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Marine Pollution Bulletin

Variations of suspended particulate matter concentrations of the Mackenzie River plume (Beaufort Sea, Arctic Ocean) over the last two decades

--Manuscript Draft--

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Corresponding Author:	Anastasia Tarasenko, PhD FRANCE
First Author:	Anastasia Tarasenko, PhD
Order of Authors:	Anastasia Tarasenko, PhD David Doxaran Bernard Gentili
Abstract:	<p>This work addresses the last 20 years' evolution of the suspended particulate matter (SPM) concentrations in the Beaufort Sea (Canadian Arctic Ocean) directly influenced by the Mackenzie River discharge. The SPM variations in the coastal zone are highlighted and related to the freshwater and solid discharges of the river measured in situ at the Arctic Red River station (150 km upstream of the river delta). The correlation between the variations of the river discharge and SPM concentration within the surface layer of the coastal waters is obvious. Rather unexpectedly, both have been slightly but significantly decreasing from 2003 to 2018-2019 and started to increase very recently (2019-2022). This change of regime could be explained by changing patterns of precipitation (especially in winter), groundwater distribution and wind-induced mixing in the coastal area.</p>
Suggested Reviewers:	Jacek Andrzej Urbanski oceju@univ.gda.pl Claireg Griffin griffin.claireg@gmail.com Joaquín Chaves joaquin.chaves@nasa.gov

We would like to thank the reviewer for his/her work, valuable comments and suggestions. Below we answer to each of them separately.

Reviewers' comments:

“The work is another study concerning the inflow of various substances (SPM, organic carbon) into the Arctic Ocean from the Mackenzie River. Due to distinct trends in air temperature in these regions and the overall ecosystem sensitivity, this topic is of significant importance. The study encompasses recent years (up to 2022), which is crucial. The paper includes all necessary elements and is clearly written. The figures are appropriate, but authors are recommended to include:

boxplots for SPM and annual discharge sums.”

The annual discharge sums were indeed not indicated in text (no number were provided explicitly), but we used them for the Pearson correlation matrix (Fig. 1 in Appendix). We added the annual discharge sums to the **legend box** of Fig.1 for each year (see below).

As for the boxplots for the SPM, we would appreciate to implement it, but unfortunately, could not understand what figure was referred. In case of Fig.1d, there are too few datapoints to add the boxplots, and in case of Fig. 2(b, d, f), there is too many datapoints, as they are daily means (Fig;2.d, f).

“My general comments are listed below.

The conducted reevaluation and use of new regression analysis methods provide a fresh perspective on the phenomena. However, the change in trend based on 3 years raises evident doubts and it might be worth discussing its significance further (especially considering that the only other study pertains to medical data).”

Thank you for this comment.

We have to add that although we do not cite other works using a segmented regression analysis, the method suggested by Muggeo, 2003 was also applied to various geophysical datasets (there are about 1900 citations of this work mentioned at

https://scholar.google.com/scholar?cites=16460629970310398020&as_sdt=2005&scioldt=0,5&hl=f), e.g.:

Zhang, S., Gan, T. Y., Bush, A. B., Liu, J., Zolina, O., & Gelfan, A. (2023). Changes of the streamflow of northern river basins of Siberia and their teleconnections to climate patterns. International Journal of Climatology.

Lee, E., Epstein, J. M., & Cohen, M. J. (2023). Patterns of Wetland Hydrologic Connectivity Across Coastal-Plain Wetlandscapes. Water Resources Research, e2023WR034553.

Börgel, F., Neumann, T., Rooze, J., Radtke, H., Barghorn, L., & Meier, H. E. (2023). Deoxygenation of the Baltic Sea during the last millennium. Frontiers in Marine Science.

Indeed, the positive trends for the last three years are not statistically significant, and the monitoring should be continued. We discussed the significance of trends in section 3, and mentioned that the positive trends for the river discharge are not significant in paragraph #6:

“The positive trend from 2019 to the present, $slope_{2Q} = 640.05x + const$, **although not statistically significant**, can indicate a progressive increase in the minimum flow impacted by a “mobilization of ground waters” as discussed by Yang et al. 2015.”

And in paragraph #7 of section 3 we say that the SPM (in situ and satellite) positive trends are questionable:

“The breakpoint of $SPM_{in situ}$ trends is slightly shifted to 2018, and CIs contain zero, thus indicating that the *segmented* model is **less reliable** for this dataset. <...>

The positive trends $slope_{2sat}$ and $slope_{2sat200}$ **are not statistically significant**, and their SEs are of the order of the estimated coefficient of linear regression (Tab.1). Nevertheless, this result is interesting for further discussion”

We added in the last paragraph of the conclusion an additional indication about statistical significance of the positive trend.

I also have reservations about the presentation of the hypothetical mechanism explaining the observed negative trend in river discharge and SPM concentration, which suddenly switches to a rapid positive trend. Since this is an original aspect of the study, the discussion should be supported by literature describing the mentioned mechanisms.

Thank you for this comment. We discussed the results with our colleagues hydrologists, and indeed changed the optics for the possible explanation of observed trends. Please, see the last three paragraphs of the results and discussion section, together with the updated Conclusions section.

The authors should include a discussion on the dynamics of the coastal zone. For instance, storms could influence the determined concentrations in the coastal area. Perhaps some seasons were more dynamic, while others were less so?

Thank you for this suggestion, we added to the analysis the wind data from ERA5 (marine area similar to that chosen for the MODIS study box).

The hypothesis that the wind stress will induce the vertical mixing and affect the SPM concentration in water is clear and reasonable. Nevertheless, there is one fundamental limitation for this analysis due to the nature of ocean color data retrieval from satellite: the highest wind speed (and thus mixing events) will mostly occur under the cloud cover with the passage of cyclones, and SPM concentrations cannot be retrieved from ocean color satellite under such conditions (presence of clouds).

To include the effect of storms to the Pearson correlation matrix (where the yearly parameters are compared), we included the *number of days of storms per year* as a new parameter (ndays_storm in Fig. 1Appendix). The day was considered stormy if over the study area there was a wind vector with a wind speed over 15 m/s. The negative correlation between days of storm and cumulative annual SPM is weak (-0.19), any conclusion can hardly be done based on this result.

Then we took the *wind speed* (the absolute value of wind vector) for each moment of time, when SPM retrieved from MODIS satellite data was available, and *averaged for the whole study area* both *Uwind_mean* and *SPM_MODIS_mean* (Fig. R2). The linear fit between the spatially averaged SPM and wind speed was low as well: $y = 0.24x + 10.71$ (where x is wind speed and y is SPM concentration).

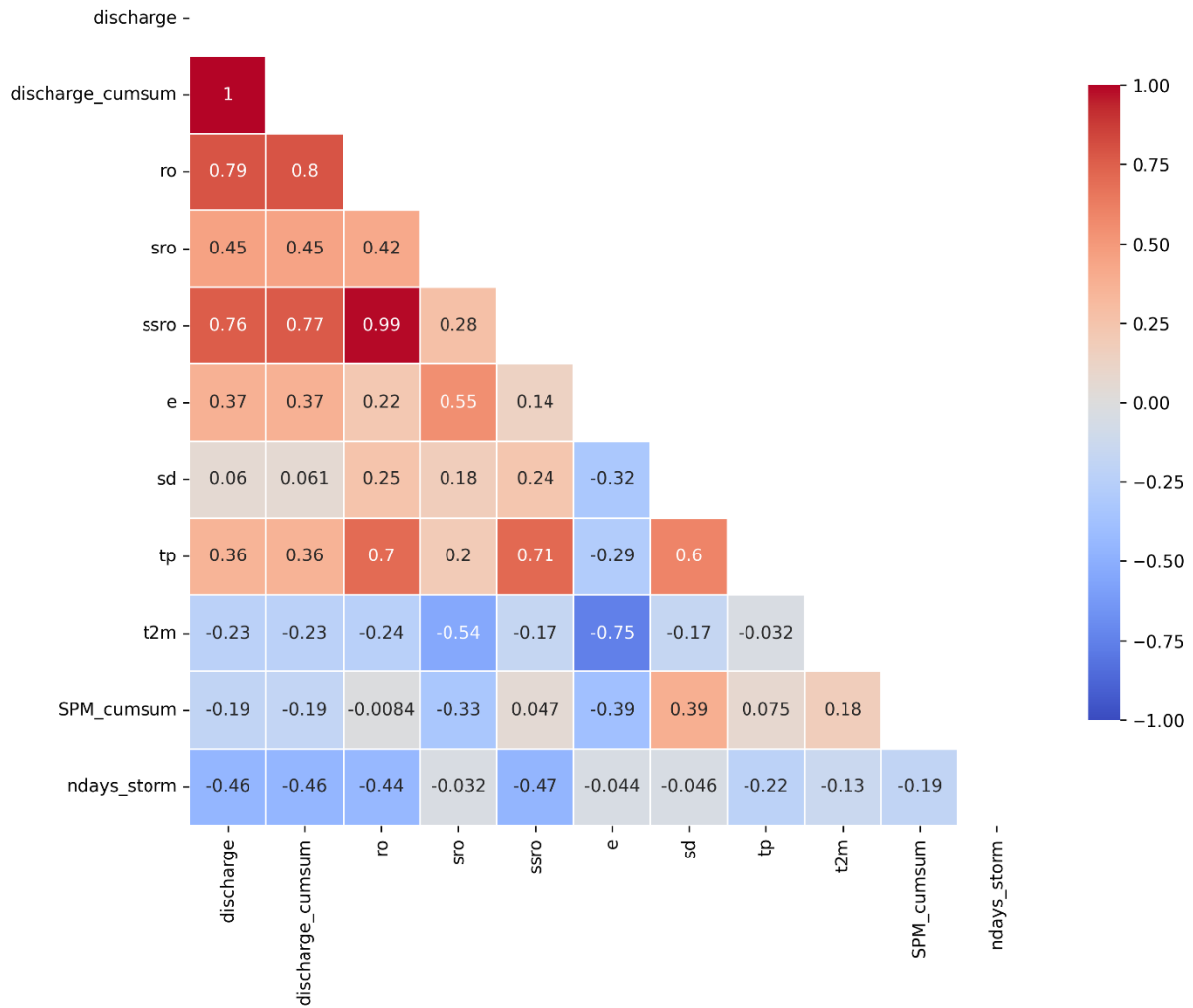


Fig.R1 Pearson correlation matrix with a new parameter referring to wind (ndays_storm).

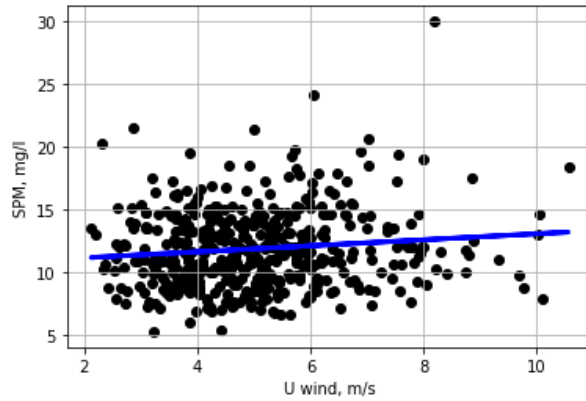
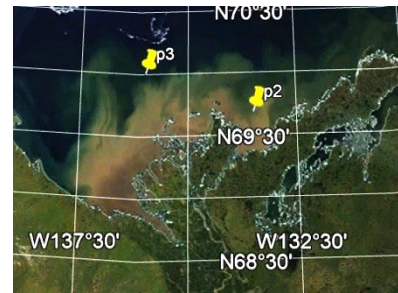


Fig.R2 Wind speed and SPM MODIS averaged over the study area for each moment of time. The blue line shows the fit

Then we tried to apply the segmented correlation method to the wind data over two sample points close to the regular river plume frontal position: $69.7^{\circ}N, -133.5^{\circ}W$ and $70^{\circ}N, 136W$.



The mean wind speed force varies from 8.5 to 10.5 m/s overall, and there is no distinguishable particular pattern (Fig. R3, red line shows the linear fit).

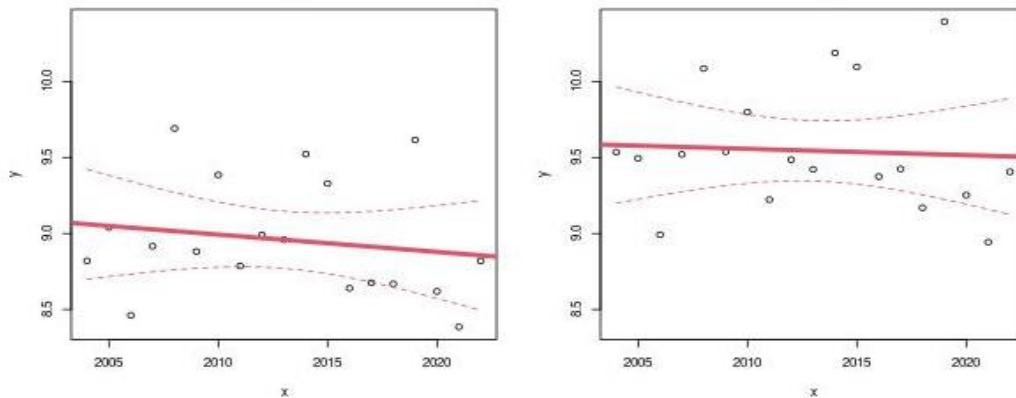


Fig. R3. Regression analysis applied to the wind force in 2 study points (left: $69.7^{\circ}N, -133.5^{\circ}W$, right: $70^{\circ}N, 136W$). The x axis is time, y axis is wind force in m/s

The only parameter that changes slightly over the 2004-2022 period is the wind direction (Fig. R4): According to linear fits, in both points, from 2004 to ~ 2020 the wind direction turned from SWW to SSW (from -100° to -130° from north).

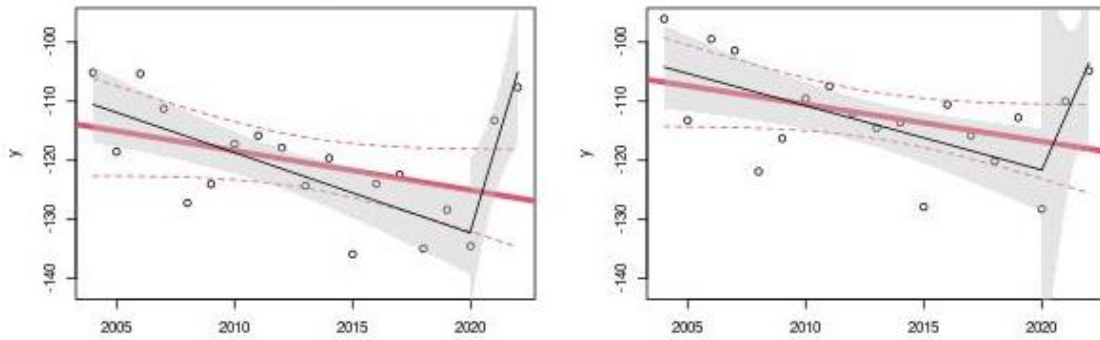


Fig. R4. Regression analysis applied to the wind direction in 2 study points (left: 69.7°N, -133.5°W, right: 70°N, 136°W). The x axis is time, y axis is wind direction in degrees from North. Red line is a linear regression result, black line is segmented regression method



Fig. R5. Illustration of possible effect of wind direction change (yellow and orange arrows), and the Ekman effect (green arrow).

This change in direction suggests rather complex consequences because of the coastal system of currents.

While the SWW wind direction (Fig.R5, left) is changing to SSW (Fig.R5; right), the Ekman effect might:

- bring the river plume into a small coastal vortex in the Mackenzie bay. In this case the SPM has more chances to recirculate within the delta area (higher SPM concentration?), but also be sedimented on the floor (lower SPM concentration)
- bring the SPM further offshore into the westward or eastward system of currents. In both cases, the SPM concentration reduces.

