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# Modelling the Frequency-Dependent Effective Excess Charge Density in Partially Saturated Porous Media

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#### **Key Points:**

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# We present a novel flux-averaging approach to compute the dynamic effective excess charge density in partially saturated porous media. The model accounts for the pore size distribution of the medium and permits to estimate the dynamic electrokinetic coupling coefficient. The proposed approach has an excellent capability for reproducing previous mod-

els and experimental measurements in the literature.

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#### 15 Abstract

In the context of seismoelectric and self-potential surveying, the effective excess charge 16 density and the electrokinetic coupling coefficient are key parameters relating the mea-17 sured electrical potential and the hydraulic characteristics of the explored porous me-18 dia. In this work, we present a novel flux averaging approach that permits to estimate 19 the frequency-dependent effective excess charge density in partially saturated porous me-20 dia. For this, we conceptualize the porous medium as a partially saturated bundle of cap-21 illary tubes under oscillatory flux conditions. We account for the pore size distribution 22 (PSD) to determine the capillary-pressure saturation relationship of the corresponding 23 medium, which, in turn, permits to determine the pore scale saturation. We then solve 24 the Navier-Stokes equations within the saturated capillaries and, by means of a flux-averaging 25 procedure, obtain upscaled expressions for: (i) the effective excess charge density, (ii) the 26 effective permeability, and (iii) the electrokinetic coupling coefficient, which are functions 27 of the saturation and the probing frequency. We analyze and explain the characteristics 28 of these functions for three different PSDs: fractal, lognormal, and double lognormal. It 29 is shown that the PSD characteristics have a strong effect on the corresponding electroki-30 netic response. The proposed flux-averaging approach has an excellent capability for re-31 producing experimental measurements and models in the literature, which are otherwise 32 based on well-known empirical relationships. The results of this work constitute a use-33 ful framework for the interpretation of electrokinetic signals in partially saturated me-34 dia. 35

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#### Plain Language Summary

Seismic waves travel throughout the Earth deforming the rocks in their passage. 37 If rocks are porous, permeable, and contain fluids in their pores, as is the case in many 38 geological formations, the wave's passage may induce oscillatory fluid flow. Minerals com-39 posing rocks are commonly electrically charged and, thus, the flowing fluid can result in 40 an electrical field. Interestingly, measuring this electrical field at the Earth's surface may 41 permit to characterize the hydromechanical properties of geological formations of inter-42 est, motivating the so-called seismoelectric method. The effective excess charge that is 43 mobilized by the fluid motion depends on the frequency content of the wave, and mod-44 els exist to estimate this dependence in terms of the rock and fluid properties. However, 45 in many scenarios in Earth sciences, rocks contain two immiscible fluid phases, such as, 46

water and air, for which frequency-dependent effective excess charge density models based
on pore-scale physics are missing in the literature. In this paper, we derive such a model
and show that it is able to reproduce previous estimates and experimental data.

#### 50 1 Introduction

The remote characterization of partially saturated geological formations using non-51 invasive techniques remains, to date, a challenging task within the field of applied and 52 environmental geophysics. Given its inherent sensitivity to flow dynamics and pore fluid 53 characteristics, the seismoelectric method can provide highly valuable information for 54 studying this type of environments (Grobbe et al., 2020; Revil et al., 2015). The phys-55 ical principles upon which seismoelectric prospecting is based on have been used in con-56 text of groundwater management and remediation (e.g., Dupuis et al., 2007; Han et al., 57 2004; Monachesi et al., 2018), exploration and production of hydrocarbons (e.g., Revil 58 & Jardani, 2010), and  $CO_2$  geosequestration operations (e.g., Zyserman et al., 2015). Novel 59 approaches addressing the complex processes behind the seismic-to-electric conversion 60 are of great interest, as they may help to better interpret seismoelectrical signatures in 61 partially saturated environments. 62

The seismic-to-electric conversion occurs when a seismic wave propagates through 63 a fluid saturated and charged porous medium, generating fluid displacements relative to 64 the pore walls (e.g., Pride, 1994). Given that, in general, the surfaces of wet minerals 65 composing porous rocks are electrically charged, an electrical double layer (EDL) arises 66 within the saturating pore fluid which counterbalances the net charge present in the min-67 erals. The EDL contains an excess of charge that is distributed in two layers: (i) the Stern 68 layer, where charges are virtually immobile, and (ii) the diffuse layer, where charges have 69 the capacity to move freely (e.g., Revil & Mahardika, 2013). Whenever a passing seis-70 mic wavefield induces flow, the excess charge located in the diffuse layer is dragged into 71 motion, generating a streaming current which, in turn, results in an electrical potential 72 distribution. The associated electrical field, which can be surveyed remotely, either at 73 the Earth's surface or at boreholes, contains valuable information regarding the hydrome-74 chanical properties of the probed geological formation. Laboratory and borehole mea-75 surements evidence that seismoelectric signals are sensitive to, for example, the poros-76 ity and permeability of porous media (e.g., Zhu et al., 2008; Wang et al., 2015), and to 77 salt concentration and dielectric permittivity of the saturating fluid (e.g., Zhu & Tok-78

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soz, 2013; Garambois & Dietrich, 2001). Seismoelectric signals measured in surface surveying or borehole logging have been used, for example, to explore earthquake rupture
characteristics (e.g., Gao et al., 2016), to identify formation boundaries associated with
lithological changes (e.g., Butler, 1996; Garambois & Dietrich, 2001), and to detect saturation changes in permeable geological formations (e.g., Thompson & Gist, 1993).

The seismoelectric conversion is traditionally modeled using of the electrokinetic 84 coupling coefficient  $C_{EK}(\omega)$ , which is a frequency-dependent parameter relating the elec-85 trical potential difference (i.e., the electrical field) and the pore fluid pressure gradient 86 driving the fluid flow. In this context, the most frequently used models to estimate  $C_{EK}(\omega)$ 87 are based on the works of: (i) Pride (1994) and (ii) Packard (1953). On the one hand, 88 Pride's (1994) model is based on volume averaging principles and on the dynamic per-89 meability model proposed by Johnson et al. (1987). On the other hand, the pioneering 90 model of Packard (1953) considers a capillary tube of a unique radius and computes the 91 streaming potential difference associated with an oscillatory flux. This model has been 92 widely applied to porous media with a certain success (e.g., Reppert, 2001). Recently, 93 Thanh et al. (2021) extended the work of Packard (1953) to take into account different 94 pore size distributions (PSD), thus showing the effects of the porous structure on  $C_{EK}(\omega)$ . 95 An alternative approach for studying the seismoelectric conversion is to compute the ex-96 cess charges that are effectively dragged in the diffuse layer, that is, the effective excess 97 charge density  $\hat{Q}_{\nu}$ , which can be subsequently used to estimate  $C_{EK}$  (e.g., Jackson, 2010; 98 Jougnot et al., 2012; Revil & Mahardika, 2013). In the literature, many studies were per-99 formed considering this effective excess charge density but neglecting its frequency-dependence, 100 that is, considering its low-frequency limit (e.g., Jougnot et al., 2013; Rosas-Carbajal et 101 al., 2020). Recently, Jougnot and Solazzi (2021) extended the definition of  $\hat{Q}_{\nu}$  to the en-102 tire frequency range  $Q_{\nu}(\omega)$ , thus allowing to compute  $C_{EK}(\omega)$ . For this, the authors in-103 tegrated the charges that are effectively dragged along individual pores across the probed 104 medium, accounting for inertial effects associated with the oscillatory pressure forcing 105 generated by a passing seismic wavefield. We remark that the latter work reconciled both 106 Pride (1994) and Packard (1953) approaches by integrating flux averaging principles and 107 the dynamic permeability concept. All of the above described works deal with the frequency-108 dependence of the coupling coefficient  $C_{EK}(\omega)$  and/or the effective excess charge  $Q_{\nu}(\omega)$ 109 under fully saturated conditions and, thus, modifications are needed if one wishes to em-110 ploy the corresponding approaches in partially saturated porous media. 111

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Evidence indicates conclusively that water content variations in porous media have 112 preeminent effects on the associated seismoelectric signatures (e.g., Bordes et al., 2015; 113 Zyserman et al., 2017). When exploring partially saturated media using the seismoelec-114 tric method, one can also use either the coupling coefficient  $C_{EK}$  and/or the effective 115 excess charge density  $\hat{Q}_{\nu}$  to study the electro-kinetic process. Warden et al. (2013) ex-116 tended the electrokinetic coupling coefficient  $C_{EK}(S_w)$  definition to address partially sat-117 urated conditions, highlighting the key influence of water content on the seismoelectric 118 conversion. The coupling coefficient in partially saturated conditions is generally obtained 119 by scaling its fully saturated counterpart by the wetting phase saturation (Bordes et al., 120 2015; Revil & Mahardika, 2013; Warden et al., 2013; Zyserman et al., 2017). Later on, 121 Revil and Mahardika (2013) proposed a simple model to compute the saturation- and 122 frequency-dependent effective excess charge density of partially saturated porous media 123  $\hat{Q}_{\nu}(\omega, S_w)$  and through it, to estimate  $C_{EK}(\omega, S_w)$ . For this, Revil and Mahardika (2013) 124 rely on concept of dynamic permeability, using a Debye approximation, and on empiric 125 and broadly used scaling laws, thus extending the approach proposed by Pride (1994) 126 to partially saturated media. As far as we know, to date, a model deriving the saturation-127 and frequency-dependent effective excess charge density  $\hat{Q}_{\nu}(\omega, S_w)$  from first principles, 128 that is, from flux averaging the pore scale physics, is lacking in the specific literature. 129 Such derivation is of fundamental importance, as it would permit to: (i) couple flux, elec-130 trokinetic properties, and the pore size distribution characteristics of porous media; (ii) 131 validate the approach proposed by Revil and Mahardika (2013). 132

In this work, we propose a novel flux averaging approach to estimate the effective 133 excess charge density as a function of saturation and frequency  $\hat{Q}_{\nu}(\omega, S_w)$ . The paper 134 is structured as follows. First, we resume the theory behind the frequency-dependence 135 of the effective excess charge density  $Q_{\nu}(\omega)$  and other parameters, such as, the dynamic 136 permeability  $\kappa(\omega)$  and the electrokinetic coupling coefficient  $C_{\rm EK}(\omega)$ . Then, we propose 137 a model to account for different saturation states in the latter. We evaluate the satu-138 ration and frequency response of the medium considering fractal, lognormal, and dou-139 ble lognormal PSDs. Finally, we compare the proposed approach with the model pro-140 posed by Revil and Mahardika (2013) and with published experimental data. 141

#### 142 2 Theory

In this section, we resume the theory of dynamic (frequency-dependent) permeability and effective excess charge density in fully saturated media. Then, we extend these definitions to the partially saturated state, considering that the pore fluids are immiscible and that their distribution throughout the pore space is determined by capillary forces. In the case of the dynamic permeability, we follow the work of Solazzi et al. (2020), who derived frequency- and saturation-dependent effective permeability estimates in partially saturated porous media.

