the Chl-Fluo indicator at this period is delayed with respect to the c_p . Regarding the b_{bp}
- 530 to c_p relationship, the former is in advance with respect to the latter at the beginning of summer,
- while at the end of summer, there is an opposite situation of delay. Between 2 and 8 days, Chl-Fluo
- is not in phase with both c_p and b_{bp} signals (*e.g.* start of summer), while the optical parameters are in phase (Figure 11).
- A possible explanation of these mid-term cycles could also be given by episodic phytoplankton 534 535 biomass increases in response to stochastic events that potentially impact the mixed layer dynamics 536 which are typical of the winter period at the BOUSSOLE site. For instance, as highlighted by Winder et al., (2010), extreme wind events could generate biomass oscillations by temporarily 537 538 reducing the phytoplankton bloom. In addition, Nezlin et al. (2002) showed, using WA, that short-539 period variations (less than 100 days) of remotely-sensed chlorophyll during spring seasons 540 correlated with surface water, air temperatures and wind stress, that have rapidly varied (i.e. at the 541 intermediate scales that we analyze here). The mechanism of these variations was an intensification 542 of phytoplankton growth resulting from mixing of the water column by wind stress and entrainment 543 of cold, rich in nutrients, water into the euphotic layer. Monteiro et al. (2015) demonstrated the 544 presence of an intra-seasonal variability that is always driven by wind stress and has a temporal scale from daily up to 14 - 20 days. These orders of magnitude are consistent with the maxima 545 reported in the AWP plots, for both winter and summer 2012 (Figures 7, 8; Table 3 and 4). 546
- 547 The general Chl-Fluo and c_p WPS also show a clear diel cycle footprint, particularly evident during 548 high biomass periods (winter and spring in Figure 5a and c). On the other hand, the diel signal is 549 less evident for b_{bp} (Figures 4 and 5b and Table 2). The correlation between c_p and Chl-Fluo at diel 550 scale (AC greater than 0.7 in Figure 6) is well known during winter and spring, when the particles
- abundance is dominated by phytoplankton cells (*Kheireddine et al.* 2014; *Bellacicco et al.*, 2016).
 The diel changes are generally considered to be driven by specific forcings (*Oubelkheir et al.*, 2005;
- 553 Oubelkheir and Sciandra 2008; Gernez et al., 2011, Loisel et al., 2011; Barnes et al., 2014;
- *Kheireddine et al.*, 2014). At diel cycle, the correlation between Chl-Fluo and b_{bp} periodicities is lower ($\cong 0.6$) than to Chl-Fluo vs. c_p (>0.7). A good correlation is found in the b_{bp} vs c_p (>0.7).
- 556 By applying the specific WA over a three-month period both in winter and summer seasons, more 557 information can be gathered about short-term variability, especially on the diel cycle.
- 558 During the winter, Chl-Fluo signal is in advance to c_p without any phases between parameters 559 despite AC showing high values (Figure 10). Reversely, the AC between Chl-Fluo and b_{bp} is lower 560 (less than 0.5) below diel cycle. The AC between b_{bp} and c_p is always lower than 0.6 (Figure 10).
- 561 In the winter, the period of strong mixing and start of the bloom (Barnes et al., 2014, Kheireddine
- 562 *et al.* 2014; *Bellacicco et al.*, 2016), c_p can be increasingly influenced by diel variations of
- abundance of phytoplankton cells (*Oubelkheir et al.*, 2005, *Oubelkheir and Sciandra*, 2008), even if
- with a marginal impact on total variability (Table 3). In contrast to c_p , b_{bp} diel cycle is not marked

by a significant seasonal variability. This would confirm that phytoplankton makes a lower contribution to b_{bp} than to c_{p} , so their seasonal changes are poorly reflected in overall seasonal changes. b_{bp} is influenced more by the presence of sub-micrometer particles such as detrital particles or heterotrophic bacteria (*Morel and Ahn*, 1991; *Stramski and Kiefer*, 1991, *Stramski et al.*, 2004), that do not have a periodical diel cycle. Recently, *Organelli et al.*, (2018) found that another main source of variability on b_{bp} is due to particles with equivalent diameters between 1 and 10 µm giving thus new insight into the b_{bp} coefficient and particles in seawater.