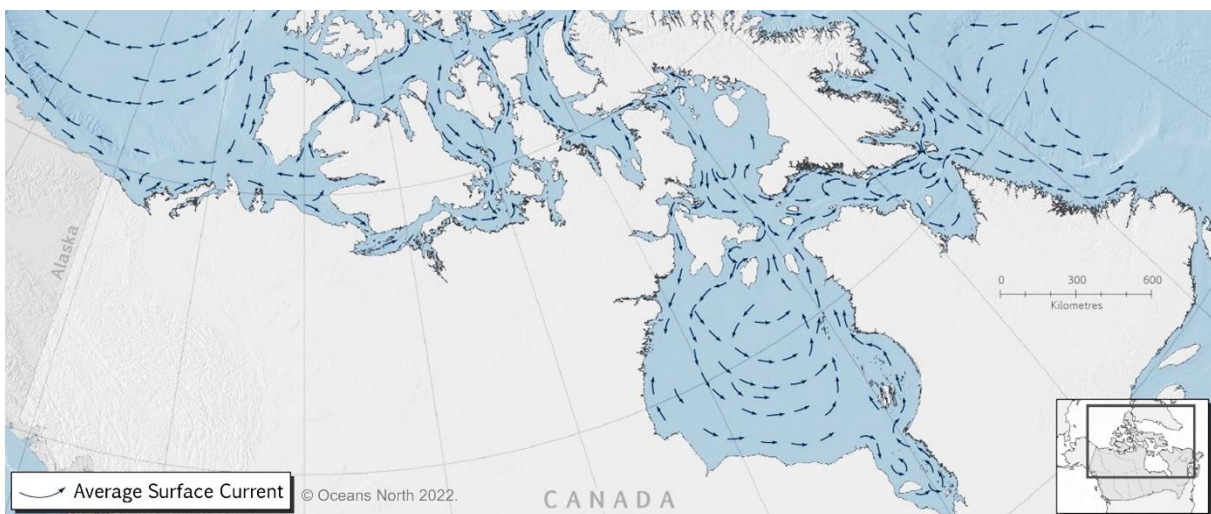


Fig. R6. Surface currents in Canadian Arctic basin (from <https://www.oceansnorth.org/wp-content/uploads/2018/07/en-02-canadas-arctic-marine-atlas-chapter-two-physical-oceanography.pdf>)

After this analysis, we conclude that understanding the impact of wind on the SPM concentration is beyond the focus of this study and should be investigated additionally.

We included the updated Pearson correlated matrix in the Appendix, and added this resumé to the methods and discussion sections:

Section 2:

“To have idea of the large-scale estimates of the wind impact on the annual cumulated SPM, we included the number of days of storms parameter.

Section 3:

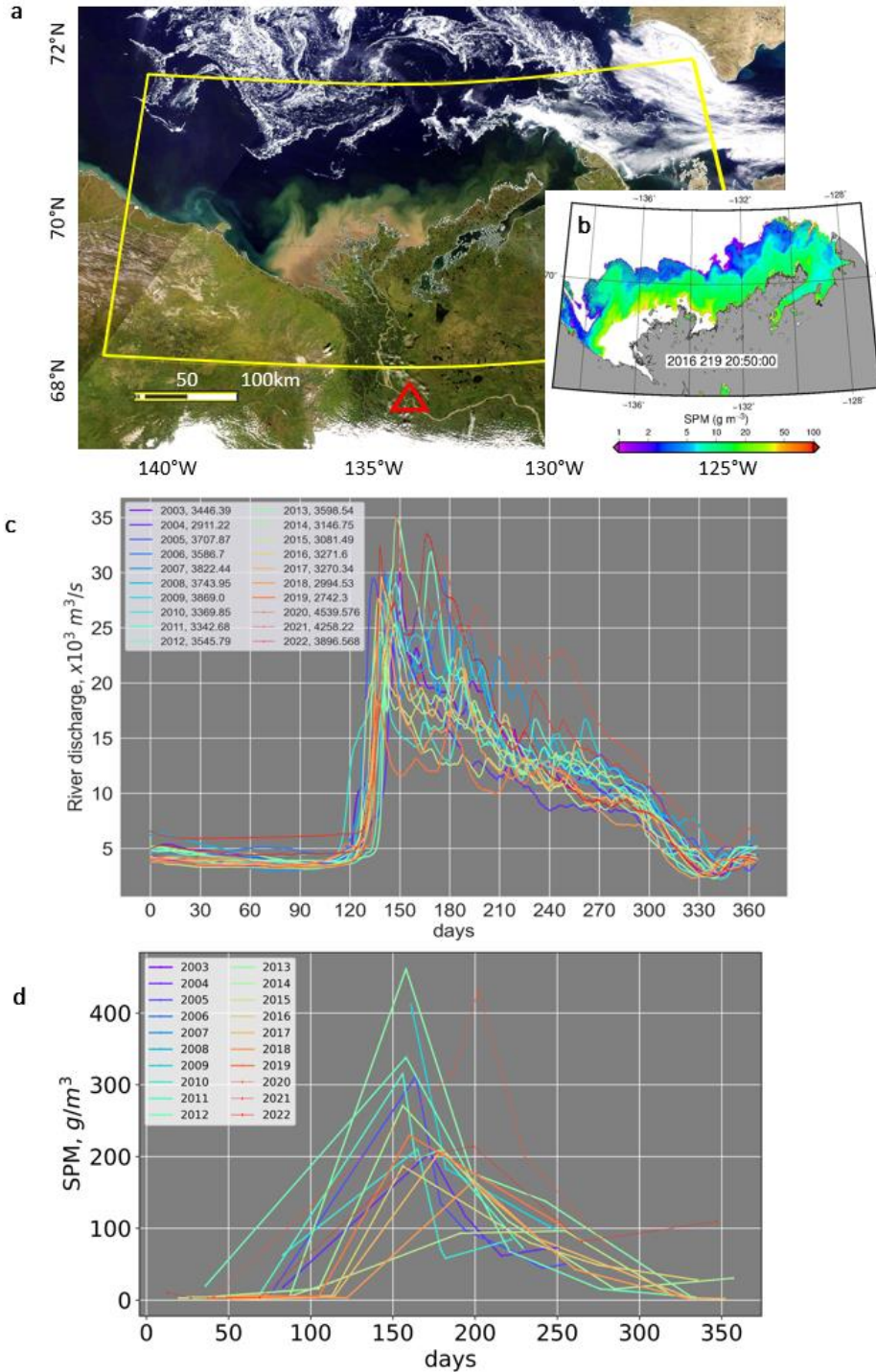
“Another question that was not addressed in our study is the impact of the wind-induced mixing on the offshore SPM concentration. The simple hypothesis proposes that the higher is the wind speed, the more mixing occurs in a shallow area, which reduces the surface SPM concentrations. At the same time, an additional hypothesis may suggest that more mixing means more re-suspension of particulate matter from the bottom sediments (increasing the SPM concentration). To verify these controversial hypotheses, we have to compare quasi simultaneous wind and SPM observations. In case of this study, this problem faces a fundamental limitation due to the nature of a satellite-retrieved SPM dataset: the highest wind speeds are observed during the cyclone passage, and SPM concentrations cannot be retrieved from satellite data under the clouds (as clouds mask the water-leaving signal).

We found a weak negative correlation (correlation coefficient is -0.19) between the number of days of storm and annual cumulated SPM, which confirms the first simple hypothesis (more mixing, less SPM), but this question should be addressed additionally with other tools, like modeling.”

In conclusion, the issue is quite complex, and the statistical significance of the (linear) regressions is not very strong. As for the detailed comments:

1. Fig. 1: The map and satellite image should have a scale.

Thank you, we added a scale in Fig.1a:



2. The link to ERA5-LAND is incorrect (it should end with "...documentation").

Thank you for your suggestion, the new link was changed to <https://confluence.ecmwf.int/display/CKB/ERA5%3A+data+documentation> and <https://confluence.ecmwf.int/display/CKB/ERA5-Land%3A+data+document>

We were following the instruction found at <https://confluence.ecmwf.int/display/CKB/How+to+acknowledge+and+cite+a+Climate+Data+Store+%28CDS%29+catalogue+entry+and+the+data+published+as+part+of+it>

3. Fig. 2: Subfigures c and e, as well as d and f, have the same subtitle, but they describe different things. This is explained in the figure caption, but it's a bit confusing - please revise.

The titles were changes, thank you for your suggestions.

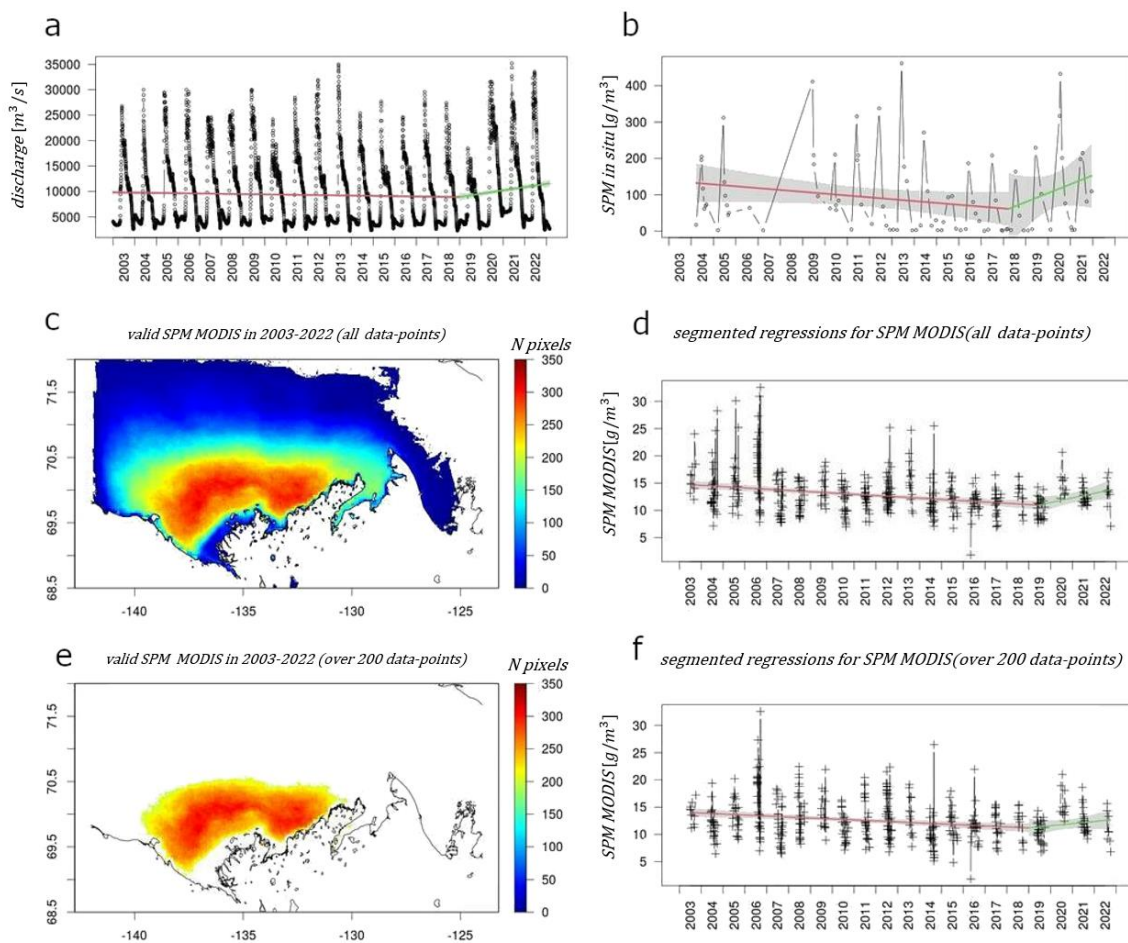


Fig.2

Highlights

Variations of suspended particulate matter concentrations of the Mackenzie River plume (Beaufort Sea, Arctic Ocean) over the last two decades

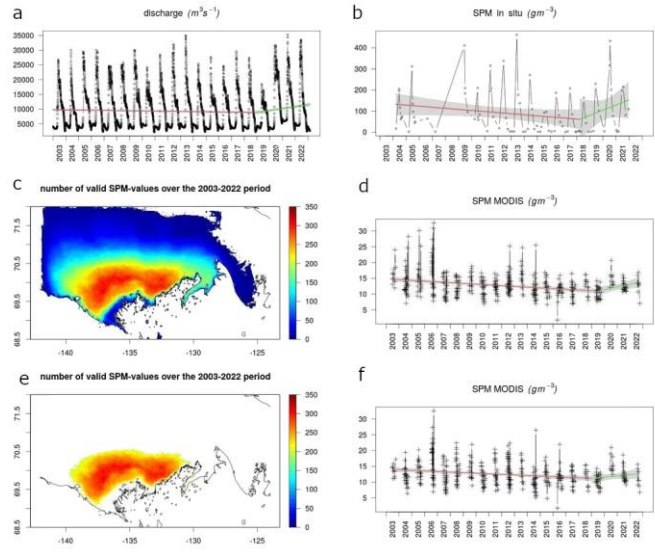
Anastasia Tarasenko, David Doxaran, Bernard Gentili

- Interannual variations of suspended particulate matter (SPM) in coastal waters influenced by the Mackenzie River were investigated over the last 20 years (2003-2022) using MODIS/Aqua satellite data
- The offshore SPM variations are related to the Mackenzie River hydrological regime, air temperature, amount of snow, and permafrost state over its draining basin
- Over the last twenty years, a statistically significant negative trend was highlighted over the period 2003-2018 for both SPM and river discharge, with a positive trend starting from 2019 up to present

Graphical Abstract

Variations of suspended particulate matter concentrations of the Mackenzie River plume (Beaufort Sea, Arctic Ocean) over the last two decades

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Variations of suspended particulate matter concentrations of the Mackenzie River plume (Beaufort Sea, Arctic Ocean) over the last two decades

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Abstract

This work addresses the last 20 years' evolution of the suspended particulate matter (SPM) concentrations in the Beaufort Sea (Canadian Arctic Ocean) directly influenced by the Mackenzie River discharge. The SPM variations in the coastal zone are highlighted and related to the freshwater and solid discharges of the river measured in situ at the Arctic Red River station (150 km upstream of the river delta). The correlation between the variations of the river discharge and SPM concentration within the surface layer of the coastal waters is obvious. Rather unexpectedly, both have been slightly but significantly decreasing from 2003 to 2018-2019 and started to increase very recently (2019-2022). This change of regime could be explained by changing winter precipitation and groundwater distribution, progressively accumulating sediments within the thawing permafrost layer and its recent release into the groundwater together with thermokarst lakes' rapid drainage.

Keywords: suspended particulate matter, ocean optics, Arctic Ocean, MODIS

1. Introduction

Climate change occurs faster in polar regions than at lower latitudes Arias et al. (2021). Global warming is usually associated in high latitudes with rising air temperature, precipitations, and permafrost thaw Miner et al. (2022); Pörtner et al. (2022). Consequently, the freshwater discharged by rivers into the Arctic Ocean is expected to increase, so as the discharge of terrestrial substances with enhanced erosion along drainage basins Doxaran et al. (2015); Matsuoka et al. (2022); Juhls et al. (2022). This would result in increasing water turbidity in coastal areas directly affected by river inputs but also boosted primary production due to higher nutrient loads and reduced sea ice cover (i.e., increasing solar light within the water column).

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4 The Mackenzie River is the fourth largest Arctic river in terms of river discharge (7% of
5 freshwater inflow to the Arctic Ocean) and is the primary source of sediment discharge
6 Carson et al. (1998); Wagner et al. (2011); Yang et al. (2015), so the changes in the
7 Mackenzie River regime will have a significant impact on the whole Arctic region Juhls et al.
8 (2022). The Mackenzie has a prominent seasonal cycle with winter lows and a summer
9 maximum of river discharge, which is related to the water cycle over its large basin with its
10 75% permafrost area and the importance of snowmelt in spring Yang et al. (2015); Grotheer
11 et al. (2020). The delta of the Mackenzie River is a complex area where the freshwater
12 massively inflows into the southern Beaufort Sea, highly impacted by a long presence of sea
13 ice. The southern Beaufort Sea is frozen most part of the year; the polynia between the fast
14 ice of the delta starts opening in May, then the water surface stays mostly icefree from July
15 to October. Over the last two decades, the sea ice conditions have become softer, and the sea
16 ice concentration (SIC) has diminished by -4-8% per decade from 2003 to 2019 Hilborn
17 and Devred (2022).