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## 2.1 Frequency-Dependent Effective Excess Charge Density in Fully Saturated Media

#### 2.1.1 Fluid Flow and Dynamic Permeability

Let us consider a cylindrical representative elementary volume (REV) of a porous 153 material of length L (m) and radius  $R_{\rm REV}$  (m). We conceptualize the fluid flow of a sin-154 gle phase across the REV using a bundle of aligned capillary tubes, oriented along the 155 axis of the cylindrical REV, comprising radii R (m) whose sizes vary from  $R_{\min}$  to  $R_{\max}$ . 156 The pore-size distribution (PSD) is such that the number of capillaries with radii between 157 R and R+dR is given by f(R)dR. Note that this conceptualization of a porous medium 158 under fluid flow is based on similar concepts as the classic model of Kozeny (1927), which 159 is broadly used in permeable soils (e.g., Mavko et al., 2009). Let us also consider that 160 an incompressible Newtonian fluid characterized by a shear viscosity  $\eta$  (Pa.s) and den-161 sity  $\rho$  (kg/m<sup>3</sup>) saturates the porous medium, whose solid matrix is assumed to be rigid 162 (Johnson et al., 1987). Note that the fluid incompressibility assumption is valid at the 163 pore scale as long as the wavelengths of possible acoustic waves traveling in the fluid are 164 much larger than the characteristic pore size (Johnson et al., 1987; Zhou & Sheng, 1989). 165 Finally, we consider that the fluid flow within the pore space is of laminar-type associ-166 ated with a small Reynold's number (Auriault et al., 1985; Smeulders et al., 1992). 167

The REV structure is then subjected to an oscillatory pore fluid pressure difference  $\Delta \hat{p} = \Delta p e^{-i\omega t}$  (Pa) along its axis., with  $\omega$  denoting the angular frequency (rad/s). Solving the incompressible Navier-Stokes equations under the assumptions mentioned above, the fluid velocity  $v^{f}$  (m/s) in a capillary of internal radius  $0 \le r \le R$  responds to (Solazzi et al., 2020)

 $v^f(r,\omega) = -\frac{1}{\tau\eta k^2} \left[ \frac{J_0(kr)}{J_0(kR)} - 1 \right] \frac{\Delta p}{L},\tag{1}$ 

where  $k^2 = i\omega\rho/\eta$  and  $J_{\nu}$  are Bessel functions of the first kind of order  $\nu$ . The tortuosity is given by  $\tau = l^*/L$ , where  $l^*$  is the actual flow path length. Note that we have dropped the harmonic term  $e^{-i\omega t}$  for ease of notation. Integrating equation (1) over the cross-sectional area of the pore, the corresponding volumetric flow rate (m<sup>3</sup>/s) through a single capillary is given by (e.g., Johnson et al., 1987)

$$q(R,\omega) = -\frac{\pi R^2}{\tau \eta k^2} \left[ \frac{2}{kR} \frac{J_1(kR)}{J_0(kR)} - 1 \right] \frac{\Delta p}{L}.$$
 (2)

The volumetric flow rate  $Q_{\text{flow}}^{\text{sat}}$  (m<sup>3</sup>/s) at the fully-saturated REV-scale can be obtained by integrating equation (2) over the entire range of pore sizes within the REV

$$Q_{\text{flow}}^{\text{sat}} = \int_{R_{\min}}^{R_{\max}} q(R,\omega) \,\mathfrak{f}(R) \,\mathrm{d}R. \tag{3}$$

The effective Darcy velocity at the REV scale v<sup>sat</sup> (m/s) is obtained by scaling the volumetric flow rate by the corresponding cross-sectional area, that is, v<sup>sat</sup> =  $Q_{\text{flow}}^{\text{sat}}/\pi R_{\text{REV}}^2$ .

If one increases the frequency of the oscillatory pressure forcing, a transition from viscous- to inertia-dominated flow occurs. For a given critical angular frequency  $\omega_c$ , the viscous skin depth  $\delta = \sqrt{2\eta/\rho\omega}$  (m) becomes comparable to the radii of the largest saturated pores (Johnson et al., 1987)

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$$\omega_{\rm c} \simeq \frac{2\eta}{\breve{R}^2 \rho},\tag{4}$$

with  $\tilde{R}$  being the characteristic radius of the saturated porous medium. For frequencies higher than  $\omega_c$  the fluid motion becomes viscously decoupled. In this context, the fluid flow and the underlying fluid pressure forcing become out of phase and the fluid flow amplitude decreases.

The dynamic (frequency-dependent) permeability  $\kappa(\omega)$  (m<sup>2</sup>) is then computed using Darcy's law, that is, relating the fluid flow and the pressure gradient along the REV (Solazzi et al., 2020)

$$\kappa(\omega) = \frac{1}{\tau R_{\rm REV}^2 k^2} \int_{R_{\rm min}}^{R_{\rm max}} \left[ \frac{2}{kR} \frac{J_1(kR)}{J_0(kR)} - 1 \right] R^2 \mathfrak{f}(R) \, \mathrm{d}R.$$
(5)

Equation (5) can be solved numerically provided that f(R),  $R_{\min}$ , and  $R_{\max}$  are known. One of the consequences of equation (5) is that the pore size distribution has an impact on the dynamic permeability characteristics, such as, the value of  $\omega_c$  (Solazzi et al., 2020). We remark that, following a different approach than the one proposed by Solazzi et al.

202 (2020), Li et al. (2021) arrived to the very same conclusion.

The low-frequency limit of equation (5) is the (Poiseuille-type) permeability of the medium (Blunt, 2017)

$$\kappa^0 = \frac{1}{\tau 8 R_{\text{REV}}^2} \int_{R_{\text{min}}}^{R_{\text{max}}} R^4 \mathfrak{f}(R) \, \mathrm{d}R.$$
(6)

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#### 2.1.2 Effective Excess Charge Density

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Let us now consider that the capillaries of the previously described porous medium 207 are saturated by a binary symmetric electrolyte (e.g., NaCl) with ionic concentration  $C_i^w$ 208 (mol m<sup>-3</sup>) and valence  $z_i = \pm 1$ , with *i* denoting the considered ion. Minerals compos-209 ing the pore walls of rocks normally exhibit surface charges when in contact with wa-210 ter. As an example, silicate and aluminosilicate minerals present negative charges un-211 der natural conditions. Let us then denote co-ions the ions that present the same charge 212 as the minerals constituting the pore walls (e.g.,  $Cl^{-}$ ) and, counter-ions those charged 213 with an opposite valence (e.g., Na<sup>+</sup>). For the system to be electrically neutral, surface 214 charges are balanced by an excess charge in the pore water. The latter are distributed 215 in the EDL. Within the EDL, the diffuse layer comprises co-ions and counter-ions that 216 are able to move and, also, is characterized by a net excess of charge. Hereafter we con-217 sider that the shear plane, that is, the plane that separates the stationary fluid and the 218 moving fluid, corresponds to the interface between the Stern layer and the diffuse layer. 219 The electrical potential along this plane is referred to as Zeta potential. 220

The distribution of the excess charges in the diffuse layer within a single capillary is governed by the Poisson-Boltzmann equation

$$\nabla^{2}\varphi\left(r\right) = -\frac{Q_{\nu}\left(r\right)}{\varepsilon_{r}\varepsilon_{0}},\tag{7}$$

where  $\varphi(r)$  (V) is the electric potential and  $\overline{Q}_{\nu}(r)$  (C m<sup>-3</sup>) is the excess charge density in the liquid at a distance  $0 \le r \le R$  from the pore-centre. The relative permittivity of the fluid and the dielectric permittivity of vacuum are given by  $\varepsilon_r$  and  $\varepsilon_0 = 8.854 \times$  $10^{-12}$  F m<sup>-1</sup>, respectively. Under the above conditions, the effective charge density responds to (e.g., Jougnot et al., 2012)

$$\overline{Q}_{\nu}(r) = N_A e_0 C_{NaCl}^w \left[ e^{\left(-\frac{e_0\varphi(r)}{k_BT}\right)} - e^{\left(\frac{e_0\varphi(r)}{k_BT}\right)} \right].$$
(8)

Generally, equation (7) is solved assuming: (i) a Debye-Hückel linear approximation, that is,  $e_0\varphi(r)/k_BT <<1$ ; (ii) that the pore size is considerably larger than thickness of the double layer. In this context, the two exponential terms in Eq. (8) can be expressed through the sinh function and, then, one can make use of the fact that for sufficiently small arguments the sinh function tends to its corresponding argument, that is,  $\sinh[e_0\varphi(r)/k_BT] \simeq$  $e_0\varphi(r)/k_BT$ . Consequently, the electric potential is given by

$$\varphi(r) = \zeta e^{\frac{r}{l_D}},$$

where  $l_D$  is the Debye length characterizing the electrical double layer thickness given

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by

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$$l_D = \sqrt{\frac{\varepsilon_0 \varepsilon_r k_B T}{2N_A e_0^2 C_{NaCl}^w}}.$$
(10)

(9)

The dependence of the  $\zeta$  potential on the ionic concentration is hereby estimated following (Pride & Morgan, 1991)

$$\zeta(C_{NaCl}^w) = a + b \log_{10}(C_{NaCl}^w). \tag{11}$$

The fitting parameters a and b are taken as a = -6.43 mV and b = 20.85 mV, as estimated by (Jaafar et al., 2009) for NaCl brine and silicate-based materials.