572 In summer, Chl-Fluo and b_{bp} have an AC of 0.8 but b_{bp} is largely in anti-phase with Chl-Fluo, as expected for this period of year of low productivity. Inversely, cp and Chl-Fluo diurnal signals have 573 an AC value of approximately 1. During this season of absence of high abundance of phytoplankton 574 575 (Kheireddine et al. 2014; Bellacicco et al., 2016), cp can be due to the daily variations of coupled 576 heterotrophic bacteria and particles pool (i.e. phytoplankton cells) which remain within the upper layer caused by the strong stratification of the water column. However, Chl-Fluo diel signal can be 577 578 determined due to the quenching effect (Xing et al., 2017). In addition, at periods of 0.5 (i.e. 12 579 hours), cp and bbp signals are in anti-phase with Chl-Fluo. Due to this limited productivity, cycles greater than 1 day are not the most impactful of the intra-seasonal variability, while conversely the 580 581 diurnal signal is well pronounced. Oubelkheir and Sciandra (2008) argued that the diel cycle of c_p, and the particles pool, is strictly influenced by changes in the properties of particles (e.g. size, 582 583 refractive index, shape and internal structure) and also by external environmental and biological agents, as reported by Binder and Durand (2002). During daytime, phytoplankton cells fix external 584 585 inorganic carbon into organic molecules determining an increase of their diameter and refractive index (Siegel et al., 1989; Stramski and Revnolds, 1993; Walsh et al., 1995), and as a consequence 586 587 an increase of scattering and attenuation cross section. This was also confirmed by laboratory experiments on a few phytoplankton species (Stramski et al., 1995; Durand and Olson, 1998; 588 Claustre et al. 2002; Poulin et al., 2018). During night-time, phytoplankton cells divide into 589 590 smaller cells with lower intra-cellular carbon content, as a result of an uptake of water during 591 division or of a loss due to respiration (Durand and Olson, 1998). Another element to consider on 592 the c_p diel variation is the relative dynamics of algal and non-algal stocks (*i.e.* heterotrophs, viruses 593 and detritus) that could vary in relation to the season and trophic regimes. Oubelkheir and Sciandra 594 (2008) showed that the c_p is strictly related to heterotrophic particles abundance coupled with 595 phytoplankton cells and associated detritus.

To summarize, through the specific three-month WA, more information about detection and power of these additional sources of variability can be obtained. The mid-term cycles (> 10 days) dominate intra-seasonal variability during the winter (period of mixing and bloom), while they are of limited impact in the case of summer which is the period of low biological production and absence of extreme meteorological events (*Gernez et al.*, 2011, *Kheireddine et al.*, 2014) indicating where the diel cycle is the most recurrent and important source of signal.



602

603 Figure 10: Cross-wavelet coherence spectra between Chl-Fluo and $b_{bp}(a)$, Chl-Fluo and $c_p(b)$, and b_{bp} and c_p 604 (c) at BOUSSOLE site for the Winter 2012. Color indicates the level of covariability between the time-series (values between 0 and 1); arrows denote the relative phase between the time-series (right: in phase; left: 605 606 antiphase; up or down: one series leads the other by 90°) and significance covariability (> 95% and thin 607 white contours). Shaded area has not to be considered as it might provide false periodic events (Torrence and 608 Compo, 1998). AC of each parameter are in Panel d-f. In red are highlighted periods where the AC is with 609 low statistical significance (less than 95%). The significance test is computed following Roesch and 610 Schmidbauer (2014).





612 Figure 11: Cross-wavelet coherence spectra between Chl-Fluo and $b_{bn}(a)$, Chl-Fluo and $c_n(b)$, and b_{bn} and c_n 613 (c) at BOUSSOLE site for the Summer 2012. Color indicates the level of covariability between the time-614 series (values between 0 and 1); arrows denote the relative phase between the time-series (right: in phase; 615 left: antiphase; up or down: one series leads the other by 90°) and significance covariability (> 95% and the 616 thin white contours). Shaded area has not to be considered as it might provide false periodic events 617 (Torrence and Compo, 1998). AC for each parameter are in Panel d-f. In red are highlighted periods where the AC is with low statistical significance (less than 95%). The significance test is computed following 618 619 Roesch and Schmidbauer (2014).