23 The open discussion of the Mackenzie River regime and its adaptation to climate change
24 depends on the studied period Woo and Thorne (2003); Doxaran et al. (2015); Yang et al.
25 (2015); Matsuoka et al. (2022); Zolkos et al. (2022). Doxaran et al. 2015 claimed a 22%
26 increase of the river discharge from 2003 to 2013. Yang et al. 2015 showed that over the
27 1973-2011 period, the positive linear trend for the Mackenzie discharge was very weak ($y =$
28 $17.21x + 8947.4$), and Woo and Thorne 2003 found no significant trend for the annual
29 discharge over the period 1968-1999, nor did Matsuoka et al. 2022 for the 1992-2018
30 period. Meanwhile, several studies agreed that the yearly amount of river discharge has
31 changed its seasonal distribution: cold season discharge becomes slightly higher, but spring
32 flows are lower (as higher air temperatures shift the snowmelt earlier) Woo and Thorne
33 (2003); Yang et al. (2015). The question of the river water discharge, thus, should be
34 regularly reevaluated.

39 In such a remote and changing environment, satellite observations have been proven to
40 be an efficient tool to monitor the evolution of Arctic coastal zones, and compensate the
41 lack of field measurements Doxaran et al. (2012); Hill et al. (2013); Juhls et al. (2022). SPM
42 field measurements, although extremely valuable, are usually too sparse for the long-term
43 variability analysis, as most of them are associated with specific summer expeditions (e.g.
44 the recent MALINA Massicotte et al. (2021) and Nunataryuk Lizotte et al. (2023) field
45 campaigns). Overall, during the ice-free season, satellite-derived suspended particulate
46 matter (SPM) concentrations typically vary from 30 g/m^3 in the delta region to 0.5 g/m^3
47 offshore over the Canada Basin Hilborn and Devred (2022).

51 Several studies analyzed satellite-derived SPM concentration variability in the southern
52 Beaufort Sea for long time periods (over 10 years), and their results are rather
53 controversial. From 2003 to 2013, Doxaran et al. 2015 found a linear trend with a significant
54 increase in monthly- averaged SPM concentrations at the mouth of Mackenzie River (+46%
55 in 10 years over the river mouth area and +71% in the river delta south to 70°N). Hilborn
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4 and Devred 2022 applied *self-organizing maps* method to identify six different regions in
5 the eastern Beaufort Sea using 17 years of MODIS (Moderate Resolution Imaging
6 Spectroradiometer) data and found only one statistically significant trend (negative) for the
7 annual SPM concentration in the deepwater Canada basin area, far offshore the Mackenzie
8 delta. At the same time, Matsuoka et al. 2022 showed a statistically significant increase of
9 dissolved and particulate organic matter concentrations (but not fluxes) in late summer:
10 0.019 g/m^3 per year and 0.069 g/m^3 per year, accordingly).

11 The present study is a continuation of Doxaran et al.2015, where we analyze the
12 evolution of the SPM concentration in the Beaufort Sea directly influenced by the discharge
13 of the Mackenzie River over the last two decades (2003-2022), using the methodology
14 developed by Doxaran et al. (2012, 2015), this time excluding the complex river delta zone.
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21 2. Materials and Methods

22 Fig.1a shows the limits of the study area (yellow rectangular) on a quasitrue color daily
23 MODIS composite for August 8, 2016. In this image, the SPM-rich Mackenzie River plume
24 propagates along the coast and northward, partly into the MIZ (marginal ice zone). The
25 variations of SPM concentrations (Fig.1b) from the delta to offshore waters are obvious,
26 which highlights complex processes along this land-sea interface. The field and satellite
27 datasets used in this study, so as the methods used to retrieve and analyze time series of
28 river discharge and SPM concentrations, are detailed hereafter.
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33 2.1. Field data sets

34 In situ measurements of the river water discharge (2003-2022) were provided by the
35 Water Survey of Canada via the ArcticGRO project McClelland et al. (2023) at the Arctic Red
36 River gauge station (station ID: 10LC014, 67.45°N 133.74°W, red triangle in Fig.1a). River
37 discharge measurements, Q , are available daily (Fig.1c). The measurements of suspended
38 matter are distributed by the ArcticGRO as *total suspended solids (TSS)* concentrations,
39 which is the same as SPM, so hereafter we will refer to in situ TSS as the
40 SPM_{insitu} . The SPM_{insitu} data were collected at most once per month, but not during all months
41 (on average, 4 to 6 measurements per year, Fig.1d). Over the studied period, some years are
42 poorly represented (one measurement per year in 2003, 2006, and 2007; no SPM
43 measurement in 2008). The lack of SPM in situ data implies the necessity to use satellite
44 data.
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50 The ArcticGRO in situ datasets pass the quality control procedure and indicate less
51 reliable measurements as "provisional data". Over the 2003-2022 period, it corresponds to
52 river discharge data in 2020-2022 and SPM_{insitu} data in 2019-2022 (shown with thinner
53 lines in Fig.1c, d).
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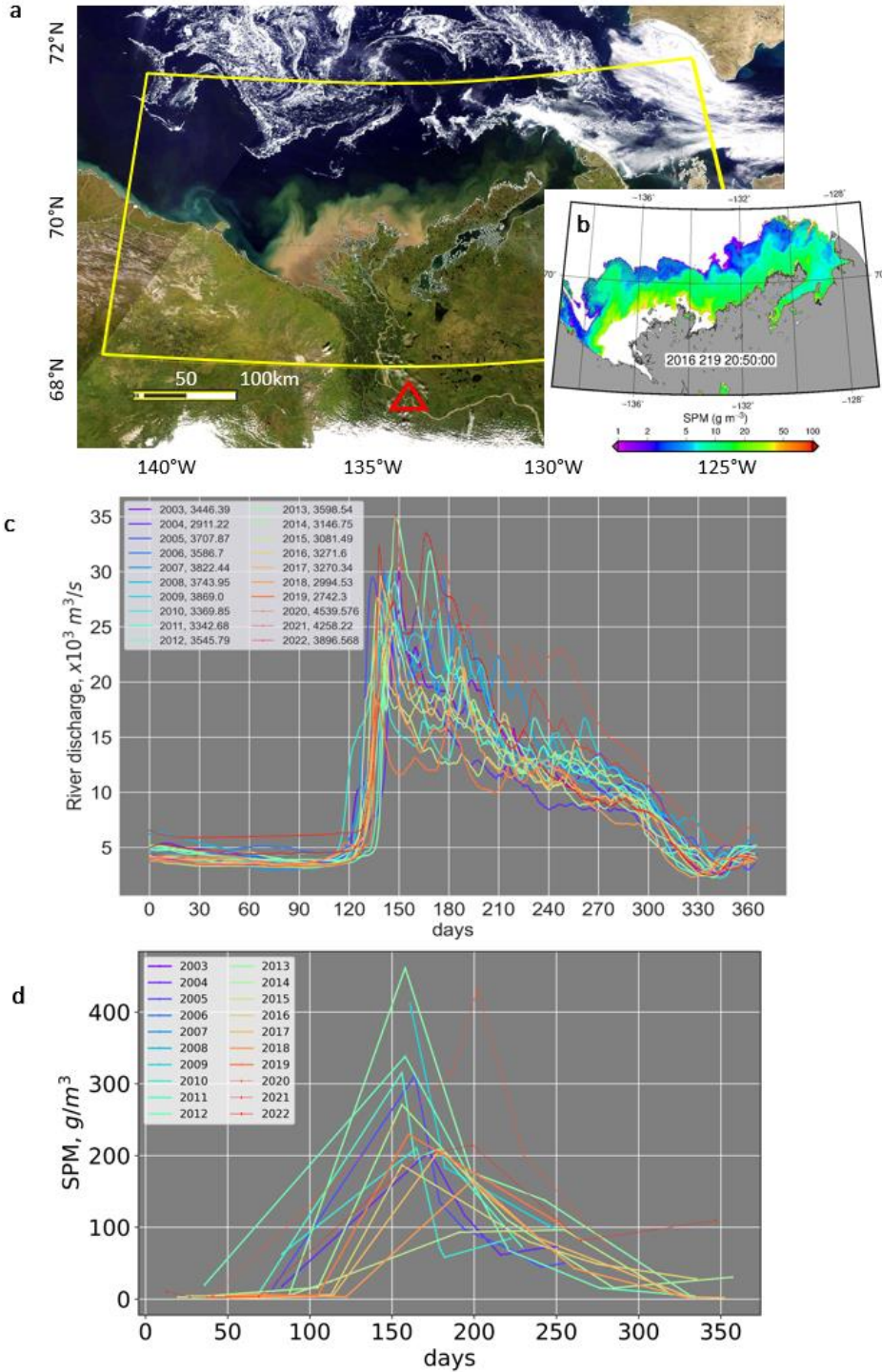


Figure 1: (a) MODIS image (August 6, 2016) locating the Arctic Red River gauge station (Tsiigehtchic) where in situ measurements are carried out (red triangle) and the coastal waters studied using satellite data (yellow rectangle); (b) SPM concentrations: Interannual variations of water discharge (the year of measurement and the annual cumulative sum of discharge is indicated in legend), (c) and SPM concentrations (d) at the Arctic Red River station. The inserted SPM color map shows the result of MODIS satellite data processing

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7 *2.2. Ocean color satellite data*

8 As in Doxaran et al. 2015, a single satellite sensor (MODIS/Aqua) was considered for
9 this study in order to avoid any bias between different sensor products when detecting
10 temporal trends. Since 2003, MODIS provides high-quality ocean color observations at a
11 good temporal resolution: several images per day of the study area during the Arctic
12 summer months Doxaran et al. (2012, 2015); Hilborn and Devred (2022); Matsuoka et al.
13 (2022). The region of interest was defined as 68.5°N-72°N, 142°W - 124°W (yellow
14 rectangle in Fig.1a). The initial dataset contained 10608 swaths collected between May 1
15 and October 31 each year over the 2003-2022 period.
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18 *Processing of satellite data.* MODIS L1A satellite data were processed using the SeaWiFS
19 Data Analysis System (SeaDAS 8.1.0) software (<http://seadas.gsfc.nasa.gov/>) and its l2gen
20 function. The atmospheric correction was performed using the NIR-SWIR algorithm of
21 Wang and Shi 2007); this method was proved to be the most appropriate for the highly
22 turbid waters at the mouth of the Mackenzie River to the less turbid waters offshore
23 Doxaran et al. (2012).
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26 SeaDAS l2gen flags were used to mask clouds and glint. In the work of Doxaran et al.
27 2015, two techniques were used to mask clouds: the default one in the coastal waters and
28 an increased cloud albedo threshold value (0.4 instead of 0.027) over the specific area of
29 the river delta zone to avoid masking the highly turbid water pixels. This procedure was
30 time-consuming and required an additional inspection of every image in the river delta
31 zone. Here, the cloud-masking method preconized by Ody et al. 2022 for river mouths was
32 used to process each satellite image only once: the 2130 nm shortwave-infrared waveband
33 was used with a cloud threshold of 0.018. The sea ice mask was computed from daily sea
34 ice concentration (SIC) data at 3.125°spatial resolution distributed by the University of
35 Bremen Spreen et al. (2008). This dataset contains AMRS-E and AMSR2 (Advanced
36 Microwave Scanning Radiometer -for Eos and -2) images. Due to several gaps in data
37 acquisition over the 2003-2022 time period, the dataset was completed with the sea ice
38 concentration from the SSMIS (Special Sensor Microwave Imager/Sounder) instrument.
39 The ice mask was created using the SIC above 0%.
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42 All satellite data (Rayleigh-corrected reflectances and masks) were reprojected on a
43 regular 250 m grid. The following criteria were then applied to exclude potentially
44 contaminated images:
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- 46 • the number of valid pixels should exceed 15% of the marine area of the study area in
47 the downloaded MODIS image;
 - 48 • the distance between the central point of the study area and that of the MODIS image
49 should be less than 1000 km to avoid aberrations on the border of the image;
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- only areas with a zenith solar angle lower than 74° were conserved.

This automatic filtering reduced the MODIS collection to 700 images. Finally, all images were inspected visually to filter out the remaining contaminated images because of not detected MIZ, clouds or cloud shadows, etc.; 651 MODIS images remained after this last quality check step.

The SPM concentrations were computed from the ratio (in %) between the remote sensing reflectance signals, R_{rs} (in sr^{-1}) in the near-infrared and green wavebands using the set of relationships established based on field measurements Doxaran et al. (2012, 2015):

$$SPM_{sat} = 0.8386 \times R_{RS}(748 : 555) \quad (1)$$

if $R_{RS}(748 : 555) < 87\%$

$$SPM_{sat} = 70 + 0.1416 \times R_{RS}(748 : 555) + 2.9541 \times e^{0.2092 \times (R_{RS}(748:555) - 87)} \quad (2) \text{ if } 87\% \leq R_{RS}(748 : 555) \leq 94\%$$

$$SPM_{sat} = 3.922 \times R_{RS}(748 : 555) - 285.4 \quad (3)$$

if $94\% < R_{RS}(748 : 555)$, where SPM_{sat} is the SPM concentration in g/m^3 and $R_{RS}(748 : 555)$ is the ratio of remote-sensing reflectances at 748 and 555 nm.

Negative $R_{RS}(555)$ and $R_{RS}(748)$ values were discarded before computing SPM, as well as computed SPM concentrations higher than $1000 g/m^3$, to remove atmospheric correction failures and residual contaminations (typically encountered along borders of clouds or sea ice).

The processed SPM images were averaged as daily, monthly, and yearly composites as simple mean averages for every pixel. Only daily composites and their mean values were used further, as monthly composites were sometimes computed using only 1 to 3 cloud-free images, which is not representative of mean monthly concentrations.

2.3. Reanalysis data

A dataset from ERA5 Land reanalysis was extracted to better understand how environmental factors did impact the Mackenzie River discharge and SPM concentrations in the adjacent coastal waters. This dataset was selected as the most suitable tool for the studies of river discharge variability Winkelbauer et al. (2022). ERA5 Land has a monthly temporal resolution and 9-km spatial resolution and is provided by Copernicus data center Munõz-Sabater et al. (2021). The area of extraction approximately corresponds to the Mackenzie drainage basin (52-70°N, 100-140°W). The following parameters were used: river runoff (ro), surface river runoff (sro), evaporation (e), total precipitation (tp), snow depth (sd), and air temperature at 2m (t2m), (see detailed description at <https://confluence.ecmwf.int/display/CKB/ERA5-Land%3A+data+document> The

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4 Pearson correlation matrix for the mean annual values of all described parameters
5 (Fig.A.3).
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7 To include a possible effect of storms on the SPM concentrations, we added to the
8 analysis the wind data from ERA5 (marine area similar to that chosen for the MODIS study
9 box). The Pearson correlation matrix (where the yearly parameters are compared) contains
10 an additional parameter "the number of days of storms per year" (*ndays_storm* in Fig. A.3).
11 The day was considered stormy if over the study area there was a wind vector with a wind
12 speed over 15 m/s.
13

14 2.4. Temporal variability

15 As previous works have shown, the linear regression model provides relatively little
16 information on interannual variability of river water discharge and related parameters,
17 such as SPM, because, over the long term, the discharge appears as stable Yang et al. (2015);
18 Matsuoka et al. (2022). Over the last 20 years, the linear trend for the in situ Mackenzie
19 discharge time series calculated with the Ordinary Least Squares (OLS) model is $y =$
20 $17.1x + 9554$ with a standard error (SE) of $63.6 \text{ m}^3/\text{s}$, and confidence intervals (CI) [0.025
21 0.975] equal to -107.5 and $141.7 \text{ m}^3/\text{s}$ for the coefficient term. Although these results are
22 very close to that of Yang et al. 2015, the statistical metrics confirm that the OLS model is
23 not suitable for the analysis in this case.
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25 For this reason, we used an estimating regression model with unknown break-points,
26 which allows describing several temporal trends Muggeo (2003). To apply this model to
27 our datasets, we used the *segmented* function in the R package *segmented*.
28

29 To compute the trends, the in situ data time series of daily river discharge (Q_{insitu}) and
30 SPM concentration (SPM_{insitu}) described above were used (Fig.2a, b). For the satellite data,
31 the spatial sparsity of observations was taken into account. Thus, to obtain SPM_{sat} time
32 series, we calculated: • one daily spatial mean SPM value for all available SPM pixels over
33 the study area, which resulted in SPM_{sat} time series (Fig.2 c-d); • one daily spatial mean SPM
34 value for the area with the highest data density (where valid satellite pixels appear more
35 than 200 times over the 2003-2022 period), which resulted in SPM_{sat200} (Fig.2 e-f)
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45 Then the segmented regression analysis technique was applied to all time series: Q_{insitu} ,
46 SPM_{insitu} , SPM_{sat} , and SPM_{sat200} . After several tests, the most statistically significant results
47 were obtained with one breaking point and two segment slopes. The following statistical
48 parameters were computed for each segment: estimated coefficient of linear regression
49 (Est.), standard error, lower and upper 95% confidence intervals (CI(95%).l, CI(95%).u,
50 respectively) (Tab.1).
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3. Results and Discussion

The seasonal cycles of both the river discharge, Q , and SPM_{insitu} are very pronounced, with high summer and low winter values (Fig.1c,d). These parameters also have a strong interannual variability. During the 2003-2022 period, the river discharge varied from $2.2 \times 10^3 \text{ m}^3/\text{s}$ (winter) to $35.2 \times 10^3 \text{ m}^3/\text{s}$ (summer), with a mean value of $9.5 \pm 6.8 \times 10^3 \text{ m}^3/\text{s}$. Typically, the discharge increases very rapidly in 2 weeks from 5 to $25 \times 10^3 \text{ m}^3/\text{s}$ in late May, when the main summer peak occurs, then slowly decreases to its winter values by November.