In this context, the effective excess charge density  $\hat{Q}^R_{\nu}$  carried by the water flow in a single capillary of radius *R* responds to (Jougnot & Solazzi, 2021)

$$\hat{Q}_{\nu}^{R}(\omega) = \frac{\int_{0}^{R} \overline{Q}_{\nu}(r) v^{f}(r,\omega) r \mathrm{d}r}{\int_{0}^{R} v^{f}(r,\omega) r \mathrm{d}r}.$$
(12)

The effective excess charge density  $\hat{Q}_{\nu}^{R}$  is different from the simple excess charge density  $\bar{Q}_{\nu}$ , since  $\hat{Q}_{\nu}^{R}$  is the excess charge that is effectively dragged by the water flow, which is smaller than the total amount of excess charge present in the diffuse layer  $(\bar{Q}_{\nu}: \bar{Q}_{\nu} >> \hat{Q}_{\nu}^{R})$ . For further details on this particular topic, we refer the readers to the discussion sections of Jougnot et al. (2019, 2020).

The effective excess charge carried by the water flow in the fully saturated REV can be obtained by integrating  $\hat{Q}^R_{\nu}(\omega)$ , weighted by the corresponding fluxes, over the entire range of pore sizes

$$\hat{Q}_{\nu}^{\text{sat,REV}}(\omega) = \frac{\int_{R_{\min}}^{R_{\max}} \hat{Q}_{\nu}^{R}(\omega) q(R,\omega) \mathfrak{f}(R) \, \mathrm{d}R}{\int_{R_{\min}}^{R_{\max}} q(R,\omega) \mathfrak{f}(R) \, \mathrm{d}R},\tag{13}$$

where  $q(R,\omega)$  is the volumetric flow rate through a single capillary of radius R given by equation (3). We remark that the supra-index "sat" denotes that the medium is fully <sup>259</sup> saturated and helps to discriminate this parameter from its partially saturated counter-

<sup>260</sup> part, defined in the next subsection of this paper.

Finally, based on the above described expressions, it is possible to define a relative excess charge density (Jougnot & Solazzi, 2021),

$$\hat{Q}_{\nu}^{\text{sat,rel}}(\omega) = \frac{\hat{Q}_{\nu}^{\text{sat,REV}}(\omega)}{\hat{Q}_{\nu}^{\text{sat,0}}},\tag{14}$$

where  $\hat{Q}_{\nu}^{\text{sat, 0}} = \lim_{\omega \to 0} \hat{Q}_{\nu}^{\text{sat, REV}}(\omega)$  is the steady-state (low frequency) excess charge density of the fully saturated medium.

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#### 2.1.3 Electrokinetic Coupling Coefficient

- At the REV scale, the electrokinetic coupling is usually quantified by means of the electrokinetic coupling coefficient (e.g., Jaafar et al., 2009)
  - $C_{EK}(\omega) = \left(\frac{\partial\varphi}{\partial p}\right)_{\mathbf{J}=\mathbf{0},\,\mathbf{\ddot{u}}_s=0} = \frac{\Delta V}{\Delta p},\tag{15}$

which is the ratio of the electrical potential difference  $\Delta V$  and the pressure difference  $\Delta p$  measured at the boundaries of a probed rock sample in the absence of total current densities  $\mathbf{J} = \mathbf{0}$  and solid frame accelerations  $\ddot{\mathbf{u}}_s = 0$ . Through a simple variable change, the frequency dependent coupling coefficient for a fully saturated medium can be expressed as (e.g., Jougnot et al., 2020; Jougnot & Solazzi, 2021; Revil & Mahardika, 2013)

$$C_{EK}^{\text{sat}}(\omega) = -\frac{Q_{\nu}^{\text{sat},\text{REV}}(\omega)\kappa(\omega)}{\eta_w \sigma^{\text{sat}}(\omega)},\tag{16}$$

where  $\kappa_w(\omega)$  and  $\hat{Q}_{\nu}^{\text{sat,REV}}(\omega)$  respond to equations (5) and (13), respectively. We re-276 mark here that the electrical conductivity  $\sigma^{\rm sat}(\omega)$  may, as well, present a frequency de-277 pendence. For a detailed derivation of equation (16), we refer the reader to, for exam-278 ple, the work of Revil and Mahardika (2013) (specifically to equations 34 to 38). Note 279 that, for steady-state conditions (low-frequency limit), Jougnot et al. (2019) showed that 280 this equation is valid for any kind of pore space geometry (pore shape and connectiv-281 ity) and that the geometrical information is carried through the permeability. In this sense, 282 as long as the thin double layer assumption is respected, permeability effects are can-283 celled in the coupling coefficient as the effective excess charge density depends on the 284 inverse of the permeability (see discussion in Jougnot et al. (2019, 2020)). However, such 285 simplification does not hold for the whole frequency range (e.g., Jougnot & Solazzi, 2021), 286 as the relationship between permeability and effective excess charge density is more com-287 plex when considering frequency-dependent effects (see equation 13). 288

The relative electrokinetic coupling coefficient can be expressed as (Jougnot & Solazzi, 2021)

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$$C_{EK}^{\text{sat, rel}}(\omega) = \frac{C_{EK}^{\text{sat, }(\omega)}}{C_{EK}^{\text{sat, }0}},\tag{17}$$

where  $C_{EK}^{\text{sat, 0}} = \lim_{\omega \to 0} C_{EK}^{\text{sat}}(\omega)$  is the steady-state electrokinetic coupling coefficient of the fully saturated medium.

# 2.2 Frequency-Dependent Effective Excess Charge Density in Partially Saturated Media

#### 2.2.1 Fluid Flow and Effective Dynamic Permeability

In the context of fluid flow in partially saturated porous media, the wetting phase flows through a fraction of the corresponding medium. Thus, Darcy's equation in partiallysaturated media is (e.g., Bear, 1972)

$$\mathbf{v}_w = -\frac{\kappa^{\text{eff}}}{\eta_w} \nabla p_w. \tag{18}$$

In equation (18),  $\mathbf{v}_w$  is Darcy's velocity of the wetting phase,  $\kappa^{\text{eff}}$  is the effective permeability of the wetting phase which responds to

$$\kappa^{\text{eff}}(\omega) = \kappa(\omega)\kappa^{\text{rel}}_{w}(p_{c},\omega), \qquad (19)$$

with  $\kappa_w^{\text{rel}}(p_c, \omega)$  denoting the frequency dependent relative permeability of the wetting phase, and  $p_c$  (Pa) the capillary pressure.

The Young-Laplace equation permits us to obtain the capillary pressure in partially saturated capillary of radius  $R_{\rm p}$  (e.g., Bear, 1972)

$$p_{\rm c} = \frac{2\gamma\cos(\beta)}{R_{\rm p}},\tag{20}$$

where  $\gamma$  (N/m) is the interfacial tension and  $\beta$  (rad) is the contact angle between the 309 solid walls and the saturating immiscible fluid phases. At the REV scale,  $p_c$  normally 310 presents a functional relationship with the saturation of the medium (e.g., Van Genuchten, 311 1980; Brooks & Corey, 1964). If the medium is at capillary pressure equilibrium, all cap-312 illaries presenting radii  $R > R_p(p_c) = \frac{2\gamma \cos(\beta)}{p_c}$  are to be saturated by the non-wetting 313 phase (e.g., Mualem, 1976) and those satisfying  $R \leq R_p(p_c)$  are to be saturated by the 314 wetting phase. It is then straightforward to compute the associated effective wetting phase 315 saturation  $S_{ew}(p_c)$ , which yields (e.g., Blunt, 2017) 316

$$S_{\rm we}(p_{\rm c}) = \frac{\int_{R_{\rm min}}^{R_{\rm p}(p_{\rm c})} R^2 \mathfrak{f}(R) \,\mathrm{d}R}{\int_{R_{\rm min}}^{R_{\rm max}} R^2 \mathfrak{f}(R) \,\mathrm{d}R}, \quad \text{with} \quad p_{\rm c,min} \le p_{\rm c} \le p_{\rm c,max}, \tag{21}$$

with  $p_{c,max} = 2\gamma \cos\beta/R_{min}$  and  $p_{c,min} = 2\gamma \cos\beta/R_{max}$ . When capillary pressures are such that  $p_c < p_{c,min}$  we have  $S_{we} = 1$  and, alternatively, when  $p_c > p_{c,max}$  we have  $S_{we} = 0$ . We remark that equation (21) assumes that the partially saturated porous medium is characterized by fully connected fluid phases, which saturate particular subsets of the probed porous medium (Blunt, 2017). The effective saturation  $S_{we}$  is related to the total  $S_w$  saturation by  $S_w = S_{we}(1-S_{wr}) + S_{wr}$ , with  $S_{wr}$  denoting the wetting fluid residual saturation.

The effective volumetric flow rates for the wetting phase can be obtained by integrating equation (3) between  $R_{\min}$  and  $R_p(p_c)$ , respectively. Then, employing equation (18), the frequency-dependent dynamic effective permeability for the wetting phase is (Solazzi et al., 2020)

$$\kappa^{\text{eff}}(p_{\text{c}},\omega) = \frac{1}{\tau R_{\text{REV}}^2 k_w^2} \int_{R_{\text{min}}}^{R_{\text{p}}(p_{\text{c}})} \left[ \frac{2}{k_w R} \frac{J_1(k_w R)}{J_0(k_w R)} - 1 \right] R^2 \mathfrak{f}(R) \, \mathrm{d}R, \tag{22}$$

where  $k_w^2 = i\omega \rho_w / \eta_w$ . Note that equation (22) is the extension of equation (5) to partially saturated media, as  $p_c = p_c(S_w)$ . As expected, in the low-frequency limit, this expression converge to its Poiseuille-type counterpart (e.g., Blunt, 2017)

$$\kappa_w^{\text{eff},\,0}(p_c) = \frac{1}{\tau 8R_{\text{REV}}^2} \int_{R_{\text{min}}}^{R_p(p_c)} R^4 \mathfrak{f}(R) \,\mathrm{d}R.$$
(23)

Please note that, in the derivation of equation (22), a no-slip condition is assumed to prevail at the interface between the saturating fluid and the pore walls. In presence of a non-negligible flow velocity at the fluid-pore wall boundary (slip condition), which may arise due to wettability effects, the dynamic permeability estimates (Li et al., 2020) and the electrokinetic response of the medium (Collini & Jackson, 2022) are expected to change. Such boundary effects are, however, beyond the scope of this work.