620 5. Conclusions

621 In the last decades, the development of fixed observation sites, such as the BOUSSOLE buoy, as 622 well as remote sensing advancement has helped to study the phenology of phytoplankton and 623 optical properties. Several works have studied temporal variability of bio-optical properties in 624 various oceanic regimes using both field and satellite data (Behrenfeld et al., 2009, Antoine et al., 625 2011; Gernez et al., 2011; Barnes et al., 2014; Kheireddine et al., 2014; Behrenfeld et al. 2015; 626 Sammartino et al., 2015; Di Cicco et al., 2017). However, most of these phenological studies focus 627 on the annual and seasonal cycles (*i.e.* long-term cycles), while there is limited literature on the mid- and short-term cycles, and their recurrence, which characterizes the temporal variability of 628 629 bio-optical properties.

- 630 In this study, we have focused on bio-optical properties using, for the first time, a statistical *a priori*
- 631 method, as the wavelet analysis (WA) is, on three-year high frequency observations. The main
- 632 goals are to determine the intra-annual dominant temporal patterns of the bio-optical parameters,
- 633 the changes of these cycles over time, the characteristics and recurring strength at those periods and
- to define the temporal relationship between the cycles of the bio-optical properties.
- 635 The WA applied here reveals a persistent annual cycle for Chl-Fluo which explains the largest
- amount of its variability. On the other hand, and unexpectedly, the 6-month cycle is the most

- 637 important and dominant temporal pattern of the b_{bp} and c_p time-series, with respect to Chl-Fluo, which accounts for half of the variability in respect to the annual cycle. Together with 6-month, the 638 4-month cycle is the third source of variability for all the parameters with different strengths. It has 639 a particular recurrence and magnitude in the case of b_{bp}, with respect to Chl-Fluo and c_p and intra-640 641 seasonal variability is driven by mid- and short-terms cycles. During the winter (season of mixing), 642 the mid-term cycles (> 10 days) are the most important. Episodic bloom events can determine these 643 cycles, as viewed in the case of bio-optical coefficients, and as also found by Winder et al. (2010). 644 During summer, the diel cycle is the most important and the main source of variability, especially for c_p and Chl-Fluo. Considering the entire time-series, at diel scale, the coherence between spectra 645 of bio-optical coefficients diminishes in respect to annual and seasonal cycles although remaining 646 647 high. Chl-Fluo - c_p and $b_{bp} - c_p$ periodicities have a strong temporal correlation in respect to Chl-Fluo- b_{bp}. At diurnal scale, Chl-Fluo depends specifically on intra-cellular and physiological 648 processes in relation to physical forcing, and on phytoplankton cells abundance. The cp at the same 649 scale has its own temporal pattern: in summer, c_p is driven mostly by heterotrophic bacteria coupled 650 651 with low phytoplankton abundances, while in winter, c_p is more associated with phytoplankton 652 particles (Oubelkheir et al. 2005, Oubelkheir and Sciandra, 2008). In the case of b_{bp}, the correlation between the periodicities of Chl-Fluo is lower with respect to c_p because the b_{bp} is influenced more 653 654 by small particles (Stramski et al., 2004; Kheireddine et al., 2014). In such a context, the recent 655 findings of Organelli et al. (2018) open challenges in understanding the complexity of marine 656 particles structure as sources of variability of the open-ocean b_{bp} signal that have to be addressed in 657 the next future in order to better constraint the use of b_{bp} observations for investigating the 658 biological carbon pump and phytoplankton phenology studies.
- This work thus highlights the need to develop in situ technologies as well as new satellite sensors at 659 660 higher temporal resolutions (e.g. geostationary satellite) for biogeochemical/bio-optical 661 measurements that have been widely recognized as a priority in the optical and oceanographic community. Indeed, high frequency observations could help to study, both in space and time, these 662 663 mid- and low-term cycles, poorly known, that currently are not taken into account in the ocean 664 color algorithms despite dominating the bio-optical variability at reduced time-series length (i.e. within the season). Lastly, as it is demonstrated here, the importance of the use of WA as a 665 powerful instrument for studying both long or short time-series of bio-optical parameters and their 666 relationships in oceanography (Winder et al., 2010, Damarcq et al. 2012, Ampe et al., 2014, 667 668 Corredor-Acosta et al., 2015, Carey et al. 2016, Sala et al., 2018).
- 669

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