The years 2013 and 2021 were exceptional with extreme *yearly maxima* of river discharge over $35 \times 10^3 \text{ m}^3/\text{s}$ on May 28 (both), while the lowest yearly maximum river discharge, $18 \times 10^3 \text{ m}^3/\text{s}$, was registered on May 18, 2019. Previously, Yang et al. 2015 also observed similar prominent interannual variations in the river discharge during the summer (e.g., in 1992 and 1995 for the 1972-2011 period). They concluded that a negative anomaly in precipitation-evaporation balance over the river basin in summer (hot and dry weather) usually results in a lower discharge the next year with an earlier maximum, while the opposite (cold and wet) weather is responsible for a higher discharge. This statement does not explain the extreme peak of discharge recorded in 2013: in 2012 the summer was “hot and dry”, so 2013 should have been a year of extremely low summer river discharge. Overall, calculated correlations between in situ river discharge and ERA5 Land total precipitation, as well as air temperature were weak (0.36 and -0.23, respectively). However, the river discharge is well correlated with snow depth cover (correlation coefficient between Q_{insitu} and sd is 0.67): the 2011-12 and 2012-13 winters were snowy (+10% and +3% sd anomaly compared to 20years median values), as well as the 2020 winter (+7% sd anomaly), but 2018-19 snow cover was weak (-7% sd anomaly). Interestingly, in 2013 the yearly means of river discharge and air temperature were close to their 20years median values, indicating the overall stability of the system during this period.

The river discharge regime in 2006-2008 and 2012 was special with two summer Q maxima, the first one in late May and the second between the end of June and mid-July. The reason for this change in river regime might be related to the precipitation seasonal pattern and the river ice opening (several ice jams crushes, creating the second peak). To investigate these particularities, a higher temporal resolution reanalysis data should be used. We also observe local maxima (values higher than the previous year) in SPM_{sat} in 2006, 2008, and 2012.

As explained in section 2.4, the OSL model does not highlight a statistically significant trend in the river discharge time series, but the *segmented* model does. Fig.2 and Table 1 present the results of calculated segmented regressions, where the river discharge and SPM concentrations demonstrate very similar trends: a negative trend from 2003 to 2018, then a recent positive trend from 2019 (2018 for SPM_{insitu}) to 2022 (Fig.2). Based on the

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4 confidence intervals, we conclude that negative trends for $slope1_Q$, $slope1_{SPM_{sat}}$, and
5 $slope1_{SPM_{sat200}}$ are statistically significant (CIs [0.025 0.975] are all negative).
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7 This negative trend of $slope1_Q = -57.21x + const$ in river discharge for the 2003-2018
8 period is opposite to that of Doxaran et al. 2015 for the 2003-2013 period, but the latter
9 study slightly overestimated the river discharge in 2013 (as the data was not yet fully
10 quality-checked) resulting in an erroneous positive (+22%) trend. As already mentioned,
11 other studies, e.g. Yang et al. (2015); Matsuoka et al. (2022) which used an OSL model did
12 not find any significant trend in the Mackenzie River discharge. The positive trend from
13 2019 to the present, $slope2_Q = 640.05x + const$, although not statistically significant, can
14 indicate a progressive increase in the minimum flow impacted by a "mobilization of ground
15 waters" as discussed by Yang et al. 2015.
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19 In situ data show that the SPM seasonal cycle generally follows the river discharge cycle,
20 with the summer maxima in late May - beginning of June. The SPM_{sat} values vary from 0.8
21 to 461 g/m^3 with a median value of $67.7 \pm 11 \text{ g/m}^3$. The breakpoint of SPM_{isitu} trends is
22 slightly shifted to 2018, and CIs contain zero, thus indicating that the *segmented* model is
23 less reliable for this dataset. At the same time, the SPM_{isitu} time series has the fewest amount
24 of data, which is not homogeneous in time and is mainly available in summer. It makes it
25 more difficult to interpret with any statistical model. SPM_{sat} and SPM_{sat200} trend analyses
26 show similar results: their negative $slope1$ values and corresponding SE are close to each
27 other, which gives confidence in the observed SPM negative trend. The positive trends
28 $slope2_{sat}$ and $slope2_{sat200}$ are not statistically significant, and their SEs are of the order of the
29 estimated coefficient of linear regression (Tab.1). Nevertheless, this result is interesting for
30 further discussion.
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35 Over the 2003-2019 period, Hilborn and Devred 2022 obtained results in good
36 agreement with a negative trend (although not significant) for SPM concentrations over the
37 southern Beaufort Sea (from -0.16 to -0.46 g/m^3 per decade) except for the Mackenzie River
38 delta, where they found an increase of SPM of 0.36 g/m^3 per decade. The difference in
39 negative coefficients' values obtained in the present study and that of Hilborn and Devred
40 (2022) comes from (1) the difference of SPM-retrieval methods and different MODIS bands
41 used; and (2) the regions: our study region of SPM_{sat200} corresponds to 3 regions over the
42 shelf identified by Hilborn and Devred (2022).
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46 In the Mackenzie delta zone, Doxaran et al. 2015 also found a significant increase of SPM
47 (+71% from 2003 to 2013) over the "Mackenzie River delta" region of Hilborn and Devred
48 (2022). The delta zone seems to play an important role of "SPM filter" between the river
49 and the coastal waters. Based on satellite observations, SPM apparently settles massively
50 in this shallow area, resulting in the formation of temporary maximum turbidity zones
51 where resuspension of bottom sediments may occur depending on the river discharge, tidal
52 currents, and wind stress (Wegner et al., 2005; Grotheer et al., 2020). Observations at high
53 spatial and temporal resolutions are required to further investigate SPM dynamics in this
54 delta zone.
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4 However, how to explain the negative trends in both river discharge and SPM
5 concentrations in the riverbed and offshore delta zone over the 2003-2018 period? Based
6 on the correlation matrix (Fig. A.3), we recognize that the river runoff and, thus, river
7 discharge variability depend mostly on the amount of total precipitation ($r = 0.7$) and the
8 snow depth ($r = 0.63$, which is another estimate of solid precipitation in winter)
9 parameters. This trivial conclusion leads us to a simple suggestion that the Mackenzie River
10 discharge slightly declined while the precipitation pattern has changed over this period.

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13 As for the SPM, the recent study of Zolkos et al. 2022 reveals a similar decrease in SPM
14 loads in most of the Siberian rivers between 1970 and 2010 and explains it by natural and
15 anthropogenic factors. The "natural factors" of sediments erosion over the river basin are
16 the physical and chemical denudation. For the Mackenzie River, the physical denudation
17 rates exceed the chemical denudation about several orders of time: mechanical denudation
18 rate at the Arctic Great River was reported as $844 \text{ t} \cdot \text{km}^{-2} \cdot \text{y}^{-1}$ in 1997, while the
19 chemical weathering occurs at a rate of $25 \text{ mm} \cdot \text{ky}^{-1}$ with a transport time of $10\text{-}400 \text{ ky}^{-1}$
20 (Vigier et al. (2001); DePaolo et al. (2006)). On the order of two decades, it is, thus, physical
21 or mechanical factors related to the river discharge and the atmospheric conditions in the
22 delta-adjacent areas that control SPM concentrations. A simultaneous decline of Q_{insitu}
23 and SPM_{insitu} in 2003–2018 means that the lower the river discharge, the less suspended
24 particulate matter it transports.

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29 Another source of variability for the "marine" SPM might be the wind mixing. The simple
30 hypothesis proposes that the higher the wind speed, the more mixing occurs in a shallow
31 area, which reduces the surface SPM concentrations. At the same time, an additional
32 hypothesis may suggest that more mixing means more re-suspension of particulate matter
33 from the bottom sediments (increasing the SPM concentration). To verify these
34 controversial hypotheses, we must compare quasi simultaneous wind and SPM
35 observations. In the framework of this study there is yet one fundamental limitation for
36 this analysis due to the nature of ocean color data retrieval from satellite: the highest wind
37 speed (and thus mixing events) will mostly occur under the cloud cover with the passage
38 of cyclones, and SPM concentrations cannot be retrieved from ocean color satellite under
39 such conditions (presence of clouds). We found a weak negative correlation (correlation
40 coefficient is $r = -0.19$) between the number of days of storm and annual cumulated SPM,
41 which confirms the first simple hypothesis (more mixing, less SPM), but this question
42 should be addressed additionally with other tools, like modeling.

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48 There is yet a question about the recent positive trends in Q and SPM over the last 3-5
49 years of observations. Although the trends are statistically not significant, can we suggest
50 any important processes that might have impacted the river discharge and sedimentation
51 transport recently?

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53 The presence of permafrost over the river basin was considered as "precluding" for the
54 groundwater's contribution to general river discharge, as it prevented the water infiltration
55 through the permafrost layer, but might have helped to create underground cavities filled
56 with non-communication water reservoirs (Vigier et al. (2001)). With a progressive

permafrost thawing, this neglected role might be re-evaluated. Matsuoka et al. 2022 observed an increase of thaw depth and precipitation, but a decrease in river discharge, which is probably explained by the contribution of ground waters, also suggested by Connolly et al. (2020). The permafrost thaw also likely affects the drainage of lakes located in the permafrost area Webb and Liljedahl (2023). A recent work of Nitze et al. 2020 described, e.g., a series of extremely quick thermokarst lakes drainage in 2018 in northwestern Alaska after the unprecedented warm (with air temperature close to 0°C) and wet winter of 2017-2018. This drainage "exceeded the average drainage rate by a factor of 10", and is supposed to continue and increase the liquid and solid discharges of Arctic rivers. This situation suggests that in the Canadian Arctic, where "large areas are susceptible to hillslope thermokarst activity" (Zolkos et al. (2022)), the permafrost thawing will progressively change chemical denudation and induce the downstream redistribution of sediments over the next millennial(s).

Table 1: Statistics for the segmented linear regression analysis of in situ river discharge (Fig.2a) - $slope_Q$; trends of SPM_{insitu} : $slope_{insitu}$ (Fig.2b); trends of SPM_{sat} for all available points: $slope_{sat}$ (Fig.2d); and trends for SPM_{sat} for the area with over 200 pixels available: $slope_{sat200}$ (Fig.2f). Negative trend parameters are described with $slope1$, positive with $slope2$

	Est.	Standard error	CI(95%).l	CI(95%).u
$slope1_Q$	-57.21	19.14	-94.73	-19.68
$slope2_Q$	640.05	142.91	359.88	920.19
$slope1_{insitu}$	-5.08	3.04	-11.13	0.96
$slope2_{insitu}$	24.46	21.86	-19.02	67.93
$slope1_{sat}$	-0.23	0.04	-0.32	-0.15
$slope2_{sat}$	0.84	0.43	-0.01	1.70
$slope1_{sat200}$	-0.18	0.04	-0.27	-0.09
$slope2_{sat200}$	0.39	0.40	-0.40	1.18

4. Conclusion

Twenty years (2003-2022) of in situ measurements (river discharge and SPM concentration at the Arctic Red River station) and satellite-derived SPM concentrations were analyzed to describe the evolution of SPM inputs in the Beaufort Sea by the Mackenzie River and its impact on the adjacent coastal waters. Using the segmented regression model, we showed two opposite trends over the last 20 years for both the river freshwater discharge and SPM concentration. Over the studied period, we observe statistically significant negative trends from 2003 to 2018-2019, then a positive trend from 2019 to 2022 for both river discharge and SPM concentrations. Our results extend previous estimations of Doxaran et al. 2015 and tend to confirm other long-term observations

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showing a rather stable freshwater discharge of the Mackenzie River, increasing SPM concentrations in the delta zone and a significant decrease in SPM concentration in adjacent coastal waters Feng et al. (2021); Hilborn and Devred (2022); Matsuoka et al. (2022).

We suggest that the observed variability indicates a simultaneous decline of water discharge and, thus, suspended particulate matter transport due to changing precipitation pattern, especially the amount of snow over the river basin with a possible role of wind-induced mixing in the marine area. We also discuss long-term effects of climate change and permafrost thawing on the sediment transport rate in the Mackenzie River.

These processes of SPM release into the Arctic Ocean should be studied next on a pan-Arctic scale: can we expect a similar behavior based on the regional variations of river discharges reported by ArcticGRO? A recent study of Zolkos et al. 2022 used only in situ data, and further analysis of SPM distribution into the Arctic Ocean with satellite data will be beneficial. Studying the seasonal variability for specific years (e.g., 2006-2008, 2012) is also required using higher resolution data, including satellite imagery, to better understand changing river regimes in Arctic regions.

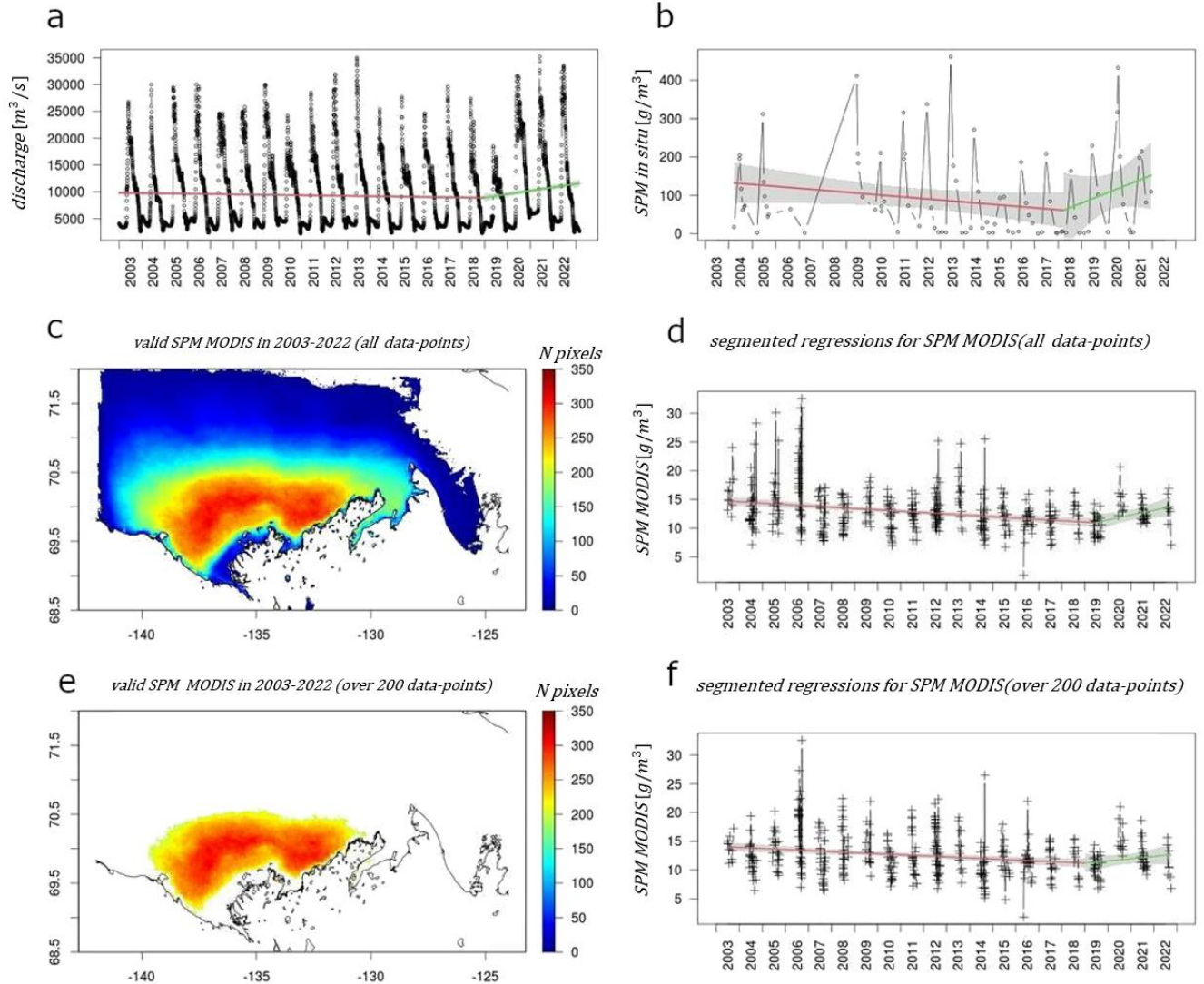


Figure 2: Interannual variations of in situ and satellite data (river discharge and SPM) with their segmented linear regression slopes (red and green colors show negative and positive regressions, respectively): (a) river discharge Q_{insitu} , (b) SPM_{insitu} time series; (c) maximum number of valid SPM_{sat} pixels in 2003-2022 (d) satellite SPM_{sat} time series; (e-f) similar to (c-d), but for the area with at least 200 valid pixels