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#### 2.2.2 Effective Excess Charge Density

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The effective excess charge carried by the water flow a the partially saturated medium is then obtained by integrating the excess charge along the pores that are effectively saturated with water for a given capillary pressure  $p_c$ , weighted by the corresponding flow rates

 $\hat{Q}_{\nu}^{\text{REV}}(p_c,\omega) = \frac{\int_{R_{min}}^{R_p(p_c)} \hat{Q}_{\nu}^R(\omega)q(R,\omega)\mathfrak{f}(R)\,\mathrm{d}R}{\int_{R_{min}}^{R_p(p_c)} q(R,\omega)\mathfrak{f}(R)\,\mathrm{d}R}.$ (24)

Note that since the capillary pressure  $p_c$  is related to the water saturation  $S_w$ , we consider  $\hat{Q}_{\nu}^{\text{REV}}(p_c(S_w),\omega) \equiv \hat{Q}_{\nu}^{\text{REV}}(S_w,\omega)$ , without loss of generality. Notably, it is possible to define a frequency- and saturation-dependent relative excess charge density

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$$\hat{Q}_{\nu}^{\text{rel}}(S_w,\omega) = \frac{\hat{Q}_{\nu}^{REV}(S_w,\omega)}{\hat{Q}_{\nu}^0(S_w)},$$
(25)

where 
$$\hat{Q}^0_{\nu}(S_w) = \lim_{\omega \to 0} \hat{Q}_{\nu}(S_w, \omega)$$

#### 2.2.3 Electrokinetic Coupling Coefficient

By means of the above defined parameters, we extend the dynamic electrokinetic coupling definition to partially saturated conditions as (Revil & Mahardika, 2013)

$$C_{EK}(S_w,\omega) = -\frac{\hat{Q}_{\nu}^{\text{REV}}(S_w,\omega)\kappa^{\text{eff}}(S_w,\omega)}{\eta_w \sigma(S_w,\omega)},$$
(26)

where  $\kappa^{\text{eff}}(S_w, \omega)$  and  $\hat{Q}_{\nu}^{\text{REV}}(S_w, \omega)$  respond to equations (22) and (24), respectively. The electrical conductivity (S/m), in its low-frequency limit, responds to (Waxman & Smits, 1968)

$$\sigma^0(S_w) = \frac{S_w^n}{F} \left( \sigma_w + \frac{\sigma_s}{S_w} \right), \tag{27}$$

where  $\sigma_s$  (S/m) is the surface conductivity,  $F = \tau/\phi$  is the formation factor, and n the second Archie's coefficient. Even though the electrical conductivity can be considered as frequency dependent, for simplicity, we hereafter consider  $\sigma(S_w, \omega) \approx \sigma^0(S_w)$ . For more information about the frequency dependence of the electrical conductivity, we refer the readers to pertinent literature on the subject (e.g., Jougnot et al., 2010; Revil, 2013).

Finally, the relative electrokinetic coupling coefficient for partially saturated media responds to

$$C_{EK}^{\text{rel}}(S_w,\omega) = \frac{C_{EK}(S_w,\omega)}{C_{EK}^0(S_w)},\tag{28}$$

where  $C_{EK}^0(S_w) = \lim_{\omega \to 0} C_{EK}(S_w, \omega).$ 

Equations (24) and (26) are the central methodological result of this paper, as they define the saturation- and frequency-dependent effective excess charge density and electrokinetic coupling coefficient at the REV scale by means of a flux-averaging upscaling procedure. We remark that both  $\hat{Q}_{\nu}^{\text{REV}}(S_w,\omega)$  and  $C_{EK}(S_w,\omega)$  depend on the PSD of the probed medium, a characteristic that is included in the corresponding expressions via the  $\mathfrak{f}(R)$  function.

#### 376 **3 Results**

In this section, we analyze the effects of frequency and saturation on the effective excess charge density  $\hat{Q}_{\nu}^{\text{REV}}(S_w,\omega)$  and the electrokinetic coupling coefficient  $C_{EK}(S_w,\omega)$ in porous media. We assess the effects of the pore size distribution in the corresponding response by considering: (i) fractal, (ii) lognormal, and (iii) double-lognormal pore size distributions.

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#### 3.1 Pore Size Distribution

3.1.1 Fractal Distribution Function

As a first case, we consider a cumulative size distribution of pores whose radii are greater than or equal to R that obeys the following fractal law (e.g., Guarracino et al., 2014; Tyler & Wheatcraft, 1990; Yu et al., 2003)

$$N(R) = \left(\frac{R_{\max}}{R}\right)^D,\tag{29}$$

where D is the fractal dimension of the pore size with 1 < D < 2 and  $R_{\min} < R < 2$ 

 $R_{\text{max}}$ . The total number of pores, from  $R_{\text{min}}$  to  $R_{\text{max}}$ , is given by

$$N_t(R_{\min}) = \left(\frac{R_{\max}}{R_{\min}}\right)^D.$$
(30)

On the other hand, differentiating N(R) with respect to R, we obtain the number of pores whose radii are between R and R + dR:

$$-\mathrm{d}N = DR_{\mathrm{max}}^D R^{-D-1} \mathrm{d}R = \mathfrak{f}(R) \mathrm{d}R.$$
(31)

#### Dividing equations (31) and (30), we obtain the probability density function $f_r(R)$

$$-\frac{\mathrm{d}N}{N_t} = DR_{\min}^D R^{-D-1} \mathrm{d}R = f_r(R) \mathrm{d}R,$$

396 such that,

$$\int_{R_{\min}}^{R_{\max}} f_r(R) \mathrm{d}R = 1 - \left(\frac{R_{\min}}{R_{\max}}\right)^D \equiv 1,\tag{33}$$

(32)

which clearly holds if  $(R_{\min}/R_{\max})^D \simeq 0$ . In this sense, the condition  $R_{\max} >> R_{\min}$ must be satisfied for fractal analysis of porous media. Please note that,  $f(R) = N_t f_r(R)$ .

#### 400 3.1.2 Lognormal Distribution Function

401 The lognormal distribution probability density function responds to

$$f_r(R) = \frac{1}{\mathfrak{s}R\sqrt{2\pi}} \exp\left(-\frac{\left(\log R - \mathfrak{x}\right)^2}{2\mathfrak{s}^2}\right).$$
(34)



Figure 1. Probability density functions associated with the pore size distributions used in this work: (a) fractal (D = 1.5), (b) lognormal ( $R^* = 10 \,\mu\text{m}$  and  $\mathfrak{s} = 0.46$ ), and (c) double lognormal ( $R^*_1 = 3.1 \,\mu\text{m}$ ,  $R^*_2 = 31 \,\mu\text{m}$ ,  $\mathfrak{s}_d = \mathfrak{s}/2$ ,  $\beta_1 = 0.09$ , and  $\beta_2 = 0.91$ ).

where  $\mathfrak{x} = \log R^*$  and  $\mathfrak{s}$  denote the scale and shape parameters. Again, we consider that  $\mathfrak{f}(R) = N_t f_r(R)$ , where  $N_t$  is the total number of pores in the medium.

#### 405 3.1.3 Double Lognormal Distribution Function

The double lognormal distribution can be regarded as the sum of two lognormal distributions with the same shape parameter  $\mathfrak{s}_d$  and responds to

$$f_{r}(R) = \beta_{1} \frac{1}{\mathfrak{s}_{d} R \sqrt{2\pi}} \exp\left(-\frac{(\log R - \mathfrak{x}_{1})^{2}}{2\mathfrak{s}_{d}^{2}}\right) + \beta_{2} \frac{1}{\mathfrak{s}_{d} R \sqrt{2\pi}} \exp\left(-\frac{(\log R - \mathfrak{x}_{2})^{2}}{2\mathfrak{s}_{d}^{2}}\right), \quad (35)$$

where  $\mathfrak{x}_1 = \log R_1^*$  and  $\mathfrak{x}_2 = \log R_2^*$ , and  $\beta_1 + \beta_2 = 1$ . Again, we consider that  $\mathfrak{f}(R) = N_t f_r(R)$ .

Figure 1 shows the representative PSDs considered in this work with pore radius 411 ranging from 1  $\mu$ m to 100  $\mu$ m: (a) fractal (D = 1.5), (b) lognormal ( $R^* = 10 \mu$ m and 412  $\mathfrak{s} = 0.46$ ), and (c) double lognormal ( $R_1^* = 3.1 \,\mu\text{m}, R_2^* = 31 \,\mu\text{m}, \mathfrak{s}_d = \mathfrak{s}/2, \, \beta_1 = 0.09$ , 413 and  $\beta_2 = 0.91$ ). We remark that smaller pore radii dominate the response of the medium 414 for fractal PSD, while pores distribute more evenly throughout the given radii for the 415 lognormal and double lognormal PSDs. As shown below, the PSD characteristics result 416 in significantly different responses for the effective permeability, the effective excess charge 417 density, and, consequently, the electrokinetic coupling in porous media. 418



Figure 2. Absolute value of the effective dynamic permeability  $\kappa_w^{\text{eff}}$ , effective excess charge density  $\hat{Q}_{\nu}^{\text{REV}}$ , and effective electrokinetic coupling coefficient  $C_{\text{EK}}$  as functions of frequency for different saturation states. Each row illustrates the result for a different PSD: (a, d, g) fractal (D = 1.5), (b, e, h) lognormal  $(R^* = 10 \,\mu\text{m} \text{ and } \mathfrak{s} = 0.46)$ , and (c, f, i) double lognormal  $(R_1^* = 3.1 \,\mu\text{m}, R_2^* = 31 \,\mu\text{m}, \mathfrak{s}_d = \mathfrak{s}/2, \beta_1 = 0.09$ , and  $\beta_2 = 0.91$ ).