5. Data availability

In situ measurements described in section 1 (river water discharge and TSS) are provided at <https://arcticgreatrivers.org/data/>. MODIS data is accessible from the NASA website <https://oceancolor.gsfc.nasa.gov>. Sea ice concentrations are available at <https://seaice.uni-bremen.de/data-archive/>. ERA5 LAND reanalysis data can be found at

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4 <https://doi.org/10.24381/cds.68d2bb30> with a detailed description at
5 <https://confluence.ecmwf.int/display/CKB/ERA5%3A+data+documentation>, and ERA5
6 wind data is available at <https://doi.org/10.24381/cds.adbb2d47> with a detailed
7 description at
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17 Greenbelt, MD, USA. doi: DOI. Accessed on 05/08/2023.
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19 No conflict of interest is stated.
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23 **Appendix A. ERA5-Land**

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25 The appendix contains Fig.A.3 illustrating correlation between ERA5 parameter, the
26 Mackenzie discharge from the Arctic GRO dataset, and in situ and satellite SPM
27 concentrations.
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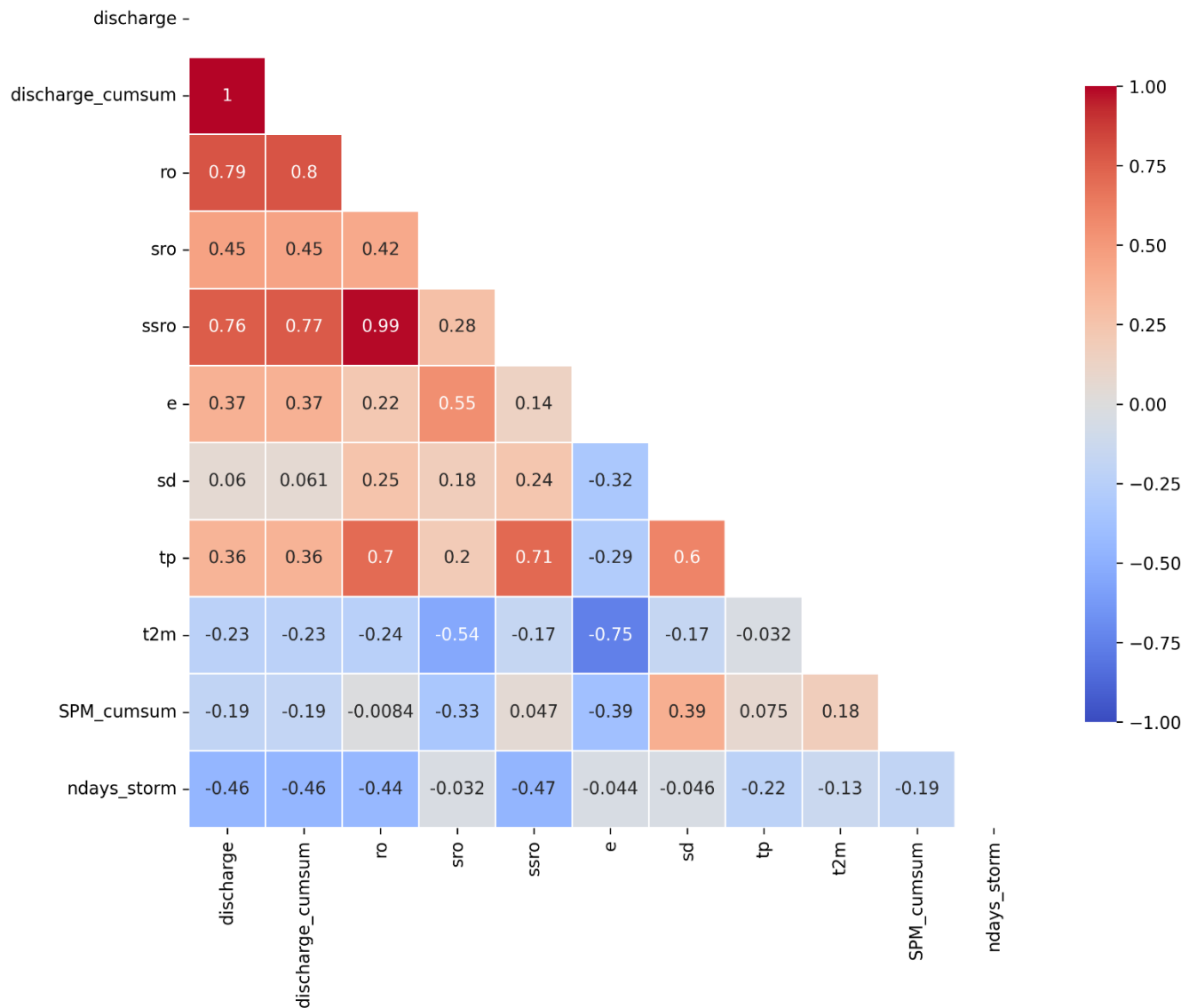


Figure A.3: Correlation matrix of all mean yearly reanalysis (ERA5 LAND) parameters over the Mackenzie basin and in situ (Arctic GRO) discharge for Mackenzie (Arctic Red station). ERA5 Land parameters are described in main text (ro - total runoff, sro surface runoff, ssro - subsurface runoff (ssro = ro-sro), e - evaporation, sd - snow depth, tp - total precipitation, t2m - air temperature at 2 m, SPM_{cumsum} - annual cumulated SPM concentrations, discharge - Arctic GRO in situ discharge).

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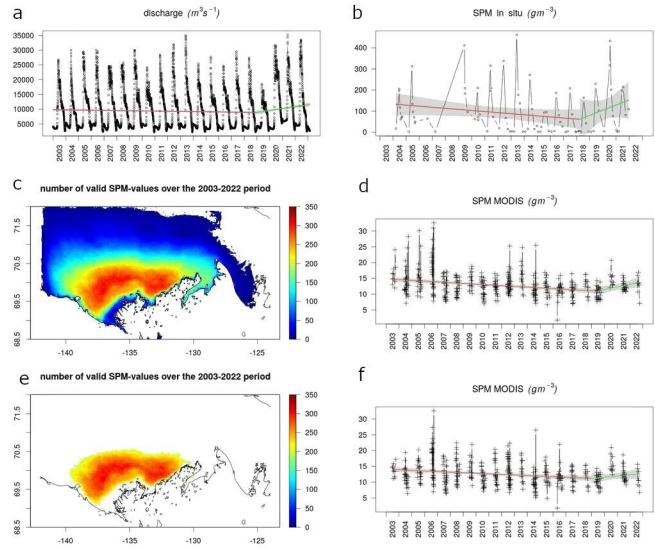
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1 Graphical Abstract

2 Variations of suspended particulate matter concentrations of the Mackenzie River
3 plume (Beaufort Sea, Arctic Ocean) over the last two decades

4 Anastasia Tarasenko, David Doxaran, Bernard Gentili



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7 Highlights

8 **Variations of suspended particulate matter concentrations of the Mackenzie River** 9 **plume (Beaufort Sea, Arctic Ocean) over the last two decades**

10 Anastasia Tarasenko, David Doxaran, Bernard Gentili

- 11 • Interannual variations of suspended particulate matter (SPM) in coastal waters
12 influenced by the Mackenzie River were investigated over the last 20 years (2003-
13 2022) using MODIS/Aqua satellite data
- 14 • The offshore SPM variations are related to the Mackenzie River hydrological regime,
15 air temperature, amount of snow, and permafrost state over its draining basin
- 16 • Over the last twenty years, a statistically significant negative trend was highlighted
17 over the period 2003-2018 for both SPM and river discharge, with a positive trend
18 starting from 2019 up to present

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5 19 Variations of suspended particulate matter concentrations of
6 20 the Mackenzie River plume (Beaufort Sea, Arctic Ocean) over
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8 21 the last two decades
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11 23 Anastasia Tarasenko^a, David Doxaran^a, Bernard Gentili^a
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14 25 Mer France
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20 27 **Abstract**

21 28 This work addresses the last 20 years' evolution of the suspended particulate matter (SPM)
22 29 concentrations in the Beaufort Sea (Canadian Arctic Ocean) directly influenced by the Mackenzie
23 30 River discharge. The SPM variations in the coastal zone are highlighted and related to the freshwater
24 31 and solid discharges of the river measured in situ at the Arctic Red River station (150 km upstream
25 32 of the river delta). The correlation between the variations of the river discharge and SPM
26 33 concentration within the surface layer of the coastal waters is obvious. Rather unexpectedly, both
27 34 have been slightly but significantly decreasing from 2003 to 2018-2019 and started to increase very
28 35 recently (2019-2022). This change of regime could be explained by changing winter precipitation
29 36 and groundwater distribution, progressively accumulating sediments within the thawing
30 37 permafrost layer and its recent release into the groundwater together with thermokarst lakes' rapid
31 38 drainage.
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35 39 *Keywords:* suspended particulate matter, ocean optics, Arctic Ocean,
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41 42 **1. Introduction**

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43 43 Climate change occurs faster in polar regions than at lower latitudes Arias et al. (2021).
44 44 Global warming is usually associated in high latitudes with rising air temperature,
45 45 precipitations, and permafrost thaw Miner et al. (2022); Pörtner et al. (2022).
46 46 Consequently, the freshwater discharged by rivers into the Arctic Ocean is expected to
47 47 increase, so as the discharge of terrestrial substances with enhanced erosion along
48 48 drainage basins Doxaran et al. (2015); Matsuoka et al. (2022); Juhls et al. (2022). This
49 49 would result in increasing water turbidity in coastal areas directly affected by river inputs
50 50 but also boosted primary production due to higher nutrient loads and reduced sea ice cover
51 51 (i.e., increasing solar light within the water column).
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4 52 The Mackenzie River is the fourth largest Arctic river in terms of river discharge (7% of
5 53 freshwater inflow to the Arctic Ocean) and is the primary source of sediment discharge
6 54 Carson et al. (1998); Wagner et al. (2011); Yang et al. (2015), so the changes in the
7 55 Mackenzie River regime will have a significant impact on the whole Arctic region Juhls et al.
8 56 (2022). The Mackenzie has a prominent seasonal cycle with winter lows and a summer
9 57 maximum of river discharge, which is related to the water cycle over its large basin with its
10 58 75% permafrost area and the importance of snowmelt in spring Yang et al. (2015); Grotheer
11 59 et al. (2020). The delta of the Mackenzie River is a complex area where the freshwater
12 60 massively inflows into the southern Beaufort Sea, highly impacted by a long presence of sea
13 61 ice. The southern Beaufort Sea is frozen most part of the year; the polynia between the fast
14 62 ice of the delta starts opening in May, then the water surface stays mostly icefree from July
15 63 to October. Over the last two decades, the sea ice conditions have become softer, and the sea
16 64 ice concentration (SIC) has diminished by -4-8% per decade from 2003 to 2019 Hilborn
17 65 and Devred (2022).

23 66 The open discussion of the Mackenzie River regime and its adaptation to climate change
24 67 depends on the studied period Woo and Thorne (2003); Doxaran et al. (2015); Yang et al.
25 68 (2015); Matsuoka et al. (2022); Zolkos et al. (2022). Doxaran et al. 2015 claimed a 22%
26 69 increase of the river discharge from 2003 to 2013. Yang et al. 2015 showed that over the
27 70 1973-2011 period, the positive linear trend for the Mackenzie discharge was very weak ($y =$
28 71 $17.21x + 8947.4$), and Woo and Thorne 2003 found no significant trend for the annual
29 72 discharge over the period 1968-1999, nor did Matsuoka et al. 2022 for the 1992-2018
30 73 period. Meanwhile, several studies agreed that the yearly amount of river discharge has
31 74 changed its seasonal distribution: cold season discharge becomes slightly higher, but spring
32 75 flows are lower (as higher air temperatures shift the snowmelt earlier) Woo and Thorne
33 76 (2003); Yang et al. (2015). The question of the river water discharge, thus, should be
34 77 regularly reevaluated.

39 78 In such a remote and changing environment, satellite observations have been proven to
40 79 be an efficient tool to monitor the evolution of Arctic coastal zones, and compensate the
41 80 lack of field measurements Doxaran et al. (2012); Hill et al. (2013); Juhls et al. (2022). SPM
42 81 field measurements, although extremely valuable, are usually too sparse for the long-term
43 82 variability analysis, as most of them are associated with specific summer expeditions (e.g.
44 83 the recent MALINA Massicotte et al. (2021) and Nunataryuk Lizotte et al. (2023) field
45 84 campaigns). Overall, during the ice-free season, satellite-derived suspended particulate
46 85 matter (SPM) concentrations typically vary from 30 g/m^3 in the delta region to 0.5 g/m^3
47 86 offshore over the Canada Basin Hilborn and Devred (2022).

51 87 Several studies analyzed satellite-derived SPM concentration variability in the southern
52 88 Beaufort Sea for long time periods (over 10 years), and their results are rather
53 89 controversial. From 2003 to 2013, Doxaran et al. 2015 found a linear trend with a significant
54 90 increase in monthly- averaged SPM concentrations at the mouth of Mackenzie River (+46%
55 91 in 10 years over the river mouth area and +71% in the river delta south to 70°N). Hilborn

and Devred 2022 applied *self-organizing maps* method to identify six different regions in the eastern Beaufort Sea using 17 years of MODIS (Moderate Resolution Imaging Spectroradiometer) data and found only one statistically significant trend (negative) for the annual SPM concentration in the deepwater Canada basin area, far offshore the Mackenzie delta. At the same time, Matsuoka et al. 2022 showed a statistically significant increase of dissolved and particulate organic matter concentrations (but not fluxes) in late summer: $0.019 \text{ g/m}^3 \text{ per year}$ and $0.069 \text{ g/m}^3 \text{ per year}$, accordingly).

The present study is a continuation of Doxaran et al. 2015, where we analyze the evolution of the SPM concentration in the Beaufort Sea directly influenced by the discharge of the Mackenzie River over the last two decades (2003-2022), using the methodology developed by Doxaran et al. (2012, 2015), this time excluding the complex river delta zone.

2. Materials and Methods

Fig. 1a shows the limits of the study area (yellow rectangular) on a quasitrue color daily MODIS composite for August 8, 2016. In this image, the SPM-rich Mackenzie River plume propagates along the coast and northward, partly into the MIZ (marginal ice zone). The variations of SPM concentrations (Fig. 1b) from the delta to offshore waters are obvious, which highlights complex processes along this land-sea interface. The field and satellite datasets used in this study, so as the methods used to retrieve and analyze time series of river discharge and SPM concentrations, are detailed hereafter.

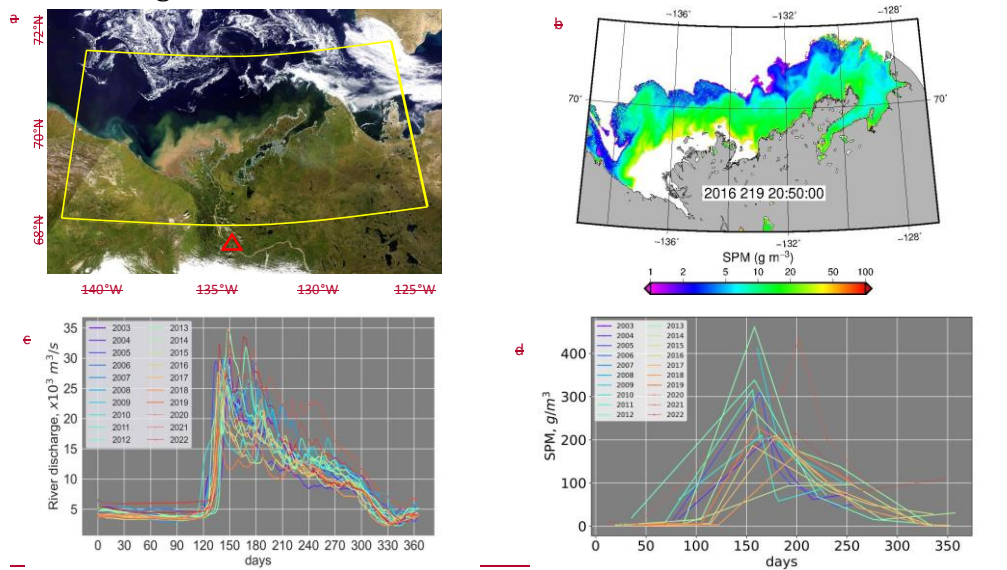


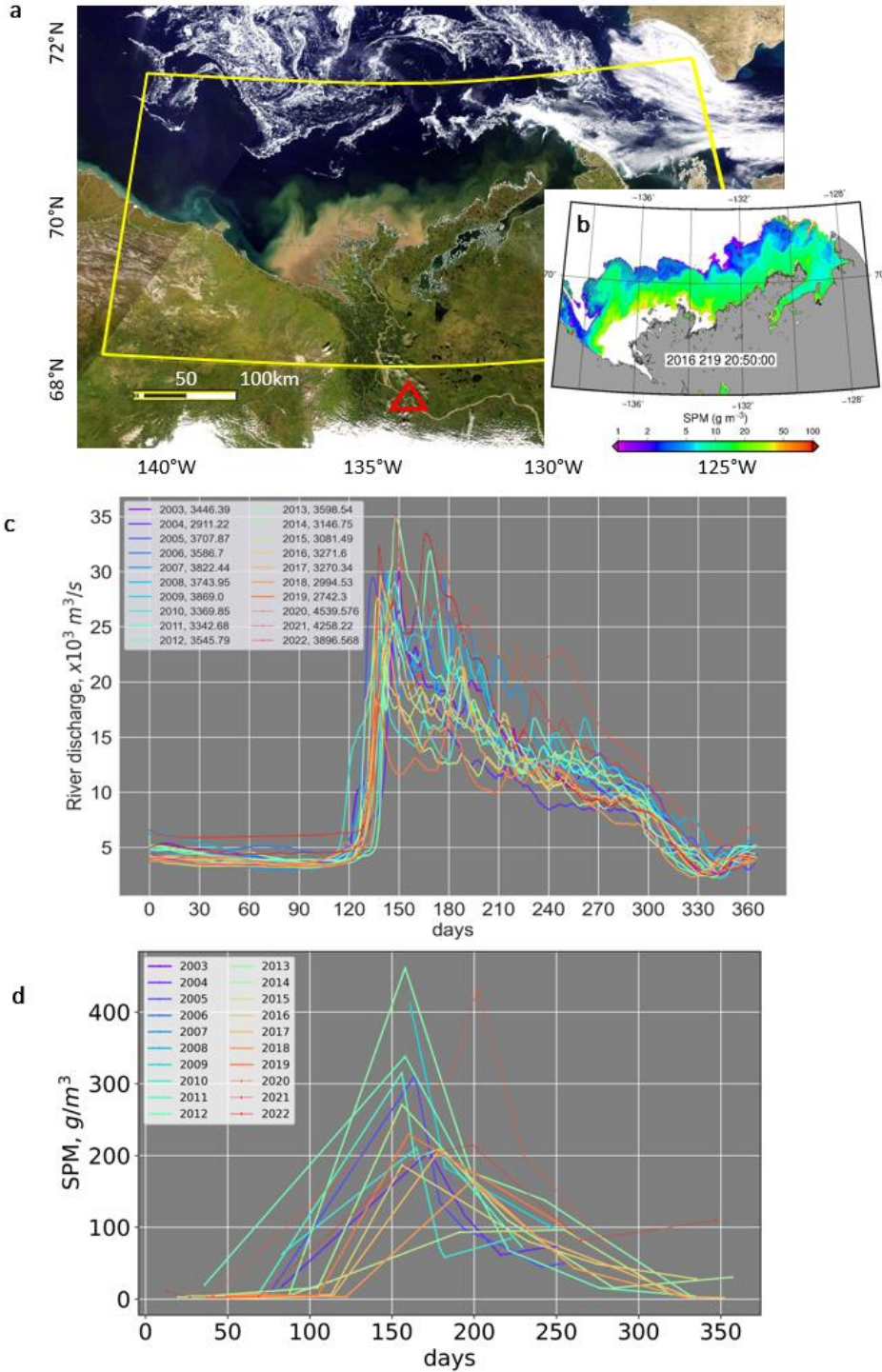
Figure 1: (a) MODIS image (August 6, 2016) locating the Arctic Red River gauge station (Tsiigehtchic) where in situ measurements are carried out (red triangle) and the coastal waters studied using satellite data (yellow rectangle); (b) SPM concentrations: Interannual variations of liquid discharge (c) and SPM concentrations (d) at the Arctic Red River station. The inserted SPM color map shows the result of MODIS satellite data processing

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2.1. Field data sets

In situ measurements of the river water discharge (2003-2022) were provided by the Water Survey of Canada via the ArcticGRO project McClelland et al. (2023) at the Arctic Red River gauge station (station ID: 10LC014, 67.45°N 133.74°W, red triangle in Fig.1a). River discharge measurements, Q , are available daily (Fig.1c). The measurements of suspended matter are distributed by the ArcticGRO as *total suspended solids (TSS)* concentrations, which is the same as SPM, so hereafter we will refer to in situ TSS as the *SPM_{insitu}*. The *SPM_{insitu}* data were collected at most once per month, but not during all months (on average, 4 to 6 measurements per year, Fig.1d). Over the studied period, some years are poorly represented (one measurement per year in 2003, 2006, and 2007; no SPM measurement in 2008). The lack of SPM in situ data implies the necessity to use satellite data.

The ArcticGRO in situ datasets pass the quality control procedure and indicate less reliable measurements as "provisional data". Over the 2003-2022 period, it corresponds to river discharge data in 2020-2022 and *SPM_{insitu}* data in 2019-2022 (shown with thinner lines in Fig.1c, d).



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136 **Figure 1: (a) MODIS image (August 6, 2016) locating the Arctic Red River gauge station (Tsiigehtchic) where**
137 **in situ measurements are carried out (red triangle) and the coastal waters studied using satellite data (yellow**
138 **rectangle); (b) SPM concentrations: Interannual variations of water discharge (the year of measurement and**

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4 139 [the annual cumulative sum of discharge is indicated in legend\), \(c\) and SPM concentrations \(d\) at the Arctic](#)
5 140 [Red River station. The inserted SPM color map shows the result of MODIS satellite data processing](#)
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9 142 2.2. Ocean color satellite data

10 143 As in Doxaran et al. 2015, a single satellite sensor (MODIS/Aqua) was considered for
11 144 this study in order to avoid any bias between different sensor products when detecting
12 145 temporal trends. Since 2003, MODIS provides high-quality ocean color observations at a
13 146 good temporal resolution: several images per day of the study area during the Arctic
14 147 summer months Doxaran et al. (2012, 2015); Hilborn and Devred (2022); Matsuoka et al.
15 148 (2022). The region of interest was defined as 68.5°N-72°N, 142°W - 124°W (yellow
16 149 rectangle in Fig.1a). The initial dataset contained 10608 swaths collected between May 1
17 150 and October 31 each year over the 2003-2022 period.

21 151 *Processing of satellite data.* MODIS L1A satellite data were processed using the SeaWiFS
22 152 Data Analysis System (SeaDAS 8.1.0) software (<http://seadas.gsfc.nasa.gov/>) and its l2gen
23 153 function. The atmospheric correction was performed using the NIR-SWIR algorithm of
24 154 Wang and Shi 2007); this method was proved to be the most appropriate for the highly
25 155 turbid waters at the mouth of the Mackenzie River to the less turbid waters offshore
26 156 Doxaran et al. (2012).

29 157 SeaDAS l2gen flags were used to mask clouds and glint. In the work of Doxaran et al.
30 158 2015, two techniques were used to mask clouds: the default one in the coastal waters and
31 159 an increased cloud albedo threshold value (0.4 instead of 0.027) over the specific area of
32 160 the river delta zone to avoid masking the highly turbid water pixels. This procedure was
33 161 time-consuming and required an additional inspection of every image in the river delta
34 162 zone. Here, the cloud-masking method preconized by Ody et al. 2022 for river mouths was
35 163 used to process each satellite image only once: the 2130 nm shortwave-infrared waveband
36 164 was used with a cloud threshold of 0.018. The sea ice mask was computed from daily sea
37 165 ice concentration (SIC) data at 3.125°spatial resolution distributed by the University of
38 166 Bremen Spreen et al. (2008). This dataset contains AMRS-E and AMSR2 (Advanced
39 167 Microwave Scanning Radiometer -for Eos and -2) images. Due to several gaps in data
40 168 acquisition over the 2003-2022 time period, the dataset was completed with the sea ice
41 169 concentration from the SSMIS (Special Sensor Microwave Imager/Sounder) instrument.
42 170 The ice mask was created using the SIC above 0%.

48 171 All satellite data (Rayleigh-corrected reflectances and masks) were reprojected on a
49 172 regular 250 m grid. The following criteria were then applied to exclude potentially
50 173 contaminated images:

- 53 174 • the number of valid pixels should exceed 15% of the marine area of the study area in
54 175 the downloaded MODIS image;

- the distance between the central point of the study area and that of the MODIS image should be less than 1000 km to avoid aberrations on the border of the image;
- only areas with a zenith solar angle lower than 74° were conserved.

This automatic filtering reduced the MODIS collection to 700 images. Finally, all images were inspected visually to filter out the remaining contaminated images because of not detected MIZ, clouds or cloud shadows, etc.; 651 MODIS images remained after this last quality check step.

The SPM concentrations were computed from the ratio (in %) between the remote sensing reflectance signals, R_{rs} (in sr^{-1}) in the near-infrared and green wavebands using the set of relationships established based on field measurements Doxaran et al. (2012, 2015):

$$SPM_{sat} = 0.8386 \times R_{RS}(748 : 555) \quad (1)$$

if $R_{RS}(748 : 555) < 87\%$

$$SPM_{sat} = 70 + 0.1416 \times R_{RS}(748 : 555) + 2.9541 \times e^{0.2092 \times (R_{RS}(748:555) - 87)} \quad (2) \text{ if } 87\% \leq R_{RS}(748 : 555) \leq 94\%$$

$$SPM_{sat} = 3.922 \times R_{RS}(748 : 555) - 285.4 \quad (3)$$

if $94\% < R_{RS}(748 : 555)$, where SPM_{sat} is the SPM concentration in g/m^3 and $R_{RS}(748 : 555)$ is the ratio of remote-sensing reflectances at 748 and 555 nm.

Negative $R_{RS}(555)$ and $R_{RS}(748)$ values were discarded before computing SPM, as well as computed SPM concentrations higher than $1000 g/m^3$, to remove atmospheric correction failures and residual contaminations (typically encountered along borders of clouds or sea ice).

The processed SPM images were averaged as daily, monthly, and yearly composites as simple mean averages for every pixel. Only daily composites and their mean values were used further, as monthly composites were sometimes computed using only 1 to 3 cloud-free images, which is not representative of mean monthly concentrations.

2.3. Reanalysis data

A dataset from ERA5 Land reanalysis was extracted to better understand how environmental factors did impact the Mackenzie River discharge and SPM concentrations in the adjacent coastal waters. This dataset was selected as the most suitable tool for the studies of river discharge variability Winkelbauer et al. (2022). ERA5 Land has a monthly temporal resolution and 9-km spatial resolution and is provided by Copernicus data center [MuñozMuñoz-Sabater et al. \(2021\)](#). The area of extraction approximately corresponds to the Mackenzie drainage basin (52-70°N, 100-140°W). The following parameters were used: river runoff (ro), surface river runoff (sro), evaporation (e), total precipitation (tp), snow

depth (sd), and air temperature at 2m (t2m), (see ~~de-detailed description at~~
<https://confluence.ecmwf.int/display/CKB/ERA5-Land%3A+data+document> The
Pearson correlation matrix for the mean annual values of all described parameters
(Fig.A.3).

~~tailed description at~~
[https://confluence.ecmwf.int/display/CKB/ERA5-](https://confluence.ecmwf.int/display/CKB/ERA5-Land%3A+data+document)
~~Land%3A+data+document~~ The similarity between the ERA5-LAND
river runoff and the Arctic GRO river discharge are shown in the
Supplementary Information section (Fig.), as well as a correlation
matrix for the mean annual values of all described parameters
(Fig.S12).

To include a possible effect of storms on the SPM concentrations, we added to the
analysis the wind data from ERA5 (marine area similar to that chosen for the MODIS study
box). The Pearson correlation matrix (where the yearly parameters are compared) contains
an additional parameter "the number of days of storms per year" (*ndays_storm* in Fig. A.3).
The day was considered stormy if over the study area there was a wind vector with a wind
speed over 15 m/s.

2.4. Temporal variability

As previous works have shown, the linear regression model provides relatively little
information on interannual variability of river water discharge and related parameters,
such as SPM, because, over the long term, the discharge appears as stable Yang et al. (2015);
Matsuoka et al. (2022). Over the last 20 years, the linear trend for the in situ Mackenzie
discharge time series calculated with the Ordinary Least Squares (OLS) model is $y =$
 $17.1x + 9554$ with a standard error (SE) of $63.6 \text{ m}^3/\text{s}$, and confidence intervals (CI) [0.025
0.975] equal to -107.5 and $141.7 \text{ m}^3/\text{s}$ for the coefficient term. Although these results are
very close to that of Yang et al. 2015, the statistical metrics confirm that the OLS model is
not suitable for the analysis in this case.

For this reason, we used an estimating regression model with unknown break-points,
which allows describing several temporal trends Muggeo (2003). To apply this model to
our datasets, we used the *segmented* function in the R package *segmented*.