#### 3.2 Numerical Analysis of the Proposed Model

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Let us consider three porous media represented by: (i) a fractal PSD with a dimension D = 1.5, (ii) a lognormal PSD characterized by  $R^* = 10 \,\mu\text{m}$  and  $\mathfrak{s} = 0.46$ , and (iii) a double lognormal PSD characterized by  $R_1^* = 3.1 \,\mu\text{m}$ ,  $R_2^* = 31 \,\mu\text{m}$ ,  $\mathfrak{s}_d = \mathfrak{s}/2$ ,  $\beta_1 = 0.09$ , and  $\beta_2 = 0.91$ . We assume that they all have  $\tau \simeq 1$ ,  $R_{\min} = 1 \,\mu\text{m}$ , and  $R_{\max} = 100 \,\mu\text{m}$  and, also, the same total number of pores  $N_t$ , which is taken from the fractal distribution characteristics (equation 30). The pore fluid properties that saturate these probed media are summarized in Table 1.

Following Solazzi et al. (2020), we numerically solve equation (22) and obtain the 427 saturation- and frequency-dependent effective permeability for the above described porous 428 media. Figures 2a, 2b, and 2c illustrate the magnitude of  $\kappa^{\text{eff}}(S_w,\omega)$  as a function of fre-429 quency for the three PSDs described above. Note that each column of Figure 2 is asso-430 ciated with one particular PSD. In this context, we plot different effective saturation states, 431 identified by different colored lines. We observe that the absolute value of  $\kappa^{\rm eff}$  decreases 432 with frequency for  $f > f_c$ , with  $f_c = \omega_c/2\pi$  denoting the critical frequency (Figures 433 2a, 2b, and 2c). Recall that  $f_c$  is determined by the PSD characteristics, specifically by 434 the largest saturated pores of the distribution. The frequency-dependent behavior of  $|\kappa^{\text{eff}}(S_w,\omega)|$ 435 is explained by the onset of the inertia effects for  $f \ge f_c$ . When inertia effects prevail, 436 the amplitude of the dynamic permeability drops and its phase increases (e.g., Zhou & 437 Sheng, 1989). As previously observed by Solazzi et al. (2020), the critical frequency  $f_c$ 438 increases with decreasing saturation, as water retreats to increasingly smaller pores. We 439 also note that  $|\kappa^{\text{eff}}|$  increases with water saturation. The corresponding response is mod-440 ulated by the PSD of the probed porous medium. The reason behind this behavior is 441 that the overall number of pores saturated by water decreases with decreasing  $S_w$ , as is 442 the case in the classic relative permeability functions (e.g., Brooks & Corey, 1964; Mualem, 443 1976). Evidently, this saturation- and frequency-dependent behavior also affects  $\hat{Q}^{\text{REV}}_{\nu}$ 444 and  $C_{EK}$  at the REV scale. Note that, as expressed in equation (26),  $C_{EK}(S_w, \omega)$  de-445 pends on  $\kappa^{\text{eff}}(S_w, \omega)$  (equation 22), which includes both the effects of  $\kappa(\omega)$  (equation 5) 446 and  $\kappa^{\text{rel}}(S_w, \omega)$  (equation 19). 447

Figures 2d, 2e, and 2f, illustrate the frequency dependence of the absolute value of the effective excess charge density  $|\hat{Q}_{\nu}^{\text{REV}}(S_w,\omega)|$  at different effective saturation states (equation 24). The parameters of the PSDs and the physical properties of the wetting

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Definition	Variable	Value	Units
Fluid shear viscosity (wetting phase)	$\eta_w$	$1 \times 10^{-3}$	Pa.s
Fluid density (wetting phase)	$ ho_w$	1000	$\mathrm{kg/m^3}$
Interfacial tension (water-air)	$\gamma$	$72  imes 10^{-3}$	N/m
Contact angle	eta	0	rad
Dielectric permittivity of vacuum	$\varepsilon_0$	$8.854\times10^{-12}$	F/m
Relative permittivity of the fluid	$\varepsilon_r$	80.1	-
Boltzmann constant	$k_B$	$1.381\times 10^{-23}$	$\rm J/K$
Avogadro number	$N_A$	$6.022\times 10^{23}$	1/Mol
Elementary charge	$e_0$	$1.6\times 10^{-19}$	$\mathbf{C}$
Ionic concentration	$C^w_{NaCl}$	$10^{-4}$	$\mathrm{Mol/L}$
Temperature	T	293.15	Κ

 Table 1. Fluid properties employed in this study.

fluid are the same as those employed in panels 2a-2c of the corresponding figure. We ob-451 serve that  $|\hat{Q}_{\nu}^{\text{REV}}(S_w, \omega)|$  increases with f for  $f > f_c$  irrespective of the saturation. (Jougnot 452 & Solazzi, 2021) explored the behavior of  $\hat{Q}_{\nu}(\omega)$  in fully saturated conditions, and ob-453 served a corresponding increase for  $f > f_c$ . By comparing panels 2a to 2c with pan-454 els 2d to 2f, we observe identical shifts in the characteristic frequency  $f_c(S_w)$ , which moves 455 towards higher frequencies for decreasing saturation. Again, this  $f_c$ -shift is different for 456 each PSD, evidencing larger change of  $f_c$  with saturation for the fractal PSD than for 457 the double lognormal PSD. Note that, as shown in previous works in fully saturated me-458 dia (e.g., Jougnot & Solazzi, 2021; Guarracino & Jougnot, 2018; Soldi et al., 2019), the 459 magnitude of the effective excess charge density increases when the characteristic cap-460 illary size decreases. Consequently, the magnitude of  $\hat{Q}_{\nu}^{\text{REV}}(S_w,\omega)$  increases with decreas-461 ing  $S_w$ , as water retreats towards relatively small pores. Both the fractal and the log-462 normal distribution characteristics, as considered in this section, present a larger num-463 ber of small pores when compared with the double lognormal PSD. This is precisely the 464 reason for a larger relative variation in  $\hat{Q}_{\nu}^{\text{REV}}(S_w,\omega)$  values for the fractal and lognor-465 mal PSDs for decreasing saturations as compared with those associated with the dou-466 ble lognormal PSD. 467

Finally, using equations (21) to (27), we predict the frequency dependence of the 468 electrokinetic coupling coefficient  $C_{EK}$  at different saturation states (Figs 2g, 2h, and 469 2i). We consider the previously described PSDs and pore fluid properties, together with 470 representative values of  $\sigma_s = 3 \times 10^{-3}$  S/m,  $S_{wr} = 0.2$ , n = 1.7 and F = 5 to model 471 the variation of electrical conductivity of porous media as a function of water saturation 472  $S_w$ . To infer the electrical conductivity  $\sigma_w$  from  $C^w$ , we employ the relation  $\sigma_w = 10 \times$ 473  $C^w$  for a NaCl solution (Sen & Goode, 1992). The results show that the magnitude of 474  $C_{EK}$  decreases with increasing frequency for  $f > f_c$  irrespective of  $S_w$  (Figs 2g, 2h, and 475 2i). This behavior is in good match with published works for the case of full saturation 476 (e.g., Jougnot & Solazzi, 2021; Pride, 1994; Revil & Mahardika, 2013). Even though the 477 amplitude of  $C_{EK}$  appears to decrease with decreasing  $S_w$  for lognormal (Fig. 2h) and 478 double lognormal PSDs (Fig. 2i), this is not the case for the fractal PSD (Fig. 2g). In 479 the latter case, we note that the amplitude of  $C_{EK}$  increases and, then, decreases with 480 saturation, a behavior that is more clearly illustrated below (Figure 3). 481

For completeness, we illustrate the behavior of  $\kappa_w^{\text{eff}}$ ,  $\hat{Q}_{\nu}^{\text{REV}}$ , and  $C_{\text{EK}}$  as functions 482 of the effective saturation  $S_{we}$ , for different frequencies (Figure 3). Black circled lines 483 denote the so-called low frequency limit for  $\kappa^{\text{eff}}$ ,  $\hat{Q}_{\nu}^{\text{REV}}$ , and  $C_{\text{EK}}$ , while colored lines de-484 pict the responses for  $f = 10^2$  Hz,  $f = 10^3$  Hz, and  $f = 10^4$  Hz. We observe that all 485 curves tend to the same value for sufficiently small  $S_{we}$  values, irrespective of the prob-486 ing frequency. This is expected, as  $f_c$  shifts towards higher frequencies for decreasing sat-487 urations (see Figure 2). Hence, the probing frequencies became smaller than  $f_c$  for suf-488 ficiently low saturations and  $\kappa_w^{\text{eff}}$ ,  $\hat{Q}_{\nu}^{\text{REV}}$ , and  $C_{\text{EK}}$  tend to their low-frequency counter-489 parts. Conversely, for increasing  $S_{we}$ , the overall responses experience a departure from 490 the low-frequency behavior. Figure 3 evidences the control that the PSD has on  $\kappa_w^{\text{eff}}(S_w)$ , 491  $\hat{Q}_{\nu}^{\text{REV}}(S_w)$ , and  $C_{\text{EK}}(S_w)$  for different probing frequencies, as we note different slopes 492 and inflections for different PSDs. 493

#### 4 Discussion 494

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In this section, we compare the  $\hat{Q}_{\nu}^{\text{REV}}(\omega, S_w)$  estimates obtained by means of the proposed flux-averaging approach with respect to those predicted by the pioneering model 496 of Revil and Mahardika (2013). Then, we address the capability of the proposed model 497 to predict experimental measurements of  $C_{EK}^{rel}(\omega, S_w)$  which, to date, have been only per-498 formed under fully saturated conditions  $(S_w = 1)$ . 499

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Figure 3. Absolute value of the effective dynamic permeability  $\kappa_w^{\text{eff}}$ , effective excess charge density  $\hat{Q}_{\nu}^{\text{REV}}$ , and effective electrokinetic coupling coefficient  $C_{\text{EK}}$  as functions of saturation for different frequencies. Each row illustrates the result for a different PSD: (a, d, g) fractal (D = 1.5); (b, e, h) lognormal ( $R^* = 10 \,\mu\text{m}$  and  $\mathfrak{s} = 0.46$ ); and (c, f, i) double lognormal ( $R_1^* = 3.1 \,\mu\text{m}$ ,  $R_2^* = 31 \,\mu\text{m}$ ,  $\mathfrak{s}_d = \mathfrak{s}/2$ ,  $\beta_1 = 0.09$ , and  $\beta_2 = 0.91$ ).