To compute the trends, the in situ data time series of daily river discharge (Q_{insitu}) and
SPM concentration (SPM_{insitu}) described above were used (Fig.2a, b). For the satellite data,
the spatial sparsity of observations was taken into account. Thus, to obtain SPM_{sat} time
series, we calculated: • one daily spatial mean SPM value for all available SPM pixels over
the study area, which resulted in SPM_{sat} time series (Fig.2 c-d); • one daily spatial mean SPM
value for the area with the highest data density (where valid satellite pixels appear more
than 200 times over the 2003-2022 period), which resulted in SPM_{sat200} (Fig.2 e-f)

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4 247 Then the segmented regression analysis technique was applied to all time series: Q_{insitu} ,
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6 248 SPM_{insitu} , SPM_{sat} , and SPM_{sat200} . After several tests, the most statistically significant results
7 249 were obtained with one breaking point and two segment slopes. The following statistical
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9 250 parameters were computed for each segment: estimated coefficient of linear regression
10 251 (Est.), standard error, lower and upper 95% confidence intervals (CI(95%).l, CI(95%).u,
11 252 respectively) (Tab.1).
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14 253 3. Results and Discussion

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16 254 The seasonal cycles of both the river discharge, Q , and SPM_{insitu} are very pronounced,
17 255 with high summer and low winter values (Fig.1c,d). These parameters also have a strong
18 256 interannual variability. During the 2003-2022 period, the river discharge varied from
19 257 $2.2 \times 10^3 \text{ m}^3/\text{s}$ (winter) to $35.2 \times 10^3 \text{ m}^3/\text{s}$ (summer), with a mean value of $9.5 \pm 6.8 \times 10^3 \text{ m}^3/\text{s}$.
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21 258 Typically, the discharge increases very rapidly in 2 weeks from 5 to $25 \times 10^3 \text{ m}^3/\text{s}$ in late May,
22 259 when the main summer peak occurs, then slowly decreases to its winter values by
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24 260 November.
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26 261 The years 2013 and 2021 were exceptional with extreme *yearly maxima* of river
27 262 discharge over $35 \times 10^3 \text{ m}^3/\text{s}$ on May 28 (both), while the lowest yearly maximum river
28 263 discharge, $18 \times 10^3 \text{ m}^3/\text{s}$, was registered on May 18, 2019. Previously, Yang et al. 2015 also
29 264 observed similar prominent interannual variations in the river discharge during the
30 265 summer (e.g., in 1992 and 1995 for the 1972-2011 period). They concluded that a negative
31 266 anomaly in precipitation-evaporation balance over the river basin in summer (hot and dry
32 267 weather) usually results in a lower discharge the next year with an earlier maximum, while
33 268 the opposite (cold and wet) weather is responsible for a higher discharge. This statement
34 269 does not explain the extreme peak of discharge recorded in 2013: in 2012 the summer was
35 270 "hot and dry", so 2013 should have been a year of extremely low summer river discharge.
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37 271 Overall, calculated correlations between in situ river discharge and ERA5 Land total
38 272 precipitation, as well as air temperature were weak (0.36 and -0.23, respectively).
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41 274 However, the river discharge is well correlated with snow depth
42 275 cover
43 276 (correlation coefficient between Q_{insitu} and sd is 0.67): the 2011-12 and ~~2012-13~~ 2012-13
44 277 winters were ~~very~~ snowy (+10% and +3% sd anomaly compared to 20years median
45 278 values), as well as the 2020 winter (+7% sd anomaly), but 2018-19 snow cover was weak
46 279 (-7% sd anomaly). Interestingly, in 2013 the yearly means of river discharge and air
47 280 temperature were close to their 20years median values, indicating the overall stability of
48 281 the system during this period.
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53 281 The river discharge regime in 2006-2008 and 2012 was special with two summer Q
54 282 maxima, the first one in late May and the second between the end of June and mid-July. The
55 283 reason for this change in river regime might be related to the precipitation seasonal pattern
56 284 and the river ice opening (several ice jams crushes, creating the second peak). To investigate
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4 285 these particularities, a higher temporal resolution reanalysis data should be used. We also
5 observe local maxima (values higher than the previous year) in SPM_{sat} in 2006, 2008, and
6 286
7 287 2012.

8 288 As explained in section 2.4, the OSL model does not highlight a statistically significant
9 trend in the river discharge time series, but the *segmented* model does. Fig.2 and Table 1
10 289 present the results of calculated segmented regressions, where the river discharge and SPM
11 290 concentrations demonstrate very similar trends: a negative trend from 2003 to 2018, then
12 291 a recent positive trend from 2019 (2018 for SPM_{insitu}) to 2022 (Fig.2). Based on the
13 292 confidence intervals, we conclude that negative trends for $slope1_Q$, $slope1_{SPM_{sat}}$, and
14 293 $slope1_{SPM_{sat200}}$ are statistically significant (CIs [0.025 0.975] are all negative).
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16 295 This negative trend of $slope1_Q = -57.21x + const$ in river discharge for the 2003-2018
17 296 period is opposite to that of Doxaran et al. 2015 for the 2003-2013 period, but the latter
18 297 study slightly overestimated the river discharge in 2013 (as the data was not yet fully
19 298 quality-checked) resulting in an erroneous positive (+22%) trend. As already mentioned,
20 299 other studies, e.g. Yang et al. (2015); Matsuoka et al. (2022) which used an OSL model did
21 300 not find any significant trend in the Mackenzie River discharge. The positive trend from
22 301 2019 to the present, $slope2_Q = 640.05x + const$, although not statistically significant, can
23 302 indicate a progressive increase in the minimum flow impacted by a "mobilization of ground
24 303 waters" as discussed by Yang et al. 2015.

25 304 In situ data show that the SPM seasonal cycle generally follows the river discharge cycle,
26 305 with the summer maxima in late May - beginning of June. The SPM_{sat} values vary from 0.8
27 306 to 461 g/m^3 with a median value of $67.7 \pm 11 \text{ g/m}^3$. The breakpoint of SPM_{insitu} trends is
28 307 slightly shifted to 2018, and CIs contain zero, thus indicating that the *segmented* model is
29 308 less reliable for this dataset. At the same time, the SPM_{insitu} time series has the fewest amount
30 309 of data, which is not homogeneous in time and is mainly available in summer. It makes it
31 310 more difficult to interpret with any statistical model. SPM_{sat} and SPM_{sat200} trend analyses
32 311 show similar results: their negative $slope1$ values and corresponding SE are close to each
33 312 other, which gives confidence in the observed SPM negative trend. The positive trends
34 313 $slope2_{sat}$ and $slope2_{sat200}$ are not statistically significant, and their SEs are of the order of the
35 314 estimated coefficient of linear regression (Tab.1). Nevertheless, this result is interesting for
36 315 further discussion.

37 316 Over the 2003-2019 period, Hilborn and Devred 2022 obtained results in good
38 317 agreement with a negative trend (although not significant) for SPM concentrations over the
39 318 southern Beaufort Sea (from -0.16 to -0.46 g/m^3 per decade) except for the Mackenzie River
40 319 delta, where they found an increase of SPM of 0.36 g/m^3 per decade. The difference in
41 320 negative coefficients' values obtained in the present study and that of Hilborn and Devred
42 321 (2022) comes from (1) the difference of SPM-retrieval methods and different MODIS bands
43 322 used; and (2) the regions: our study region of SPM_{sat200} corresponds to 3 regions over the
44 323 shelf identified by Hilborn and Devred (2022).
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4 324 In the Mackenzie delta zone, Doxaran et al. 2015 also found a significant increase of SPM
5 325 (+71% from 2003 to 2013) over the "Mackenzie River delta" region of Hilborn and Devred
6 326 (2022). The delta zone seems to play an important role of "SPM filter" between the river
7 327 and the coastal waters. Based on satellite observations, SPM apparently settles massively
8 328 in this shallow area, resulting in the formation of temporary maximum turbidity zones
9 329 where resuspension of bottom sediments may occur depending on the river discharge, tidal
10 330 currents, and wind stress (Wegner et al., 2005; Grotheer et al., 2020). Observations at
11 331 high spatial and temporal resolutions are required to further investigate SPM dynamics in
12 332 this delta zone.

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16 333 However, how to explain the negative trends in both river discharge and SPM
17 334 concentrations (in the riverbed and offshore the delta zone) derived from field and satellite
18 335 observations over the 2003-2018 period? ~~The Based on the correlation matrix (Fig. A.3),~~
19 336 ~~we recognize that the river runoff and, thus, river discharge variability depend mostly on~~
20 337 ~~the amount of total precipitation (r = 0.7) and the snow depth (r = 0.63, which is another~~
21 338 ~~estimate of solid precipitation in winter) parameters. This trivial conclusion leads us to a~~
22 339 ~~simple suggestion that the Mackenzie River discharge slightly declined while the~~
23 340 ~~precipitation pattern has changed over this period.~~

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27 341 As for the SPM, the recent study of Zolkos et al. 2022 reveals a similar decrease in SPM
28 342 loads in most of the Russian Arctic-Siberian rivers between 1970 and 2010, and explains it
29 343 by natural and anthropogenic factors. At the same time, the rivers' discharge (mean and
30 344 peak values) The "natural factors" of sediments erosion over the river basin are the physical
31 345 and chemical denudation. For the Mackenzie River, the physical denudation rates exceed
32 346 the chemical denudation about several orders of time: mechanical denudation rate at the
33 347 Arctic Great River was reported as 844 t * km⁻² * y⁻¹ in 1997, while the chemical
34 348 weathering occurs at a rate of 25 mm*ky⁻¹ with a transport time of 10-400 ky⁻¹ (Vigier et
35 349 al. (2001); DePaolo et al. (2006)). On the order of two decades, it is, thus, physical or
36 350 mechanical factors related to the river discharge and the atmospheric conditions in the
37 351 delta-adjacent areas that control SPM concentrations. A simultaneous decline of Q *in situ*
38 352 and SPM *in situ* in 2003-2018 means that the lower the river discharge, the less suspended
39 353 particulate matter it transports.

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44 354 Another source of variability for the "marine" SPM might be the wind mixing. The simple
45 355 hypothesis proposes that the higher the wind speed, the more mixing occurs in a shallow
46 356 area, which reduces the surface SPM concentrations. At the same time, an additional
47 357 hypothesis may suggest that more mixing means more re-suspension of particulate matter
48 358 from the bottom sediments (increasing the SPM concentration). To verify these
49 359 controversial hypotheses, we must compare quasi simultaneous wind and SPM
50 360 observations. In the framework of this study there is yet one fundamental limitation for
51 361 this analysis due to the nature of ocean color data retrieval from satellite: the highest wind
52 362 speed (and thus mixing events) will mostly occur under the cloud cover with the passage
53 363 of cyclones, and SPM concentrations cannot be retrieved from ocean color satellite under
54 364 such conditions (presence of clouds). We found a weak negative correlation (correlation
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4 365 coefficient is $r=-0.19$) between the number of days of storm and annual cumulated SPM,
5 366 which confirms the first simple hypothesis (more mixing, less SPM), but this question
6 367 should be addressed additionally with other tools, like modeling.

7 368 There is yet a question about the recent positive trends in Q and SPM over the last 3-5
8 369 years of observations. Although the trends are statistically not significant, can we suggest
9 370 any important processes that might have impacted the river discharge and sedimentation
10 371 transport recently?

11 372 The presence of permafrost over the river basin was considered
12 373 mostly stable, so the link between the river discharge and SPM load is
13 374 not straightforward. Zolkos et al. suggest that the decrease of SPM is
14 375 due to the decrease of the sediments input “and/or increased
15 376 sedimentation within the fluvial network”. Nonetheless, the Kolyma
16 377 yearly SPM flux has doubled over the same period, which was
17 378 associated with the geological stratification of more erodible
18 379 quaternary sediments altering with ice. In the Canadian Arctic “where
19 380 large areas are susceptible to hillslope thermokarst activity”, the
20 381 permafrost thaw can induce the downstream redistribution of
21 382 particular sediments over the next “centuries or millennial”. We might
22 383 observe the effect of the sediment accumulation during a “precluding”
23 384 for the groundwater’s contribution to general river discharge, as it
24 385 prevented the water infiltration through the permafrost layer, but
25 386 might have helped to create underground cavities filled with non-
26 387 communication water reservoirs (Vigier et al. (2001)). With a
27 388 progressive permafrost thawing and its recent release in SPM
28 389 measurements.

29 390 The permafrost thaw likely controls as well the drainage of lakes located in the
30 391 permafrost area Webb and Liljedahl (2023)., this neglected role might be re-evaluated.
31 392 Matsuoka et al. 2022 observed an increase of thaw depth and precipitation, but a decrease
32 393 in river discharge, which is probably explained by the contribution of ground waters, also
33 394 suggested by Connolly et al. (2020). Finally, The permafrost thaw also likely affects the
34 395 drainage of lakes located in the permafrost area Webb and Liljedahl (2023). A recent work
35 396 of Nitze et al. 2020 described, e.g., a series of extremely quick thermokarst lakes drainage
36 397 in 2018 in northwestern Alaska after the unprecedented warm (with air temperature close
37 398 to 0°C) and wet winter of 2017-2018. This drainage “exceeded the average drainage rate by
38 399 a factor of 10”, and is supposed to continue and increase the liquid and solid discharges of
39 400 Arctic rivers. This situation suggests that in the Canadian Arctic, where “large areas are
40 401 susceptible to hillslope thermokarst activity” (Zolkos et al. (2022)), the permafrost thawing
41 402 will progressively change chemical denudation and induce the downstream redistribution
42 403 of sediments over the next millennial(s).

43 404 Based on these results, we conclude that the following mechanism ex-
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Table 1: Statistics for the segmented linear regression analysis of in situ river discharge (Fig.2a) - $slope_Q$; trends of SPM_{insitu} : $slope_{insitu}$ (Fig.2b); trends of SPM_{sat} for all available points: $slope_{sat}$ (Fig.2d); and trends for SPM_{sat} for the area with over 200 pixels available: $slope_{sat200}$ (Fig.2f). Negative trend parameters are described with $slope1$, positive with $slope2$

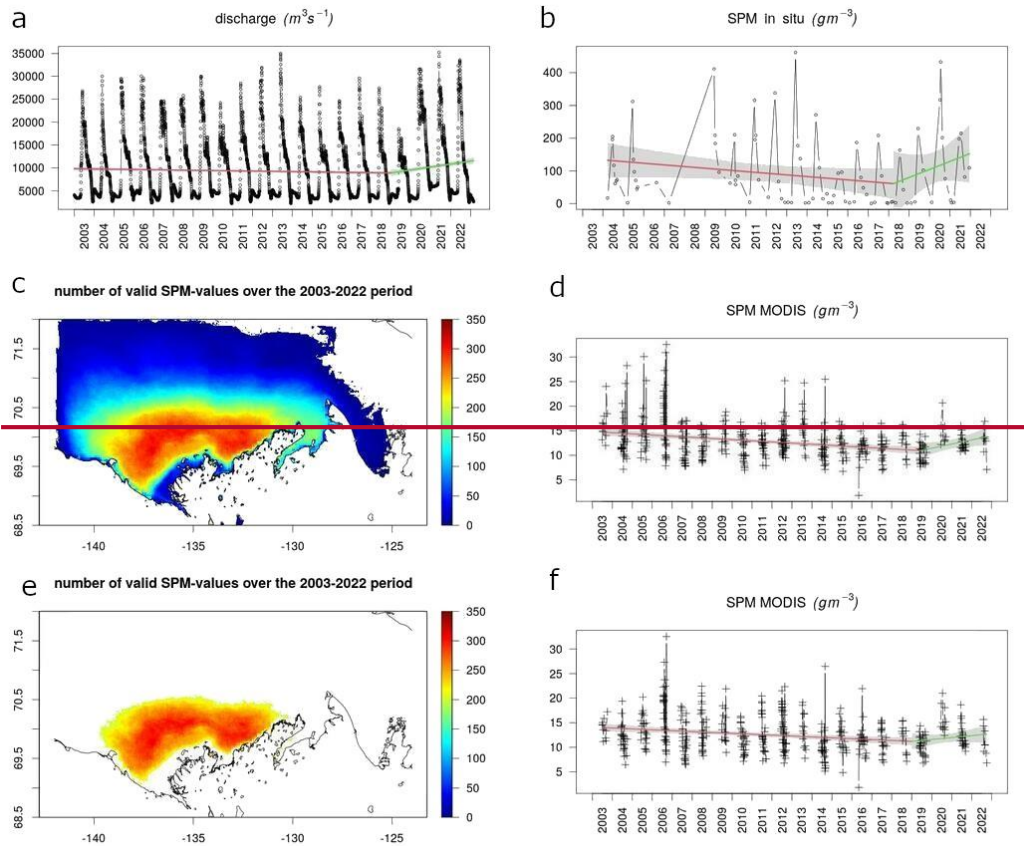
	Est.	Standard error	CI(95%).l	CI(95%).u
$slope1_Q$	-57.21	19.14	-94.73	-19.68
$slope2_Q$	640.05	142.91	359.88	920.19
$slope1_{insitu}$	-5.08	3.04	-11.13	0.96
$slope2_{insitu}$	24.46	21.86	-19.02	67.93
$slope1_{sat}$	-0.23	0.04	-0.32	-0.15
$slope2_{sat}$	0.84	0.43	-0.01	1.70
$slope1_{sat200}$	-0.18	0.04	-0.27	-0.09
$slope2_{sat200}$	0.39	0.40	-0.40	1.18

~~plains the observed negative trend in river discharge and SPM concentrations switched by a rapid positive trend: the progressive thaw of the permafrost helped to accumulate sediments in the ground inner impermeable layers up to the certain moment (appr. 2018), then these additional sediments arrived into the groundwater, and together with the drainage of lakes, the accumulated sediments started its quick release with groundwater into the main surface flow and the Beaufort Sea.~~

4. Conclusion

Twenty years (2003-2022) of in situ measurements (river discharge and SPM concentration at the Arctic Red River station) and satellite-derived SPM concentrations were analyzed to describe the evolution of SPM inputs in the Beaufort Sea by the Mackenzie River and its impact on the adjacent coastal waters. Using the segmented regression model, we showed two opposite trends over the last 20 years for both the river freshwater discharge and SPM concentration. Over the studied period, we observe a statistically significant negative ~~trend~~trends from 2003 to 2018-2019, then a positive trend from 2019 to 2022 for both river discharge and SPM concentrations. Our results extend previous estimations of Doxaran et al. 2015 and tend to confirm other long-term observations showing a rather stable freshwater

442 discharge of the Mackenzie River, increasing SPM concentrations in the
 443 delta zone and a significant decrease in SPM concentration in adjacent



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 445 Figure 2: Interannual variations of in situ and satellite data (river discharge and SPM) with their segmented
 446 linear regression slopes (red and green colors show negative and positive regressions, respectively): (a) river
 447 discharge $Q_{in\ situ}$, (b) $SPM_{in\ situ}$ time series; (c) maximum number of valid SPM_{sat} pixels in 2003-2022 (d) satellite
 448 SPM_{sat} time series;
 449 (e-f) similar to (c-d), but for the area with at least 200 valid pixels

450 coastal waters Feng et al. (2021); Hilborn and Devred (2022); Matsuoka et al. (2022).