#### 4.1 Comparison with Previous Models

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<sup>501</sup> In their classical work, Revil and Mahardika (2013) proposed the following empir-<sup>502</sup> ical model for the frequency- and saturation-dependent effective excess charge density

$$\hat{Q}_{\nu}^{\text{REV}}(S_w,\omega) \simeq \hat{Q}_{\nu}^0(S_w) \sqrt{1 - i\omega\tau_k(S_w)},\tag{36}$$

where  $\hat{Q}^0_{\nu}(S_w)$  denotes the low-frequency value of the effective excess charge density and  $\tau_k$  denotes the relaxation time of the frequency-dependent behavior, which is given by

$$\tau_k(S_w) = \kappa^{\text{eff}}(S_w) \frac{\rho_w F S_w^{1-n}}{\eta_w}.$$
(37)

In order to compute equation (36), Revil and Mahardika (2013) use the volume averaging model of Linde et al. (2007)

$$\hat{Q}^{0}_{\nu}(S_{w}) \simeq \frac{\hat{Q}^{\text{sat, 0}}_{\nu}}{S_{w}}.$$
 (38)

where  $\hat{Q}_{\nu}^{\text{sat, 0}}$  is the low-frequency effective excess charge density in fully saturated conditions. To estimate  $\kappa^{\text{eff}}(S_w)$ , the authors take the Brooks and Corey (1964) model

$$\kappa^{\text{eff}}(S_w) = \kappa_0 S_w^{\frac{2+3\lambda}{\lambda}},\tag{39}$$

with  $\lambda$  a fitting parameter that is determined by the pore space characteristics of the probed medium.

It is important to remark that equation (36) is based on a linear and low-frequency 515 approximation of the dynamic permeability, which is commonly used to deal with  $\kappa(\omega)$ 516 in the space-time domain (e.g., Revil & Mahardika, 2013). Given that the proposed flux-517 averaging approach (equation 24) is developed in the space-frequency domain, our es-518 timates are not limited by such assumption. More importantly, in equation (36), Revil 519 and Mahardika (2013) assume that equations (38) and (39) hold for the probed medium. 520 If we wish to compare our approach with such model, it is important to analyze first the 521 validity of equations (38) and (39) for the considered PSDs. 522

Figure 4 shows a comparison between equations (38) and (39) with the corresponding ones considered in this work, which respond to equations (6) and (24) (in the low frequency). We observe that equation (39) correctly reproduces the tendencies of the effective permeability associated with the fractal PSD (Figure 4a), which is in agreement with the observations of Soldi et al. (2017) for fractal media. Considering typical  $\lambda$  values, we note that equation (39) tends to underestimate  $\kappa^{\text{eff},0}(S_w)$  for lognormal and double lognormal PSDs (Figures 4b and 4c). We remark that the considered pore-structure



Figure 4. (a, b, c) Effective permeability and effective excess charge density as a function of the effective saturation in the low-frequency limit for: (a, d) fractal (D = 1.8,  $R_{\min} = 23$  nm and  $R_{\max} = 4.7 \ \mu\text{m}$ ), (b, e) lognormal ( $R^* = 1.4 \ \mu\text{m}$  and  $\mathfrak{s} = 0.15$ ), and (c, d) double lognormal ( $R_1^* = 1.0 \ \mu\text{m}$ ,  $R_2^* = 1.5 \ \mu\text{m}$ ,  $\mathfrak{s}_d = \mathfrak{s}/2$ ,  $\beta_1 = 0.4$ , and  $\beta_2 = 0.6$ ) PSDs. Panels (a, b, c) illustrate the behavior of equation (39) for  $\lambda = \{1, 4\}$  (magenta solid lines) and that of equation (6) (black circled lines). Panels (d, e, c) illustrate the behavior predicted by equation (38) (magenta solid lines) compared with that of (24) in the low-frequency range (black circled lines).

is highly idealized and these differences can be a source of mismatch, as equation (39)530 is known to provide reliable predictions of  $\kappa^{\text{eff},0}(S_w)$  in siliciclastic rocks. On the other 531 hand, Figures 4d to 4f allow us to test the assumption expressed in equation (38). In-532 terestingly, we note that this equation provides with a fair representation of  $\lim_{\omega \to 0} |\hat{Q}_{\nu}^{\text{REV}}|$ 533 when considering a lognormal PSD. However, it tends to give biased representations of 534 the corresponding variable for fractal and double lognormal PSDs. Particularly, equa-535 tion (38) results in estimations that significantly differ from those predicted by the low-536 frequency limit of equation (24) for low saturations. We conclude that, when compar-537 ing the proposed approach with that of Revil and Mahardika (2013) (equation 36), dif-538 ferences associated with the estimates given by equations (38) and (39) may be a source 539 of mismatch. In order to circumvent this issue and, also, given that performing low-frequency 540 measurements of  $\kappa^{\text{eff},0}(S_w)$  and  $\hat{Q}_{\nu}^{\text{REV},0}(S_w)$  is feasible in laboratory setups, in the fol-541 lowing, we propose to perform the comparison of equations (24) and (36) assuming that 542  $\kappa^{\text{eff},0}(S_w)$  and  $\hat{Q}_{\nu}^{\text{REV},0}(S_w)$  are known and, in this case, given by those resulting from 543 the flux-averaging approach proposed in this work. As such, below, we concentrate solely 544 on comparing the frequency-dependent response predicted by our model and that of Revil 545 and Mahardika (2013). 546

Figure 5 shows a comparison between the results from equation (24) (solid lines) 547 and (36) (dashed lines) for the proposed PSDs, at each column. The first row depicts 548 the absolute value of  $|\hat{Q}_{\nu}^{\text{REV}}(S_w,\omega)|$  as a function of frequency (for different saturation 549 states) and the second row the corresponding behavior as a function of saturation (for 550 different probing frequencies). We observe that, in general, the proposed flux-averaging 551 model and the model of Revil and Mahardika (2013) provide with similar tendencies. In 552 particular, estimates are largely similar for saturations approaching unity (Figures 5a 553 to 5c) and at relatively low frequencies (Figures 5d to 5f). Given that the proposed model 554 obtained the corresponding estimates, for the first time in the literature, averaging the 555 physical processes from the pore to the REV scale, we are thus validating the model of 556 Revil and Mahardika (2013) and proving its consistency and robustness despite its sim-557 plicity. 558

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#### 4.2 Comparison with Experimental Data

To date, measurements of  $C_{EK}^{rel}(\omega)$  for different probing frequencies have been performed only under fully saturated conditions. The proposed model should have the ca-

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Figure 5. Effective excess charge density as a function of (a, b, c) frequency and (d, e, f) of the effective saturation. We consider: (a, d) a fractal PSD ( $D = 1.8, R_{\min} = 23 \,\mu\text{m}$  and  $R_{\max} = 4.7 \,\mu\text{m}, m = 1.19$ ); (b, e) a lognormal PSD ( $R^* = 1.4 \,\mu\text{m}$  and  $\mathfrak{s} = 0.15, m = 1.06$ ), and (c, d) a double lognormal PSD ( $R_1^* = 1.0 \,\mu\text{m}, R_2^* = 1.5 \,\mu\text{m}, \mathfrak{s}_d = \mathfrak{s}/2, \beta_1 = 0.4$ , and  $\beta_2 = 0.6, m = 1$ ) PSDs. Solid lines illustrate the behavior of the proposed flux-averaging model (equation 24) and dashed lines illustrate the behavior predicted by Revil and Mahardika (2013) model (equation 36).



Figure 6. (a) Real and (b) imaginary parts of the relative electrokinetic coupling coefficient as functions of frequency. We illustrate results of the proposed model for different saturations (solid lines) using a lognormal PSD. Red squares depict the experimental measurements of Glover, Walker, and Jackson (2012) for an Ottawa sand at full saturation. We also illustrate the predictions of the proposed model for other saturations (colored lines). The dashed blue arrows indicate the direction in which saturation decreases.

pability of representing experimental measurements in such end-member scenario of sat uration and, also, to predict the partially saturated response of the corresponding me dia.

Figure 6 shows the frequency dependence of the real and imaginary parts of  $C_{EK}^{rel}(\omega)$ 565 at fully saturated conditions as reported by Glover, Walker, Ruel, and Tardif (2012) for 566 an Ottawa sand sample. The Ottawa sand is characterized by a mean grain radius of 235 567  $\mu$ m. Glover, Walker, Ruel, and Tardif (2012) used a 10<sup>-3</sup> mol/L NaCl electrolyte. We 568 employed a lognormal PSD in combination with equation (17), that is equation (28) with 569  $S_w=1$ , to model the behavior of the measured data. We take  $R_{min}=1.05\,\mu{\rm m}$  and  $R_{min}=1.05\,\mu{\rm m}$ 570  $105\,\mu\text{m}$ , as well as  $R^* = 60\,\mu\text{m}$  and s = 0.15. Note that  $R^*$  is close to the effective pore 571 radii of  $r_p$  = 67  $\mu {\rm m},$  as reported by Glover, Walker, Ruel, and Tardif (2012) for the cor-572 responding sample. The pore fluid properties are summarized in Table 1. We observe 573 that the proposed model is able to reproduce experimental data (Figure 6, black lines). 574 We also illustrate variations predicted by the proposed model for  $C_{EK}^{rel}$  for saturations 575 of  $S_w = 0.6, 0.3$  and 0.1 (Figure 6, colored lines). The dashed blue arrows indicate the 576 direction in which saturation decreases in Figures 6a and 6b, evidencing an increase of 577



Figure 7. Amplitude of the electrokinetic coupling coefficient as function of frequency for different saturations. Each column displays the results for a different PSDs: (a) fractal (D = 1.8,  $R_{\min} = 23$  nm and  $R_{\max} = 4.7 \mu$ m), (b) lognormal ( $R^* = 1.4 \mu$ m and  $\mathfrak{s} = 0.15$ ), and (c) double lognormal ( $R_1^* = 1.0 \mu$ m,  $R_2^* = 1.5 \mu$ m,  $\mathfrak{s}_d = \mathfrak{s}/2$ ,  $\beta_1 = 0.4$ , and  $\beta_2 = 0.6$ ). Red squares depict the experimental measurements of Peng et al. (2020) for a fully saturated sandstone. We also illustrate the predictions of the proposed model for other saturations (colored lines). The dotted blue arrows indicate the direction in which saturation decreases.

 $f_c$  with decreasing  $S_w$ , a fact that is also observed in the effective permeability curves for lognormal distributions (Figure 2b).

Figures 7 shows the frequency dependence of  $|C_{EK}^{rel}|$  at full saturation for a sand-580 stone sample, as measured by Peng et al. (2020) (red squares). We show that the pro-581 posed model is capable of fitting the main trend of experimental data by means of the 582 three PSD described in this study, by using: (a) fractal  $(D = 1.8, R_{\min} = 23 \text{ nm and})$ 583  $R_{\rm max} = 4.7 \ \mu {\rm m}$ ), (b) lognormal ( $R^* = 1.4 \mu {\rm m}$  and  $\mathfrak{s} = 0.15$ ), and (c) double lognor-584 mal  $(R_1^* = 1.0 \,\mu\text{m}, R_2^* = 1.5 \,\mu\text{m}, \mathfrak{s}_d = \mathfrak{s}/2, \beta_1 = 0.4, \text{ and } \beta_2 = 0.6)$ . We remark that, 585 because of computational restrictions involved with the numerical integrations performed 586 in this work, we do not carry out a full inversion of the parameters but empirically find 587 those which provide a relatively good fit with experimental data. Nevertheless, these pa-588 rameters are similar to those reported by Thanh et al. (2021) for the same sample to model 589 the frequency dependence of the electrokinetic coupling coefficient, that is directly ex-590 pressed via the Zeta potential rather than the effective excess charge density. Once again, 591 we illustrate the predicted variations of  $|C_{EK}^{rel}|$  for different saturations (colored lines). 592

<sup>593</sup> Finally, Figure 8 shows the frequency dependence of  $|C_{EK}|$  measured by Zhu and <sup>594</sup> Toksoz (2013) for a sample of Berea sandstone (squared curves) for different electrical <sup>595</sup> conductivities. We employ the proposed model considering  $R_{\min} = 0.13 \,\mu\text{m}$  and  $R_{\max} =$ 



Figure 8. Amplitude of the electrokinetic coupling coefficient as functions of frequency under saturated conditions for different electrical conductivities. Colored squares denote measurements taken by Zhu and Toksoz (2013) for a Berea sandstone. We illustrate the predictions of the proposed model for three different PSDs: (a) fractal (D = 1.65,  $R_{\min} = 0.13 \ \mu\text{m}$  and  $R_{\max} = 30 \ \mu\text{m}$ ), (b) lognormal ( $R^* = 6.3 \ \mu\text{m}$  and  $\mathfrak{s} = 0.15$ ,  $R_{\min} = 0.13 \ \mu\text{m}$  and  $R_{\max} = 30 \ \mu\text{m}$ ), and (c) double lognormal ( $R_1^* = 1.0 \ \mu\text{m}$ ,  $R_2^* = 3.0 \ \mu\text{m}$ ,  $\mathfrak{s}_d = \mathfrak{s}/2$ ,  $\beta_1 = 0.4$ , and  $\beta_2 = 0.6$ ).

- $_{596}$  30  $\mu$ m for the different PSD (see solid lines), that is, (a) fractal (D = 1.65), (b) lognor-
- <sup>597</sup> mal  $(R^* = 6.3 \mu \text{m and } \mathfrak{s} = 0.15)$  and (c) double lognormal  $(R_1^* = 1.0 \mu \text{m}, R_2^* = 3.0 \mu \text{m},$
- $\mathfrak{s}_{d} = \mathfrak{s}/2, \ \beta_{1} = 0.4, \ \mathrm{and} \ \beta_{2} = 0.6).$  The values of  $C_{EK}^{0}$  are reported to be  $0.3 \times 10^{-6}$ ,
- $_{599}$  0.15×10<sup>-6</sup>, 0.065×10<sup>-6</sup>, 0.035×10<sup>-6</sup> and 0.024×10<sup>-6</sup> V/Pa for 0.012, 0.048, 0.095, 0.18
- and 0.32 S/m, respectively (Zhu & Toksoz, 2013). Using equation (17), we are able to
- obtain  $C_{EK}^{rel}$  and, hence,  $C_{EK}$ . It is seen that the proposed approach using three con-
- sidered PSDs is capable of reproducing the experimental data very well.

#### 5 Conclusions

We proposed a flux averaging approach to compute the frequency- and saturation-604 dependent effective excess charge density in partially saturated porous media. For this, 605 we conceptualized the pore space as a bundle of capillary tubes with a given pore size 606 distribution (PSD). We modeled the frequency dependence of the effective excess charge 607 density by solving the Navier-Stokes equations under oscillatory flow conditions within 608 the capillaries that are effectively saturated for a given capillary pressure state. Over-609 all, we derived expressions for: (i) the capillary pressure-saturation relationship of the 610 probed medium; and for the saturation- and frequency-dependent (ii) effective perme-611 ability  $\kappa^{\text{eff}}(S_w,\omega)$ , (iii) effective excess charge density  $\hat{Q}_{\nu}^{REV}(S_w,\omega)$ , and (iv) electroki-612 netic coupling coefficient  $C_{\rm EK}(S_w, \omega)$ . 613

The variation of  $\kappa^{\text{eff}}$ ,  $\hat{Q}_{\nu}^{\text{REV}}$  and  $C_{\text{EK}}$  with frequency at different saturation states 614 are analyzed and explained for three different PSDs (fractal, lognormal and double log-615 normal PSDs). It is shown that the PSD has strong effect on the critical frequency  $\omega_c$ 616 and the characteristics of  $\kappa^{\text{eff}}$ ,  $\hat{Q}_{\nu}^{\text{REV}}$  and  $C_{EK}$  as functions of frequency and saturation. 617 Namely, the critical frequency  $\omega_c$  increases with decreasing water saturation  $S_w$  for a given 618 PSD. The reason is that when the water saturation decreases, only the smaller radius 619 pores are saturated by water, leading to a decrease of the characteristic radius represen-620 tative  $\hat{R}$  of the saturated pores. This process affects the effective excess charge density 621 at the REV scale. The proposed model is compared with previous models in the liter-622 ature and, in the case of full saturation, it is also compared with published data. We found 623 that the proposed model is capable of reproducing the frequency-dependence of  $\hat{Q}^{\text{REV}}_{\nu}$ 624 as predicted by previous models, which do not rely in a flux-averaging approach, pro-625 vided that the low-frequency estimates of the effective excess charge and effective per-626 meability are correct. On the other hand, our approach was able to represent experimen-627 tal measurements of the coupling coefficient  $C_{\rm EK}$  for different frequencies, conductivi-628 ties, and rock properties. The proposed approach is valid for practically any PSD and 629 constitutes a practical framework for the interpretation of seismoelectric signatures of 630 partially saturated media. 631

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#### <sup>632</sup> Data Availability Statement

The data and computational codes associated with this paper are available online from https://doi.org/10.5281/zenodo.6620462.

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#### 640 References

- Auriault, J.-L., Borne, L., & Chambon, R. (1985). Dynamics of porous saturated
   media, checking of the generalized law of Darcy. J. Acoust. Soc. Am., 77(5),
   1641–1650.
- Bear, J. (1972). Dynamics of Fluids in Porous Media. N. Y.: Elsevier.
- Blunt, M. J. (2017). Multiphase Flow in Permeable Media: A Pore-Scale Perspective.
  Cambridge University Press.
- Bordes, C., Sénéchal, P., Barrière, J., Brito, D., Normandin, E., & Jougnot, D.
- (2015). Impact of water saturation on seismoelectric transfer functions: a laboratory study of coseismic phenomenon. *Geophys. J. Int.*, 200(3), 1317-1335.
- Brooks, R., & Corey, A. (1964). Hydraulic properties of porous media. Colorado
  State University.
- Butler, K. E. (1996). Seismoelectric effects of electrokinetic origin. PhD thesis, University of British Columbia.
- <sup>654</sup> Collini, H., & Jackson, M. D. (2022). Relationship between zeta potential and
   <sup>655</sup> wettability in porous media: insights from a simple bundle of capillary tubes
   <sup>656</sup> model. J. Colloid Interface Sci., 608, 605–621.
- <sup>657</sup> Dupuis, J. C., Butler, K. E., & Kepic, A. W. (2007). Seismoelectric imaging of the <sup>658</sup> vadose zone of a sand aquifer. *Geophysics*, 72(6), A81–A85.
- Gao, Y., Harris, J. M., Wen, J., Huang, Y., Twardzik, C., Chen, X., & Hu, H.
- (2016). Modeling of the coseismic electromagnetic fields observed during
  the 2004 Mw 6.0 Parkfield earthquake. *Geophys. Res. Lett.*, 43(2), 620–627.
- Garambois, S., & Dietrich, M. (2001). Seismoelectric wave conversions in porous