451 We suggest that the observed variability indicates a ~~recent progressive accumulation of~~
 452 ~~sediments during the permafrost thaw, its further release into the drainage system,~~
 453 ~~Mackenzie flow, and finally into the Beaufort Sea since 2018. An important snow~~
 454 ~~accumulation followed by a spring thawing with extreme air temperatures aggravates the~~
 455 ~~drainage and sediment accumulation and release on the medium-term timescale.~~
 456 ~~simultaneous decline of water discharge and, thus, suspended particulate matter transport~~
 457 ~~due to changing precipitation pattern, especially the amount of snow over the river basin~~
 458 ~~with a possible role of wind-induced mixing in the marine area. We also discuss long-term~~
 459 ~~effects of climate change and permafrost thawing on the sediment transport rate in the~~
 460 Mackenzie River.

These processes of SPM release into the Arctic Ocean should be studied next on a pan-Arctic scale: can we expect a similar behavior based on the regional variations of river discharges reported by ArcticGRO? A recent study of Zolkos et al. 2022 used only in situ data, and further analysis of SPM distribution into the Arctic Ocean with satellite data will be beneficial. Studying the seasonal variability for specific years (e.g., 2006-2008, 2012) is also required using higher resolution data, including satellite imagery, to better understand changing river regimes in Arctic regions.

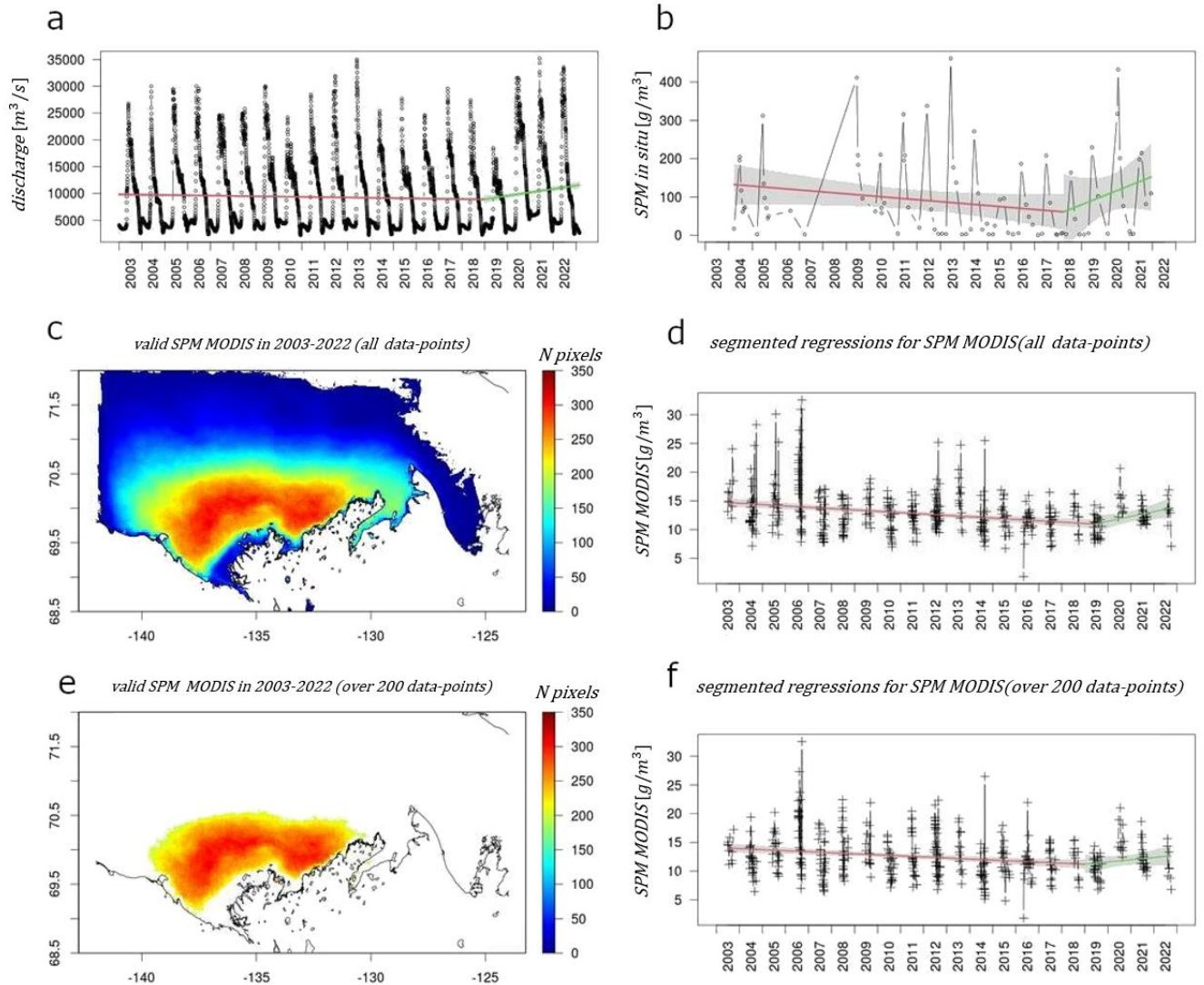


Figure 2: Interannual variations of in situ and satellite data (river discharge and SPM) with their segmented linear regression slopes (red and green colors show negative and positive regressions, respectively): (a) river discharge Q_{insitu} , (b) SPM_{insitu} time series; (c) maximum number of valid SPM_{sat} pixels in 2003-2022 (d) satellite SPM_{sat} time series; (e-f) similar to (c-d), but for the area with at least 200 valid pixels

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7 476 **5. Data availability**
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9 477 In situ measurements described in section 1 (river water discharge and TSS) are
10 478 provided at <https://arcticgreatrivers.org/data/>. MODIS data is accessible from the NASA
11 479 website <https://oceancolor.gsfc.nasa.gov>. Sea ice concentrations are available at
12 480 <https://seaice.uni-bremen.de/data-archive/>.
13 481 <https://seaice.uni-bremen.de/data-archive/>.
14 482 [ERA5 LAND reanalysis data can be found at https://doi.org/10.24381/cds.68d2bb30](https://doi.org/10.24381/cds.68d2bb30) with
15 483 a detailed description at
16 484 <https://confluence.ecmwf.int/display/CKB/ERA5%3A+data+documentation>, and ERA5
17 485 wind data is available at <https://doi.org/10.24381/cds.adbb2d47> with a detailed
18 486 description at
19 487 [ERA5 LAND reanalysis data can be found at https://doi.org/10.24381/cds.68d2bb30](https://doi.org/10.24381/cds.68d2bb30).
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25 488 **6. Acknowledgments**
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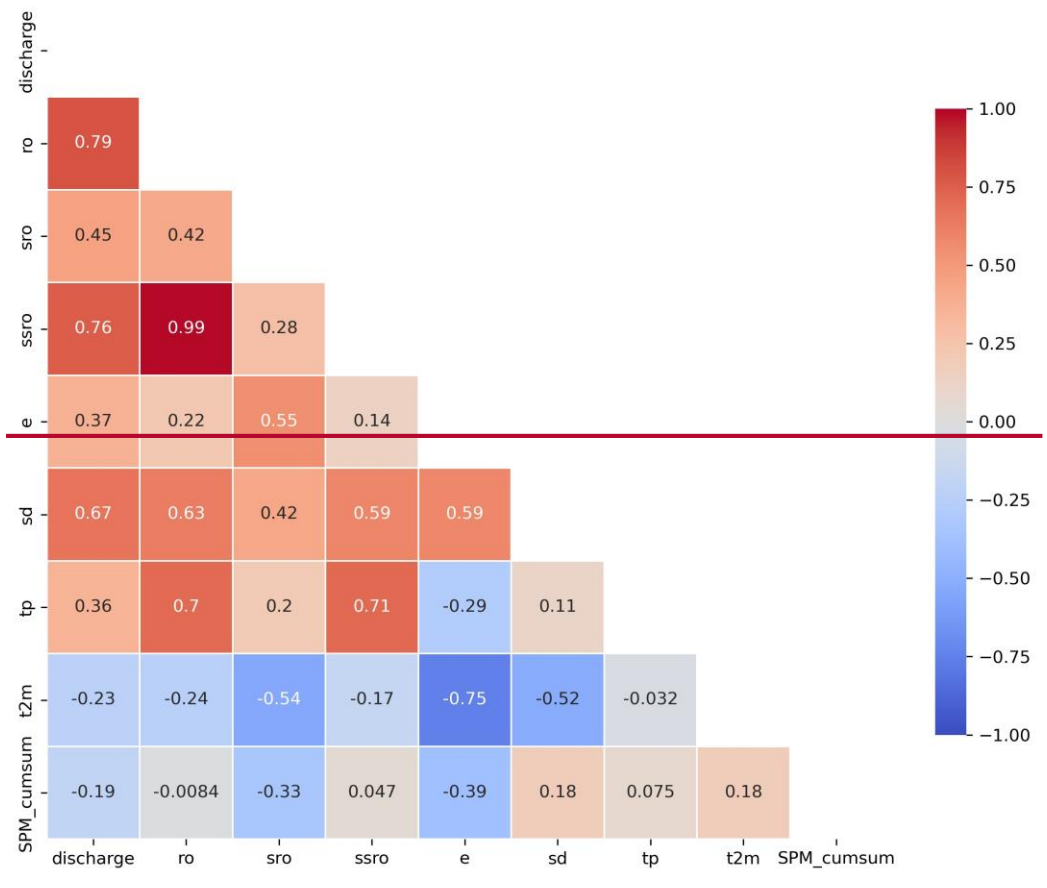
27 489 This study is part of the Nunataryuk project. The project has received funding under the
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31 493 Flight Center, Ocean Ecology Laboratory, Ocean Biology Processing Group.
32 494 Moderateresolution Imaging Spectroradiometer (MODIS) Aqua Data; NASA OB.DAAC,
33 495 Greenbelt, MD, USA. doi: DOI. Accessed on 05/08/2023.
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36 496 No conflict of interest is stated.
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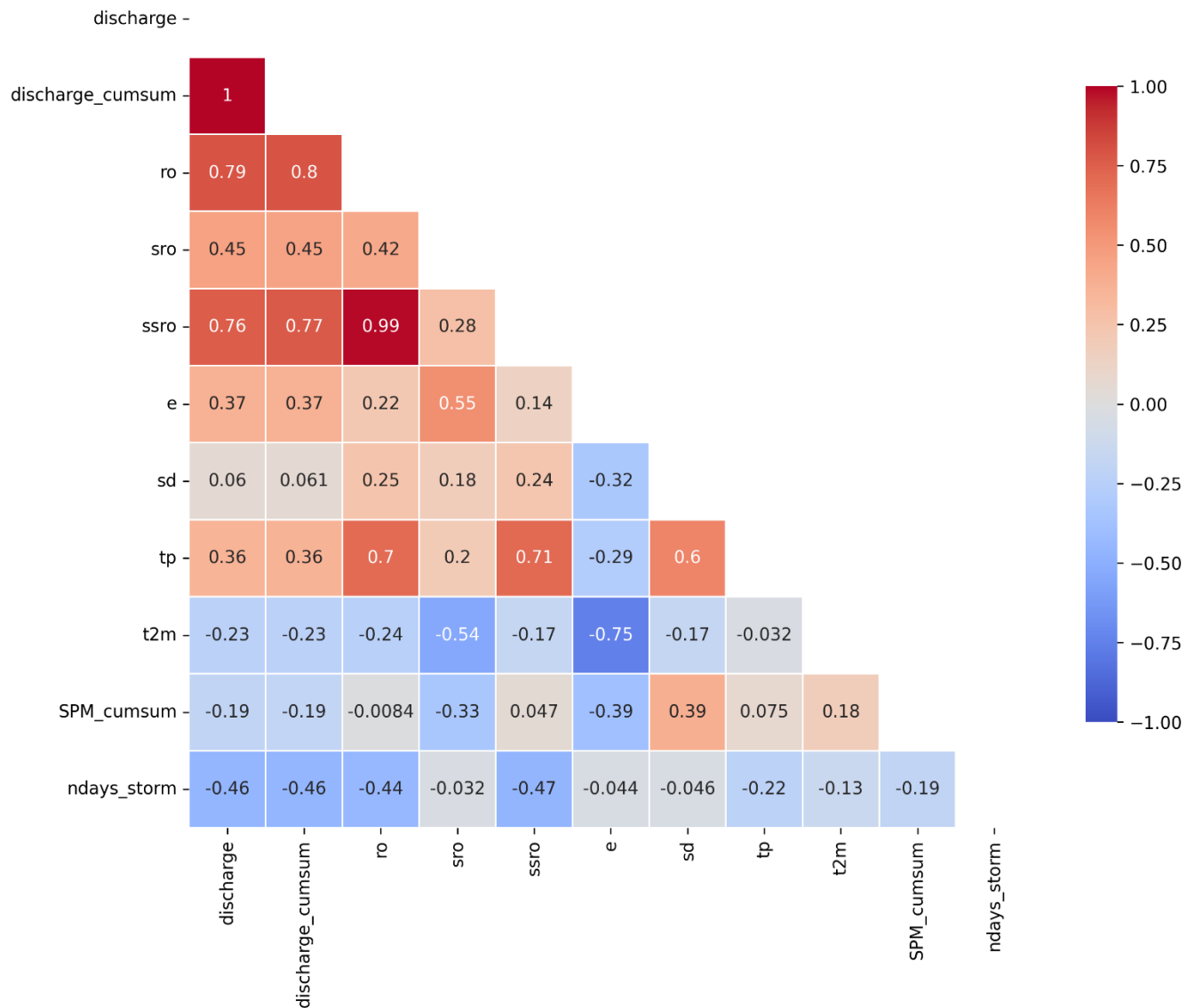
39 498 **Appendix A. ERA5-Land**
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41 499 The appendix contains Fig.A.3 illustrating correlation between ERA5 parameter, the
42 500 Mackenzie discharge from the Arctic GRO dataset, and in situ and satellite SPM
43 501 concentrations.
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Figure A.3: Correlation matrix of all mean yearly reanalysis (ERA5 LAND) parameters over the Mackenzie basin and in situ (Arctic GRO) discharge for Mackenzie (Arctic Red station). ERA5 Land parameters are described in main text (ro - total runoff, sro surface runoff, ssro - subsurface runoff (ssro = ro-sro), e - evaporation, sd - snow depth, tp - total precipitation, t2m - air temperature at 2 m, SPM_{cumsum} - annual cumulated SPM concentrations, discharge - Arctic GRO in situ discharge).

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Declaration of interests

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

AT has written the original and revised manuscript, participated in the investigation, visualization and data analysis. DD did the conceptualization, project administration, supervision and manuscript revision and editing. BG did the data curation, formal analysis, visualization and manuscript editing.