-29-

663	media: Field measurements and transfer function analysis. $Geophysics, 66(5),$
664	1417–1430.
665	Glover, P. W. J., Walker, E., & Jackson, M. (2012). Streaming-potential coefficient
666	of reservoir rock: A theoretical model. $Geophysics$ , $77(2)(2)$ , D17-D43.
667	Glover, P. W. J., Walker, E., Ruel, J., & Tardif, E. (2012). Frequency-dependent
668	streaming potential of porous media-part 2: Experimental measurement of
669	unconsolidated materials. Int. J. Geophys., 2012.
670	Grobbe, N., Revil, A., Zhu, Z., & Slob, E. (2020). Seismoelectric exploration: The-
671	ory, experiments, and applications (Vol. 252). John Wiley & Sons.
672	Guarracino, L., & Jougnot, D. (2018). A physically based analytical model to
673	describe effective excess charge for streaming potential generation in water
674	saturated porous media. J. Geophys. Res.: Solid Earth, 123(1), 52-65.
675	Guarracino, L., Rötting, T., & Carrera, J. (2014). A fractal model to describe the
676	evolution of multiphase flow properties during mineral dissolution. $Adv Water$
677	Resour, 67, 78–86.
678	Han, SJ., Kim, SS., & Kim, BI. (2004). Electroosmosis and pore pressure
679	development characteristics in lead contaminated soil during electrokinetic
680	remediation. $Geosci. J., 8(1), 85-93.$
681	Jaafar, M. Z., Vinogradov, J., & Jackson, M. D. (2009). Measurement of streaming
682	potential coupling coefficient in sandstones saturated with high salinity nacl
683	brine. Geophys. Res. Lett., 36(L21306), doi:10.1029/2009GL040549.
684	Jackson, M. D. (2010). Multiphase electrokinetic coupling: Insights into the impact
685	of fluid and charge distribution at the pore scale from a bundle of capillary
686	tubes model. J. Geophys. Res.: Solid Earth, 115(B7).
687	Johnson, D. L., Koplik, J., & Dashen, R. (1987). Theory of dynamic permeability
688	and tortuosity in fluid-saturated porous media. J. Fluid Mech., 176, 379-402.
689	doi: 10.1017/S0022112087000727
690	Jougnot, D., Linde, N., Revil, A., & Doussan, C. (2012). Derivation of soil-specific
691	streaming potential electrical parameters from hydrodynamic characteristics of
692	partially saturated soils. Vadose Zone J., 11(1), 272-286.
693	Jougnot, D., Mendieta, A., Leroy, P., & Maineult, A. (2019). Exploring the effect
694	of the pore size distribution on the streaming potential generation in saturated
695	porous media, insight from pore network simulations. J. Geophys. Res.: Solid

696	$Earth, \ 124(6), \ 5315-5335.$
697	Jougnot, D., Revil, A., Ghorbani, A., Leroy, P., & Cosenza, P. (2010). Spectral
698	induced polarization of partially saturated clay-rocks: a mechanistic approach.
699	Geophys. J. Int., 180(1), 210-224.
700	Jougnot, D., Roubinet, D., Guarracino, L., & Maineult, A. (2020). Modeling stream-
701	ing potential in porous and fractured media, description and benefits of the
702	effective excess charge density approach. In Advances in Modeling and Inter-
703	pretation in Near Surface Geophysics (pp. 61–96). Springer.
704	Jougnot, D., Rubino, J. G., Carbajal, M. R., Linde, N., & Holliger, K. (2013).
705	Seismoelectric effects due to mesoscopic heterogeneities. Geophys. Res. Lett.,
706	$4\theta(10), 2033-2037.$
707	Jougnot, D., & Solazzi, S. G. (2021). Predicting the frequency-dependent effective
708	excess charge density: A new up-scaling approach for seismoelectric modelling.
709	Geophysics, 86(3), 1-10.
710	Kozeny, J. (1927). Über kapillare Leitung des Wassers im Boden. Akad. Wiss,
711	136(2a), 271-306.
712	Li, J. X., Rezaee, R., & Müller, T. M. (2020). Wettability effect on wave propaga-
713	tion in saturated porous medium. J. Acoust. Soc, $147(2)$ , $911-920$ .
714	Li, J. X., Rezaee, R., Müller, T. M., & Sarmadivaleh, M. (2021). Pore size
715	distribution controls dynamic permeability. $Geophys. Res. Lett., 48(5),$
716	e2020GL090558.
717	Linde, N., Jougnot, D., Revil, A., Matthäi, S., Arora, T., Renard, D., & Doussan, C.
718	(2007). Streaming current generation in two-phase flow conditions. <i>Geophys.</i>
719	Res. Lett., $34(3)$ .
720	Mavko, G., Mukerji, T., & Dvorkin, J. (2009). The rock physics handbook: Tools for
721	seismic analysis of porous media. Cambridge University Press.
722	Monachesi, L. B., Zyserman, F. I., & Jouniaux, L. (2018). An analytical solution
723	to assess the sh seismoelectric response of the valoes zone. $Geophys. J. Int.,$
724	213(3), 1999-2019.
725	Mualem, Y. (1976). A new model for predicting the hydraulic conductivity of unsat-
726	urated porous media. Water Resour. Res., 12(3), 513–522.
727	Packard, R. G. (1953). Streaming potentials across glass capillaries for sinusoidal
728	pressure. J. Chem. Phys., 21(2), 303-307.

729	Peng, R., Di, B., Glover, P. W. J., Wei, J., Lorinczi, P., Liu, Z., & Li, H. (2020).
730	Seismo-electric conversion in shale: experiment and analytical modelling. Geo-
731	phys. J. Int., 223(2), 725-745.
732	Pride, S. R. (1994). Governing equations for the coupled electromagnetics and
733	acoustics of porous media. Phys. Rev. B, $50(21)$ , 15678.
734	Pride, S. R., & Morgan, F. D. (1991). Electrokinetic dissipation induced by seismic
735	waves. <i>Geophysics</i> , 56(7), 914-925.
736	Reppert, P. M. (2001). Frequency-dependent streaming potentials. J. Colloid and
737	Interf. Sci., 234(1), 194 - 203. doi: 10.1006/jcis.2000.7294
738	Revil, A. (2013). Effective conductivity and permittivity of unsaturated porous ma-
739	terials in the frequency range 1 mHz–1GHz. Water Resour. Res., 49(1), 306-
740	327.
741	Revil, A., & Jardani, A. (2010). Seismoelectric response of heavy oil reservoirs: the-
742	ory and numerical modelling. Geophys. J. Int., 180(2), 781–797.
743	Revil, A., Jardani, A., Sava, P., & Haas, A. (2015). The seismoelectric method: The-
744	ory and applications. John Wiley & Sons.
745	Revil, A., & Mahardika, H. (2013). Coupled hydromechanical and electromag-
746	netic disturbances in unsaturated porous materials. Water Resour. Res., 49,
747	744-766.
748	Rosas-Carbajal, M., Jougnot, D., Rubino, J. G., Monachesi, L., Linde, N., & Hol-
749	liger, K. (2020). Seismoelectric signals produced by mesoscopic heterogeneities:
750	Spectroscopic analysis of fractured media. Seismoelectric Exploration: Theory,
751	Experiments, and Applications, 269–287.
752	Sen, P. N., & Goode, P. A. (1992). Influence of temperature on electrical conductiv-
753	ity on shaly sands. $Geophysics$ , $57(1)$ , 89-96.
754	Smeulders, D., Eggels, R., & Van Dongen, M. (1992). Dynamic permeability: refor-
755	mulation of theory and new experimental and numerical data. J. Fluid Mech.,
756	245, 211-227.
757	Solazzi, S. G., Rubino, J. G., Jougnot, D., & Holliger, K. (2020). Dynamic perme-
758	ability functions for partially saturated porous media. Geophys. J. Int., $221(2)$ ,
759	1182–1189.
760	Soldi, M., Guarracino, L., & Jougnot, D. (2017). A simple hysteretic constitutive
761	model for unsaturated flow. Transport Porous Med., 120(2), 271–285.

- Soldi, M., Guarracino, L., & Jougnot, D. (2019). An analytical effective excess
  charge density model to predict the streaming potential generated by unsaturated flow. *Geophys. J. Int.*, 216(1), 380-394.
- Thanh, L. D., Jougnot, D., Solazzi, S. G., Van Nghia, N., & Van Do, P. (2021). Dy namic streaming potential coupling coefficient in porous media with different
   pore size distributions. *Geophys. J. Int.*. doi: 10.1093/gji/ggab491
- Thompson, A., & Gist, G. (1993). Geophysical applications of electrokinetic conversion. Lead. Edge, 12(12), 1169–1173.
- Tyler, S. W., & Wheatcraft, S. W. (1990). Fractal processes in soil water retention.
   Water Resour. Res., 26(5), 1047–1054.
- Van Genuchten, M. T. (1980). A closed-form equation for predicting the hydraulic
  conductivity of unsaturated soils. Soil. Sci. Soc. Am. J., 44(5), 892–898.
- Wang, J., Hu, H., & Guan, W. (2015). Experimental measurements of seismoelectric
  signals in borehole models. *Geophys. J. Int.*, 203(3), 1937–1945.
- Warden, S., Garambois, S., Jouniaux, L., Brito, D., Sailhac, P., & Bordes, C. (2013).
  Seismoelectric wave propagation numerical modelling in partially saturated
  materials. *Geophys. J. Int.*, 194 (3), 1498–1513.
- Waxman, M. H., & Smits, L. (1968). Electrical conductivities in oil-bearing shaly
  sands. Soc. Pet. Eng. J., 8(02), 107–122.
- Yu, B., Li, J., Li, Z., & Zou, M. (2003). Permeabilities of unsaturated fractal porous
   media. Int. J. Multiph. Flow, 29(10), 1625–1642.
- Zhou, M.-Y., & Sheng, P. (1989). First-principles calculations of dynamic permeability in porous media. *Phys. Rev. B*, 39(16), 12027.
- Zhu, Z., & Toksoz, M. N. (2013). Experimental measurements of the streaming potential and seismoelectric conversion in berea sandstone. *Geophys. Prospect.*, 61, 688-700.
- Zhu, Z., Toksöz, M. N., & Burns, D. R. (2008). Electroseismic and seismoelectric
   measurements of rock samples in a water tank. *Geophysics*, 73(5), E153-E164.
   doi: 10.1190/1.2952570
- Zyserman, F. I., Jouniaux, L., Warden, S., & Garambois, S. (2015). Borehole seis moelectric logging using a shear-wave source: possible application to CO<sub>2</sub>
   disposal? Int. J. Greenh. Gas Control., 33, 89–102.
- <sup>794</sup> Zyserman, F. I., Monachesi, L. B., & Jouniaux, L. (2017). Dependence of shear

-33-

- <sup>795</sup> wave seismoelectrics on soil textures: a numerical study in the vadose zone.
- <sup>796</sup> Geophys. J. Int., 208(2), 918–935.